Antecedent Synoptic Environments Most Conducive to North American Polar/Subtropical Jet Superpositions

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Maps of tropopause pressure help to identify the location of the jets.

While each jet occupies its own climatological latitude band, substantial meanders are common.

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Occasionally, the latitudinal separation between the jets can vanish resulting in a vertical **jet superposition**.

Modified from Defant and Taba (1957)



The pole-to-equator baroclinicity is combined into a much narrower zone of contrast in the vicinity of a jet superposition.

Intensified frontal structure is often attended by a strengthening of the superposed jet's transverse circulation.

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Christenson et al. (2017) highlight three locations that experience the greatest frequency of jet superpositions:

1) Western Pacific

2) North America

3) Northern Africa

Climatological frequency of Northern Hemisphere jet superposition events per cold season (Nov–Mar) 1960–2010



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Jet superpositions can be an element of high-impact weather events

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 - Jet superposition enhanced the poleward moisture transport via its ageostrophic circulation (Winters and Martin 2014; 2016).



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How do these structures develop?



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Jet Superposition Event Identification and Classification

0000 UTC 27 April 2010



Isolated grid points over North America in the CFSR (Saha et al. 2014) characterized by a polar and subtropical jets during Nov–Mar 1979–2010.



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0000 UTC 24 October 2010



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Isolated grid points over North America in the CFSR (Saha et al. 2014) characterized by a jet superposition during Nov–Mar 1979–2010.



0000 UTC 24 October 2010



1000

В

Winters and Martin (2014); Christenson et al. (2017)

B

1. Isolated grid points over North America in the CFSR (Saha et al. 2014) characterized by a jet superposition during Nov–Mar 1979–2010.

Jet Superposition Frequency – All Times



- 1. Isolated grid points over North America in the CFSR (Saha et al. 2014) characterized by a jet superposition during Nov–Mar 1979–2010.
- 2. Retained analysis times that rank in the top 10% in terms the number of grid points characterized by a jet superposition.

Jet Superposition Frequency – Top 10% Times



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- 2. Retained analysis times that rank in the top 10% in terms the number of grid points characterized by a jet superposition.
- 3. Filtered retained analysis times to group together jet superpositions that are < 30 h and < 1500 km of one another.

Jet Superposition Frequency – Top 10% Times



Jet Superposition Event Classification

1. Determined the mean position of the 2-PVU contour on the 320-K and 350-K surfaces at each analysis time in the CFSR.



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Polar Dominant



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- 2. Compared the position of the jet superposition centroid at the start of each event against the climatological position of the 2-PVU contour.
 - Polar Dominant
 - Subtropical Dominant



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- 2. Compared the position of the jet superposition centroid at the start of each event against the climatological position of the 2-PVU contour.
 - Polar Dominant
 - Subtropical Dominant
 - Hybrid

















Number of events

Frequency of Subtropical Dominant Jet **Events**



Frequency of Subtropical Dominant Jet **Superposition Events**



Frequency of Subtropical Dominant Jet Superposition Events



 Legend
Centroid of all events
Composite movement of superposition

Jet Superposition Event Composites:

Polar Dominant vs. East Subtropical Dominant

Polar Dominant Jet Superposition Events

2 Days Prior to Jet Superposition



N=80

Polar Dominant Jet Superposition Events

1 Day Prior to Jet Superposition



Anom.

N=80





m s⁻¹

mm



Polar Dominant Jet Superposition Events



Polar Dominant Jet Superposition Events



Polar Dominant Jet Superposition Events





1 Day Prior to Jet Superposition

N=76







250-hPa Jet, Geo. Height, & Geo. Height Anom., & 300-hPa Geo. Temp Adv.





250-hPa Jet, Geo. Height, & Geo. Height Anom., & 300-hPa Geo. Temp Adv.


















Future Work

- Apply piecewise PV inversion (e.g., Davis and Emanuel 1991) to quantify the influence that polar cyclonic and tropical anticyclonic PV anomalies have on deforming the tropopause during each type of superposition event.
- Examine the impact that each type of jet superposition event has on the evolution of the downstream large-scale flow pattern.
- Utilize numerical simulations of jet superposition events to examine the sensitivity of jet superposition to diabatic processes.
- Further illuminate the connection between jet superposition events and high-impact weather events (e.g., severe weather, cyclogenesis, floods).

Supplementary Slides

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Sample Jet Superposition Centroid Calculation

Calculated the centroid of each jet superposition based on all valid grid points at a particular analysis time.

To calculate the centroid, there must exist a group of 18 superposition grid points, of which no superposition grid point is >1000 km away from another superposition grid point.





Frequency of Dominant Jet Superposition Events



Frequency of West Subtropical Dominant Jet Superposition Events



Jet Superposition Event Classification



Polar Dominant Jet Superposition Events

3 Days Prior to Jet Superposition



N=80

E. Subtropical 250-hPa Jet, Geo. Height, & Geo. Height Anom., & 300-hPa Geo. Temp Adv. **Dominant Jet Superposition** -10080 **Events 3 Days** m s⁻¹ **Prior to Jet** 250-hPa Jet, MSLP Anom., PWAT Anom., & OLR Anom. **Superposition** -4 _4 **N=76**

Δ

mm

Jet Superposition Conceptual Model



Jet Superposition Conceptual Model



Ageostrophic Transverse Jet Circulations

Upper Troposphere



Traditional four-quadrant model

Geo. cold-air advection (CAA)

along the jet axis promotes **subsidence** through the jet core

Geo. warm-air advection (WAA)

along the jet axis promotes ascent_through the jet core

Lang and Martin (2012)

Background

Insight into how the tropopause can be restructured from a PV perspective can be found by consulting Wandishin et al. (2000)

Two processes can account for "foldogenesis":

- Differential vertical motions can <u>vertically</u> <u>steepen</u> the tropopause.
- 2) Convergence or a vertical shear can produce a differential horizontal advection of the tropopause surface.



These same mechanisms are also likely to play an important role in superpositions.

Background



FIG. 2. (a) Cold season average of zonally averaged Δy (km) for 5-K isentropic layers ranging from 300–305 to 365–370 K. The 315–330- and 340–355-K layers are highlighted in light gray shading. (b) The average frequency of occurrence of grid points with a maximum wind speed value within the 5-K isentropic layers along the abscissa per cold season. The 315–330- and 340–355-K layers are shaded in blue and red, respectively.