Rossby wave breaking and extreme precipitation events in the central and eastern United States

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Motivation

- Widespread extreme precipitation events (EPEs) in central and eastern U.S. during cool season can result in highimpact flooding
- Evidence from case studies indicates widespread EPEs occur in conjunction with baroclinic Rossby wave breaking (RWB)
- Climatological and dynamical linkages between RWB and EPEs in U.S. have not yet been examined



Data source: NCEP Stage IV

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320-K PV (PVU, gray shading), 1000–200-hPa IVT (kg m⁻¹ s⁻¹, vectors and tan shading) 6-h Stage-IV precip (mm, color shading)

Data sources: NCEP Stage IV and NCEP CFSR 0600 UTC 27 Dec 2015

Science questions

- 1. Do widespread EPEs in the central/eastern U.S. preferentially occur in conjunction with RWB?
- 2. How does RWB supply the ingredients for EPEs?

This presentation will focus on the nexus between RWB and moisture transport linked to EPEs.

RWB: An aspect of baroclinic wave life cycles





ARs form as an aspect of the dynamical evolution of baroclinic waves that establishes favorable conditions for heavy precipitation

Data and methods

Climatology of widespread EPEs

- Use 24-h (ending 1200 UTC) 0.25° gauge-based precipitation analyses for 1979–2015 from NOAA CPC Unified Precipitation Dataset
- Define top 5% (299) of days with ≥1 extreme precip value in domain as widespread EPEs
- Retain only days during Sep–May without a tropical cyclone in domain
- Consider consecutive days as one event; retain only largest-scale day for statistical analysis
- Examine final sample of 201 widespread EPEs
- Define t₀ as start time of 24-h period of EPE



Data and methods

breaking Rossby waves

Adaptation of method from Wernli and Sprenger (2007)

- Potential vorticity (PV) streamers are manifestations of RWB
- PV streamers identified on 2-PVU contour on 310-, 320-, and 330-K surfaces in 0.5° NCEP CFSR
- PV streamers identified as pairs of points along 2-PVU contour separated by distance *d* < 1000 km and by contour length *l* > 3000 km
- Classify streamers based on orientation angle relative to meridional baseline through midpoint of interval *d*:
 - > 15°: anticyclonic (LC1) wave breaking
 - \circ < -15°: cyclonic (LC2) wave breaking
 - o all others: "meridional"



Climatology of



320-K PV streamer frequency displayed as an anomaly relative to the climatological frequency (%, shading; only statistically significant values shown) and composite 320-K PV (PVU, black)



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Preferred regions for PV streamer occurrence associated with EPEs



320-K PV streamer frequency anomaly (%, shading; only statistically significant values shown) averaged between $t_0 - 12$ h and $t_0 + 12$; climatological frequency contoured in black every 0.5%

Approach

- Identify streamers that overlap ¼ of area of 10° × 10° box
- Consider EPE linked to RWB if streamer identified within 24-h period centered on t₀





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Results for 320-K PV streamers

- ~48% (97 of 201) of EPEs linked to RWB
 - **LC1: ~**47% (46)
 - **LC2:** ~26% (25)
 - meridional: ~27% (26)

320-K PV streamer frequency anomaly (%, shading; only statistically significant values shown) averaged between $t_0 - 12$ h and $t_0 + 12$; climatological frequency contoured in black every 0.5%



Results when analysis repeated to include PV streamers identified on at least one of three isentropic surfaces (i.e., 310, 320, 330 K)

- ~76% (153 of 201) of EPEs linked to RWB
 - **LC1: ~**49% (75)
 - LC2: ~23.5% (36)
 - meridional: ~27.5% (42)

Approach:

- Construct composites for EPEs linked to PV streamers identified on 320-K surface
- Examine only LC1 and LC2 cases to highlight distinct EPE scenarios

320-K PV streamer frequency anomaly (%, shading; only statistically significant values shown) averaged between $t_0 - 12$ h and $t_0 + 12$; composite 320-K PV (PVU, black) at t_0 for PV streamers overlapping box



60°N 60°N 017 LC1 40°N 40°N N = 4620°N 20°N 120[']W 120[']W 90°W 150°W 90[°]W 60°W 150°W 60°W 60°N 60°N Do 0 \bigcirc 40°N LC2 N = 25 20°N 20°N 576 0 300 kg m 120[°]W 150°W 120°W 90°W 60°W 150°W 90°W 60°W -2 -2 -10 -6 2 10 hPa -16 -8 2 6 8 16 mm

320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red) 320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg $m^{-1} s^{-1}$, green and vectors)

t_o – 36 h



320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red) 320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg $m^{-1} s^{-1}$, green and vectors)

t_o – 24 h



320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red) 320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg $m^{-1} s^{-1}$, green and vectors)

12 h

60°N 60°N I C1 40°N-N = 4620°N 20° 1014 120[']W 150°W 90[°]W 60°W 120°W 90[°]W 150°W 60°W 020 60°N 60°N 40°N LC2 N = 2520°N 20°N 300 kg m 150°W 120°W 90°W 60°W 150°W 120°W 90°W 60°W -2 -2 -10 -6 2 10 hPa -16 -8 2 6 8 16 mm

320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red) 320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg $m^{-1} s^{-1}$, green and vectors)

0 h





320-K PV (PVU, thick black), SLP (hPa, thin black), SLP anomaly (hPa, shading; only statistically significant values shown), 1000–500-hPa thickness (dam, red) 320-K PV (PVU, black), PW anomaly (mm, shading; only statistically significant values shown), IVT (kg $m^{-1} s^{-1}$, green and vectors)





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Time series of ingredients for precipitation in precipitation domain



Summary

- Climatological and dynamical linkages between RWB and EPEs over portions of the central/eastern U.S. during 1979–2015 examined
- Large majority (~76%) of EPEs examined found to occur in connection with RWB; LC1 dominant relative to LC2
- PV streamers associated with EPEs occur over discrete regions centered over the western U.S.
- RWB linked to formation of high-amplitude, slow-moving wave pattern that establishes persistent corridor of strong water vapor transport (i.e., AR)
- Water vapor transport supports EPE occurrence in presence of dynamical forcing for ascent

Extra slides

Data and methods

Climatology of widespread EPEs





factor of increase in probability of extreme precipitation relative to climatology (shading; only statistically significant values shown) for days during Sep–May on which a streamer overlaps 1/4 of the area of the box

Lagrangian perspective



Trajectory density for 120-h backward trajectories released during EPE from 5° × 5° box centered on maximum precipitation location that exhibited >5 g kg⁻¹ decrease in specific humidity in final 24 h; time-mean composite IVT vectors for t_0 – 72 h and t_0 + 24 overlaid

Linkage of EPEs to Rossby waves



Hovmöller of 250-hPa merid. wind anomalies (m s⁻¹, shading), statistical significance at 95% confidence level (black contours)