Transverse Jet Circulations and their Impact on the Production of Sensible Weather

Andrew C. Winters

1 March 2018











1) What are the characteristics of the Jet Stream?

1) How was the Jet Stream "discovered"?

1) How do transverse jet circulations impact the production of sensible weather?

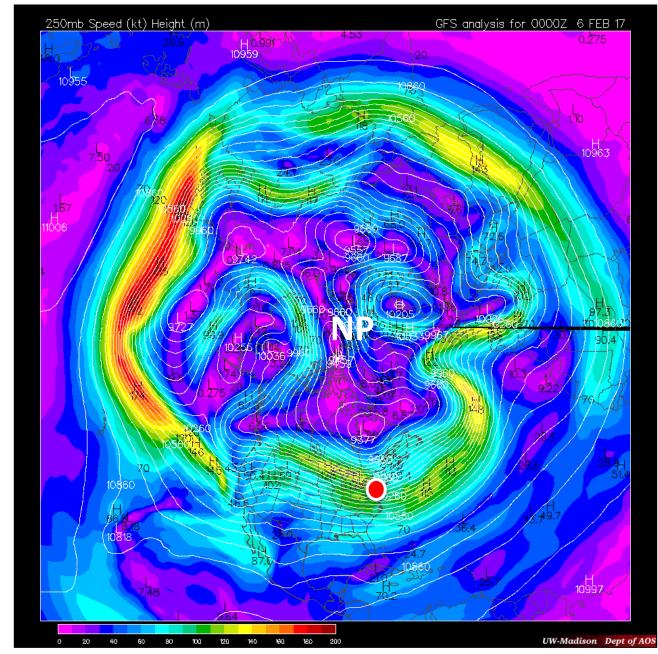
250-hPa Wind Speed

0000 UTC 6 Feb 2017



State College, PA

 \mathbb{NP} North Pole



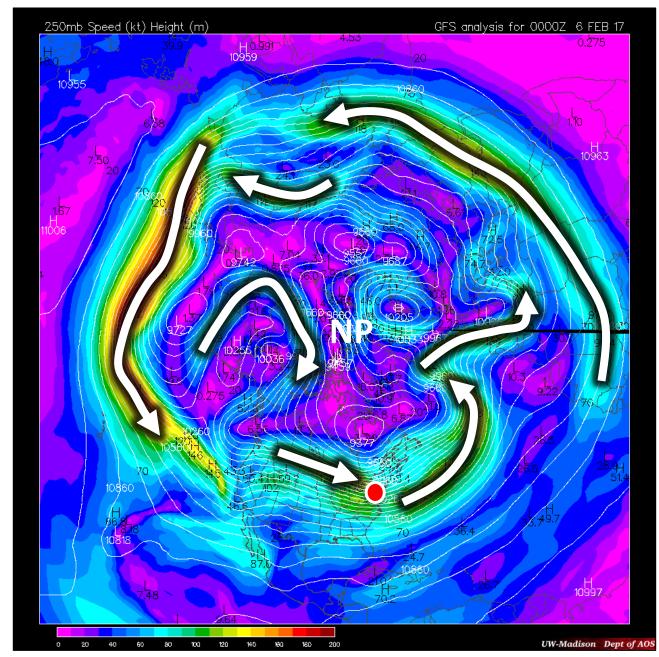
250-hPa Wind Speed

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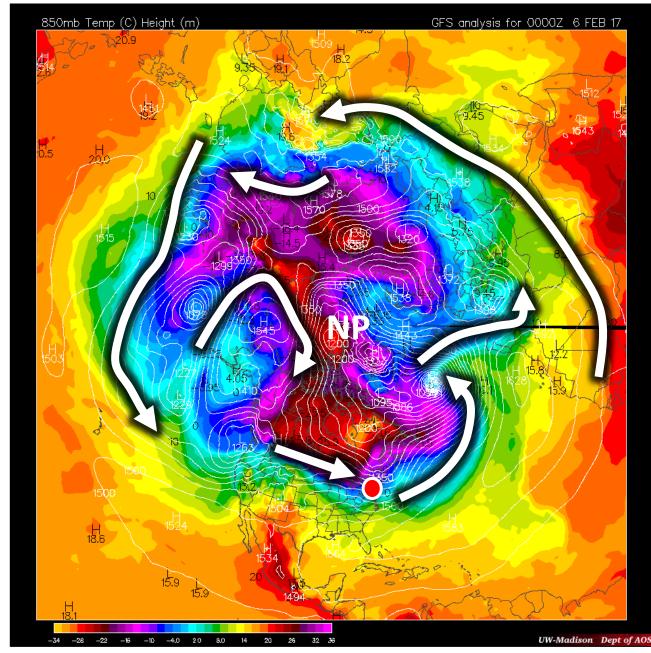
850-hPa Temperature

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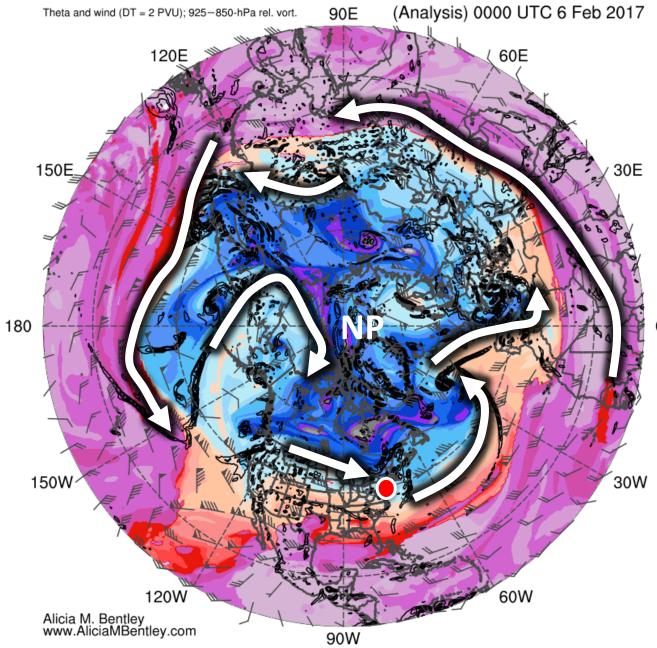
Tropopause Potential Temperature

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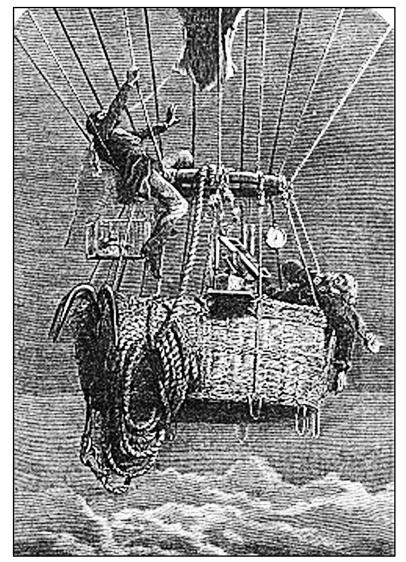
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Coxwell and Glaischer (1862)

The Flight of a Lifetime!

Manned balloon ascent to ~29000 feet.



Illustrated London News

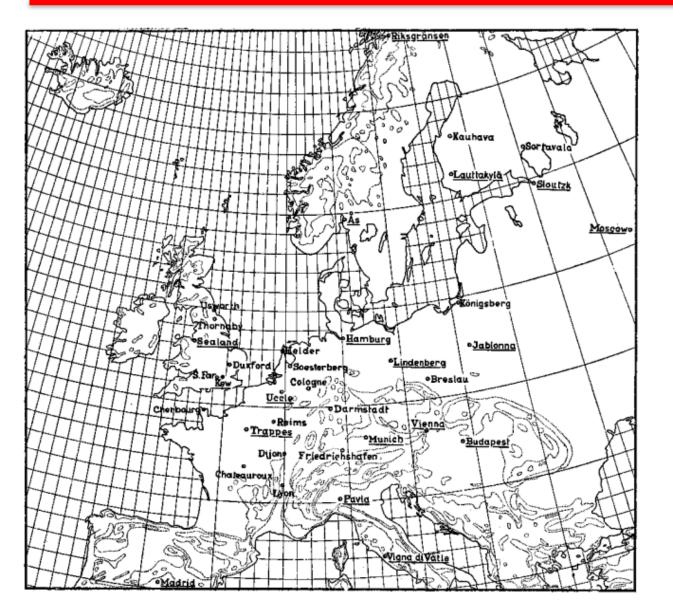
Teisserenc de Bort (1902)

Discovery of the stratosphere

Temperature stops decreasing at a particular distance above the Earth's surface.



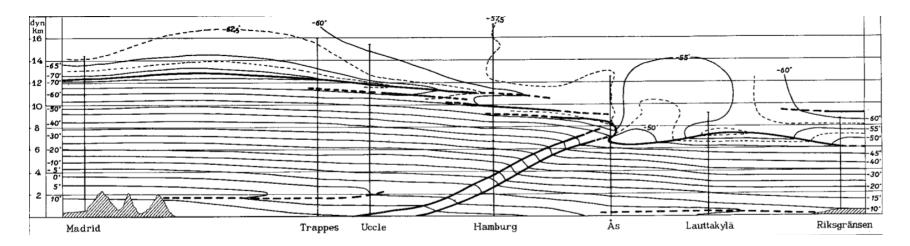
Nature



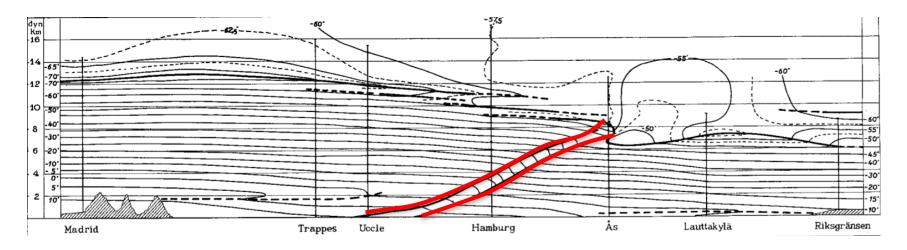
Bjerknes and Palmén (1937)

Coordinated **"swarm ascents"** at 18 different locations across Europe.

Bjerknes and Palmén (1937)

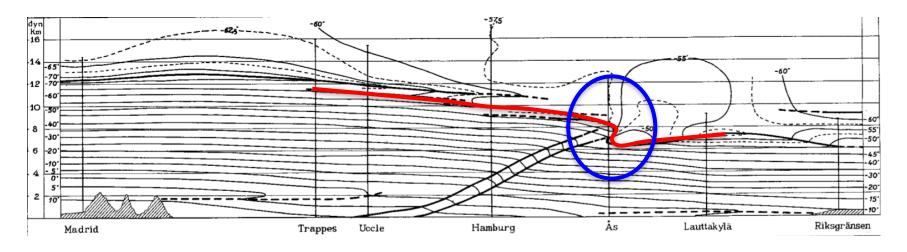


Bjerknes and Palmén (1937)



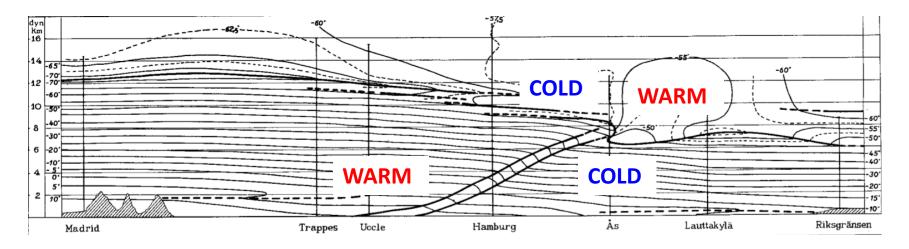
• The front is a **transition zone** across which the temperature gradient is discontinuous.

Bjerknes and Palmén (1937)



- The front is a **transition zone** across which the temperature gradient is discontinuous.
- The tropopause **abruptly lowers** at the location where the polar front intersects the tropopause.

Bjerknes and Palmén (1937)



- The front is a transition zone across which the temperature gradient is discontinuous.
- The tropopause **abruptly lowers** at the location where the polar front intersects the tropopause.
- The meridional temperature gradient reverses directly above the tropopause break.

Reid Bryson and Bill Plumley – Weather Officers in the

Pacific during World War II (1944) (Bryson 1994).



CCR

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MIT

Carl-Gustaf Rossby – First to refer to the phenomenon as the "jet stream" (1947).

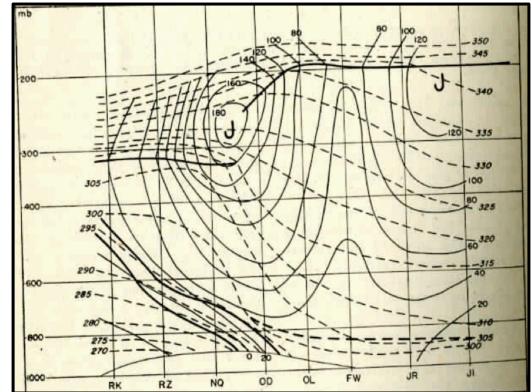
University of Chicago (1947)

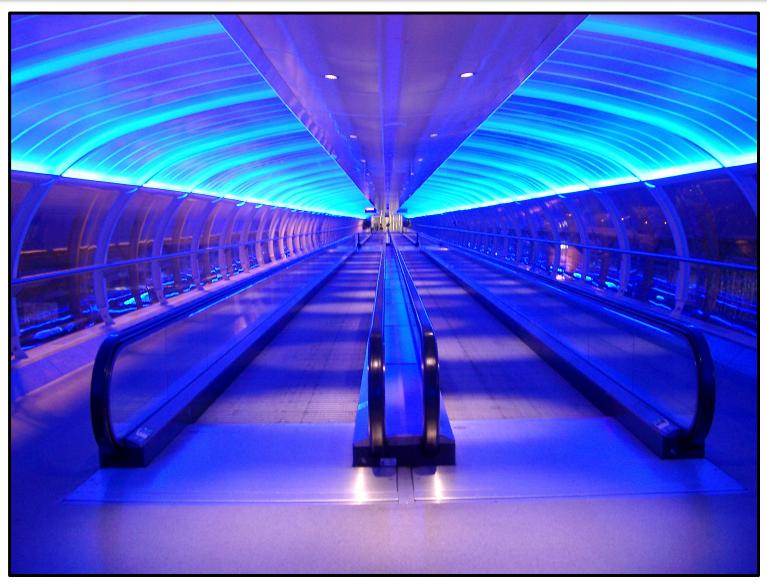
One of the first hemispheric examinations of the midlatitude circulation.

1) The jet was characterized by a nearly continuous band of strong zonal wind speeds.

2) The jet sat atop the strongly baroclinic polar front.

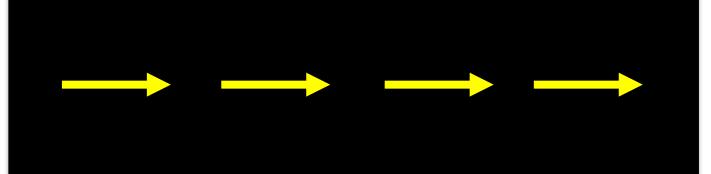
3) The jet was nestled squarely in a tropopause break.



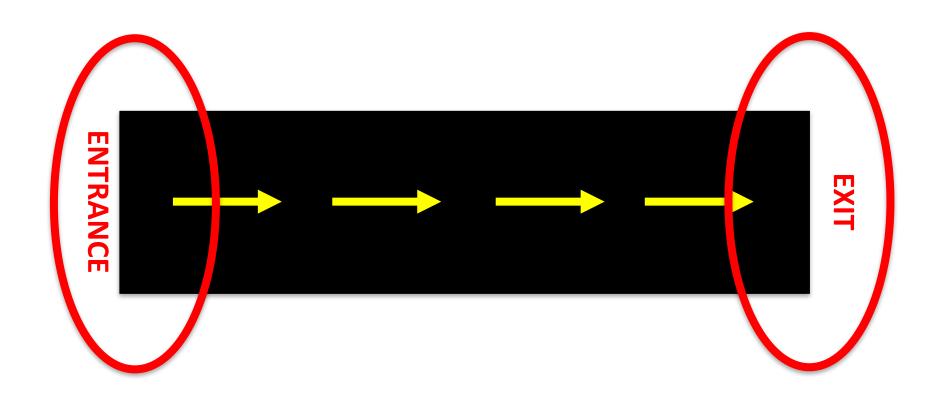


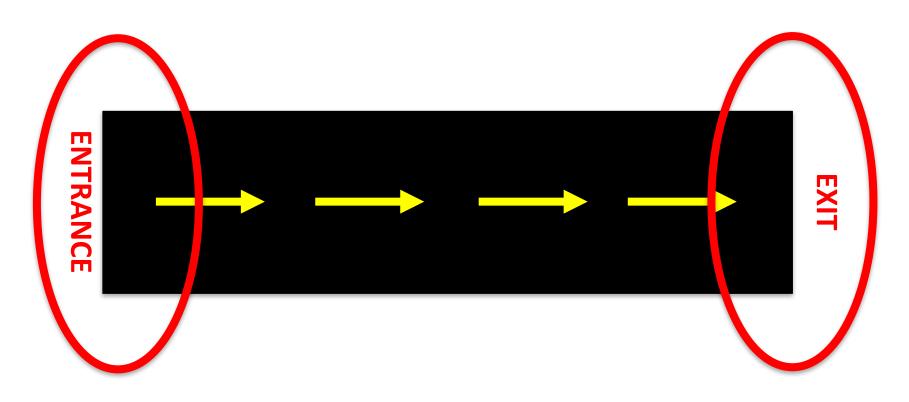
Photoeverywhere.co.uk

ENTRANCE



EXIT





Areas where you accelerate or decelerate with respect to the walkway are important for generating clumsiness

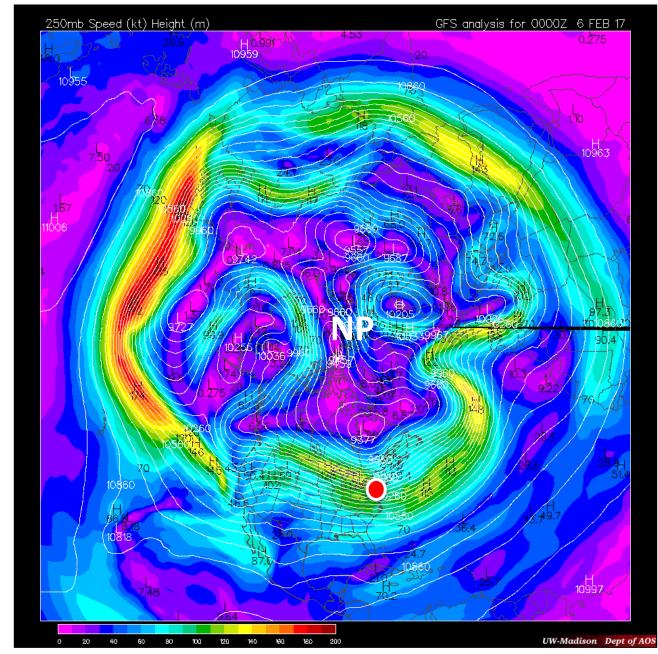
250-hPa Wind Speed

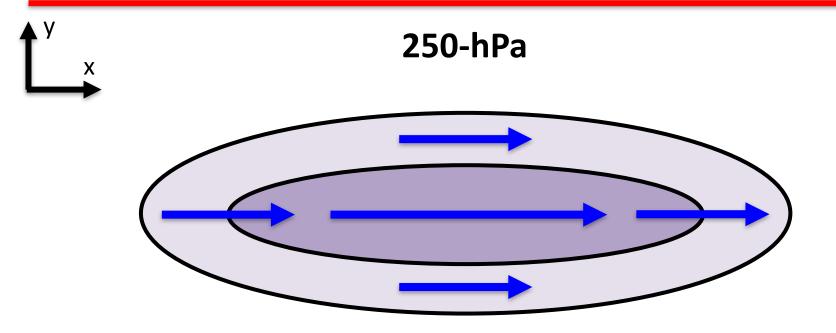
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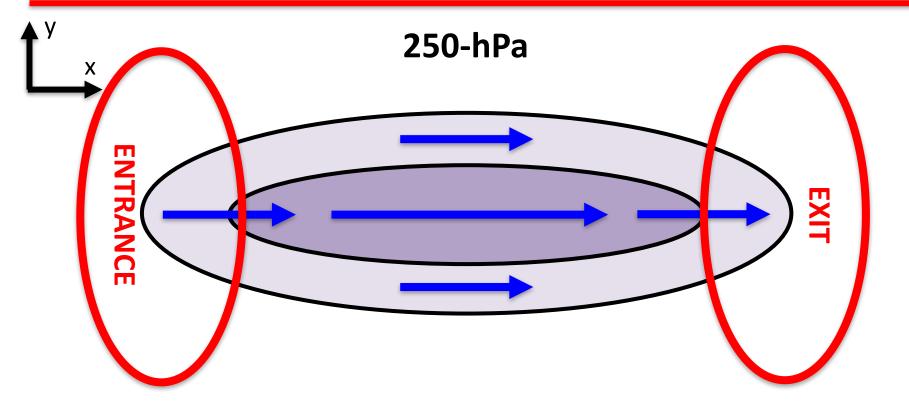
State College, PA

 \mathbb{NP} North Pole

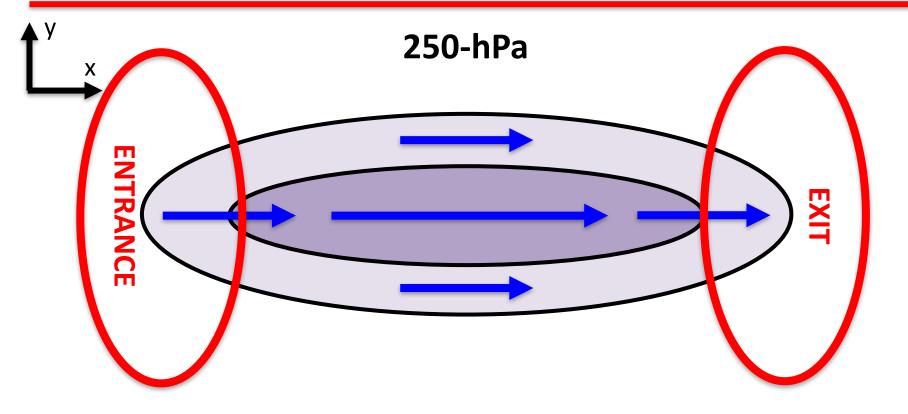


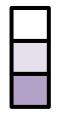


No wind speed Slow wind speed Fast wind speed



No wind speed Slow wind speed Fast wind speed

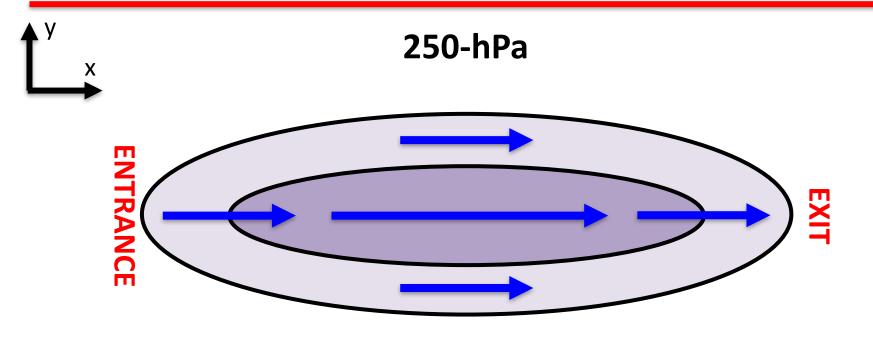




No wind speed Slow wind speed Fast wind speed

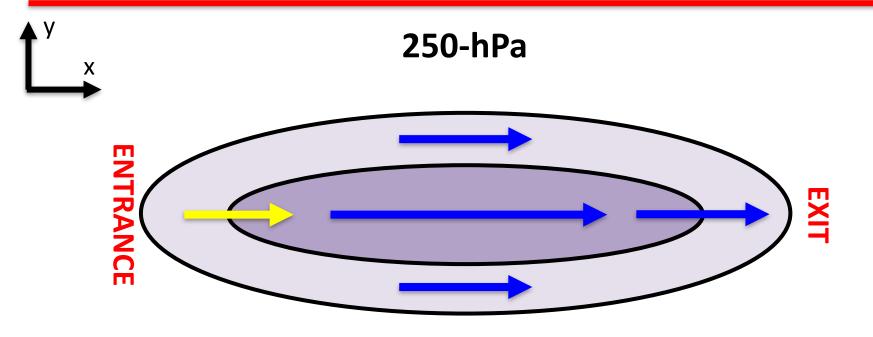
Wind Vectors

Areas where the wind accelerates or decelerates are important for generating weather

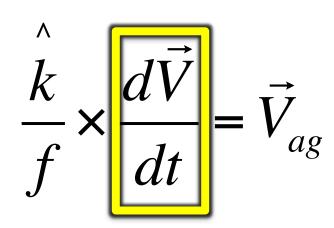


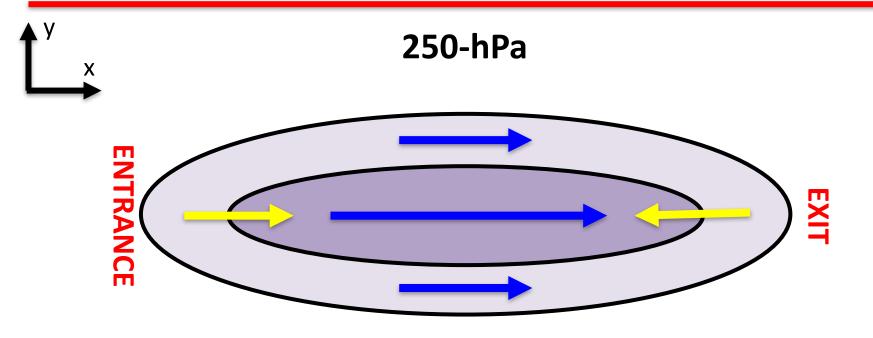
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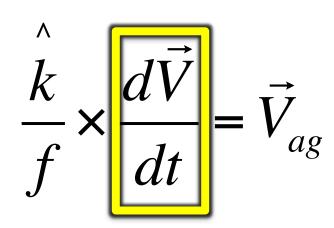


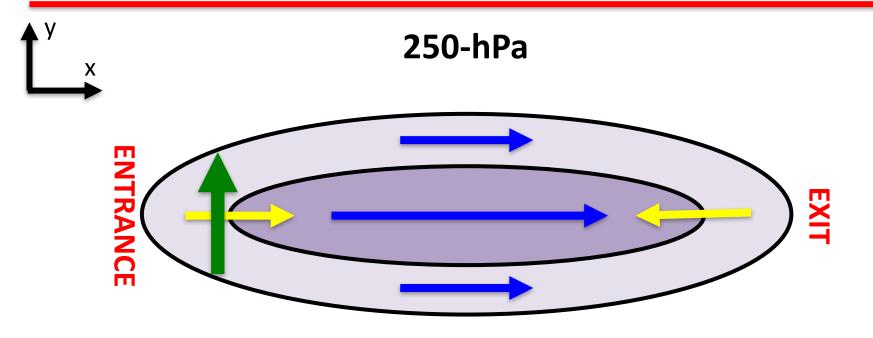
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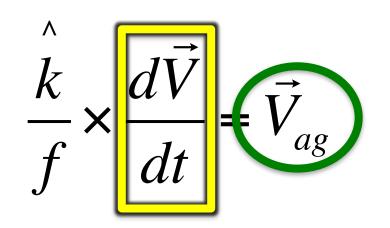


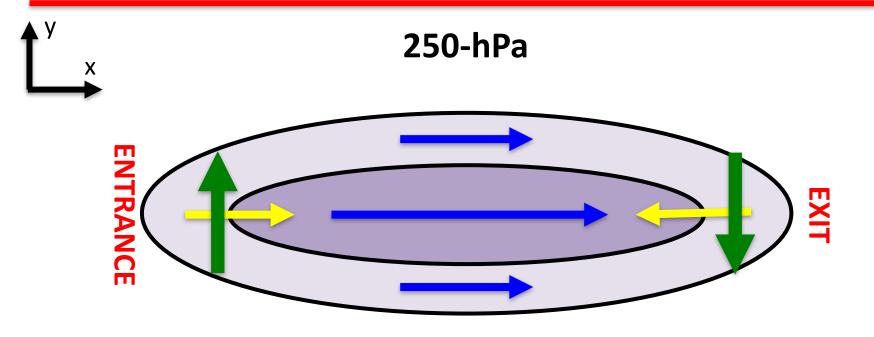
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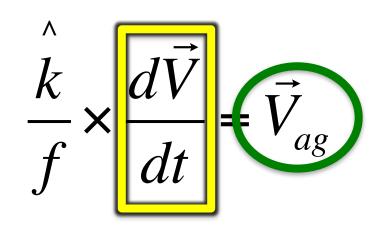
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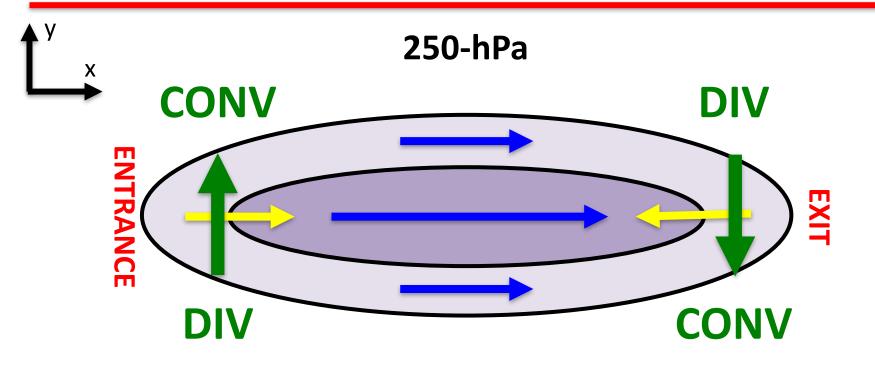




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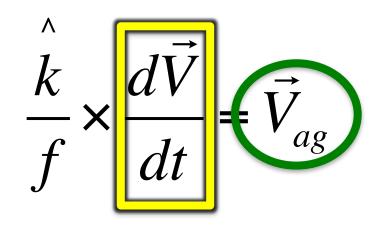
Wind Vectors

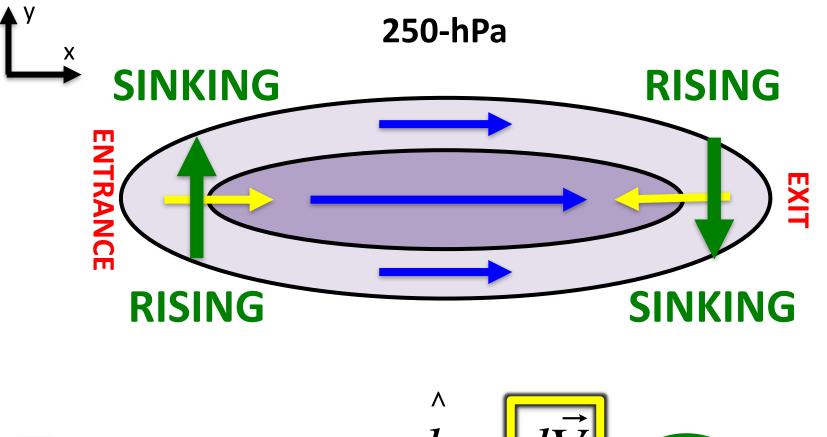




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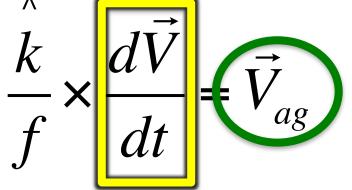
Wind Vectors

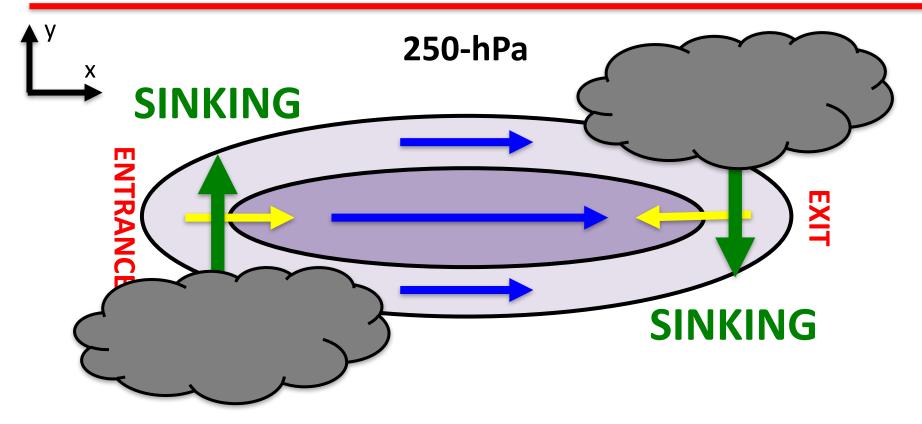


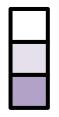


No wind speed Slow wind speed Fast wind speed

Wind Vectors







No wind speed Slow wind speed Fast wind speed

Wind Vectors

Areas where the wind accelerates or decelerates are important for generating weather

Transverse Jet Circulations

The **Sawyer (1956)–Eliassen (1962) Circulation Equation** provides a way to diagnose the transverse circulations associated with active fronts.



Arnt Eliassen

Norwegian Encyclopedia

$$(-\gamma \frac{\partial \theta}{\partial p})\frac{\partial^2 \psi}{\partial y^2} + (2\gamma \frac{\partial \theta}{\partial y})\frac{\partial^2 \psi}{\partial p \partial y} + (-\frac{\partial u_g}{\partial y} + f)\frac{\partial^2 \psi}{\partial p^2} = Q_g - \gamma \frac{\partial}{\partial y}(\frac{d\theta}{dt})$$

Where:

$$\omega = \frac{\partial \psi}{\partial y}$$

$$v_{age} = -\frac{\partial \psi}{\partial p}$$

$$\Delta z_1' \Delta z_2' \Delta z_3' \Delta z_4'$$

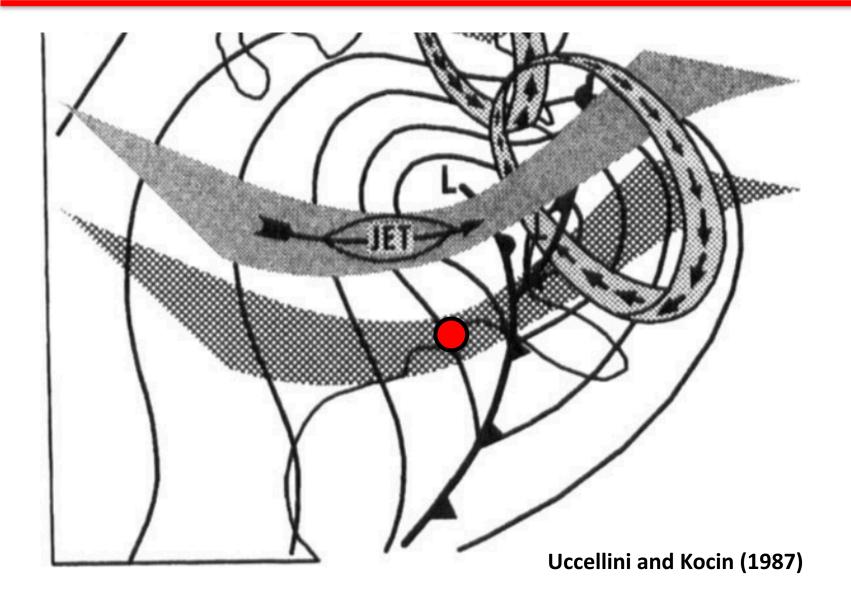
$$(-\gamma \frac{\partial \theta}{\partial p})\frac{\partial \psi}{\partial y^2} + (2\gamma \frac{\partial \theta}{\partial y})\frac{\partial \psi}{\partial p \partial y} + (-\frac{\partial u_g}{\partial y} + f)\frac{\partial \psi}{\partial p^2} = Q_g - \gamma \frac{\partial}{\partial y}(\frac{d\theta}{dt})$$

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$$\left(-\gamma\frac{\partial\theta}{\partial p}\right)\frac{\partial^{2}\psi}{\partial y^{2}} + \left(2\gamma\frac{\partial\theta}{\partial y}\right)\frac{\partial^{2}\psi}{\partial p\partial y} + \left(-\frac{\partial u_{g}}{\partial y} + f\right)\frac{\partial^{2}\psi}{\partial p^{2}} = Q_{g} - \gamma\frac{\partial}{\partial y}\left(\frac{d\theta}{dt}\right)$$

Where:

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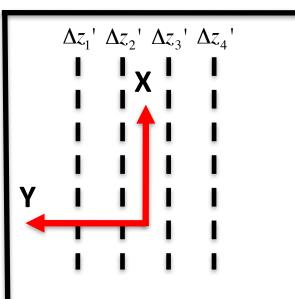
$$\left(-\gamma\frac{\partial\theta}{\partial p}\right)\frac{\partial^{2}\psi}{\partial y^{2}} + \left(2\gamma\frac{\partial\theta}{\partial y}\right)\frac{\partial^{2}\psi}{\partial p\partial y} + \left(-\frac{\partial u_{g}}{\partial y} + f\right)\frac{\partial^{2}\psi}{\partial p^{2}} = Q_{g} - \gamma\frac{\partial}{\partial y}\left(\frac{d\theta}{dt}\right)$$

Static Stability

$$\Delta z_1' \Delta z_2' \Delta z_3' \Delta z_4'$$

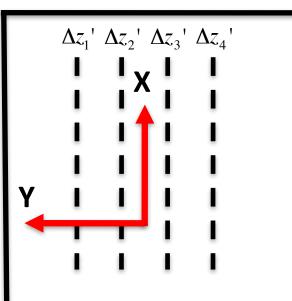
$$\left(-\gamma\frac{\partial\theta}{\partial p}\right)\frac{\partial^{2}\psi}{\partial y^{2}} + \left(2\gamma\frac{\partial\theta}{\partial y}\right)\frac{\partial^{2}\psi}{\partial p\partial y} + \left(-\frac{\partial u_{g}}{\partial y} + f\right)\frac{\partial^{2}\psi}{\partial p^{2}} = Q_{g} - \gamma\frac{\partial}{\partial y}\left(\frac{d\theta}{dt}\right)$$

Static Stability Across-Front Baroclinicity



$$\left(-\gamma\frac{\partial\theta}{\partial p}\right)\frac{\partial^{2}\psi}{\partial y^{2}} + \left(2\gamma\frac{\partial\theta}{\partial y}\right)\frac{\partial^{2}\psi}{\partial p\partial y} + \left(-\frac{\partial u_{g}}{\partial y} + f\right)\frac{\partial^{2}\psi}{\partial p^{2}} = Q_{g} - \gamma\frac{\partial}{\partial y}\left(\frac{d\theta}{dt}\right)$$

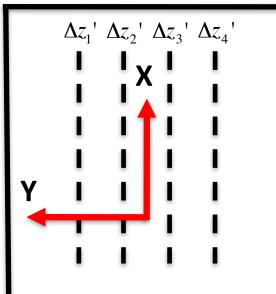
Static Stability Across-Front Baroclinicity Horizontal Absolute Vorticity

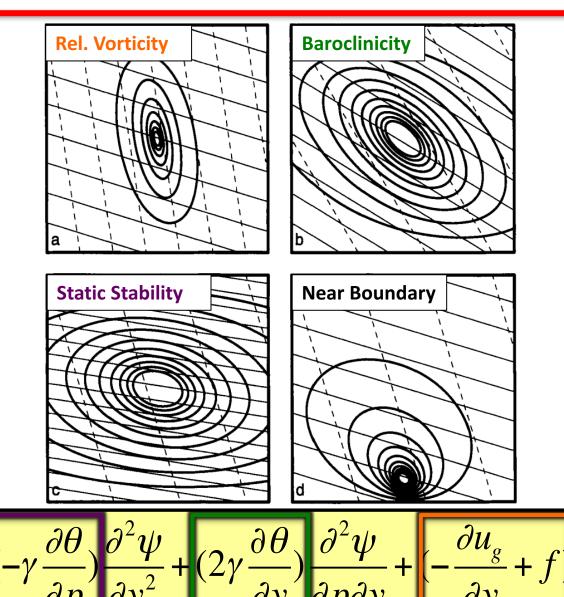


$$\left(-\gamma\frac{\partial\theta}{\partial p}\right)\frac{\partial^{2}\psi}{\partial y^{2}} + \left(2\gamma\frac{\partial\theta}{\partial y}\right)\frac{\partial^{2}\psi}{\partial p\partial y} + \left(-\frac{\partial u_{g}}{\partial y} + f\right)\frac{\partial^{2}\psi}{\partial p^{2}} = Q_{g} - \gamma\frac{\partial}{\partial y}\left(\frac{d\theta}{dt}\right)$$

Static Stability Across-Front Baroclinicity Horizontal Absolute Vorticity

Frontal Characteristics





Hakim and Keyser (2001)

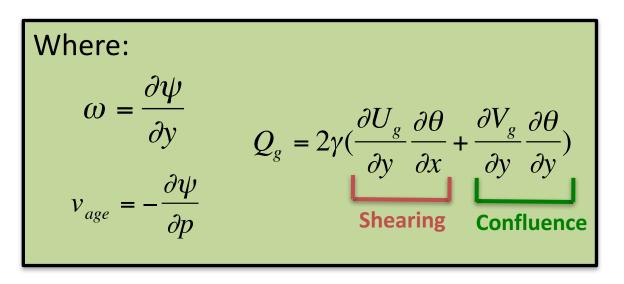
How do the coefficients of the Sawyer–Eliassen Equation modulate the resultant circulation?

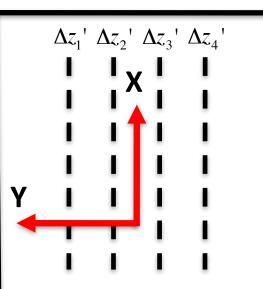
 $\partial^2 \psi$

$$\left(-\gamma\frac{\partial\theta}{\partial p}\right)\frac{\partial^{2}\psi}{\partial y^{2}} + \left(2\gamma\frac{\partial\theta}{\partial y}\right)\frac{\partial^{2}\psi}{\partial p\partial y} + \left(-\frac{\partial u_{g}}{\partial y} + f\right)\frac{\partial^{2}\psi}{\partial p^{2}} = Q_{g} - \gamma\frac{\partial}{\partial y}\left(\frac{d\theta}{dt}\right)$$

Static Stability Across-Front Baroclinicity Horizontal Absolute Vorticity

Frontal Characteristics Geostrophic and Diabatic Forcing





$$(-\gamma \frac{\partial \theta}{\partial p})\frac{\partial^2 \psi}{\partial y^2} + (2\gamma \frac{\partial \theta}{\partial y})\frac{\partial^2 \psi}{\partial p \partial y} + (-\frac{\partial u_g}{\partial y} + f)\frac{\partial^2 \psi}{\partial p^2} = Q_g - \gamma \frac{\partial}{\partial y}(\frac{\partial \theta}{\partial t})$$
$$Q_g = 2\gamma(\frac{\partial U_g}{\partial p}\frac{\partial \theta}{\partial p} + \frac{\partial V_g}{\partial p}\frac{\partial \theta}{\partial p})$$

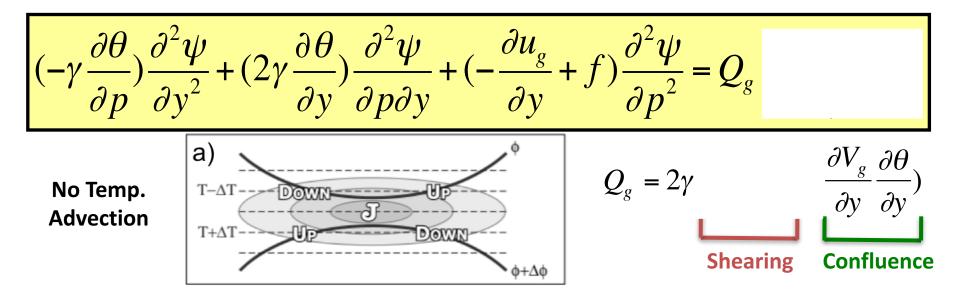
$$Q_{g} = 2\gamma \left(\frac{\partial O_{g}}{\partial y} \frac{\partial O}{\partial x} + \frac{\partial V_{g}}{\partial y} \frac{\partial O}{\partial y}\right)$$

Shearing Confluence

$$(-\gamma \frac{\partial \theta}{\partial p})\frac{\partial^2 \psi}{\partial y^2} + (2\gamma \frac{\partial \theta}{\partial y})\frac{\partial^2 \psi}{\partial p \partial y} + (-\frac{\partial u_g}{\partial y} + f)\frac{\partial^2 \psi}{\partial p^2} = Q_g$$

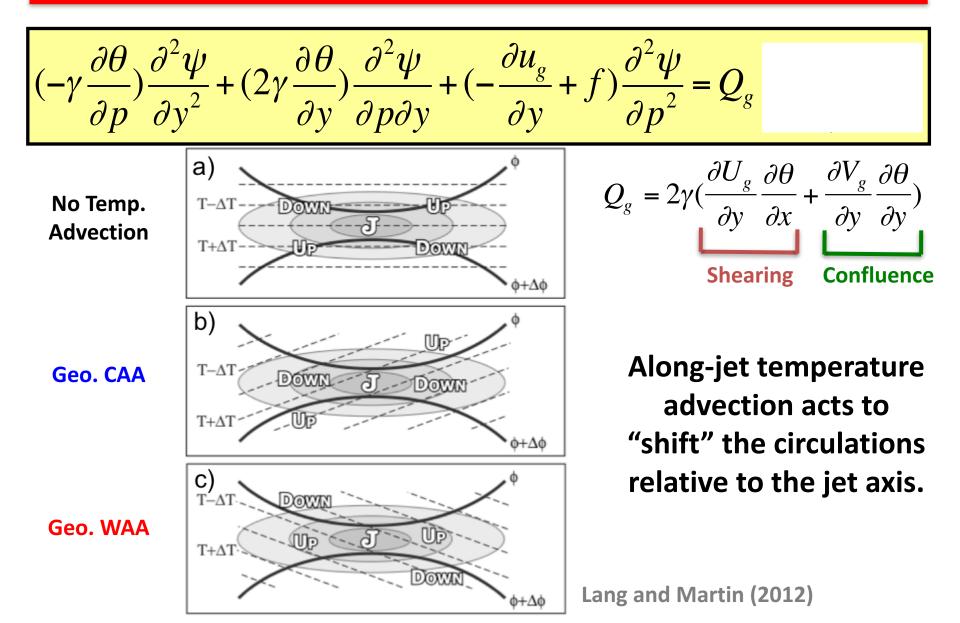
$$Q_{g} = 2\gamma \left(\frac{\partial U_{g}}{\partial y} \frac{\partial \theta}{\partial x} + \frac{\partial V_{g}}{\partial y} \frac{\partial \theta}{\partial y}\right)$$

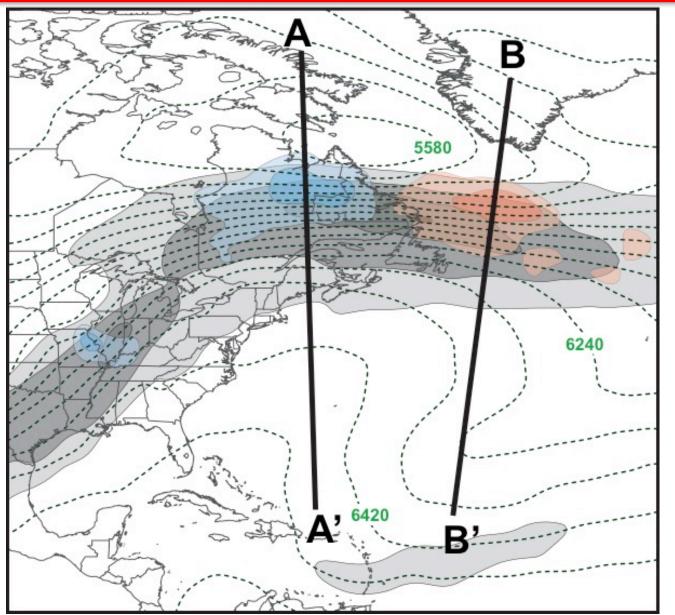
Shearing Confluence



The absence of any along-jet temperature advection returns the traditional fourquadrant model.

Lang and Martin (2012)

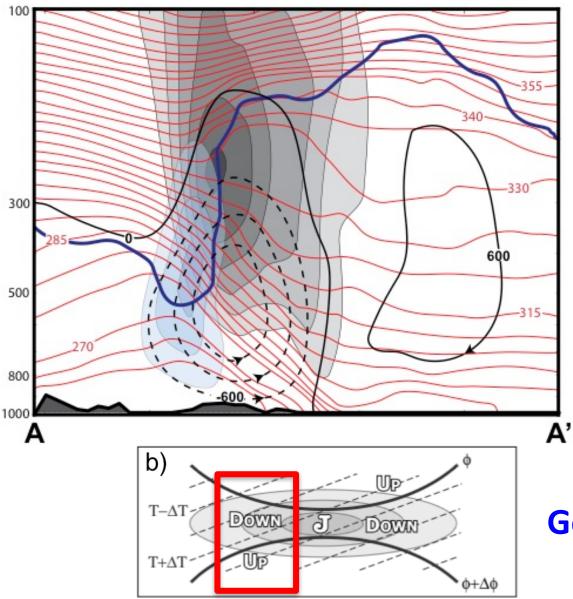




1200 UTC 22 Dec 2013

Geo. CAA in the jet entrance region

Geo. WAA in the jet exit region

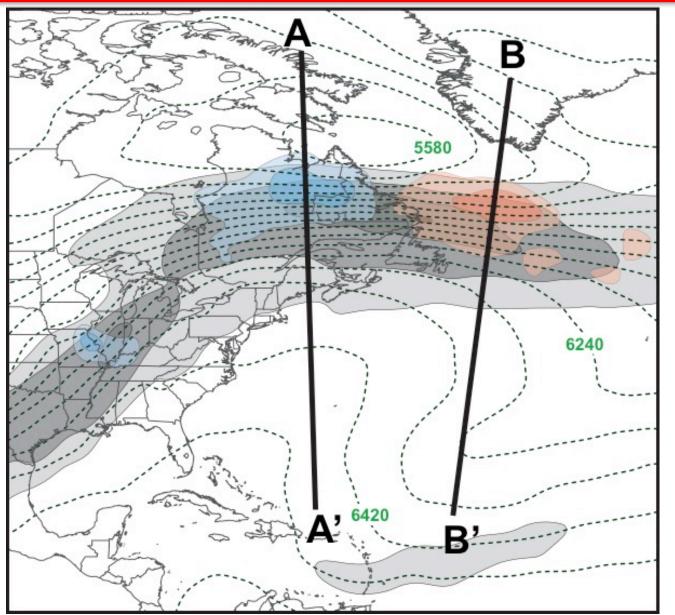


1200 UTC 22 Dec 2013

The result is a **thermally direct circulation**.

Subsidence is present slightly poleward of the jet core.

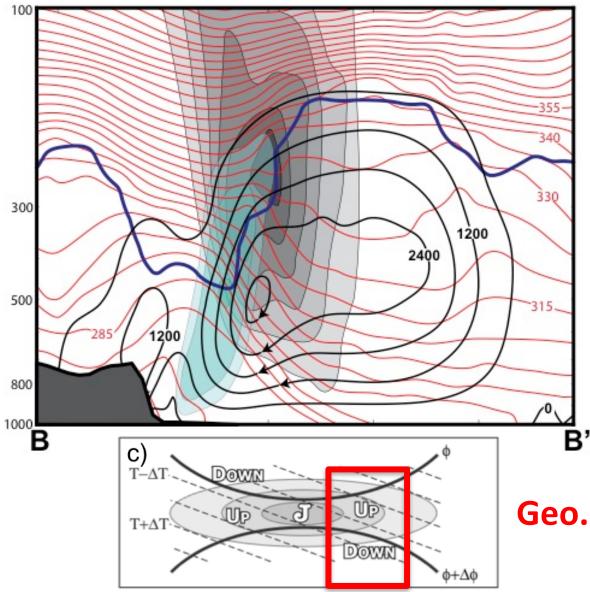
Geo. CAA



1200 UTC 22 Dec 2013

Geo. CAA in the jet entrance region

Geo. WAA in the jet exit region



1200 UTC 22 Dec 2013

The result is a **thermally indirect circulation**.

Ascent is present slightly poleward of the jet core.

Geo. WAA

Impacts of Transverse Jet Circulations on the Production of Sensible Weather

- Severe Weather Outbreaks

(e.g., Omoto 1965; Uccellini and Johnson 1979; Hobbs et al. 1990; Martin et al. 1993)

Impacts of Transverse Jet Circulations on the Production of Sensible Weather

- Severe Weather Outbreaks

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- Cyclogenesis

(e.g., Uccellini et al. 1984; Uccellini et al. 1985; Uccellini and Kocin 1987; Whitaker et al. 1988; Barnes and Colman 1993; Lackmann et al. 1997)

Impacts of Transverse Jet Circulations on the Production of Sensible Weather

- Severe Weather Outbreaks

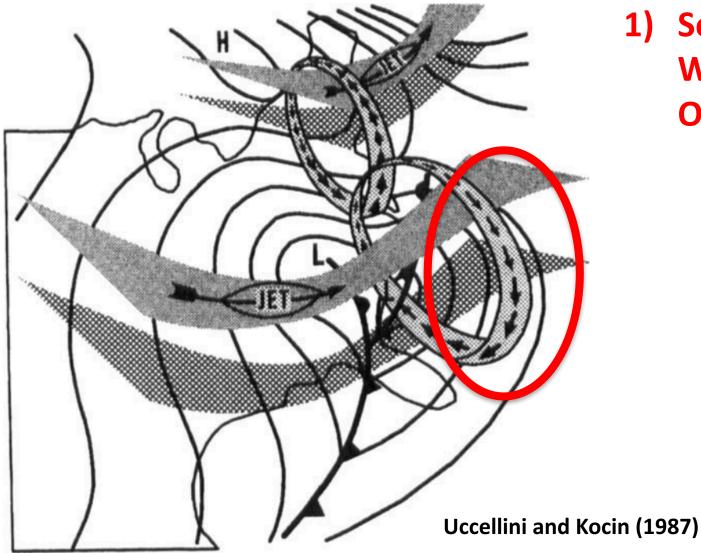
(e.g., Omoto 1965; Uccellini and Johnson 1979; Hobbs et al. 1990; Martin et al. 1993)

- Cyclogenesis

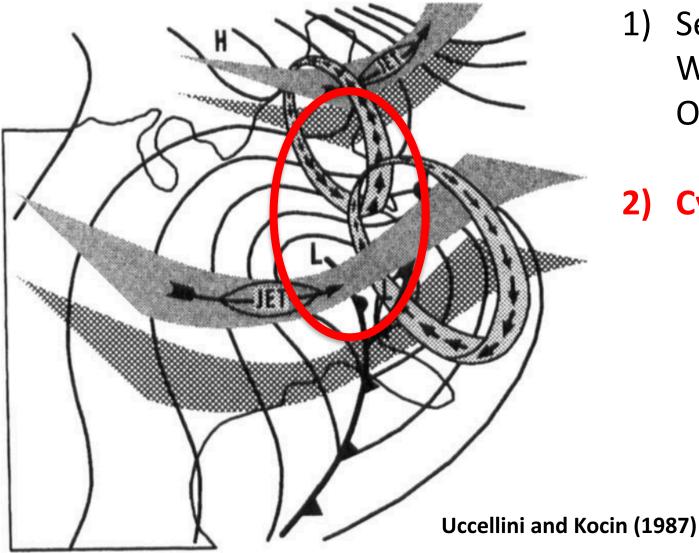
(e.g., Uccellini et al. 1984; Uccellini et al. 1985; Uccellini and Kocin 1987; Whitaker et al. 1988; Barnes and Colman 1993; Lackmann et al. 1997)

- Moisture Transport

(e.g., Uccellini and Johnson 1979; Uccellini et al. 1984; Uccellini and Kocin 1987; Winters and Martin 2014)

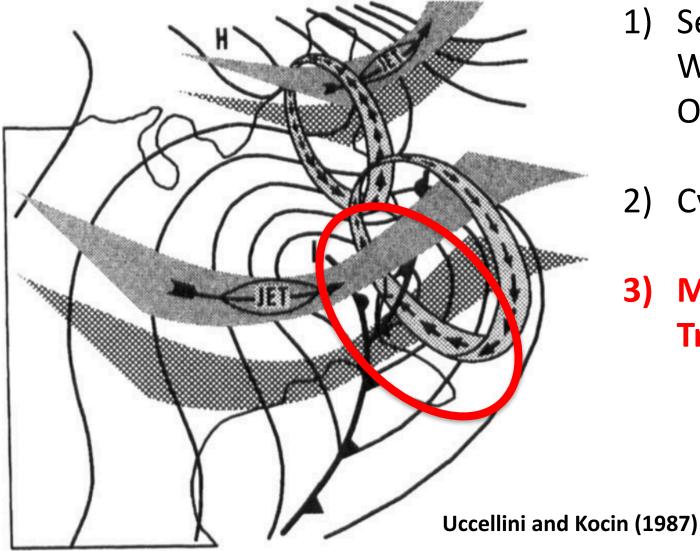


1) Severe Weather Outbreaks



Severe
 Weather
 Outbreaks

2) Cyclogenesis



) Severe Weather Outbreaks

2) Cyclogenesis

8) Moisture Transport

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