

CONTINENTAL COLLISIONS ANALAGOUS TO THAT  
FORMING THE QINGHAI-XIZANG PLATEAU

Kevin Burke, J.F. Dewey, W.S.F. Kidd and A.M.C. Şengör  
Department of Geological Sciences  
State University of New York at Albany  
Albany, New York 12222 U.S.A.

INTRODUCTION

Papers presented at this symposium have provided, for the first time, a coherent picture of the present state and geological development of the Qinghai-Xizang plateau as a response to the Cenozoic collision of India with Asia along the Indus-Yarlung Zangbo suture zone.

The importance of the collisional process in tectonic evolution has been appreciated since Wilson (1) pointed out that recognition of the plate structure of the lithosphere indicated that, through much of earth history, oceans have been opening in some places and closing in others.

The process of ocean closing is complex and diverse and the variety of phenomena associated with collision are well illustrated in the Alps and the Mediterranean. A distinctive set of collisional phenomena developing after suturing involves continental thickening, partial melting and both intrusive and extrusive magmatic activity (2). In the Qinghai-Xizang plateau we interpret the extent of these processes to be very generally indicated by: (I) the area of very high ground (> 5 km) and (II) the envelope of areas of geothermal activity in the Himalaya and on the plateau (Tong Wei and Zhang Mingtao, this volume).

The exact mechanism of continental thickening and partial melting is as yet poorly understood and future studies of the Qinghai-Xizang plateau are most likely to throw light on the processes that operate. It is clear that thickening the continent in response to convergence will bring its lower part into a realm where isotherms are normally hot enough to induce partial melting but the thickened continent will be cold and conduction alone could only warm it very slowly. Some combination of conduction and convective heat transfer seems essential and the role of the mantle is as yet uncertain. Some workers (e.g. Toksöz, this volume) have suggested that the occurrence of an unusually hot uppermost mantle may influence partial melting of the continent after collision.

ANALOGUES OF TIBET

There are numerous areas on earth where igneous rocks whose formation has involved partial continental melting are distributed over a wide area on one side of a suture zone that marks a former continental collision. In some of these areas the time of suturing can be shown to have preceded by a few million years the episode of widespread igneous activity. We have

interpreted such areas as analogues of the present Tibetan plateau (2, 3) Table 1 lists a selection of these areas in order of increasing age. No effort has been made to be exhaustive in the list. Although some of the areas are better known than others, the main aim of the compilation is to show that Tibet-style continental reactivation phenomena are and have been (at least since the Proterozoic) a relatively common element in earth history.

Figure 1 is a map of the world illustrating the distribution of the sutures and Tibet-style terrains listed in the table. These areas constitute a substantial proportion of all continental crust, and we suggest that Tibetan style reactivation represents a significant stage in continental evolution involving the separation of high-level rocks enriched in water and large ion lithophile elements from anhydrous granulitic lower crust.

#### References

- (1) Wilson, J.T., Static or mobile earth: the current scientific revolution, in *Gondwanaland Revisited: New evidence for continental drift*. Proc. American Philos. Soc., 112, 309-320 (1968)
- (2) Dewey, J.F., Burke, K.C.A., Tibetan, Variscan and Precambrian basement reactivation: Products of continental collision, *Jour. Geology* 81, 783-792 (1973)
- (3) Burke, K., Dewey, J., The lost oceans of Africa, *Jour. Min. Geol. Nigeria Min. Geol. Met. Soc.* 6, 75 (1971)
- (4) Şengör, A.M.C., Kidd, W.S.F., Post-collisional tectonics of the Turkish-Iranian Plateau and a comparison with Tibet, *Tectonophysics* 55, 361-376 (1979)
- (5) Trümpy, R., et al., *Geology of Switzerland, A Guidebook*, Wepf and Co. Publishers, Basel (1980)
- (6) Bögel, H., Schmidt, K., *Kleine Geologie der Ostalpen*, Ott Verlag, Thun (1976)
- (7) Fujita, K., Pre-Cenozoic tectonic evolution of northeast Siberia, *Jour. Geology* 86, 159-172 (1978)
- (8) Şengör, A.M.C., Mid-Mesozoic closure of Permo-Triassic Tethys and its implications, *Nature* 279, 590-593 (1979)
- (9) Şengör, A.M.C., Yilmaz, Y., Ketin, I., Remnants of a pre-late Jurassic ocean in northern Turkey: fragments of Permo-Triassic Palaeo-Tethys? *Geol. Soc. America Bull.* 91 (Part I), 599-609 (1980)
- (10) Ivanov, S.N., Per filiev, A.S., Poutchkov, V.N., Les traits principaux de la structure géologique de l'Oural, in *La Chaîne Varisque d'Europe moyenne et occidentale*, CNRS Paris, 571-582.
- (11) Hatcher, R.D., Jr., Tectonics of the western Piedmont and Blue Ridge southern Appalachians: Review and speculation, *Am. Jour. Sci.* 278, 276-304 (1978)
- (12) Wickham, J., Roeder, D., Briggs, G., Plate tectonics models for the Ouachita foldbelt, *Geology* 4, 173-176 (1976)
- (13) King, P.B., Marathon revisited, in *Symposium on the geology of the Ouachita Mountains* 1, 41-69 (Arkansas Geological Commission (1977))
- (14) Burke, K., Dewey, J.F., Orogeny in Africa, in *African Geology*, Ibadan Univ. Press, Ibadan, 583-608 (1972)
- (15) Şengör, A.M.C., Butler, J.C., The Llano Uplift, central Texas: a Proterozoic example of continental collision? *Geol. Soc. America Abst. Progs.* 9, 72-73 (1977)
- (16) Hoffman, P., Dewey, J.F., Burke, K., Aulacogens and their genetic relation to geosynclines with a Proterozoic example from Great Slave Lake, Canada, in *Soc. Econ. Paleontologists and Mineralogists Spec. Pub.* 19, 38-55 (1974)

TABLE 1

## CONTINENTAL COLLISIONS WITH TIBETAN-STYLE REACTIVATION

|    | COLLISIONAL<br>EVENT                | SUTURE<br>ZONE                           | AGE  | KEY REFERENCES   |
|----|-------------------------------------|--|--|--|
| 1  | Arabia-<br>Eurasia                  | Bitlis-Zagros                            | Began in Mid-<br>Miocene,<br>continues                         | (4) Şengör and Kidd<br>(1979)  |
| 2  | Alps                                | Arosa-Matrei<br>zones and<br>equivalents | Late Middle<br>Eocene  | (5) Trümpy <i>et al.</i><br>(1980)<br>(6) Bögel and Schmidt<br>(1976)              |
| 3  | India-<br>Eurasia                   | Indus-Yarlung<br>Zangbo                  | Middle to<br>Late Eocene                                       | (2) Dewey and Burke<br>(1973)<br>Numerous papers in<br>this volume                 |
| 4  | Verkoyansk                          | Chersky                                  | End Jurassic   | (7) Fujita (1978)  |
| 5  | Cimmerian<br>Continent-<br>Laurasia | Various Palaeo-<br>Tethyan sutures       | Late Triassic<br>to Late<br>Jurassic                           | (8) Şengör (1979)<br>(9) Şengör <i>et al.</i><br>(1980)<br>Şengör (this<br>volume) |
| 6  | Hercynian                           | Lizard-Harz Orsteigal-<br>Massif Central | Upper Carb.<br>Mid-Carb.                                       | (2) Dewey and Burke<br>(1973)  |
| 7  | Eurlaurentia-<br>Asia               | Urals                                    | Early Permian  | (10) Ivanov <i>et al.</i><br>(1977)  |
| 8  | Eurlaurentia-<br>Gondwanaland       | S. Appalachian-<br>Ouachita-Marathon     | Late Carb.   | (11) Hatcher (1978)<br>(12) Wickham <i>et al.</i><br>(1976)<br>(13) King (1977)    |
| 9  | Pan African                         | Numerous sutures                         | Latest Pre-<br>cambrian<br>Locally into<br>Lowest<br>Paleozoic | (14) Burke and Dewey<br>(1972)   |
| 10 | Grenville                           | Llano (Texas)                            | $1 \times 10^9$ years  | (2) Dewey and Burke<br>(1973)<br>(15) Şengör and<br>Butler (1977)                  |
| 11 | Coronation                          | Wopmay                                   | $1.8 \times 10^4$ years  | (16) Hoffman <i>et al.</i> ,<br>(1974)   |

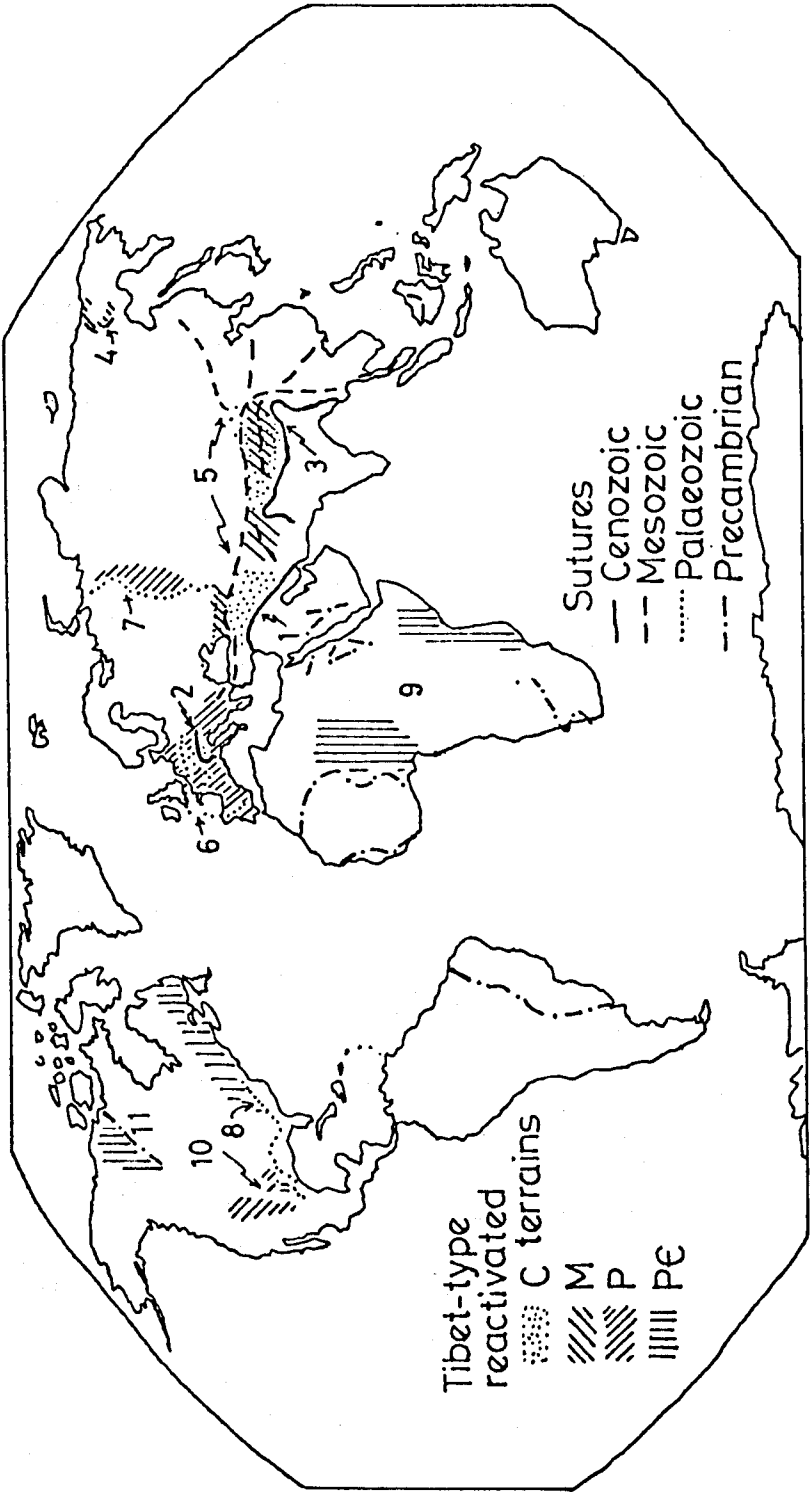


Figure 1. Location of selected Tibetan-type reactivated terrains and suture zones with ages ranging from Cenozoic to Proterozoic. Numbers correspond to those in Table 1.