GRENVILLE BASEMENT IN THE PIEDMONT EAST OF THE PRE-APPALACHIAN (PRE-CALEDONIAN) EDGE(?) OF THE NORTH AMERICAN CRATON

20502

FARRAR, Stewart S. and GLOVER, Lynn, III, Orogenic Studies Laboratory, Department of Geological Sciences, Virginia Polytechnic Institute and State University, Blacksburg, VA 24061 Several recent tectonic models of the Appalachians interpret the western edge of the central Piedmont gravity high as the pre-Appalachian edge of the North American craton (Rankin, 1975; Hatcher and Zietz, 1980; Hatcher, 1982). The Piedmont to the east is variously interpreted as part of Africa, Avalonia, or some "suspect terrane" formed elsewhere.

Mapping of the eastern Piedmont has defined the Grenville age Goochland granulite terrane, 20-30 km in width, running from near Raleigh, N. C. to Fredricksburg. Va. and continuing under Coastal Plain cover to the north. This terrane consists of Grenville garnet-pyroxene and two-pyroxene granulites and sillimanite-K-feldspar pelitic gneisses in a terrane which has been retrogressively metamorphosed by an upper greenschist to middle amphibolite facies Alleghanian event. The Goochland terrane is in pre-Alleghanian thrust contact with the late Precambrian to early-Paleozoic Eastern slate belt and Chopawamsic volcanics. In the north it lies on and in the south it lies well to the east of the central Piedmont gravity high, and is interpreted to have been thrust from farther east. The simplest explanation for the origin of this terrane is that it is a microcontinent which was rifted from North America in the late-Precambrian and reattached in the mid-Paleozoic.

PRELIMINARY RESULTS OF A DETAILED STUDY OF PART OF THE SAURATOWN MOUNTAINS BASEMENT MASSIF, NORTH CAROLINA

No 20504

HATCHER, Robert D., Jr., HEYN, Teunis, MCCONNELL, Keith I., Department of Geology, University of South Carolina, Columbia, SC 29208, and GODFREY, Stephen C., Route 7, 96 Hathaway Circle, Greenville, SC 29609.

Detailed geologic mapping in the southwestern portion of the Sauratown Mountains anticlinorium basement massif has yielded some interesting preliminary results. A sequence of Grenville basement (?) orthogneiss is separated from high grade cover metasedimentary rocks by a mylonitic interval suggesting existence of a premetamorphic detachment thrust. These rocks either grade stratigraphically upward into or metamorphic grade drops upward (southeastward, northwestward) into a kyanite or low er grade assemblage of metasandstone, quartzite, pelitic schist (graphitic to the northwest), and amphibolite (formerly called Ashe Formation) Cover sequences are distinctively different of the NW and SE flanks of the anticlinorium suggesting that if they are the same, considerable shortening has occurred. All the rocks and contacts in this area are polydeformed in the usual sequence of two or more sets of isoclinal folds superceded by crenulation cleavage, then more brittle structures The entire sequence is cut by brittle Mesozoic faults (Stony Ridge fault zone with small displacement and diabase dikes. Most rock units are superbly exposed in creeks but poorly exposed along roads, which, because most mapping in this region has been reconnaissance, has led to the common misconception of no exposure in this area and the rest of the To date we have made significant changes in locations of contacts on published USGS 2-degree reconnaissance maps raising questions regarding the utility of this program in complex terranes with poor road exposure and little existing detailed geologic mapping.

EVIDENCE AGAINST PRE-GRENVILLE METAMORPHISM IN THE READING PRONG OF SOUTHEASTERN NEW YORK

Nº 21005

MOSE, Douglas G. and NAGEL, Susan, Department of Geology, George Mason University, Fairfax, VA 22030, JACKSON, Richard J., Exploration, Exxon Company, Denver, CO 80201, and TUNSOY, Ahmet, Department of Geology, Florida State University, Tallahassee, FLA 32306 A few reports using Rb-Sr whole-rock and U-Pb zircon data show that the Reading Prong, Manhattan Prong and Blue Ridge contain rocks older than about 1050-1150 m.y., the time of the Grenville high-grade meta-morphic event. However, the meaning of the pre-Grenville ages is in doubt. A compilation of Rb-Sr whole-rock isochron ages is shown below (1 sigma error is less than 5% for the ages and less than 0.2% for the initial 875r/865r ratios):

ROCK UNIT	AŘEA	AGE	I.R.	INDEX TO AREAS
Fordham Gneiss- outcrop	ave. A	1350	0.7021	A= Manhattan Prong
Fordham Gneiss- site A	Α	1113	0.7248	(GSA v. 93, p. 391)
Fordham Gneiss- site B	Α	1100	0.7138	(d3A v. 33, p. 331)
Fordham Gneiss- site C	Α	1326	0.7025	B≕ Western Side of
Highland Metavolcanics	В	1139	0.7029	the Hudson River
Highland Paragneiss	В	1147	0.7046	C= Eastern Side of
Storm King Pluton	В	1106	0.7066	the Hudson River
Canopus Pluton	В	1070	0.7038	che nadson kivel
Canada Hill Pluton	В	913	0.7186	D= New Milford and
Reservoir Granite Gneiss	С	1225	0.7025	Housatonic Highland
Pink Granite Gneiss	D	1045	0.7040	

An examination of these and other data plus a consideration of the behavior of stratified rocks (volcanic and sedimentary) vs. crystalline rocks (plutonic and metamorphic) suggests that the Grenville event was not preceded by a significantly earlier metamorphic event.

A THREE-STAGE MODEL OF ADIRONDACK TECTONIC HISTORY

Nº 14266

MHITNEY, Philip R., Geological Survey, NY State
Museum, Rm. 3140 CEC, Albany, NY 12230

A three-stage tectonic history is proposed to account for the principal features of Adirondack geology. (1) Prior to "Grenvillian" tectonic activity, Proterozoic rocks in the Adirondacks consisted of platformtype sediments with some volcanics, resting unconformably on older basement. About 1.1 ga, mantle upwelling developed along a NNE trend from the southern Adirondacks into the Central Granulite Terrain in Quebec. A large cushion of basaltic magma accumulated between subcrustal lithosphere and the lower crust causing partial melting and intrusion of the sphere and the lower crust causing partial metring and intrusion of the managerite-charnockite suite into higher crustal levels. Fractionation of the basaltic magma led to high-Al basalt which intruded the crust and differentiated further to yield the anorthosite suite (Emslie, 1978). Anorthosite was then diapirically emplaced at relatively shallow depths. (2) Delamination or convective thinning of subcrustal lithosphere was followed by intracontinental subduction, formation of large NE-trending To llowed by intracontinental subduction, formation of large NE-trending folds and nappes, and SE-over-NW thrusting. Doubling of the continental crust over the Adirondack highlands and the Central Granulite Terrain led to development of the presently observed 8 ± 1 kb metamorphic assemblages. (3) Renewed vertical movements of large, low density and relatively rigid anorthosite bodies produced further intense deformation of earlier structural (3) Renewed vertical movements of large, low density and tion close to the anorthosite, and deflection of earlier structural trends. Generalized uplift of the zone of doubled crust led to NW tilting of the Carthage-Colton mylonite zone which marks the boundary between the upper (NW lowlands) and lower (Adirondack highlands) crustal slabs. Evidence from numerous recent Adirondack studies is presented in support of this model, and tests of the validity of the model are

LINEATIONS, SHEATH FOLDS, AND IMPLICATIONS FOR ADIRONDACK TECTONIC HISTORY

No 17326

MCLELLAND, James, Colgate Univ., Hamilton, NY
Quartzofeldspathic gneisses of the southern Adirondacks exhibit pronounced E-W lineations consisting principally of lithic ribbons; mullion-like corrugations; mafic rich streaks; and oriented elongate minerals. The most prominent linear fabric occurs within gneisses of the Piseco Group and the overlying Sacandaga Fm. and is best developed where these units are exposed in the Piseco anticline. Here the rocks resemble pencil gneisses consisting of ribbons of quartz and feldspar. The average dimension of these ribbons is (0.05x0.25x20) cm. The origin of the ribbons has been established by tracing lineated units into regions of lesser deformation where they appear to be megacrystic quartz monzonitic gneisses. Lineation of these rocks is thought to have been caused by regional simple shear accompanied by flattening and resulting in ductile stretching giving rise to EW ribbons, mullions, corrugations, etc.. These lineations parallel the axes of the two earliest recognized fold sets of the region. The first of these is isoclinal-recumbent and the second is open and upright. Recently discovered sheath folds are elongated E-W and appear to consist of early isoclinal axes rotated into their present orientation from an intial N-S alignment. It is believed that the initial isoclinal folding took place in response to E-W simple shear responsible for the strong regional lineation. Perturbations in the flow field resulted in rotation of the early fold axes, bring them into parallelism with the stretching lineations, Upright E-W folds of the second generation may result from convergences in the flow field. The assymetry of feldspar augen and their tails suggest that displacement in the simple shear field was east side up and to the west. The Carthage-Colton mylonite zone may represent a zone of high strain parallel to the plane of regional simple shear. The regional shear couple probably resulted from plate interaction and crustal thickening during the Grenville Orogeny.

MYLONITE ZONES, THRUSTS, AND TECTONICS OF THE GREN-VILLE IN THE ADIRONDACKS

No 19883

ANDERSON, S.L., BURKE, K.C., KIDD, W.S.F., and PUTMAN, G.W., Dept. of Geological Sciences, SUNY-Albany, Albany, N.Y. 12222 Relationships between granitic and metasedimentary rocks in the Adirondacks have been explained by two models, neither of which are much concerned with the great variations in strain evident in the rocks: intrusion of sills, or a folded basement-cover assemblage. In the latter. huge isoclinal folds have been postulated and used to explain layerparallel foliation and small-scale isoclinal folds. We suggest that, as in the Alps and Himalayas, such nappes are better viewed as thrust slices. The foliation is most intense in discrete zones parallel with large-scale lithologic layering; these high-strain zones contain much mylonite and at least one contains pseudotachylite. One in the south-central Adirondacks, the Sacandaga "Formation" of McLelland, is an extensive, mylonitic high strain zone containing interleaved metapelitic and mafic rocks. Before reorientation by major folds, it was regionally gently dipping. We suggest that it is a major thrust; a strong E-W elongation lineation widespread in the southern Adirondacks gives the movement direction. Other major, originally gently inclined mylonitic zones exist elsewhere in the Adirondacks and we also identify them as significant thrusts. Without younging or fossil evidence it is difficult to show that a deformed lithologic succession is a depositional sequence and major thrusts make it improbable that more than partial depositional sequences are present. Wilson-cycle interpretation of Grenville tectonics permits a major continental collision to explain the thrusts and later folding (Himalayas) but more complex models including earlier tectonics along an active convergent margin (e.g. Tibet) are perhaps more probable.

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