SEDIMENT CHARACTERISTICS AROUND THE KENYON ISLAND GROUP, GREAT SACANDAGA LAKE (NY): ECONOMIC POTENTIAL OF DREDGING AND LAND RECLAMATION

A thesis presented to the Faculty

of the University at Albany, State University of New York

in partial fulfillment of the requirements

for the degree of

Master of Science

College of Arts and Sciences

Department of Earth and Atmospheric Sciences

Arthur M. Ambrosino 2001

University at Albany, State University of New York

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The thesis for the master's degree submitted by

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under the title

SEDIMENT CHARACTERISTICS AROUND THE KENYON ISLAND GROUP, GREAT SACANDAGA LAKE (NY): ECONOMIC POTENTIAL OF DREDGING AND LAND RECLAMATION

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ABSTRACT

The Kenyon Island Group lies within the Great Sacandaga Lake's largest and most unnavigable shallow water shoal. The shoal measures approximately 5.0 mi². The Kenyon Islands, including Mead and Deer Islands, occupy approximately ³/₄ mi² during the months of annual high lake level. During the late shallow water season of 1998, a NE to SW diagonal transect of nine core samples were collected by a Geoprobe coring device, with recoveries between 8' and 16' deep sections. In this study these cores were used to identify the stratigraphy, classify the sediment grain sizes, evaluate the abundance of economic minerals, and calculate the most cost effective and environmentally sound method of deepening the lakebed for navigation and recreational purposes and reconfiguring the new real property created.

Here I present evidence for the existance of a previously unidentified large moraine field, of approximately 12 mi^2 , located between two different end moraines of the Pleistocene Epoch. Today, about 1/3 of that moraine field surrounds the Kenyon Island Group, within the shorelines of the Great Sacandaga Lake. The south and east end moraine is visible in an arcuate path which includes the two major Kenyon Islands and Deer Island. The south and east end moraine is approximately $\frac{3}{4}$ mile wide and is confined by the valley walls of the Sacandaga Basin. This end moraine must have been at least 300 feet high during the Pleistocene Epoch. The north and west moraine was deposited before the south and east end moraine and is also visible in exposures around the lake. This forms moraine rock fields in arcuate bands, which are confined by the valley walls of the Sacandaga Basin. The north and west end moraine was originally

½ miles wide and must have stood at least 120 feet high.

The moraine field contains all the classic landmarks including kames, kettle lakes, eskers, drumlins, flutes and fossil streambeds. The moraine field sediments that exist are no more than 45 feet thick beneath the lakebed and 45 to 120 feet thick outside the shorelines and above bedrock. The bedrock may be Cambrian (570-510 Ma), including Little Falls Dolomite, Theresa Dolomite and/or Potsdam Sandstone.

This thesis proposes a remedy involving dredging and earth-moving heavy equipment to permanently deepen part of the lakebed of the Great Sacandaga Lake that is currently of little use because the area is typically too shallow for regular lake navigation. If there are minerals within the sediments of the study area that are of current economic importance and have values significant enough to pay for a lakebed deepening, the geochemistry, mineralogy and mineral chemistry will reveal them. These minerals can then be identified and evaluated as the mechanism that will support the tremendous costs associated with such a major deepening effort. The long term outcome of a Great Sacandaga Lake deepening, beside improving the navigational, recreational, and fisheries of the lake will create room for an additional 200 billion gallons of water, which could be priceless in the next frontier.

The moraine field sediment profile in the study area may be unique to the whole of the Great Sacandaga Lake. The question of dredging the lakebed versus simpler earthmoving heavy equipment to reconfigure the lakebed, anywhere else in the Great Sacandaga Lake, will most certainly require further research and could yield different results.

Acknowledgements

I would particularly like to thank Dr. John I. Garver for all the effort he extended this project. He gave of himself and the splendid Union College facilities, far more than would have been expected and his insight and counsel was greatly appreciated.

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CHAPTER 1

INTRODUCTION

Purpose of the study

During the last several decades, there has been an increasing dispute between the downstream "interests" and the "recreational users" of the water of the Great Sacandaga Lake in upstate New York. The downstream "interests" include power generation dams, Hudson River water levels, ocean water/salt front concerns at Poughkeepsie and flood control at Albany, among others.

Large annual water-level variations to accommodate the downstream interests cause large areas of very shallow water in the Great Sacandaga Lake (GSL), especially in midsummer, which is a situation the landowners and boating public find unacceptable. In an effort to reach accommodations to all, an Upper Hudson-Sacandaga Comprehensive Settlement Negotiation Team is presently studying water level remedies (See Appendix A-Permits and History). They have proposed raising the GSL high water mark from 768' to 770' over a twenty-year time span. Any remedy they agree upon will be a compromise between upstream and downstream interests.

This thesis proposes a remedy involving dredging and earth-moving heavy equipment to permanently deepen part of the lakebed of the Great Sacandaga Lake that is currently of little use because the area is typically too shallow for regular lake navigation. To accomplish this task it is necessary to identify the stratigraphy of the lakebed sediments at various points within the lake, classify the sediment grain sizes,

evaluate the potential economic mineral wealth, and determine the most cost effective and environmentally sound method of deepening the lakebed and reconfigurating existing Islands and shorelines.

The Kenyon Island Group, perhaps the largest and most unnavigable shallow water shoal in the Great Sacandaga Lake, is the ideal location for investigating such a project. The Kenyon Island Group contains the former and now submerged community of Osborne Bridge. To address the above issues a transect of core samples were collected by Geoprobe in the Kenyon Island Group using gravel roads which were serviceable during the late shallow water season of 1998. Nine Geoprobe 1 3/4" diameter cores were collected, between 8' and 16' in length.

Location of study area

The Kenyon Island Group is located in the north central portion of the Great Sacandaga Lake, in the town of Northampton, in northeast Fulton County, New York State (Figure 1.1). The Kenyon Island Group is just over two miles south of Northville, New York. Some portions of the Great Sacandaga Lake are located in Northwest Saratoga County, but are outside this project review. The whole of the lake lies within the perimeter ("Blue Line") of the Adirondack State Park (Figures 1.2).

The 1908 New York State Water Supply Commission Maps, numbers 20 and 21, were digitized, merged and enhanced to show the GPS and coring sites (Figure 1.3).

These maps depict the area before the Conklingville Dam was built in 1929, which subsequently flooded the area creating the Sacandaga Reservoir.

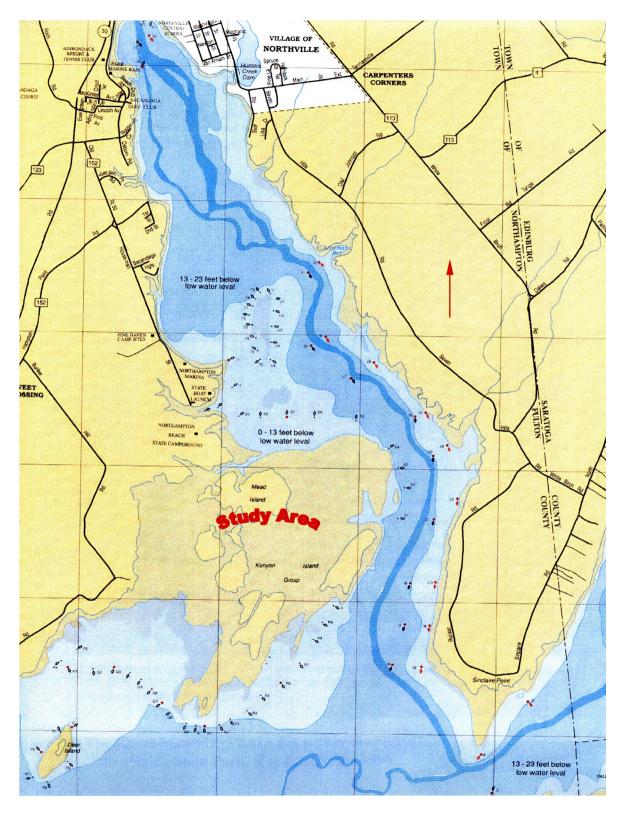


Figure 1.1 The Kenyon Island Group, Great Sacandaga Lake Navigation Map, Jimapco, 1991.

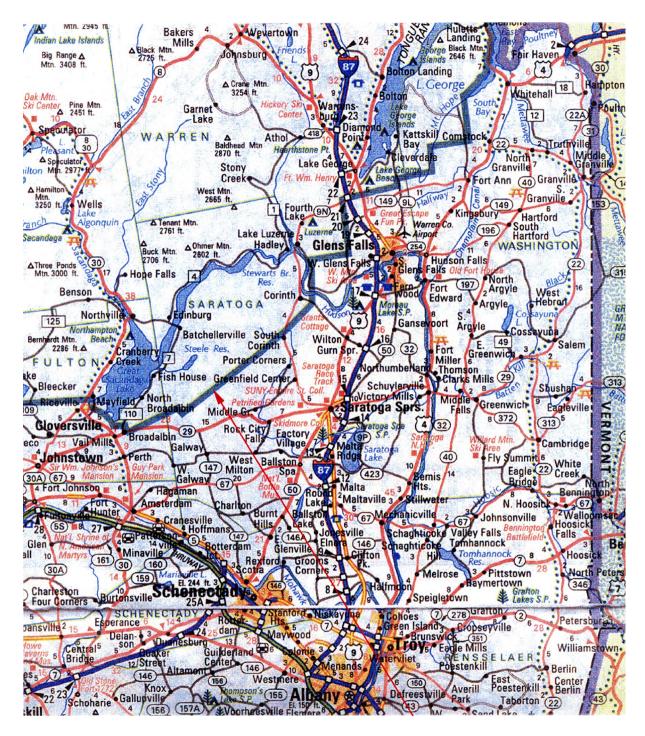


Figure 1.2 The Great Sacandaga Lake lies within the "Blue Line" of the Adirondack State Park, Rand McNally, 1999.

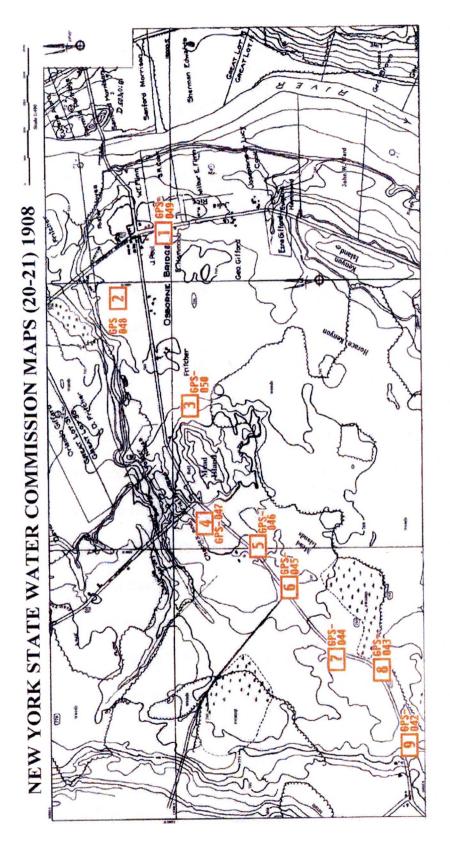


Figure 1.3 The 1908 New York State Water Commission Maps, Numbers 20 and 21, were Digitized, Merged and Enhanced to show the GPS and Coring Sites in red.

Regional geology

The geologic structure of the Sacandaga Basin lies on a great sunken block between two Precambrian crystalline blocks of the Mohawk Valley fault series. In the lower Paleozoic Era, during the Taconic Orogeny (~ 450 MA), the region was faulted due to flexure (downwarping). The bedrock across the Mohawk Valley was cut by ten or more northeast trending faults, which have greatly influenced modern topography. The faults are almost vertical with the downdropped block to the east (Figure 1.4). One exception is the fault located in the Sacandaga basin study area (Arnow, 1951).

The Noses Fault crosses Fulton County southwest of Johnstown (New York), trending northeastwardly, crossing the county line northwest of Northville. The upthrown block, at the northern most end, is on the west side of the fault. This side forms the southern-most crystalline rock escarpment of the Adirondack Mountains. Less than 6.5 *mi* to the east of the Noses, the Batchellerville fault forms another escarpment of crystalline rock extending from Batchellerville (in Saratoga County) southward past Northampton (Fish House), and then southeastward.

The Batchellerville fault also has a displacement of about 1500' but the upthrown block side is to the East. The Sacandaga Reservoir and all the nearby communities, including the submerged town of Osborne Bridge and the Kenyon Islands lie on the plain of this downdropped block, between these two Precambrian crystalline masses (Figure 1.5), (Fisher, Isachsen and Rickard, 1995).

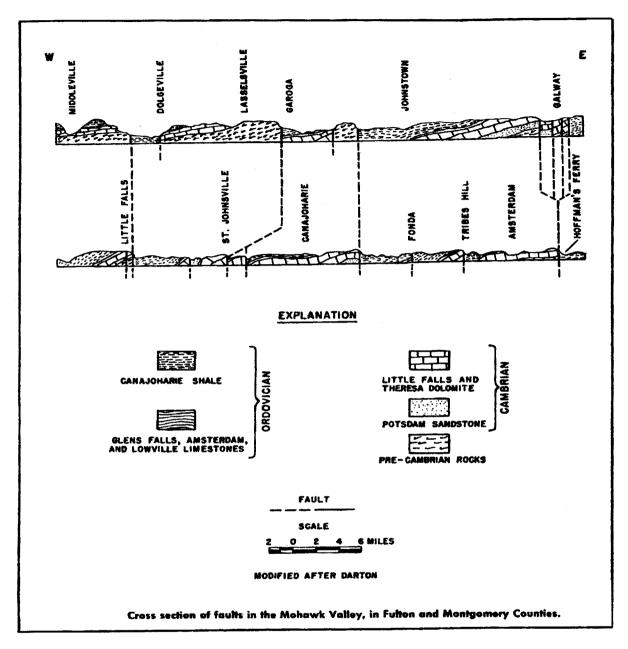


Figure 1.4 The Ground Water Resources of Fulton County, New York, U.S. Geological Survey, Bulletin GW – 24, Arnow, 1951.

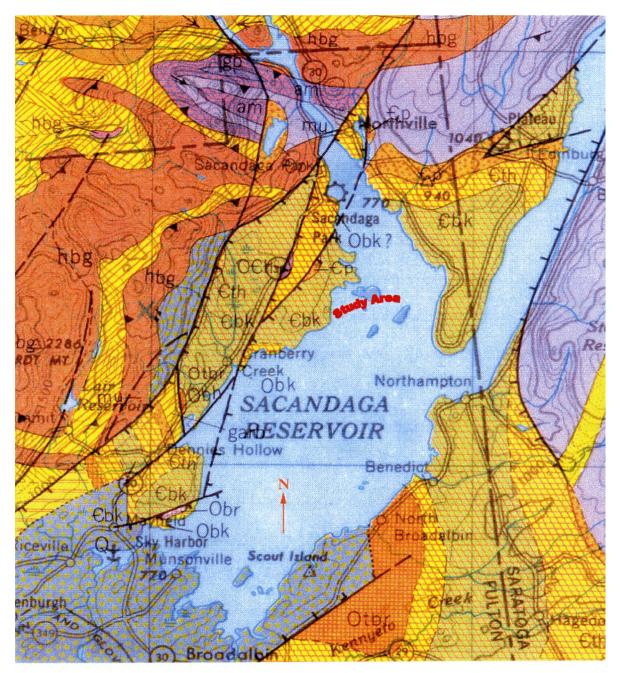


Figure 1.5 Bedrock Geologic Map of New York State, Hudson-Mohawk Sheet #15, New York State Museum and Science Service, 1995.

Cbk-Upper Cambrian Beekmantown Group, Tribes Hill Limestone.

Cp--Cambrian Potsdam Sandstone.

Cth-Cambrian Theresa Galway Formation Dolostone, Sandstone & Shale.

Obk-Ordovician Beekmantown Group Tribes Hill, Limestone, Dolostone.

Q--Quaternary-Glacial and Alluvial deposits.

The bedrock of the Sacandaga Basin range from Precambrian crystalline rocks on the east and west escarpments to Cambrian and Ordovician beneath the overburden of the down-dropped block, within which the Kenyon Island Group resides. Perhaps the best characterization of the bedrock comes from the water well logs obtained by the 1951 U.S. Geological Survey of Fulton County Water Resources (Arnow, 1951). The wells that exist within a 2 *mi* circumference suggest that bedrock is from 10-45′ beneath the overburden at the Kenyon Island site and no more than 30-120′ deep closer to the valley walls of the Sacandaga Basin. Since 1951, legislation has been passed such that water well drilling companies are not required to have bedrock tailings examined and only depth and Health Department requirements exist today. No outcrop exposures are available in the study site area.

The 1987 Surficial Geology Map of New York by Donald H. Cadwell indicates that the whole of the Kenyon Island Group is fluvial deltaic sand. Some glacial tills form boundaries existing above the fluvial deltaic sand and north of Northville where outwash sands and gravels exist (Figure 1.6). In 1911 The University of the State of New York, State Museum, Topographic Map, Bulletin #153, by Miller, describes the Kenyon Islands landscape as Quaternary glacial and postglacial sediments. In 1929 The University of the State of New York, State Museum, Glacial/Geographic Map, Bulletin #280, by Brigham, describes the sediments of the Kenyon Island Group and Osborne Bridge as water laid drift (Figure 1.7).

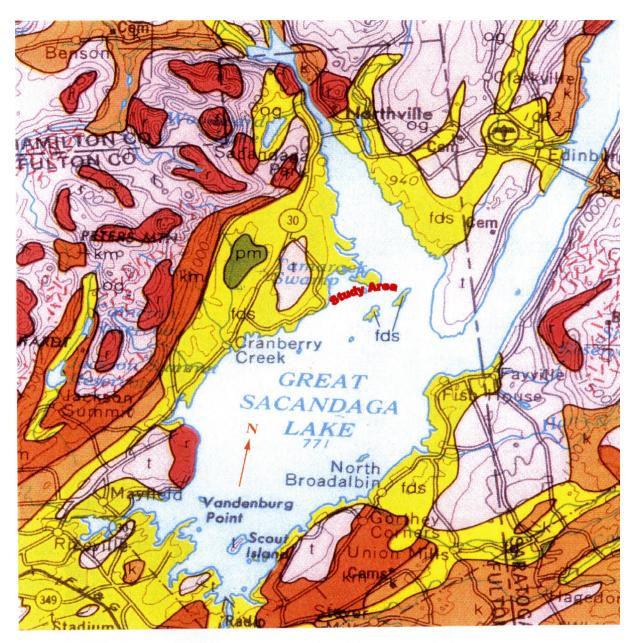


Figure 1.6 Surficial Geologic Map of New York State, Hudson-Mohawk Sheet #40, New York State Museum and Science Service, 1987.

t--till
fds--fluvial deltaic sand

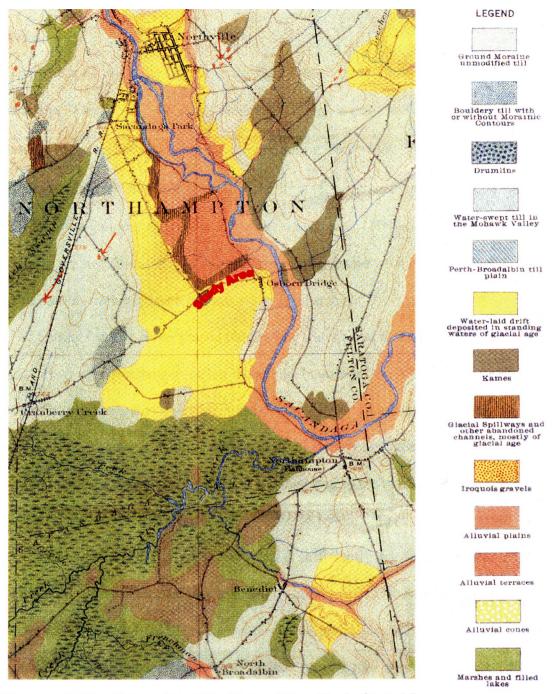


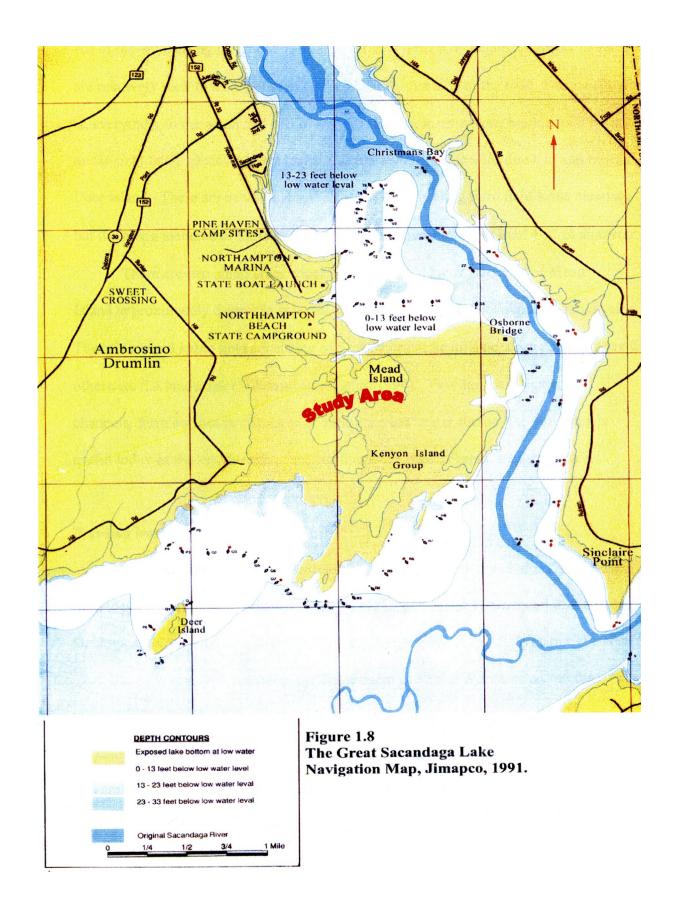
Figure 1.7 The University of the State of New York, State Museum, Glacial/Geographic Map, Bulletin No. 280, with Cropped legend, by Brigham, 1929.

Setting

The spillway at the Conklingville Dam, which was built in 1929 to create the Sacandaga Reservoir is 771' above sea level. At high water the Kenyon Island Group including Mead and Deer Islands have shorelines at about 768'. The lake level begins to drop in (months). Water is discharged from the reservoirat an average rate of 1 1/2" per day (~1 billion gallons) and by the low water season over 150 billion gallons of water have been drained from the reservoir. It takes approximately 6 months for the transformation from high water to low water level status.

During the low-water season the shallow water shoal that surrounds the Kenyon Island Group becomes dangerous to recreational boaters on the lake. The situation becomes so serious that almost nine miles of shallow warning buoys are strung around the greater shoal, warning boaters not to enter. This buoy line extends from Deer Island to beyond Northampton Marina. Boaters are relegated to the marked Sacandaga River Channel for much of the boating season in order to by-pass the shallows (Figure 1.8).

The exposed lake bottom at low water is sufficiently dry during the summer season to travel across the old farm roads of the old town of Osborne Bridge. Dozens of remnant foundations where the old farmhouses and barns once stood are visible. Except for the occasional sandbar from anomalous lake currents like the one at core site #8 and ice-rafted sediments from just west of core site #2 to beyond core site #1, there is almost no appreciable new sedimentation across the whole shoal since the Conklingville dam flooded the area in 1929.



All of the lands along the gravel road used for the NE to SW diagonal transect of cores, are relatively rock free for between 400' and 800' either side of the road. The exception is, everything to the south and east of Osborne Bridge in an arcuate boulder belt (rocks up to car size) that includes the two major Kenyon Islands, southward down to and beyond Deer Island. There are occasional small rock piles indicating farm field stone clearing, but the preponderance of these rocks are small boulders no bigger than 4-8" in diameter.

There are two small areas consistent with kettle lakes, one west of Mead Island approximately 6 acres in size and another to its southeast of about an acre in size. There is a small kame field consisting of 5 or 6 moundlike hills no more than 4' above the otherwise flat landscape (Lillesand and Kiefer, 1994). Two unusual relict stream channels, form backwater depressions during the low water season, the north one is useful today as the boat launch of the Northampton Beach State Campgrounds.

Previous work

The first surveying of the Broadalbin Quadrangle, which included the proposed Sacandaga Reservoir, was begun in 1899 by the United States Department of the Interior Geological Survey and completed in 1911, by the geographer, H. M. Wilson. The New York State Survey, represented by the Department of Public Works, surveyed the triangulations and W. H. Lovell surveyed the topography. Because construction of the Sacandaga Reservoir was anticipated, the proposed lake was stippled in and the shoreline was ambiguous. No attempt was made to depict the geology and neither the Noses fault nor the Batchellerville fault were delineated.

The first surveying and mapping of the Sacandaga terrain was undertaken by decree of the New York State Legislature in 1902. It was performed by John R. Freeman, C. E. and was finished in 1908. This study was undertaken for the purpose of creating a reservoir and only focused on the topography of the region and setting the proposed high water mark ("Flow Line") and the dam site at Conklingville, New York. The projects legacy is the substantial body of work known as the New York State Water Commission Maps of 1908.

Simultaneous with the mapping of the Broadalbin Quadrangle by the U. S. Department of the Interior Geological Survey, The University of the State of New York, State Museum, began the first geological survey of the Quadrangle. The work was done by W. H. Miller and completed by 1911 as Bulletin #153 (Miller, 1911). Miller also stippled in the proposed Sacandaga Reservoir layered over the still viable communities within the anticipated "Flow Line." Miller correctly identified the Noses fault to the west of the Sacandaga Basin but failed to identify the Batchellerville fault to the east. His description of the surficial geology was also very broad, calling it Quaternary glacial and postglacial sediments. Later on in his career when he had completed his geological survey of the Gloversville Quadrangle in 1922-23, Miller speculated that the Sacandaga River once drained to the Mohawk River rather than turning northward near Osborne Bridge to drain into the Hudson River, as it does today.

In 1929, The University of the State of New York, State Museum used the same Broadalbin Quadrangle and added the glacial and geographic features. The work was done by A. P. Brigham and presented in Bulletin #280 (Brigham, 1929). Brigham characterized the Kenyon Island Group and Osborne Bridge surficial geology as "water

laid drift" deposited in standing waters of glacial age. Brigham suggested that the preglacial Sacandaga River followed the present axis of the Sacandaga Reservoir and then flowed in a southwesterly direction, paralleling the Noses fault line, reaching the Mohawk River just west of Fonda, New York. Geophysical exploration in 1949 by the United States Geological Survey in cooperation with the State of New York, Department of Conservation, Water Power and Control Commission indicate the west edge of a buried escarpment and could possibly confirm Brigham's postulate (Arnow, 1951).

The United States Geological Survey in cooperation with the State of New York. Department of Conservation, Water Power and Control Commission work, done in 1951 by Theodore Arnow, called Bulletin #GW-24 (Arnow, 1951), is the most complete report on the geology of the region. The study began in 1946 as part of a statewide survey of the groundwater resources of the State of New York. It is called "The Groundwater Resources of Fulton County, New York". The study examined the records of 219 water wells in Fulton County, did chemical analysis on water samples, incorporated The University of the State of New York, State Museum geological work, and did the geophysical resistivity and seismic profiles of the overburden east of Gloversville, New York. It did not, however, examine the surficial geology of the County or the Sacandaga Basin, except in cooperation with local water well drilling companies whose well logs and tailings were used during the study period. It recognized in a generic way, Brigham, Miller, and others' work in identifying the Quaternary deposits that mantle the region, but limited its scope to the sources, quantity and quality of the groundwater of Fulton County (Arnow, 1951). This report, however, did correctly interpret the former work of Darton, (1897). Darton (1897) had suggested multiple faulting north of the Hoffman's Ferry

fault, east of Amsterdam. However, Darton did not identify the Batchellerville fault. It was later identified by the U. S. Geological Survey, Groundwater Branch workers who conducted the Groundwater Resources of Fulton County study (Arnow et al, 1951). This report and summarized by the work of Darton (1897), Miller (1911) and Wilson (1911), identifies the occurrence of the great sunken block between the Noses and the Batchellerville faults as being due to a regional uplift.

The Geologic Map of New York by The University of the State of New York,

State Museum, compiled by (Fisher et al, 1970) and the Surficial Geologic Map of New

York compiled by Caldwell and Dineen (1987) are both composites of previous work and
aerial photography and should be considered as good references only. Neither is
completely definitive nor intended to be.

The General Soils Report of Fulton County, New York (Davis et al,1971) was very important in arriving at the geologic perspectives this thesis espouses. The United States Department of Agriculture Soil Conservation Service prepared this report in cooperation with the Fulton County Soil and Water Conservation District. Most of the work was done using aerial photography and United States Geological Survey maps as working base maps. The report was prepared as a general planning guide, based on definitive soil characteristics, which affect the degree and type of limitations that can be expected for community development, recreational, or farmland use.

A soil association map was compiled with 50 soil-type associations across Fulton County. Each of the 50 soil characteristics were given soil classifications including their suitability for such things as agricultural use, campsites, roadbeds, athletic fields, and septic drainage fields, etc. (see Table 1.1, soil-type associations involving the study area).

ISTICS	Droinggo		Well & Moderate	Well	Moderate & Poorly	Poorly	Well & Moderate	ir Moderate
	Sound can	7	Gravel	Not Suitable	Not Suitable	Not Suitable	Not Suitable	Not Suitable
TARACTER		Septic Fields	Moderate	Severe Stoniness	Moderate, Wetness	Severe Wetness	Severe Stoniness	Moderate
GENERAL SOILS REPORT - SOIL ASSOCIATION AND CHARACTERISTICS	Soil Suitability and Limitations	Athletic Fields	Moderate and Stony	Severe Stoniness	Moderate and Stony	Severe & Wetness	Severe Stopiness	Moderate
		Streets, Roadbeds	Moderate	Severe Storiness	Moderate	Severe & Wetness	Severe Stoniness	Moderate
		Camp Sites	Slight	Severe Stoniness	Moderate	Severe & Wetness	Severe Stoniness	Moderate
		Field Crops	Slight	Severe Stoniness	Slight	Moderate	Severe Stoniness	Moderate
		Association	Charlton-Sutton Stony	Chariton, Stony and Extremely Stony	Sutton-Leicester Stony	Leicester Stony	Chariton-Sutton Extremely Stony	Sudbury
	13	qunN	8		12	14	12	32

Table 1.1 Soil Type Associations in Study Area from 1971 General Soils Report of Fulton County, New York.

Although the work is interpretive and little fieldwork was actually done (Davis et al, 1971), patterns do exist that are quite revealing (Figure 1.9). The work was done in the spring of 1971 during high water level status of the Great Sacandaga Reservoir. Doing the work in the spring distorted the Kenyon Island Group soil-type perspectives and Davis's field workers did not expend the time boating out to the islands. Thus, this report overlooked the forested terrain and the lands adjacent to the Kenyon Islands that form the focus of this thesis.

Finally, the Hudson River-Black River Regulating District (the new District created in 1959 by the New York State Legislature when it combined the Hudson River and Black River Regulating Districts) commissioned a study (Gabrielson, 1982). The purpose of the study was to determine the storage elevation curve of the reservoir to estimate how much sedimentation from stream tributaries and shoreline erosion was affecting the lake capacity. Their conclusion was that after 53 years of operations, the holding capacity of the reservoir had diminished by 1/3 of a billion cubic feet or about 2.6 billion gallons.

Format of this study

In this thesis, I try to present a logical argument for a deepening project in the Great Sacandaga Lake by looking at the subsurface soil stratigraphy using Geoprobe coring in Chapter 2. Core descriptions, access, permits, maps and GPS coordinates, methods and the questions raised are presented.

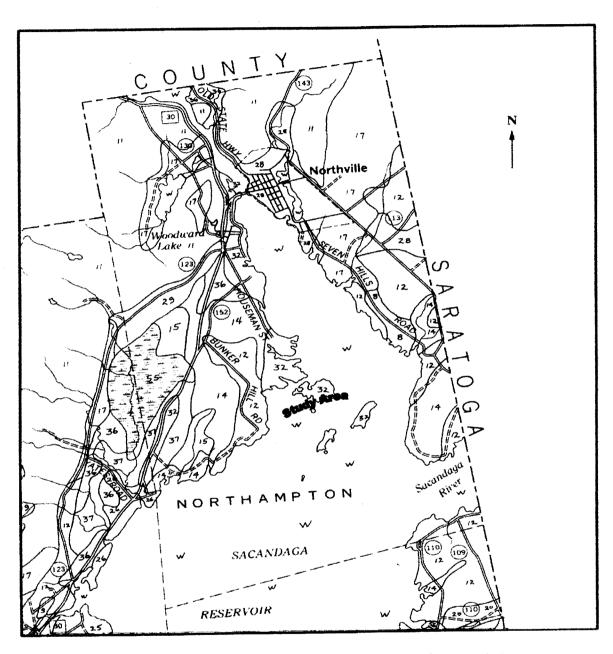


Figure 1.9 Soil Type Associations in Study Area from the General Soils Report of Fulton County, New York, Arnow et al, 1971.

In Chapter 3, a grain size analysis is conducted to try to assist in the interpretation of the cores. I have also examined the clay-sized fractions with X-Ray Diffraction and evaluate using *Ostracodes* to examine lake history. Idrisi software was used on Landsat 4, 5 imagery and isodata clustering during the late shallow water seasons of 1984 and 1988 to evaluate the bathymetry and land cover of the study area.

Some recent and historical accounts of the potential presence of heavy minerals in the sediments or bedrock around and in the reservoir are taken into account in Chapter 4. A trace metal analysis performed on different soil horizons in the geoprobe cores was used to determine which horizon minerals, if any, have value. A total heavy mineral mineralogy on stratigraphically averaged sediments from the cores was also performed. I also try to determine if the heavy minerals were entrained in glacial debris transported from many miles away, or rather, if they are beneath the overburden in the bedrock. The potential mineral resource of the heavy minerals will also be discussed.

In Chapter 5, I take a look at the issues involved in Kenyon Island reconfiguration. They include access and ownership of the lakebed sediments, the planning and geography of the reconfiguration possibilities, the costs of the undertakings, the real property values, and the timetables of the wisest approach to the deepening effort. In accordance with the new field and laboratory evidence, the potential economic impact a deepening project could have is presented and some possible future work and the conclusions are discussed.

CHAPTER 2

GEOPROBE CORING

Geoprobe Coring

Nine sediment cores were collected with a Geoprobe[™] in a transect across the Kenyon Island Group. The goal for each core site was to obtain continuous soil/sediment core samples, but this is sometimes not attainable especially if the cutting shoe hits gravel, cobbles or large rocks.

The Geoprobe used for this coring was a model #5400 operated by Environmental Maintenance Technology (Schenectady, New York). The nine Geoprobe core samples were collected along a NE to SW diagonal transect, across the shallow water shoal, which includes the Kenyon Island Group and Mead Island. Access was accomplished during the low-water season of 1998, by driving along the gravel roadbeds of the former village roads of Osborne Bridge, which is submerged during the high water season. Everywhere along the transect of core holes, the ground was generally flat, but the ability to adjust the plumb of the core holes so precisely, made the operations very smooth and effortless. Minimal disturbance to the ground is a great advantage and the probing rate is fast between sections (from 5' to 25' per minute depending on the soil/sediment type). Each probe rod section is 48" long with a drive head at the top and a cutting shoe on the bottom. The outside diameter of the stainless steel probe rod is 2.0125" and the outside diameter of the cutting shoe is 2.2" with an inside diameter of exactly 1.8125". A 44" CAB (Cellulose Acetate Butyrate) polymer liner, exactly 1 3/4" OD fits snuggly inside. The liners are approximately 1/32" thick and have a small flare on the down side to fit

over the cutting shoe on the inside of the probe casing. These liners are clear and offer a good view of newly pushed cores, at least until the soil/sediment moisture obscures the view. For clarity the 44" liners are assigned an added 4", weighted over the sediment horizons to account for the 4' penetrations of each probe section. The first 4' section is pushed and removed. Each successive core is pushed in 4' more with the aid of 1.25" solid steel probe extensions. Extensions are added and removed one at a time, so deep cores are more and more time consuming.

Environmental Maintenance Technology's Geoprobe was powerful enough to push aside small rocks, but this never became a significant problem in any of the core holes. Our purpose was to soil/sediment probe as many holes as possible into the subsurface of the shallow water shoal known as the Kenyon Island Group, in one workday. The coring was performed on November 13, 1998, and we recovered a total of nine cores (See figure 1.3).

Core descriptions

In cores #1 and #4 we encountered small rocks in gravel horizons, which were about 2" in diameter and these rocks blocked the recovery of complete cores until the cutting shoe rounded them enough to accept them (Figure 2.1). Figure 2.1 should be referred to whenever detailed descriptions of the core horizons are being discussed. None of the cores encountered any rocks of enough consequence to force us to relocate the truck for a second push. For a more complete discussion of the core horizons (See Appendix B – Core Descriptions).

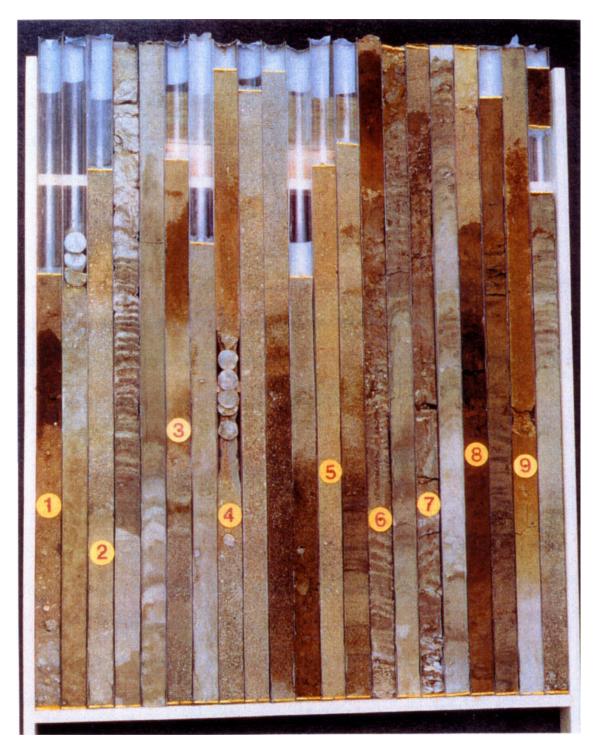


Figure 2.1 Geoprobe Core Horizons.

Core #1 was located just SSW of the Four Corners of old downtown Osborne Bridge. Its file number is 049 and its GPS (Global Positioning Satellite) coordinates are N (43°-10′-801″) and W (74°-08′-998″). The elevation was approximately 766′ above sea level. The site was also just north or on the edge of an end moraine, which was revealed further into the investigation of the study area (see page 53-55). We were probably on the former lands of either J. Paul or E. Hammond (see Figure 1.3).

Core #2 was located just NW of Osborne Bridge on the former lands of Foster Noyes (see Figure 1.3). It was located on the banks of the Sacandaga River Floodplain, which is pervasive north to the Northville Bridge, even at low water status. Its file number is 048 and its GPS coordinates are N (43°-10′-865″) and W (74°-09′-809″). The elevation was approximately 765′ above sea level.

Core #3 was located just E of Mead Island on the former lands of L. D. Fritcher (see Figure 1.3). It was at an elevation of about 765' above sea level. Its file number is 050 and its GPS coordinates are N (43°-10'-821") and W (74°-09'-791"). The site was formerly farmed and a few small stone piles existed adjacent to Mead Island.

Core #4 was located at the NW corner of Mead Island adjacent to a backwater inlet that faces the Sacandaga River floodplain to the east. Core #4's file number is 047 and its GPS coordinates are N (43°-10′-715″) and W (74°-10′-252″). Its elevation was precisely at the high water mark of 768′ above sea level. The site was on the former lands of Charles and Arietta Mead (see Figure 1.3).

Core #5 was located WSW of Mead Island on the edge of the largest eroded micro-island of a hummock field in the local greater Mead Islands group. The micro-island is usually submerged until late summer. The site was picked just on the eastern edge of the old Osborne Bridge Road and was 765' above sea level. Its file number is 046 and its GPS coordinates are N (43°-10'-527") and W (74°-10'-230"). It is unclear by the 1908 New York State Water Commission Maps who the former owner was but it could have been Horace Kenyon. In a straight line, the site was about 900' from core site #4 (see Figure 1.3).

Core #6 was located about 800' further SW along the old Osborne Bridge gravel road from site #5 (see Figure 1.3). This site is approximately 760' above sea level and it is unclear who the former owner was, but was likely to be Horace Kenyon. The site was on the East Side of the road and appears to be a former dirt road heading northwest to an anonymous unimproved road connected to the Bunker Hill road. Its file number is 045 and its GPS coordinates are N (43°-10'-431") and W (74°-10'-421").

Core #7 was located at a bend in the old Osborne Bridge Road, SW of core #6 by some 1200' (see Figure 1.3). It's not perfectly clear from the maps but this property could have been formerly the same lands of Ed Gifford. This site was also at the 760' elevation above sea level and on the West Side of the road. Its file number was 044 and its GPS coordinates were N (43°-10'-332") and W (74°-10'-487").

Core #8 was sited approximately 800' further SW along the old Osborne Bridge Road. It is on the East Side of the road at 764' above sea level and was probably the former lands of Ammon Blowers (see Figure 1.3). Its file number is 043 and its GPS coordinates are $N(43^{\circ}-10'-262'')$ and $W(74^{\circ}-10'-603'')$.

Core #9 is located at the high water mark (768') on the shoreline of the old Green's Beach Campsites. It would be just south of the old Osborne Bridge Road and on the former lands of Ammon Blowers (see Figure 1.3). Its file number is 042 and its GPS coordinates are N (43°-10′-182″) and W (74°-10′-798″). Core #9 was approximately 1200′ further SW along the old Osborne Bridge Road.

Core Analysis

After collecting the cores the ends were taped tight with duct tape to prevent the escape of sediments and moisture. All of the cores were labeled at retrieval with a file number and designated which ends were tops and bottoms. They were then transported to the Earth and Atmospheric Sciences Sediments Laboratory at the University at Albany.

Special racks were built to hold the cores and restrain them one by one for splitting each core liner lengthwise, proximally. The University machine shop built to my specifications a brass-cutting implement to which we attached two roofing knife blades. These were meant to slice the core liners only, at the centers 180° apart along the lengths. Hand handles were fabricated and 50# monofilament fishing line was attached to the handles and used to cut the sediments proximally through the length of the liners. A steel length weight was placed over the already split cores and a flat plastic shim was

pulled down through the center of the liners, separating the top and bottom liner and sediments. Once that was done, the top split could be inverted and laid cut side up. The University Graphics Photographer took pictures of the still moist core halves (see Figure 2.1). The remaining halves were covered with clear plastic strips, (44" x 1¾") and archived in the Sediment Lab at the University of Albany.

The actual core liner lengths, which are 44", were adjusted to reflect a 48" push Geoprobe depth. Each core was given a file number, which referred to the core site no matter what the depth of the pushed core was. The sediment horizons were designated by their status regarding depth. For example, a distinct horizon in core site #1, file number 049, at the bottom of the second section from 78" to 96" deep would hereafter be referred to as 0497896. An interesting horizon at core site #4, file number 047, at the bottom of the fourth section from 176" to 192" deep would hereafter be referred to as 047176192, etc.

Each distinct horizon, were split and weighed. Portions were used for analysis, of trace metals, particle grain size analysis, pH and conductivity. Read about the purpose and results of the particle grain size study in Chapter 3 and trace metals in Chapter 4. Additionally, plugs were taken at intervals in the varved banding of the fine sand and clay horizons in sites #6 and #7 for identification of the clay constituents and to evaluate the need for a more detailed study of the *Ostracode* fauna.

CHAPTER 3

PARTICLE GRAIN SIZE ANALYSIS

Purpose

The core descriptions in Chapter 2 were based on observation, texture, color and facies changes, which reflect changes in the depositional environment. Facies changes at the top 16" of core site #8 show the sandbar made by currents during the last 70 years.

The top 32" of modern (<70 years) ice rafted clasts and postglacial glaciofluvial gravels at core site #2 are Recent, due to the annual deltaic flooding at this part of the Sacandaga River as it opens-up to the newly created reservoir. Alignment of the core profiles contiguously by elevation with removal of all the blank spaces failed to allow clear correlation of sediment horizons from one site to another. Similarities can be seen between cores #6 and #7 or between cores #8 and #9, but no correlation exists between cores #5 and #6 and little exists between cores #7 and #8, even though the distances between probes was a mere 800' in each case.

The depositional environment given the nearby Sacandaga River and/or the observations of the authors of previous work in the study area ie; glaciofluvial (Brigham, 1929), fluvial (Miller, 1909) and floodplain or fluvial deltaic (Cadwell, 1987) did not explain the deposition of the study area sediments. If the deposition of the core section horizons were either glaciofluvial or fluvial, we would expect to see sediment types with several components each related to processes of different flow regimes, characteristic of a polymodal distribution. Historical accounts of the Sacandaga River, before the reservoir was created, indicate numerous floodings and drownings during lumber log

jams in the Sacandaga Basin. Even the 1908 Water Commission Maps indicate flood events at 766 feet above sea level, though the Sacandaga River elevation in the study area before the reservoir averaged 740 feet above sea level. River velocities, size and density of clasts and particle shapes would all be variable aspects of transported and entrained sediments, which should be expected to impart some degree of hydrodynamic sorting (Prothero and Schwab, 1996).

If the deposition were that of a floodplain, we would expect to see great variations over both small vertical and small horizontal distances, and the importance of the fine component, which characterizes the lowland suspended load type of rivers. Any fine suspended loads outside of the main channel of the river would have had to come about by the river overtopping its banks (Brown, 1985). Fine suspended loads would appear as layers in the sediment horizons and appear as abundances in core horizon splits.

Finally, if the deposition of the core sediment horizons were those of a lacustrine or lake environment, one would expect the preponderance of the sediments to be clays and silts, less than (0.0625mm/62.5μm) in size and showing sequences of laminae deposited in still water with annual layering reflecting lacustrine conditions (Loizeau et al., 1994). Important clues would be clay, terrestrial organic matter and freshwater crustaceans such as *Ostracodes*. The laminated layers in the clay horizons of cores #2, #6 and #7 looked to be good candidates for such conditions.

Over the last decade or so, several automatic rapid sediment analyzers based on light diffraction have been developed which are ideal for the measurement of the particle grain size of natural sediments, especially fine-grained sediments characteristic of fluvial/glaciofluvial and floodplain or lacustrine sediments. The geology department at

Union College in Schenectady, New York, has a Coulter LS-100 with a laser light source, from Coulter Electronics Ltd (USA) in their core analysis lab. Data from this insrument was used to characterize the particle grain sizes to assist in the interpretation of the depositional environments. Grain-size distributions would be presented in two ways: graphs of grain-size distributions will be plotted and compared; and statistical measures such as mean, median, mode and sorting are arithmetically calculated. In this chapter I will also identify and analyze the clay sized fraction with X-Ray Diffraction. Although an Ostracode study of the varved fine sand and clay layers of the core horizons at sites #2, #6 and #7 may not be appropriate for a chapter on grain-size, I did the work and hoped to learn whether (Brigham, 1929) was right to have postulated what he called a former "Glacial Lake Sacandaga" in the basin. The results of the Ostracode study will determine if a coherent pollen profile study needs to be undertaken. No bulk density or pore water analysis will be undertaken in the fine sand and clay-sized fraction to estimate the physical or chemical weathering effects on the sediments. All of the horizons examined either for particle grain size or economic worth (see Chapter 4) will be repeats of the most interesting and colorful horizons. References to a pH and Conductivity analysis, in this Chapter, on the core horizons interesting enough to further our overall goal of evaluating the potential economic worth of the sediments for a deepening project, will be discussed in Chapter 4.

Methods and Principles

Core horizons thought most interesting, as described in Chapter 2, page 48, were split and weighed, with approximately 30 dry grams used for particle grain size analysis

and pH and Conductivity measurements. Each 30-gram horizon split was dried in a low temperature oven at 50° C (maximum temperature, so as not to lose any volatiles that may have existed). Core splitting revealed that there were no root systems to be removed. Absence of root systems was not assumed to indicate organics were not present, but rather an indication that no aquatic plants or terrestrial plants existed at any of the core sites. Approximately 20 grams were split-out of the 30 gram charge and set aside for the pH and conductivity measurements. Of the roughly 10 grams left, enough of a portion from each core horizon was sieved-out (<2mm) and weighed to about 2 grams (g) and placed in a pan. In each core horizon exactly 0.50g of sieved sample was weighed to the nearest 0.01g and placed in a pre-labeled and weighed centrifuge tube. Labeling recorded the core horizon sample numbers and weights and was used for duplicate records. All of the remaining core horizon (<2mm) fractions were tested with weak sulfuric acid for secondary carbonates, but none whatsoever were detected. Therefore, no carbonate extractions were necessary.

Despite the lack of visible roots, the removal of organic matter is essential, especially because it binds mineral particles together; particularly clay sized particles, and hinders dispersion. For removal of organics, 3 milliliters (ml) of concentrated (30%) H_2O_2 (Hydrogen Peroxide) was added to each centrifuge tube and sample. Each tube and sample was placed under a fume hood and gloves and glasses were used with a calibrated pipette to add the H_2O_2 . Each sample was stirred with a glass rod and the rod was rinsed with deionized (DI) water (rinsed back into the centrifuge tube so that no material was lost). Organic matter is oxidized by H_2O_2 and an initial reaction or frothing indicates that organic matter is present. No organic matter was evident in any sample except for very

small amounts (as indicated by the extent of frothing) in a few of the surface sediment horizons. Still, to be absolutely sure that all of the organics would be destroyed, each centrifuge tube and sample was placed upright in beakers about half filled with water on a hotplate set to 50° C (none of the tubes were submerged more that 75% in H₂0). When the small amount of frothing in any of the tubes was complete enough to ensure that no further loss of fluid in the tubes was occurring, each tube was filled evenly (to about 1½" deep) with DI water and left overnight on the hotplate in the fume hood to thoroughly digest any remaining organics. The next morning the fluid levels in the tubes were topped off with DI water (to balance the tubes for centrifuging). Caps were placed on each set of centrifuge tubes and samples and they were centrifuged at 5000-7000 RPM for 3 minutes. Each tube supernatant was decanted carefully so that no sediment escaped. The centrifuge tubes were filled to ¼ level and the caps replaced. The tubes and samples were timed and vigorously hand shaken for 30 seconds and re-centrifuged for another 3 minutes.

The removal of biogenic silica is necessary for all lake sediment samples and surface soil developed horizons. Physiological actions of organisms, both because of the study area's history of farming and the lake microorganisms, effect the cycling of the chemical constituents of the core sediment minerals. To do this, 10ml of 1 Molar NaOH (Sodium Hydroxide) was added to the centrifuge tubes and samples. Each tube was left upright in beakers without caps in water and placed back on the hotplate, set at 50° C for at least 4 hours. The samples were not allowed to dry out or cool down during this period. The centrifuge tubes were capped, placed in racks and then in a reciprocating shaker for 30 minutes. After the required time, the centrifuge tubes were put back in the

water-filled beakers on the hotplate for another 30 minutes. Each tube was again put in groups for centrifuging at 5000-7000 RPM for another 3 minutes. Each tube supernatant was carefully decanted so as to avoid losing any sample material. Virtually no reaction occurred on any sample during the removal of the biogenic silica, so the tubes were filled evenly to ¾ full with DI water, timed and hand shaken vigorously for 30 seconds apiece and re-centrifuged one more time. Each sample was carefully decanted again, so that no sediment was lost and the tubes were filled to ¾ full.

A chemical dispersant made of non-ionic surfactant (0.3 Molar), sodium hexametaphosphate (5 Molar) and sodium carbonate (1 Molar) was added to each centrifuge tube and sample and stirred with a glass rod. The glass rods were carefully rinsed with DI water back into the tubes to prevent sample contamination and material loss. The samples were then capped and stood upright in racks and placed on the reciprocating shaker table for two hours, after which they were left overnight to await particle size analysis with the Coulter laser diffraction machine. The dispersant is supposed to accomplish three tasks: (a) adding sodium ions to increase the exchangeable sodium and cause a repulsion between particles, (b) adding hexametaphosphate which is adsorbed on to positive electrical charges on the sesquioxides and clays, so preventing attraction to negatively charged clay and (c) adding carbonate to raise the pH of the solution and so remove positive charges (Loizeau et al., 1994).

Using the Coulter Laser (LS) Particle Size Analyzer is essentially automatic.

Each centrifuge tube and sample is placed upright in a beaker of water, atop an ultrasonic vibrating machine for 5 minutes prior to run time. Vibrating ensures that dissolved air from stirring and shaking is degassed because bubbles do not interfere with the particle

size distributions. The filename and sample information is entered, the sample added after the water column is flushed clear with filtered water. After each sample the Coulter automatically saves the particle size distribution as a log-µm Cumulative Curve that can be grouped as advised (Figures 3.1a and 3.1b). The Volume Statistics generated can be seen in Appendix B. The exact same protocol, ultrasonic degassing, data entry, sample addition and chamber purging follow each sample run.

The principle of the Coulter Laser Grain size analyzer is based on the interaction between light and particles of sediment. Diffraction is the preponderant effect: particles of a given size scatter light through a given angle, the angle increasing with decreasing particle size. A laser beam of monochromatic light ($\lambda = 750$ nm) is collimated to form a 13 mm diameter beam that passes through the sample cell containing an upwards-moving suspension. After the samples are ultrasonically dispersed, the centrifuge tube and its sample contents are introduced into the fluid module that generates a high-speed suspension stream inside the sample cell. The stream is strong enough to move particles up to 900µm in diameter with a density of up to 4g cm⁻³. Size results are supplied in terms of equivalent spherical cross-sectional diameter and the Coulter resolves sizes into 72 classes from 0.4 to 900µm with an exact logarithmic progression. The Coulter Laser Particle Size Analyzer does not take into account particles calculated to be smaller than 0.4 µm, such as ultra fine clays and silts. The efficiency of their detection is proportional to the actual clay content. For ultra fine sediments containing less than 10% of clay, the efficiency is about 40-45% whereas for clay content greater than about 80% it reaches 100%.

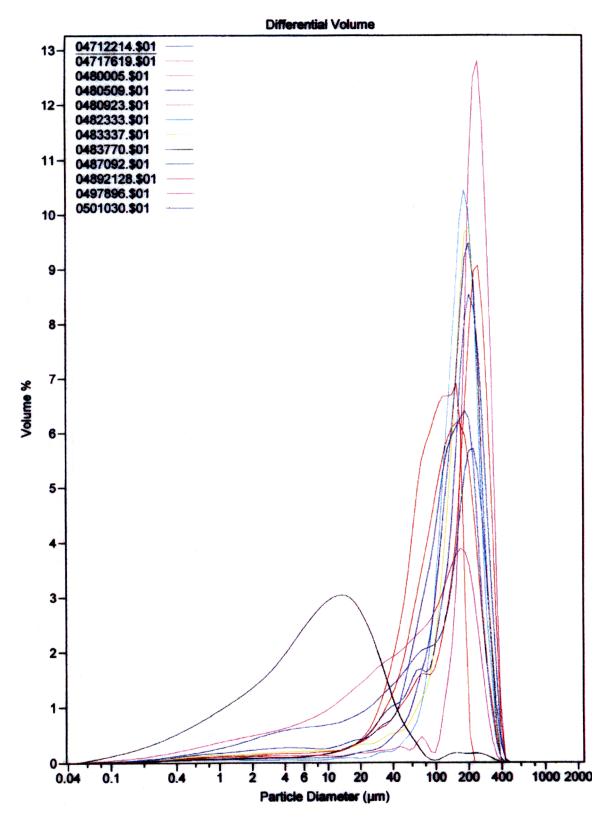


Figure 3.1a Coulter LS Particle Size Analyzer, Sediment/Core Lab, Geology Department, Union College.

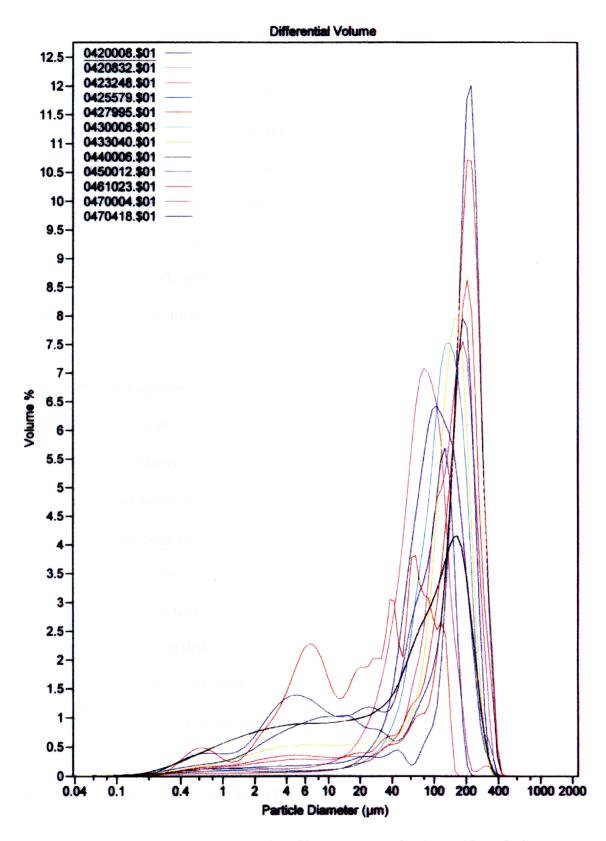


Figure 3.1b Coulter LS Particle Size Analyzer, Sediment/Core Lab, Geology Department, Union College.

The diffracted light is focused by two Fourier lenses, which direct light of identical diffraction angle toward one of three sets of detectors for low, medium and high angles (Figure 3.2 and Table 3.1). Particle location in the sample cell has no influence on the measurement. Consequently, a stream of particles can be passed through the beam to generate a stable diffraction pattern (Loizeau et al., 1994). From all accounts, the Coulter can be expected to measure the distribution of sediment sizes equally or better, and certainly much faster than can be measured by sieving. Laser diffraction computes simultaneously in the whole range of sizes, expressed in volume (or surface) units, whereas sieving results are obtained in weight units.

X-Ray Diffraction of clay sized fractions

Several of the probes showed laminated layer horizons, but apparently the laminated layers occurred mainly in sand fractions. The discovery of two cores, #6 and #7, with laminated layers of what appeared to be solid <50µm to depths of at least 8' deep and probably more, pose challenges to a potential Great Sacandaga Lake deepening project. Were these new laminated layers solid clay or a mixture of fine sand and clay and to what extent do these horizons exist?

If the clay species discovered were 100% pure and had no superfluous mineral constituents they would be a recoverable resource. Expandable clays such as smectite have a wide array of industrial applications such as catalytical activity and ion exchange. If the clays were Talc or kaolin type clays they would have applications in powders and cosmetics or as paper fillers and coatings.

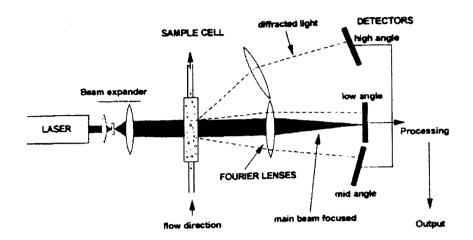


Figure 3.2 Schematic diagram showing the arrangement of the main Components of the Coulter LS-100, after Loizeau, 1994.

Table 3.1 Main characteristics of the laser diffraction grain size analyzer Coulter LS-100, after J-l Loizequ et al., 1994.

Laser type	solid, 4mW
Wave lenght of light	λ=750nm
Number of detectors	127
Size range	0.4-900 µm
Number of size classes	72
Fluid Module capacity	1-2 litres
Weight of sediment needed	0.05-2 g

If the clay has no intrinsic value, it still could have some very practical purpose, as a constituent in topsoil production for the reconstructed island or islands. Several of the smaller outlying Kenyon Islands and Mead Islands are barren with few, if any, trees. The larger islands have trees but most are stunted and are littered with dead and fallentrees. Good topsoil, depending on the source referenced, has 15-20% clay, 40-45% sand, 35-40% sesquioxides and 5-10% humus. The right proportions necessary to build a topsoil that does not need constant nutrient additions and will support trees and grasses would have to be determined by a soil scientist, but the shallow water shoal sediments (those sediments to be dredged) should have a sufficient contingent of each in the sediment mix. The ideal island that might be reconfigured in a deepening project will be discussed in Chapter 5, but would involve approximately 3-4 million cubic yards (yd³) of topsoil (12" deep). Clays of little intrinsic value (between 500,000-800,000 yd³) could be utilized in topsoil mixtures during island reconstruction.

The extent to which the fine sand and clay horizons exist could not possibly be determined without a much more extensive probe grid work. That may not be necessary though, if the X-Ray Diffraction (XRD) results reveal no potential economically valuable clay fraction and/or it contained superfluous contamination that required separation. If the clays cannot be determined to be valuable and dredging them is impractical, then earth-moving heavy equipment used to accomplish a deepening project would overcome the clay dilemma wherever necessary.

Clay plugs (1/2" marble size) were taken from the upper core sections of core #6 and #7 at 36" of depth. Core #6 includeds a light colored varve layer and core #7

includes a dark colored varve layer. The plugs were sent to Dr. Rossman F. Giese Jr. at the State University of New York at Buffalo, for X-ray diffraction analysis.

The clay fraction is mostly chlorite (>80%) and lepidolite (20%), with the balance having 12Å (angstrom) spacing, suspected of being <4µm common silicates feldspar and quartz. Fine sand and clays accounted for approximately 12% of the samples by weight. There were no expandable clays in the clay fractions. No further X-Ray diffractions were planned because both cores #6 and #7 were so consistent throughout and there were no apparent facies changes evident.

Therefore, I conclude that the <4µm fraction has no intrinsic economic value and are essentially all feldspar and quartz rendering them useful only as a constituent of topsoil and/or fill. One important result of the X-Ray Diffraction is that the clay is from the Chlorite Group and chlorite and lepidolite might help with the interpretation of the rock provenance of the sediments. This also means that the <4µm horizons would have to have been deposited as rock flour (pulverized while being transported). Rock flour is finely comminuted, chemically unweathered material, consisting of silt and clay sized angular particles of rock forming minerals. The extent to which silt and clay sized sediments exist here in the study area provides clues as to the sediment depositional environment and perhaps denudation rates, suggesting that the transport mechanism was a glacier. If the clay fraction is lacustrine, then either an *Ostracode* study or a coherent pollen profile might help with interpreting the environment of deposition.

Ostracode Study

Plugs were removed from cores #2, #6 and #7 at approximately 60", 36" and 60" deep respectively, on the premise that *ostracodes* could be preserved in the sediments. Most *Podocopina* are between 1-2mm, to ensure capture of juveniles or those of smaller stature, each plug was wet sieved at 80μm (0.07mm) and any sediments that passed through the sieve were discarded. The fraction that was saved, larger than 80μm, was washed back into a cup and a small amount of non-ionic surfactant was added. The sediments were mostly very fine sands after sieving out the smaller clay fraction, but did have clumps of clayed particles. The sediments were stirred and the surfactant aided in de-clumping the clayed particles. The cups were covered and left for several days, with occasional stirring. The sediments were dried in an oven at 50° C and examined under the microscope (x 100 magnification).

Absolutely no *Ostracodes* were discovered in any of the core horizons examined. Because of this, it was decided not to do any coherent pollen profiles, which might turn out to be of limited value in this context. The implications are telling and indicate that no surface water exposure to sunlight and photosynthesis, or eolian (wind blown) organics lent a part in living organisms becoming part of the fabric of the depositional environment of the >80µm (often refered to as slimes) core sediments. This could mean that no lake existed in the study area at anytime within the framework of the depths of our coring since the last glacial epoch. Or it could just mean that *Ostracodes* were not preserved or present in our limited analysis. One could argue that three sample plugs are not fully representative, but without *Ostracodes* it could be difficult to make a case that the subsurface sediments were viable to life and living organisms. Other interpretation

could be imagined but might be impractical given reasonable alluvial denudation rates in a valley such as the Sacandaga Basin. One would expect the sediments brought about by erosional or due to a fluvial nature would show some degree of rounding and these sediments do not. Therefore, my conclusion is that if a lake existed either now or since the last glacial epoch, life ought to have left a fingerprint, at least some organics would be detected at depth (see page 32 and 33 – methods) and that the total overburden is probably the result of a single glacial event in the late Pleistocene.

Interpretations

In the New York State Museum Bulletin No. 280, 1909, Brigham recognized what he called strong moraine, recessional moraine, outwash terraces and kames in the Sacandaga Valley, north of Northville, New York. He even points out that the Northville Quadrangle Topographic Map adequately shows these features. He was convinced though, that the Sacandaga Vly, represented a former glacial lake he called "Glacial Lake Sacandaga" (Brigham, 1909). Perhaps the basin, covered in trees, grasses and swamps, did not lend itself to easy field work and he was only able to interpret the Sacandaga River shorelines.

The aerial photographs used by the United States Department of Agriculture Soil Conservation Service in preparing the 1971 General Soils Report of Fulton County, New York, tell another story (Davis et al, 1971). Here, the mapping clearly suggest, the existance of remnant end moraine sections forming further arcuate bands south of Northville, New York, particularly at Hampton Point, Northampton Marina and across the Kenyon Island Group.

Landsat Thematic Mapper Imagery, obtained from the Eros Data Center in Sioux Falls, South Dakota also indicate the existance of morainal ramparts (morainal ramparts as described by Jackson, 1997) throughout the Great Sacandaga Lake. I used Thematic Mapper Spectral Bands 1: (0.45 to 0.52μm – blue), 4: (0.76 to 0.90μm – reflective infrared) and 7: (2.08 to 2.35μm – mid-infrared) on Idrisi 32 software to run an Unsupervised Land-Cover Examination with isodata clustering classification to study the land cover around the Great Sacandaga Lake (Jensen, 1996). I also used orthophoto's and the USGS Topographic map to improve the image classifications, obtaining close to 80% accuracy. The images used were obtained on 11/08/1984 and 09/16/1988, during exceptionally low water periods for the reservoir and were in raster format. The land cover classifications are very similar to the bathymetry of the GSL Navigation Maps.

Examination of the Coulter LS Particle Size Distribution Cumulative Curves in Figures 3.2a and 3.2b from 24 different sediment horizons shows that the mean of volume percentages is about $130\mu m$ (0.15mm). The consistency is remarkable given the possible range of depositional environments that could be imagined, such as fluvial, glaciofluvial, floodplain or lacustrine. In fact, the particle grain sizes that most reveal the parent material are somewhere between fine sand and very fine sand. The transect of core probing covered over 2 miles in length, diagonally across more than $5mi^2$ of the shallow water shoal study area. Except for horizon (0483770) in core #2, which was almost entirely fine sand with some clay, little variation existed. Glaciers, which erode and deposit sediment in a very distinctive manner may be the reason the deposition appears to be subglacial.

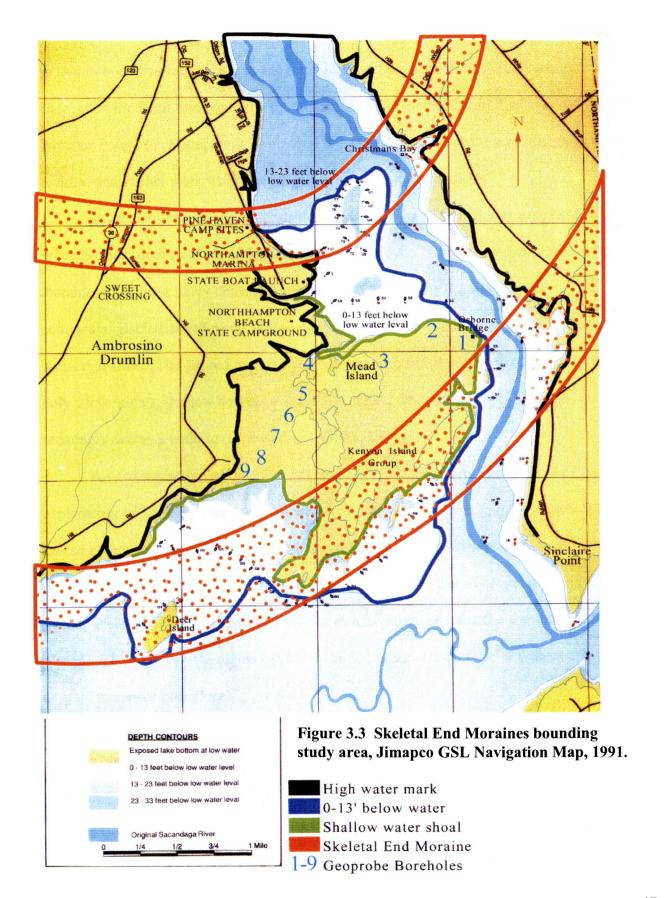
Glacial sediment is entrained in various ways and the mechanisms that transport and deposit clasts by ice make them notoriously poorly sorted, angular and heterogeneous. Clasts of a wide variety of rocks, derived from whatever lithologies lying upstream from the deposit, float in the predominantly <50 µm fine sand and clay "rock flour" matrix. The matrix is not the product of normal aqueous flow, but rather, deposited by a mode of mass transport in glacial ice. The sediment mix is generally characterized by a disrupted gravel framework, within which, stones are not generally in contact. The sediment mix is unstratified and notable for containing more matrix than gravel sized clasts (pebbles are less than 10%). All material incorporated on, within or near the base of glacial ice settles out slowly, varying with the seasons. The high viscosity and low velocity of moving ice permit no sorting and little abrasion, although grinding of the bedload against valley walls and floor can generate additional finegrained sediment. Glacial loads are plastered onto the substrate as ground moraine (lodgement till) or accumulate wherever there is a balance between rates of ice flow and melting. This accumulation produces end, recessional, and terminal moraines (Prothero and Schwab, 1996).

There must be supporting field evidence to identify ground moraine or deformed lodgement till including varved clay sequences (freeze and thaw cycles) and very large boulders forming lateral, end and/or terminal moraines (Prothero and Schwab, 1996). In fact, the shallow water shoal study area has all these features. The transect of cores lies almost entirely within the ground moraine field and its sediments are deformed around the flutes, kames and drumlins as stoss and lee topography. The kettle lake sediments created by the superincumbent glacial ice were deformed by the overlain lobe weight.

The study area may have been part of a sequence of outlet glacier (outlet glaciers as described by Jackson, 1997) advances called morainal ramparts. Each rampart looks as if it were an amphitheatre at the mouth of the valley, whose leading edge develops as an end moraine, its outside edges develop as lateral moraines and its central portion develops as ground moraine or deformed lodgement till (Prothero and Schwab, 1996).

Outlet glaciers, may explain the morainal rampart topography exhibited by the study area. The skeletal end moraine to the south and east which makes an arcuate path south of Osborne Bridge across the Kenyon Islands, Deer Island, past the Ambrosino Drumlin and Cranberry Creek and to the southwest of Tamarack Swamp forms the rampart boundary of the study area to the south. The skeletal end moraine to the north and west makes a smaller arcuate path from the Johnson Road, Christmans Bay, through Northampton Marina, underneath the Ambrosino Drumlin and to the north of the Tamarack Swamp makes up the northern boundary of the study area (Figure 3.5). The greater area ($\sim 12mi^2$) makes up the whole of the area confined by the two skeletal end moraines described (see USGS Topographic Map, Northville Quadrangle, 1970).

These are not all inclusive of the morainal ramparts produced by the Sacandaga Outlet Glacier during the late Pleistocene. At least one other skeletal end moraine can be described farther north of these two. This one occurs in an arcuate path across the valley from Gould Hill to just north of Hampton Point. Perhaps, based on topographic and navigation maps another existed just south of Northville across to the old Scout Island and to the north of the Mountain Road.



A terminal moraine most likely existed as the final rampart, south of the two skeletal end moraines bounding the study area and it ran in an arcuate band from Fish House and the Batchellerville Fault through to the southernmost end of the Mountain Road, just north of Paradise Point. No attempt will be made in this report to convince the reader of these realities and further work would be needed to confirm this hypothesis.

The glacial sequencing suggests that the Sacandaga Outlet Glacier may have preceded any continental glaciation and was overlain by it. The ground moraine and end moraines were most likely deformed and probably smeared by the weight of the continental glaciation and could explain the deformation and correlation issues in the core horizon profiles. The ground moraine is consistant with flat plains having local relief of only a few meters, shallow depressions with no natural drainage outlet and gullies seemingly draining nothing (Lillesand and Kiefer, 1994).

An outlet glacier and the morainal ramparts it produced seem the only plausible explanation of the study areas depositional environment and its topography. Microscopic investigations reveal essentially no rounding of grains and they remain sharp and angular throughout. For a good general explanation of the glaciation that affected the region in the late Pleistocene (see Brigham, 1929), which adequately explains the continental glaciation that covered the region including the Sacandaga Outlet Glacier with as much as 1000 meters of glacial ice.

CHAPTER 4

MINERAL WEALTH

Theories behind the study

Any deepening for this part of the Great Sacandaga Lake would involve dredging or earth moving by heavy equipment. Either of these approaches to permanently deepening the lakebed would be very expensive and careful consideration of many variables, especially in the case of dredging, would need to be evaluated. Earthmoving heavy equipment would be easier and less costly, but the only way to recover these costs would be through the new real property created, or Island rebuilding and reconfigured lakefront property values. The current "State of the Art" dredging machines can pump up to 500 cubic yards of material (solids up to 50% content and move objects as large as 8" in diameter) and are the method of choice in deepening projects, where "heavies" (mineral wealth) need to be recovered. Heavies minerals have a specific gravity greater (usually >2.85) than the surrounding sediments (a detrital mineral less than 1% in most sediments forming a minor constituent or accessory mineral). They typically include magnetite, ilmenite, zircon and noble metals and they may be recovered to offset the cost of operations. Simultaneous to extracting lake bottom sediments and separating heavies, the dredged slurries are pumped as much as 3000-4000' to be used to rebuild or reconfigure adjacent islands. Dredging machines are particularly ineffective in consolidated sediments such as clays. Dredging and pumping abilities are diminished and additional implements need to be added such as, auger heads with clay degraders or bucketwheel cutterheads to disaggregate the clay. Because clay sized particles remain suspended longer due to their shape and specific gravity, serious turbidity problems could

also be encountered if dredging was used. The topic of island rebuilding will be examined more fully in Chapter 5.

Bedrock studies and assays of the slope sediments in the feeder streams to the Great Sacandaga Lake have revealed bedrock units that contain interesting minerals (Miller, 1911). Glaciers in the Sacandaga Valley indicate that the deposition of the sediments of the study area must have come from an up-ice or upstream bedrock and overburden. There are known granite pegmatites (Miller, 1911) around the Sacandaga Valley and glacial erosion probably accounts for the many billions of tons of sediment atop the bedrock of the study area and adjacent morainal ramparts.

Dredging and separating the minerals of the study area is the only method that has the potential to generate profits beyond the altruism of the deepening cause. If the sediments have native elements, oxides, sulfides or even any of the pathfinder elements that could point to bedrock mineralization, I can then evaluate their economic importance. I have found mineral abundances and/or mineralizations using ICP-MS (Inductively Coupled Plasma Mass Spectroscopy), ICP-AES (Plasma Atomic Emission Spectroscopy) or standard "Fire Assay" procedures. The multi-element analysis using ICP-MS requires a tri-acid digestion procedure, which produces "near total", extended sensitivity digestion to selectively leach a much wider spectrum of elements. The Geoprobe core horizons that looked interesting and had the most colorful sediments, after cleaning-off small samples of their coatings with weak acid, were sent off for analysis. One hundred gram samples were split from the liners, marked and sent off, none of which contained any cleaned (no coatings washed-off) horizon sediments.

Multi-element analysis was performed on 26 Geoprobe samples and one nearby bedrock unit sample of the Potsdam Sandstone (from the Ridge Road by Carpenters Corners) as identified by nearby water well records FU 141, 0740900 NAD27 of the USGS, (see Appendix C). Facilities exist between the University at Albany and Rennselear Polytechnical Institute to analyze the samples, but it was decided that due to the potential economic importance of the sediments, more credibility would be attached to the results from an outside independent source. The Assaying and ICP-MS multi-element analyses were performed by ALS Chemex Labs, Inc., in Mississauga, Ontario, Canada and Reno, Nevada (see Appendix D). Forty-Four (44) elements were analyzed and the results are displayed as ppb (parts per billion), ppm (parts per million) or as % (percent).

Occurrences of minerals in much higher abundances and worthy of recovery are not commonly found in glacially derived sediments. If the Sacandaga Outlet Glacier derived its sediments from bedrock that has economically important minerals, we might have deposits of these ore minerals at our disposal. Mineral deposits of economic value in sedimentary deposits (i.e. placers) are those in which minerals are sufficiently high enough to make it economically worthwhile to extract. Because minerals are so widely distributed in many common rocks, the issue of resources and reserves is dominated by costs of recovery and selling price.

In sediments, if the elements are locked-up in mineral ores with specific gravities above 1.8-1.9 and if the particles are no smaller than 20 μ m, dredging recoveries are commonly greater than 95% after washplant operations (Hartman, 1996). The sediments in my study area have a mean size above 125 μ m (>0.15mm) as confirmed by the Coulter

Laser Grain-Size analysis. This point is central to our whole study because there is an advantage to dredging unconsolidated sediments and mining them as *placer deposits*. In fact, some mineral deposits in Quaternary sand bodies are currently of greater importance than deposits of any other type. For example, titanium oxides from shoreline deposits account for virtually all-current production and 36% of all identified resources of titanium in the United States and 55% for the entire world. The current economic mineral ore rutile (TiO₂) in placer deposits in Australia, have cutoff grades (0.1% TiO₂) below the average crustal abundances of TiO₂ (Force, 1991). Another advantage accorded to dredge mining is that it is a completely non-chemical extraction and environmentally friendly. Perhaps the greatest advantage though, is the low cost of equipment and operational costs, one-tenth of hard rock mining, making placer mining profitable even with very low-grade mineral ores (Holland and Petersen, 1995).

The trace metal analysis and assay abundances in my core horizon samples are listed in Appendix D. They represent only 44 of the more than 100 fundamental substances that consist of elements of only one kind. The absence of the remaining 66 elements, with a few exceptions, was chosen by design and meant to maximize the potential economic opportunities a specific gravity separation of the elements, such as dredge mining, would offer. Two elements that should be considered in further research include tin (Sn) and zirconium (Zr).

Heavy mineral mineralogy and assays of mineral fractions were used to identify those heavy minerals of economic importance. A thorough, microscopic study has been accorded the sediments of the study area and several elements pointed out in the trace element analysis have been specifically identified in their mineral ores. Several other

identification methods were employed in this study such as mineral associations; many minerals exist in sympathy with other minerals and sometimes only with other minerals associated with specific rock types (Chesterman, 1995). Magnetization and fluorescence studies were used to help identify or eliminate suspected mineralogy. A radiological study was used because of rare earth element enrichments discovered in the study area sediments (see Appendix E). Using a Landers, Frary and Clark Radiological Survey Meter (Geiger counter), made by the Victoreen Instrument Company of Cleveland, Ohio (air-ionization chamber) the sediments registered 0.11 R/hr (Roentgens per hour). A Roentgen is a measure of radiation intensity related to the ionization of air. The dose to soft tissue is about 10 milligray when exposure is 1 Roentgen.

Trace elements (those present in amounts less than 0.1% in general) play an important role in defining the nature of rocks and from which igneous partial melts they are derived. They are valuable indicators because the extent to which different minerals incorporate these elements varies sharply as a function of temperature, pressure or composition of the fluid from which the minerals crystallized. Of special interest are barium, rubidium, strontium, zirconium and the rare-earth elements. The incorporation of trace elements in igneous rocks also depends on the crystal structures of the minerals crystallizing from the melt. Minerals such as rubidium, strontium and the rare-earth elements (REE) tend to remain in the melt longer because their ionic radii and charges prevent them from incorporation into simpler structures (Raymond, 1995). Most of the sediments in the study area should be from igneous and metamorphic Grenville gneissic rocks of the Adirondacks. A correlation matrix was incorporated in an attempt to discover the rock provenances of these minerals.

The next important result of the multi-element analysis is the the role played in identifying or dismissing Dispersion Patterns; the patterns of distribution of chemical elements, especially trace elements, in the wall or surface rocks of an ore body or in the surface sediments surrounding it (Smee, 1983). The sediment in the study area lies atop bedrock that could contain mineralizations. A simple pH and ionic conductivity measurement on suitably prepared sediment horizon slurries may be a sufficient indicator of diffusion through the overburden. Any variations to background levels of elements in the sediment horizons over an ore zone would indicate an anomaly (Smee, 1983).

Measurements of soil pH and conductivity on study area soil solution slurries were made following methods of (Rowell, 1994). Trace element dispersion pattern anomalies, did not occur in our sediment horizon samples (see Table 4.1).

pH and Conductivity Measurements Great Sacandaga Geoprobe Boreholes

Borehole	Weight (grams)	рН*	Conductivity†
0427995	11.5	6.25	.220
0433040	9.5	6.40	.275
0480005	12.5	6.20	.275
0470418	12.0	6.25	.275
0501030	12.0	6.35	.250
04892128	14.5	6.60	.165
0491832	11.5	6.00	.475
0423248	9.5	6.55	.160
0480509	10.0	6.05	.245
0497896	15.5	6.15 + [6.10]	.235+[.005]
0425579	10.0	6.10 + [6.20]	.135+[.015]
0482333	9.5	6.10	.140
0483770	22.0	6.40 + [6.35]	.180+[.050]
0480923	13.5	6.00	.210
0450012	12.0	6.00	.270
0420832	9.5	6.10	.180
0430006	8.0	5.95 + [5.85]	.195+[.040]
0420008	10.5	5.90	.155
0461023	10.5	5.90	.180
0440006	7.0	5.45	.430
047122140	12.5	6.10	.110
0487092	12.5	6.30	.150
0483337	11.5	6.10	.230
0470004	11.5	6.00	.190
0467896	13.5	6.10	.160

Table 4.1

^{*} Hanna Instruments pH Meter (HI 9218) with Stability Indicator † Solinst Conductivity + Temperature (°C) [measured in (μ S x 1000) @ ~ 16.8 °C] + variations with agitation

Inductively Coupled Plasma (ICP) Spectroscopy developed in the early 1980's offers a way to carry out simultaneous analysis of large suites of elements including pathfinder elements associated with noble metals. Plasma Mass Spectroscopy (ICP-MS) offers the added advantage of introducing diverse selective extraction procedures for the detection of trace element suites. The object is to isolate and quantify that portion of the element in a sample horizon, which is relatively loosely bound, indicating prior mobility, either from depth (if mineralizations exist in the bedrock below the overburden) or if they were transported from short distances away, as is the case of our study area.

The secondary weathering environment is of fundamental importance and should be examined, because enough time has passed since the last glacial epoch in the study area. Secondary iron and manganese oxides occurring as coatings, concretions or as discrete particles in soils like those of the study area are significant because of their strong scavenging of important ore metals. Fe and Mn are pathfinder elements and are particularly important as scavengers of silver (Ag) and cobalt (Co). Although most sediment contains much greater amounts of Fe oxides than Mn oxides, the later possess greater sorption capability for trace elements. Such reactivity is due to characteristics of Mn such as: (1) it exists in several oxidation states (I, III, IV); (2) it forms non-stoichiometric oxides with different valencies; (3) its higher valence oxides exist in several crystalline (e.g. birnessite, lithiophorite, pyrolusite) or pseudocrystalline (manganite) forms; and (4) it forms co-precipitates and solid solutions with Fe oxides, owing to their similar chemical properties (Chao and Theobald, 1976).

There are various mechanisms of binding labile elements (water-extractable, organic and oxide-bound) in the secondary environment, including physical and chemical

sorption, precipitation, chelation and complexation. Phases present in the study area sediments, which are likely to scavenge free elements other than amorphous Mn and Fe oxides, are humic and fulvic acids of soil horizons. The characteristics of the sediment and soil profile depend not only on the parent rock material but also on climate. topography and biological activity, all issues addressed previously in this study. I know from this study that some of the Geoprobe cores had appreciable quantities of fine sands and chlorite clay. Adsorption of metals by soil constituents can be described as specific or non-specific. Specific adsorption implies exchange of cations and anions forming covalent bonds within ionic lattices. Metals most able to form hydroxyl (OH) complexes such as hydrous oxides of Fe, Mn and Al are the major constituents that are responsible for this type of adsorption. A soil's cation exchange capacity (CEC) is a measure of its ability to provide sites for non-specific adsorption. The forces involved in the adsorption of ionic species on charged surfaces are electrostatic and governed by Coulumb's Law of Attraction and Repulsion (Press and Siever, 1986). CEC's of most soils range from a few to several hundred meq/100g and far exceeds anion exchange capacity as negative rather than positive charges dominate colloidal surfaces. The distribution of an element between soil solution and solid phases at equilibrium is governed by: (1) the density of surface binding sites for each component (e.g. clay mineral, hydrous oxides); (2) the binding intensity of the metal ion to each component; (3) the abundance of each component; (4) the chemical characteristics (e.g. pH, ligands) and (5) the concentration of other ions, major or trace, competing for binding sites (Hall, 1998). ALS Chemex's quantitative determinations such as by Fire Assay and Extended Sensitivity acid digestion leaches used in ICP-MS are designed to be specific to elemental concentrations in my

study area and to exploit elements from mineral ores that have economic potential, in today's markets.

Noble metal mineralizations, besides the obvious and most important native states of these elements, are associated with pathfinder elements such as tellurium (Te), antimony (Sb), arsenic (As), bismuth (Bi) and more rarely mercury (Hg). The trace multi-element analyses were designed to evaluate their potential noble metal complexations in our sediment horizons and also to evaluate the mineral associations of noble metals with lead (Pb), copper (Cu), chromium (Cr) and nickel (Ni). There are almost 75 platinum-group minerals, often derived from deep-seated slow cooling and more fluid ferromagnesian rocks such as gabbro and peridotite, both igneous-plutonic rocks. Because platinum-group minerals include palladium, iridium, osmium and rhodium, they have a wide range of specific gravities; much of the world supply of these metals is derived from clastic placer deposits. If platinum-group metals are found, we would expect to find enrichments in nickel, iron and chromium, which would be indicative of there close association and environment together as elements that crystallize out of a melt first (Cabri, 1976). Gold, because of its general unreactivity, occurs in nature in only about 10 mineral ores, principally with tellurides and pyrites or arsenopyrites, but sometimes with antimony and bismuth (Henley, 1975). Silver is complexed with about 20 mineral ores, mostly sulfides, arsenides, tellurides and sulfosalts. Finding silver would be more important as a guide to mineral associations than for its economic potential.

Rare-earth elements (REE) were known at the turn of the century as monazite sands, but these sands are principally sources of thorium (Th) and uranium (U) for use in

mantles in gas lamps and electrodes in arc lamps. The sands are found as littoral deposits and placer mined extensively on the seacoast of Brazil (Hedrick, 1995). The specific gravity of monazite is between 4.9-5.3, its color is yellow and its luster is resinous. The mineral crystals are resistant to weathering, hardness about 5 ½, and are derived from parent rocks of granite pegmatite, and to a lesser extent in carbonatites and gneiss. Monazite possesses slight and low radioactivity due to its Th and U content, which is strong enough to affect undeveloped film and can be measured by Geiger counters (Kithil, 1915). When first discovered, REE were known only in their oxide forms and because they resembled the oxides of the alkaline earths (CaO, BaO, etc.) and did not seem to form common minerals, they were labeled rare earths. REE really do not deserve their name because they are much more abundant than many other geochemically scarce metals and today have wide uses in many diverse applications. The REE are produced mainly from two minerals today, monazite (CeYPO₄) and bastnaesite (CeFCO₃). Although the formulas of both minerals are written for cerium compounds, all of the REE substitute for cerium in their structures by atomic substitution (Boyle, 1982). There are 15 REE (see Table 4.2), starting with lanthanum and ending with lutetium. They differ only in the number of electrons in their inner electron shells and have very similar chemical properties. I designed our trace element analyses to look for cerium (Ce), lanthanum (La), thorium (Th), uranium (U), phosphorous (P-because of phosphates PO₄) and yttrium (Y-because it behaves like REE) because I was interested in heavy mineral sands of current economic significance, with specific gravities heavy enough to dredge.

Most geochemical evaluations are considered on a population basis. That is, the larger population, like the one we designed for this study, will include multiple sub-

populations. Sub-populations include for example, the noble metals, Au, Pt and palladium (Pd) or those elements that noble metals are complexed with (pathfinders) such as As, Sb, Mn, Fe, Te, etc., the REE oxides and all of the elements that are associated with individual rock environments

Table 4.2

The rare-earth elements. The relative amount of REE produced is in proportion to its geochemical abundance in the crust. (From Craig, 1988)

Name	Chemical Symbol	Atomic Number	Geochemical Abundance (% by weight)
Yttrium*	Y	39	0.0035
Lanthanum	La	57	0.005
Cerium	Ce	58	0.0083
Praseodymium	Pr	59	0.013
Neodymium	Nd	60	0.0044
Promethium	Pm	61	0.0000
Samarium	Sm	62	0.00077
Europium	Eu	63	0.00022
Gadolinium	Gd	64	0.00063
Terbium	Tb	65	0.0001
Dysprosium	Dy	66	0.00085
Holmium	Но	67	0.00016
Erbium	Er	68	0.00036
Thulium	Tm	69	0.000052
Ytterbium	Yb	70	0.00034
Lutetium	Lu	71	0.00008

^{*}Yttrium is commonly classed with the REE because its properties are so similar.

Results and discission

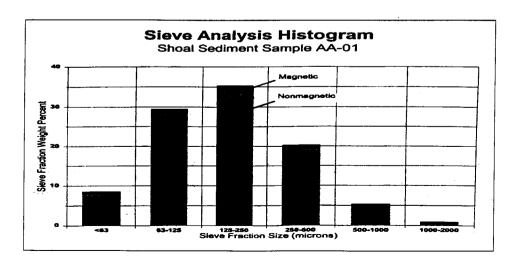
The geochemical results from the Great Sacandaga Lake study area are intriguing and promising. The fact that the sediments are related to glaciation and were locally derived means that the parent rocks deserve some further investigation. With that in mind, a hard rock sample from an adjacent outcrop, thought to be underneath the sediment overburden of the study area (USGS water well logs), was sent along with the Geoprobe core horizon samples to be analyzed by ALS Chemex. The results are listed separately in Appendix D and are not too different from the unconsolidated sediments of our study area. The rock is of course, much more ancient and metamorphosed and those elements that are more mobile like the alkalis Ca and Mg are slightly diminished. Those elements that are more labile like K and Na are slightly enriched and those elements that are more resistant or locked-up in minerals that are durable like those of the REE are about the same. Both the sediment samples of the study area and the hard rock sample would be considered weathered. Throughout this study and the results, the precise geochemical abundance backgrounds for weathered rocks in the literature is difficult to verify, but this hard rock specimen is not sufficiently enriched in any element that approaches ore grade. By incorporating otherwise "un-interesting" sample populations, though, we can get a glimpse outside of any bias the study area sediments may have introduced through secondary mobility of the elements.

The trace multi-element procedures designed for this study have given us a fair chemical analysis, but were only designed to yield the 44 elements (as in Appendix D) that would give us the strongest signals. The procedures avoid the conflicts of weaker signals, which arise using near-total tri-acid digestion. This also means that some of the

more resistant minerals and their abundances such as rutile (Ti) and REE like monazite (Ce, La, Th, and Y) may not be fully reported in this study. In further investigations, fusion and X-ray Fluorescence spectroscopy detection can be counted on for total element detection and will likely reveal somewhat higher element abundances in our sediments. Zircon (Zr) and cassiterite (Sn), also resistant minerals, are identified or indicated in the sediments and should also be determined by lithium metaborate fusion/ICP-MS (absolute total digestion) in further investigations.

Black sands are ubiquitous in my study area sediments and every horizon sample sent-off for analysis included these sands. Rudimentary magnetization tests on these sediments reveal about 10-12% of the black sands are magnetic and less than 5% of the mineral aggregates are magnetic. Sieve analysis performed by myself and by Overburden Drilling Management Limited of Nepean, Ontario, Canada (the principal mineralogy company used for this project) are nearly identical. A histogram showing weight percent plotted against size fractions reveals the extent of the heavy mineral population of the study area sediments (see Table 4.3, below).

Table 4.3



A more thorough radiological survey was done at the Radiation Safety Lab at the University of California at Los Angeles and listed in Appendix G (including a decay scheme). The element isotopes of europium (152Eu, 155Eu), Lead (210Pb), protactinium (234Pa) and thorium (234Th) were slightly elevated. Other lanthanide and actinide elements are also indicated. It was decided not to pursue a grams activity per sample analysis and leach test at this time because the gamma radiation appears to be no problem and the beta radiation has only slightly higher energy. No determination, without a leach test analysis will be speculated, regarding the danger of drinking from local water wells in the Sacandaga Basin. Investigations into dry beach sand recoveries of heavy mineral sands in Australia, containing monazite sands, indicate that air-borne particles have to be carefully damped to keep employees safe from radiological dangers. No such dangers are encountered by wet gravimetric dredge mining recoveries (Uranium Information Center, Ltd., Australia), such as the type we espouse in this report.

Based on a 20' averaged deepening of the study area total dredged sediment amounts to 92 million yd³ with all the calculations based on 3.0x10¹¹¹ lbs of sediments (92 million yd³ X 3000 lbs/yd³). The sediments actually weighed over 3300 lbs/yd³, but everywhere, quantities and potential worth were estimated downward and costs were estimated upward. If I was unsure of an element's mineralogy, occurrence or complexation, I opted to take the conservative approach and declare no value for purposes of this study. If we can identify mineral sand ore such as monazite and are unsure of its element mineralogy (REE all substitute for cerium in their structures by atomic substitution), then they were merged into one category because all of the minerals have high specific gravity, are considered quite valuable and are all easily seperated

(Craig, 1988). Some elements of high abundance that exist as constituents of minerals were classed too light, based on their known occurrences, and others whose occurrences are not known but expected to be in minerals too light to be easily seperated, were given a \$0 ore grade value until further work is done. Most elements examined, even though their ore minerals have potential economic value, are at or below average crustal abundances. Only those minerals that are clearly identified, to date, and exist in quantities appreciable enough to dredge successfully were assigned an ore grade worth. This does not mean that other dredgeable minerals are not recoverable and worth recovering.

The coatings on the sediments are identified as of this writing from the mineralogy as goethite (hydrogen iron oxide), with substantial amounts of manganese and iron. The coatings are heaviest in the heavy sand and gravely horizons and may well be due to the freer flow of groundwater and the secondary-weathering environment.

There are no dispersion pattern anomalies in any of the sediment horizons and no enrichments of pathfinder elements either. In fact, the relative abundances of noble metals are equal to the average crustal abundances for these elements and no indication of pathfinder elements would be expected in this scenario (Tilling et al, 1973). This means that the overburden does not lie atop any noble metal mineralization and that no nearby up-ice or upstream noble metal bedrock mineralization is likely to be discovered either. This would also hold true for base metal mineralizations.

The pH and conductivity readings are as expected, uneventful as well. A more thorough, much narrower (testing every 100'-rather than every 800') gridwork of tests would need to be conducted if any anomalies were detected, but that does not seem to be

indicated in these sediments. Ionic migration is likely to electrochemically transport through as much as 100-150' of glacial sediment overburden, since the last glacial epoch, especially in areas of high groundwater content and/or clay sediment (Smee, 1983) and time and thickness should have had no effect on halos or dispersion patterns in our study area sediments.

In the secondary environment, anomalous concentrations of metals, derived from some mineralized source are taken into solution in the generally reducing and slightly acidic environment below the water table. Metal cations are incorporated into the interlayer structure of the phyllosilicates (clay minerals) by ion exchange. Phyllosilicates that have a large open structure and can offer sites to metal cations, even those with 8 and 12 fold coordination such as Cs and Sr with large radius ratios can become enriched. These clays often take on colorations (plaqueing) different than in their normal structures. with increased mass (Bradshaw et al, 1974). This does not appear to have happened to our study area clays and no different or unusual element enrichments or colorations occur in any horizons with elevated clay mineralogy. Cesium (Cs), rubidium (Rb) and lithium (Li) are often associated with lepidolite ores, which were identified in the sediment horizons. Lepidolite is a phyllosilicate mineral with a specific gravity of about 2.8-2.9. Some washplants have been designed with overflow filters to collect phyllosilicates plaqued with noble metals, but for this writing, we chose to ignore this and assigned no ore grade value to them.

The trace multi-element proceedures used on the study area sediments only pointed to element enrichments, but without total heavy mineral (THM) mineralogy done on the separable heavies of the sediments of the study area, we cannot assign any mineral

value to the results. Two companies were chosen for their special reputations in the field: Overburden Drilling Management Limited (ODM) of Nepean, Ontario, Canada as mineralogy experts and QIT-Fer et Titane (QIT) of Tracy and Sorel, Ontario, Canada as industry experts. At the University I attended to heavy mineral (heavies) separations in the same manner as these two companies, to isolate the heavies from the bulk sediment. do the sieve analysis for populations size-fractions of heavies and to identify and verify microscopically as best I could, the heavy mineral fractions that they might attain. This first step is accomplished by weighing from the archived sediment cores, collected using Geoprobe, stratigraphically averaged sediments from all the lithologies. I shipped just over 8.3 kilograms to ODM and used 1.0 kilogram for my own separations. Three 9-kilogram bags were shipped to OIT from sites around the Great Sacandaga Lake. including one from the study area, for bulk processing from an industrial perspective. Methylene iodide with a specific gravity greater than 3.2 (after re-use and dilution with acetone) was used by ODM and myself to separate the "heavies". QIT used electrostatic extraction, which separates conductors (oxides) from non-conductors (garnet, pyroxene and amphibole) followed by a magnetic fractionation at different field intensities to separate the various iron and titanium oxides, such as ilmenite, titano-magnetite and rutile, etc.

The results of these studies are all remarkably consistent, withnoted differences (Appendix D). the descriptions can be seen in. Each of the three techniques measured between 8.5% and 12.5% (average 10.5%) THM in the study area sediments. My results and ODM both determined that the THM fraction resides essentially between $45\mu m$ and $1000\mu m$, with, 1-2% of the THM residing above or below these fractions. The mean

(193µm) totals approximately 35% of the THM. The estimates of the heavy mineral populations by species of minerals from the ODM mineralogy are listed in Table 4.4, below.

Table 4.4

MINERALOGY

PERCENT	MINERAL	FORMULA
3	Augite	Ca, Na (Mg, Fe, Al) (Al, Si) ₂ O ₆
3	Diopside	Ca Mg Si ₂ O ₆
32	Horneblende	$(Ca, Na, K)_{2-3} (Mg, Fe^{2+}, Fe^{3+}, Al_5) (Si Al)_8$
3	Hyperthene	(Mg, Fe) Si O ₃
2	Zircon	Zr Si O ₄
2	Titanite	Ca Ti Si O₄
7	Magnetite	Fe ₃ O ₄
7	Ilmenite	Fe Ti O ₃
35	Almandine	Fe ₃ ²⁺ Al ₂ Si ₃ O ₁₂
3	Apatite	Ca ₅ (PO ₄) ₃ (F, Cl, OH, REE)
3	Luecoxene	Ti O ₂

ODM and QIT used Scanning Electron Microscopy (SEM) and 200 grain-count methods to evaluate the THM. In addition to the above-mentioned minerals, QIT identified rutile (TiO₂ at <0.4%), monazite ((Ce, La, Y, Th)PO₄ at <0.3%) and zircon (ZrSiO₄ at ~1.0%). Potentially other minerals could be identified with continued grain-counts, but for practical purposes and given the reputations of the two companies, the results are substantive.

The most abundant mineral of economic significance is the garnet almandine, at nearly 35% of the THM; it represents more than 3.5% of the total sediments in the study area. Almandine is particularly hard (~7.5 Mohs scale) and brittle, with a specific gravity ranging from 3.56-4.32. Almandine generally occurs in Diorite of plutonic rocks (Chesterman, 1995). Examinations of the THM fractions by all the methods discussed above indicate sharp angular grains with essentially no rounding, consistent with glacial

sediments. Since the THM's are from averaged lithologies, the inference clearly indicates the transport mechanism is by entrainment in ice and not by fluvial deposition. This is a very positive circumstance for almandine because its principal use today is as an industrial abrasive. Additionally, the almandine exists in perfect size fractions and requires no crushing to be used as an industrial abrasive. In an averaged 20' deepening of the study area, 8.7 billion pounds of almandine would be recovered at 90% washplant efficiencies. No ore-grade values could be found in Standard and Poor's or Platt's Metal week, but USGS experts claim \$0.09-\$0.12 per pound Wyoming ore-grade almandine. Phone conversations with industrial mineral officials indicate \$0.05-\$0.06 per pound. Using \$0.09 per pound the mineral almandine is worth \$782,000,000 in the study area. Production costs are listed in Chapter 5.

Another of the most abundant and recoverable heavy mineral sands from the study area sediments are the oxides of titanium, principally ilmenite. Titanium minerals account for between 6.5% and as much as 20.8% according to ODM and QIT, respectively, of the THM population (expressed as TiO₂). This represents 0.69% of the total sediments in the study area. A separate assay of the THM's, done again by ALS Chemex, verified this percentage (see Appendix F). The economically important Ti minerals in the study area THM's include; Ilmenite (average 48% TiO₂), titano-magnetite (FeO mineral with up to 15% TiO₂), altered ilmenite (luecoxene, >55% TiO₂), and rutile (>85% TiO₂). Ilmenite deposits of magmatic origin (the anorthosite-ferrodiorite type) are currently the most important types of TiO₂ mineral deposits. Ilmenite from magmatic ferrodioritic sources are known to exist just 30-50 miles north of the study area in Sanford Lake and Tahawus, N.Y. (Force, 1991). In an attempt to identify heavy mineral

provenance, the inter-element relationships were examined using a correlation matrix of elements normalized as a percent of total (see Table 4.5). Trace elements such as the noble metal assay results were eliminated. Several elements revealed significant correlations, with rare earth's (likely toexist together) and mafic elements (Ti and Fe), showing significance among themselves. Significance usually means that two elements when compared to each other have coefficients above ~.60. The correlation matrix I generated was too large to fit in this document so I recoded the elements relevant to the THM's and while a fair fit to the regression line emerged, I was disappointed with the histogram and regression formula generated. The recoded correlation matrix (R² -values) are presented below in Table 4.5.

Table 4.5

Correlatiom Matrix of Selected Element Geochemistry

	Al	Ca	Fe	Ti	Cr	La	Се	Y	U	Th
Al	1.000	.218	205	309	731	.494	.555	.018	.454	.319
Ca	.218	1.000	.184	.297	049	185	053	.027	103	234
Fe	205	.184	1.000	.959	.231	.300	.370	.823	.427	.515
Ti	309	.297	.959	1.000	.382	.075	.160	.731	.332	.367
Cr	731	049	.231	.382	1.000	598	596	075	410	288
La	.494	185	.300	.075	598	1.000	.970	.662	.743	.823
Се	.555	053	.370	.160	596	.970	1.000	.685	.779	.791
Y	.018	.027	.823	.731	075	.662	.685	1.000	.719	.746
U	.454	103	.427	.332	410	.743	.779	.719	1.000	.837
Th	.319	234	.515	.367	288	.823	.791	.746	.837	1.000

It is my contention that although the essential, accessory and associated minerals of ferrodioritic rocks correlate with the THM's and should be the likely source rocks for these THM's, the correlation matrix does not substantiate my conclusions. The only explanation I have for the inconsistency is that for robust calculations of correlation coefficients, the 28 samples (and 44 element variables) are too few to generate reasonable conclusions. In addition, the sediments may not be mature enough to exhibit clays scavenging mobile ions and as such, the elements do not appear to be correlating with aluminum. Nevertheless, the main purpose of this research was not to discover the rock provenances, but rather to identify minerals with enough economic potential to pay for a deepening of the Great Sacandaga Lake.

The titanium-mineral sands in our study area, principally ilmenite, at 0.69% of the total sediments of the study area equal 1.86 billion pounds at 90% washplant efficiencies. Ore-grade values from Standard and Poor's and Platt's Metal week indicate that these mineral sands are worth \$0.45 per pound. Phone conversations with industry officials indicate that prices may be as low as \$.15 per pound, until product chemistry has determined the contaminates in the ilmenite. QIT ran the ilmenite product chemistry (see Appendix D - Mineralogy). It shows the best magnetic fractions (using MicroProbe) contain 48% TiO₂, 39% FeTi, ~0.68% Al₂O₃, 0.03% CaO, 0.51% MgO, 0.02%P₂O₅, 0.47% MnO, 0.19% SiO₂, 0.03% Cr₂O₃, 0.26% V₂O₅, and 0.02% ZrO₂.

At a very conservative minimum estimate of \$0.15 per pound the ilmenite mineral sands are worth more than \$279 million dollars in the study area. Production costs are listed in Chapter 5.

Titanium metal has unique properties (density half that of steel, strength retention and atmospheric corrosion immunity) making it an ideal construction material for engines, airframes and structural purposes. Almost all Ti metal production involves tetrachlorination and reaction with either Mg or Na to form 85% pure metal and pig iron. called sponge. Ti sponge is worth from \$2-3.00 dollars per pound (Platt's, 2000). The same beginning process is used to make microcrystalline TiO₂ used for white pigment in paint, plastic and paper and is by far the largest single use of the metal. TiO₂ forms more than 20% by weight of paints and the pigment industry consumes more than 90% of all titanium minerals mined. The two communities of Gloversville and Johnstown with their under-capacity Wastewater Treatment Facilities and failing economies could benefit tremendously from further processing of titanium mineral sands. Benjamin Moore Paint Company has a new and very large facility in Johnstown and another Paint Company is planning to move to the area soon. Again, the titanium mineral sands estimated for the study area, and/or if the whole lake is deepened, do not have to be value added to, or further processed, to be worth what I anticipate to be their in-place value. It should also be pointed out here, that the whole lake may not contain heavy mineral sands such as REE or TiO₂ throughout its sediments in the same concentrations and further research will need to be done to estimate their abundance and potential worth.

The Rare Earth Oxides (REO) in the study area sediments are certainly enriched (see Appendix D). The fact that they exist in ore minerals such as apatite, zircon and monazite (already identified), that are durable and have high specific gravity makes them perfect candidates for dredge mining. Separately, the REO are quite valuable (Hedrick, 1995), worth between \$9 to \$2500 per pound, but even as a heavy mineral sand, they are

worth \$0.45-\$0.55 per pound (Table 4.6). Although REO are separated from one another in high purity by ion exchange methods, some of them, like cerium, are separated by taking advantage of its quadrivalent character. Today, many REO applications use a very substantial amount of the mixed REO just as they are obtained from their minerals ores or with extra cerium added or their lanthanum and cerium removed.

Rare-Earth Oxide (REO) Prices—by Various Corporations in 2000

Table 4.6

Oxide	Percent Purity*	Price per pound
Cerium	96-99.50	\$8.45-16.80
Dysprosium	95-96	\$38-60
Erbium	96-98	\$60-86
Europium	99.99	\$320-450
Gadolinium	99.99	\$55-60
Holmium	99.9	\$230-480
Lanthanum	99.99	\$8.75-10.50
Neodymium	95-99.9	\$11-40
Lutetium	99.99	\$2,000-2,500
Praseodymium	96.0	\$14.50-16.80
Samarium	96.0	\$30-34
Terbium	99.9	\$311-375
Thulium	99.9	\$1,600-1,650
Ytterbium	99.0	\$100-105
Yttrium	99.99	\$40-50

^{*} Purity expressed as percent of total REO (U.S. Department of the Interior, Hedrick, 1995).

Each of the REO has distinct individual properties and they form organic chelates, which are taken advantage of in modern ion-exchange methods of separation (Parker, 1984). Domestic mine production of REO increases every year because of their increasing demand in industrial applications and production is expected to increase substantially in the years to come (Hedrick, 1995). Apatite, zircon and monazite heavy mineral sands account for over 270 million pounds at 90% washplant efficiencies in this study area. Until the pre-feasibility study is completed, however, I will not assign any value to apatite and zircon as potential REO ore-grade minerals. My results indicate that 31 million pounds of monazite heavy mineral sands should be available in the study area at 90% washplant efficiencies with a value of approximately \$13 million dollars at \$0.45 per pound. Production costs are listed in Chapter 5.

It should be noted here again that the nearby communities of Gloversville and Johnstown in Fulton County could benefit tremendously from the mineral processing of these concentrates. Because of the tremendous capacity of the Gloversville-Johnstown Wastewater Treatment Plant, which was designed to accommodate a much larger leather industry (the leather industry is essentially a metals industry using metals and metal salts to denature rawskins), mineral's processing there should be a natural fit.

The last mineral abundance of economic significance in the study area sediments evaluated in this report are the magnetic fractions of iron, magnetite and titano-magnetite. ODM and QIT estimated between 7.1 - 7.6% (average 7.35%) magnetite in the THM population. This means that 0.77% magnetite exists in the whole sediment overburden of the study area or approximately 2.1 billion pounds at 90% washplant efficiencies.

Even though magnetite does not have significant ore-grade value, Platt's Metals weekly quotes are approximately \$35 per metric ton or \$0.015 per pound. Since magnetite has a specific gravity of 4.9-5.2 and is magnetic, it is easily obtainable and separable in washplant operations. In the study area, magnetite mineral ore worth exceeds \$30 million dollars. Production costs are listed in Chapter 5.

CHAPTER 5

SUMMARY

Island rebuilding and shoreline reconfiguration

The primary goal of this study was to identify potential mineral wealth in the sediments of the Great Sacandaga Lake study area and to evaluate the economic rational for permanently deepening the lakebed as a remedy to water-level fluctuations that many find intolerable. If we chose to deepen with earth-moving heavy equipment, the lake would have to be drained and the costs would have to be covered by the new rebuilt Island property values and the shoreline reconfiguration improvements. A modest beginning project like the one proposed in the study area, with an ideal island configuration to maximize the shoreline worth, would create over 10 miles of shoreline and have a shorefront value of almost \$11 million dollars. The Hudson River-Black River Regulating District Board (HHBRRD) would be unable to hire the earth moving, roadwork and access bridges needed to access the property for many times that amount. The Board would need to issue bonds to pay for the costs of earth-moving equipment and apportion the costs along to beneficiaries, but would have to claim the improvements as assets because it couldn't sell the property to a developer at a loss. The downstream beneficiaries would gain a modest benefit due to the increase of more than 18 billion more gallons of water in the reservoir (just in this phase-20' deep x 4.5 mi²), but would not be happy with the costs associated with the altruism of a deepening cause. The new lands permit (any newly created shoreline would be leased, like the current shoreline properties) beneficiaries would love the property, but both their permits and the permit

tenants on the other side of the lake would not be happy with the tremendous permit fees associated with such a project. Water-level drawdown would have to begin early in the season and the work would have to be done and completed in the coldest part of the year, which would make everyone displeased. And finally, if mineral wealth were available, an earth-moving heavy equipment method of recovering the mineral wealth could not be feasibly completed in a single season in the study area.

The water levels of the Great Sacandaga Lake are fixed historically, to regulate the flow of waters during times of high water and times of drought. The spillway at Conklingville is the point of regulation and it is unlikely that building a newer dam to raise water levels would be wise and in the interests of all involved. The remedy of choice would be to deepen the Lake and find a way to do it that benefits everyone including the HRBRD. The only prudent method of deepening is to remove some of the lakebed so that the reservoir holds more water. The only plausible way to do this is to find some mechanism that maintains the continued operation of the reservoir and pays for itself at the same time. Therefore, the only logical way to proceed is with *dredge mining* and washplant recoveries of the valuable heavies in the sediments of the lakebed.

The primary investigations of the placer deposits (valuable heavies) have been done, at least in the study area and points upstream in the Sacandaga River. Considerable enrichments in heavy mineral sands that are worth recovering include almandine, ilmenite, magnetite, and the REO (residing in monazite, apatite and zircon heavy mineral sands). There are probably smaller but recoverable quantities of other multiple oxides, but the work to verify them has not been done yet. The heavy mineral concentrate mineralogy has been done and without any further mineral sand beneficiation, the

remunerations derived from the project prove that it can be accomplished without costs to anyone, satisfy all of the beneficiaries, and pay-off handsomely (see Table 5.1).

Table 5.1

Heavy mineral sands. Estimated dredge mining worth of study area heavy mineral sands and projected worth after dredging the whole Reservoir.

Mineral	Study Area	Whole Reservoir
Almandine	\$782,000,000	\$7,042,000,000
Ilmenite	\$279,000,000	\$2,511,000,000
Magnetite*	\$30,640,000	\$827,280,000
REE	\$13,000,000	\$351,000,000
Total Potential Worth	\$1,104,640,000 Ф	\$10,731,280,000

Φ Dredge depth 6.6 yd (twenty feet deep x 4.5 square miles) in three seasons (~seven months each).

The costs to operate a successful dredge-mining operation are well known from various sites around the world. In one titanium sands operation in Australia, using 500 tph (tons per hour) cutter-head suction dredges, 12 mo/yr, 7 days/wk, 24 hr/day in unconsolidated sands like Great Sacandaga Lake sediments, it costs \$0.32/m³ without amortization, depletion, depreciation, interest or overhead. Running time is said to be ~95%. Recovery is >92%, determined by tailings, sand sampling and recovery of ore reserves, determined by drilling to be 98%. Floating washplants are used to recover, by specific gravity, rutile, zircon and ilmenite, using MDL spirals and Reichert cones (Hartman, 1996). Inflated estimated costs (in today's dollars) for recovering heavies from the sediments in the study area are listed in Table 5.2.

Heavy mineral sand dredging costs. Estimated costs to dredge mine the study area mineral sands and the projected costs to dredge the whole Reservoir.

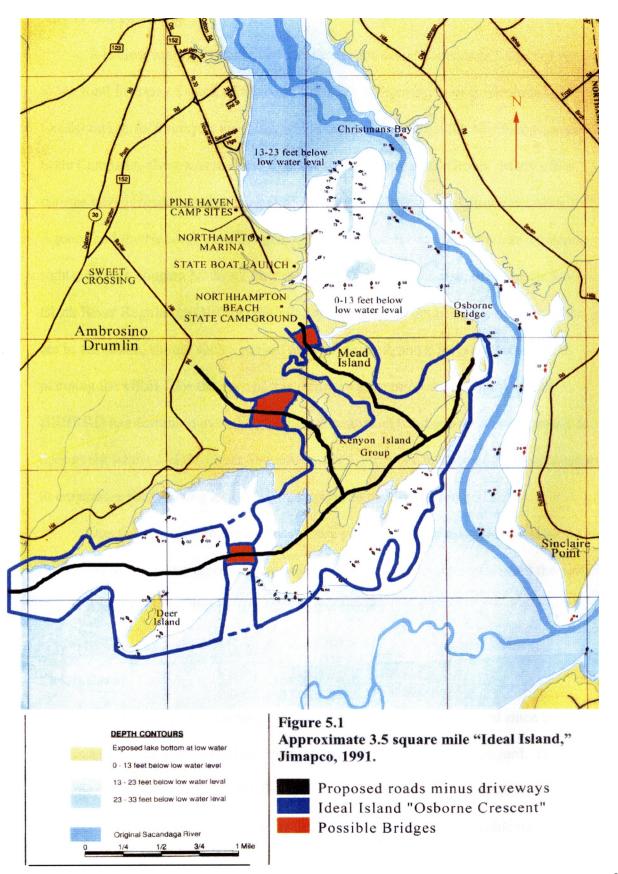
Operation	Study Area	Whole Reservoir	
Employees	\$11,000,000	\$990,000,000	
Dredging	\$45,000,000	\$405,000,000	
Dredges	\$6,000,000	\$12,000,000	
Wash Plants	\$15,300,000	\$30,600,000	
Transportation	\$40,000,000	\$360,000,000	
Executive Operations	\$1,500,000	\$13,500,000	
Outside Contracting	\$500,000	\$4,500,000	
Amortization & Interest	\$1,550,000	\$13,950,000	
Insurances	\$3,000,000	\$27,000,000	
Overhead	\$4,750,000	\$47,250,000	
Tax Planning	\$5,800,000	\$52,200,000	
Bridges (two in study area)	\$6,000,000	\$9,000,000	
Total Potential Costs	\$140,400,000Ф	\$1,965,000,000	

 $[\]Phi$ Study area based on three seasons (7 month seasons x 24 hours per day).

The base of the rebuilt island would be shaped in an arcuate band across the islands of Deer and the Kenyon's, which is the skeletal end moraine to the east and south in the study area (see Figure 3.5). The remnants of the north and west skeletal end moraine extending from the Northampton Marina and the New York State Boat Launch would be dredged and removed to the greatest extent possible. End moraines are loaded with large rock debris and a late season heavy earth-moving deepening may be necessary to remove the shallow water threat of the north and west skeletal end moraine. Everything in between the two skeletal end moraines is dredgeable and worth a recovery consideration.

Ideal Island

The ideal island would have more than 10 miles of shoreline and be about 3.5 mi^2 in area. This arcuate ideal island would have as many appendages as practicable to provide as much shoreline, bays and inlets needed to give the new real property charm and character (Figure 5.0). Everything to the north and west would be dredged, the heavies removed and the tailings used to rebuild the island. Three scenarios for bridges could be imagined, no bridges, one bridge and/or two bridges to connect the roads to the building sites. In Table 5.2, the cost of building two bridges in the study area under the ideal island scheme, are included. Anything practical could be built on the ideal island, including a golf course, but it might also include a marina, a water filtration plant, wastewater treatment plant and homesites. A suitable name would be Osborne Crescent. There are 2240 acres in 3.5 mi^2 . The shoreline properties, over 530 of them, at 100' of beachfront and selling at over \$200 per beachfront foot would be worth near \$11,000,000. If the lots are $100' \times 200'$ in length ($\sim 1/2$ acre), that would use-up about 265 acres of real estate. Figuring about 10 miles of road right-of-ways and driveways (60' wide), uses-up only another 75 acres of real estate. The balance of the interior lands, some 1900 acres, can be used for anything, including restaurants and shops, parks and trees or even some kind of amusement park or resort. My preference would be to keep as much open space as possible, but I would expect the local people to have ideas of their own about how to successfully manage this newly created real property. The size of the ideal island could be somewhat different than proposed in the study area.



Access and ownership

It is clear multiple entities have a stake in the Great Sacandaga Lake. According to the Real Property Bureau of the New York State Department of Environmental Conservation, for example, the DEC, which owns and operates the Northampton Beach State Campsite, also owns the land rights to the Kenyon Island Group. Many other entities have jurisdiction over the use of the lands, including the Adirondack State Park Agency and the National Park Service, while many others have interests in the water rights, such as Niagara Mohawk Power Corporation and of course, the Hudson River-Black River Regulating District (HRBRRD). There is no doubt that enormous obstacles are in the future, should such a project ever be realized, and a great deal of work in planning the effort. For the sake of this writing and in my personal opinion, the HRBRRD has dominion over the planning, management and execution of a project to deepen the lakebed of the Great Sacandaga Lake. No work has been done in this report to emphasize the planning aspects of this proposed deepening project.

One thing is clear to the success of any project with the magnitude and scope of work such as this paper espouses; the project has to satisfy the interests of all the parties involved and be in the "Public Benefit" for everyone.

Conclusions

So many years of wrangling and threats of litigation have passed since the vocal opponents of shallow water in the Great Sacandaga Lake began being heard. During all this time, many citizens from both the upstream and downstream interests have tried to come up with a palatable remedy to the navigation and recreational problems at the lake.

In this report, I have tried to show that a deepening project could actually be practical, environmentally sound and profitable at the same time.

All dredging operations are based on 5600 yd³ per hour, with like washplant operation capacity. That translates to nearly 17,000,000 pounds of lakebed removed, heavies separated and tailings delivered to rebuild the ideal island in the study area per hour. Over 125,000 pounds of heavies would be separated per hour and over 3,000,000 pounds per day. This means over 60 truckloads of heavies will need to be transported to Amsterdam for transport along railways to a smelting plant, unless provisions are set-up to do the smelting in the Gloversville-Johnstown area. No provisions are made in this paper to build a smelting plant in the Gloversville-Johnstown communities, but the potential worth of smelted TiO₂ heavies increases to between \$1.8-2.4 billion dollars, just with the first year dredged TiO₂ heavies of the study area. The cost of transporting the heavies and the equipment to do it is estimated in Table 5.2. By the economy of scale, the cost of transportation increases only by the cost of fuel over the life of the project, if the whole reservoir is dredged for heavies. The total potential worth of smelted heavies if the whole reservoir is dredged, is \$20 billion dollars (a very conservative estimate).

The dredge pace is set to accomplish the 6.6 yd deep x 1.5 mi² task in the study area within a normal high water season, approximately seven months. The whole of the reservoir would take between 25-30 years to dredge and recover heavies. The dredging and washplant systems have been designed for proven efficiency, safety and visibility, operator comfort and particularly for quiet operations. No credit will be given the corporations who build and design this equipment because it can be built and supplied by several such companies.

The reservoir releases about 150 billion gallons of water each season. At >6 billion gallons of increased water capacity each dredging season, the removal of the lakebed will create over 180 billion gallons of drinking water capacity if the whole reservoir is dredged. This extra drinking water capacity, will be a priceless commodity in the new frontier. No attempt will be made in this report to quantify it as an asset. Certainly, a deeper Lake will have profound effects on the well-being of its fish and other fauna. If the deepening continues throughout the Lake, the disruption to the lakebed organisms (estimate 27 like-areas to be dredged) will be completely reversed in each area within a three-year period (based on discussions with local biology professors). The answer to this dilema is to understand what fauna makes a living in soils, 6 months of the year and in water the remaining 6 months of the year?

The study area sediments proved to be extremely consistent throughout the sediment overburden of the study area. The heavy minerals identified are only slightly enriched from typical crustal sediments but as discussed they do have considerable value within the morainal ramparts investigated. It can be expected that at least another area of comparable size with sediments exactly like those of our study area can also be dredged for deepening and recovering heavy mineral sands, up the Sacandaga River. These two areas alone would take 6 seasons to dredge completely and be worth \$2.3 billion dollars. If the future studies of the GSL sediments indicate what I believe to be a terminal end moraine to the south and east of the Kenyon Island end moraine, at least another \$11 billion dollars worth of heavy mineral sands could potentially be recoverable.

This study has proved that a deepening project begun like the proposed study area project and expanded to include more and more of the lakebed sediments as appropriate, would be a perfect remedy to the dilemma of shallow water in the Great Sacandaga Lake. Furthermore, the public benefit and the economic advantages and spin-off businesses this project would deliver to the upstate economy are the obligation of the HRBRRD and the other regulating entities to allow to be brought to fruition. Much work needs to be done, but this project could develop through the planning phases, begun and start solving all the beneficiaries' problems within the next 5 years. It is my contention, that for future generations and the well-being of the industry of New York State, that an extra 180 billion gallons of drinking water is reason enough to undertake this project!

The next step I intend to take, beginning in the Fall of 2001 is to undertake a prefeasibility study of the whole of the Great Sacandaga Lake. I intend to investigate the resource and take a more careful look at the methods of extraction and methods of processing the mineral sands. I want to investigate more thouroughly, the potential markets, capacities and costs and look carefully at potential environmental problems and remedies. There is a lot of work ahead to understand reasonably well about the quality and quantity of the mineral resources in the Great Sacandaga Lake. The objectives of this report have been successfully completed, though there is a long road ahead before investors might want to participate in a program to design, seek permits, start operations and begin marketing of these heavy mineral sands.

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Appendix A PERMITS AND HISTORY

New York State Department of Environmental Conservation Office of Natural Resources, Region 5

Route 86 - P.O. Box 296, Ray Brook, New York 12977

Phone: (518) 897-1276 FAX: (518) 897-1370



January 27, 1999

University of Albany c/o Arthur M. Ambrosino Dept. of Earth & Atmospheric Sciences 1400 Washington Avenue, Room 351 Albany, NY 12222

Dear Mr. Ambrosino:

Enclosed is a Temporary Revocable Permit which will allow you to take one inch soil cores with a maximum depth of 20', on a portion of the Northampton Campground. This permit is subject to the Terms and Conditions listed. The permit will expire on April 1, 1999.

Regional Forester Tom Martin, located in the Ray Brook Office, will act on behalf of the Department of Environmental Conservation as Regional Land Manager. Should you have any questions or concerns, please contact Mr. Martin at 518/897-1276.

Sincerely,

Thomas H. Wahl

Thomas Allahl

Regional Supervisor of Natural Resources

THW:mb Enclosure

cc: T. Wolfe

J. Grossman

T. Kapelewski

FR: D. Brooks

ECO: M. Trottier

File: TRP#2 File

New York State Department of Environmental Conservation Reg 5 TRP 1996 TEMPORARY REVOCABLE PERMIT FOR THE USE OF STATE LANDS (see page 2 for standard terms and conditions) Name: University of Albany Address: c/o Arthur M. Ambrosino Dept. of Earth & Atmospheric Sciences 1400 Washington Avenue, Room 351 Albany, NY 12222 Telephone Number: 518/442-4466 Town: Mayfield State Land Location: County: Fulton Land Designation: Intensive Use Patent, Tract, Lot, Etc. Northampton Beach Campground The Department of Environmental Conservation hereby grants permission to the above named permittee to use these State lands described above and shown on the attached map for the purposes stated on the attached application and as amended below subject to the standard terms and conditions listed on page 2 and to the following special terms and conditions. (Use additional sheet if necessary). This TRP will allow you to take one inch soil cores with a maximum depth of 20', on a portion of the Northampton Campground. Unless otherwise suspended or revoked, this permit shall expire on April 1, 1999. APPROVAL Regional Director JAN 15 1999 Date Central Office (If required)

NOTICE OF EXPIRATION OF TEMPORARY REVOCABLE PERMIT					
This is to inform you that this permit expired on					
and that all conditions of such permit have been satisfactorily met.					
Date	Regional Director				

STANDARD TERMS AND CONDITIONS

- 1. This permit shall at all times be subject to the approval of the Regional Land Manager and may be suspended or revoked at any time with due cause.
- 2. The permittee shall notify the Regional Land Manager at least 48 hours prior to commencing use and upon completion of use.
- 3. The activities authorized under this permit shall not interfere with normal administration of the area by the Department.
- 4. No damage will be done to State land, State facilities or boundary or survey markers.
- 5. The permittee is responsible for any inadvertent or deliberate damage caused by the exercise of this permit and will be held responsible for restoration, rehabilitation or repair at the permittee's expense.
- 6. No trees or other vegetation shall be cut, disturbed or removed unless specifically authorized by the Regional Land Manager.
- 7. The State land covered by this permit shall be kept free of litter and debris and be left in a condition satisfactory to the Regional Land Manager at the expiration of the permit or on completion of activities authorized by the permit.

THE GREAT SACANDAGA LAKE ISLAND REBUILDING PROJECT: DREDGING AND REBUILDING THE KENYON ISLAND GROUP

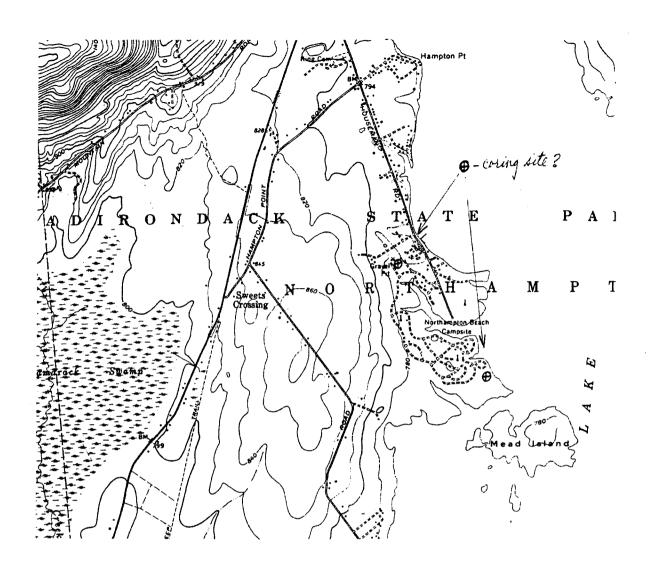
AMBROSINO, Arthur M., Department of Earth and Atmospheric Sciences, University at Albany, 1400 Washington Avenue, Albany, N.Y. 12222, aa7148@cnsvax.albany.edu

During the last decade, there has been increasing dispute between the downstream interests and the recreational users of the waters of the Great Sacandaga Lake. The downstream interests make up eleven power generation dams, Hudson River water/ Ocean water salt front concerns at Poughkeepsie, and flood control at Albany, which was the principal reason the Great Sacandaga Lake was created. The recreational users are the landowners/boat operators, around the Lake.

Large water-level variations to accommodate the downstream interests cause large areas of very shallow water in the Great Sacandaga, especially in midsummer, which is a situation the boating public finds unacceptable.

Dredging is the only remedy to deepen the Lake significantly for boating; thus, it is of interest to know the types of sediments at various points in the Lake and what the environmental concerns would be. How many meters of sediments from the erosion of the islands have accumulated since the dam created the Lake in 1930, and how has the Basin filled with sediments since the last glacial epoch?

The Kenyon Island Group, perhaps the largest and most unnavigable shallow water shoal in the Great Sacandaga Lake, is the ideal location for such a possible dredging project. The island group is uninhabited, accessable and appears to be the ideal soil for dredging. The dredged tailings will be used to rebuild the islands. Core sampling should identify the underlying sediment stratigraphy and establish the feasibility of such a project and its worthiness. In addition, sediment cores will be analyzed and interpreted with respect to Climate Change since the last glacial epoch, approximately 15,000 years B.P..



STATE OF NEW YORK



Board of Hudson River-Black River Regulating District

Sacandaga Field Office

737 Bunker Hill Rd., Mayfield, New York 12117

Phone (518) 661-5535 FAX (518) 661-5720

August 20, 1998

On this date Mr. Arthur Ambrosino was allowed to borrow the original Power Commission maps to have them scanned and will then return them immediately.

Arthur Ambrosino

STATE OF NEW YORK



Pourd of Hudson River-Black River Regulating District

Sacandaga Field Office 737 Bunker Hill Rd., Mayfield, New York 12117

Phone (518) 661-5535 FAX (518) 661-5720

September 2, 1998

On this date Mr. Arthur Ambrosino returned the Original Power Commission maps to the Sacandaga Field Office.

Marlene Junguera

STATE OF NEW YORK



Board of Hudson River-Plack River Regulating District 350 Northern Boulevard, Albany, New York 12204 Phone (518) 465-3491 FAX (518) 432-2485

FAX COVER SHEET

TO:	ARTHUR AWITHOSINO
FROM:	Tom BREWER
DATE:	11/12/98
SUBJECT:	CORE SAMPLES
	NO. OF PAGES (including cover sheet):
MESSAGE:	THIS WILL CONFIRM THAT YOU ARE
	1260 TO GOLLECT CORE SEDIMENT
	S IN THE AREA OF THE KENDON ISLANDS
	AT SACANDAGA LAKE ON 11/13/94.
	Thomas E. Brine
	CHIEF ENGINEER

IF YOU HAVE ANY PROBLEMS OR QUESTIONS RECEIVING, PLEASE CALL OUR OFFICE IMMEDIATELY.



Maps to swear by...not at!

LICENSING AGREEMENT

DATE: 8/25/98	
CUSTOMER: Archur M. Arnt LICENSE NUMBER: C.982508	1705LD - 50NY ALDERY
LICENSE NUMBER: C.982508	- bay 192-5825
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	Christina T. Fisk, Vice President

History of the study area

The Sacandaga Basin has long been an important geopolitical region in the eastcentral part of New York State, no less important to the Native Americans who named the lowlands "Sacandog," meaning "Land of Waving Grass." As early as the 1860's, the New York State Legislature considered the idea of constructing a dam and building a reservoir in the basin. The Sacandaga Basin contains approximately one third of the waters of the Upper Hudson Watershed. For many years, disastrous floods regularly wreaked havoc in the communities along the Hudson River, from Glens Falls to Albany. In the early 1900's, the Legislature gave its blessing to the New York State Water Supply Commission to survey the basin and identify where a suitable dam should be constructed. In 1908 the Water Supply Commission's task was completed and the proposed "Flow Line" and the site of the dam at Conklingville were identified. The State Legislature established the Hudson River Regulating District in 1922 as a public benefit corporation. The District was given a broad spectrum of legal powers to accomplish this mission, including the authority to build and operate the reservoir, issue bonds and apportion costs on its beneficiaries to finance construction, maintenance and for operations. The District's specific responsibilities are to reduce floods by excess run-off and to augment Hudson River water flow at times of drought or other periods when river flows are low. Work began on the dam in 1927, was completed in 1929 and the valley was flooded by 1930.

The project cost \$12,000,000, and the 1100-foot Conklingville Dam retains close to 300 billion gallons of water in the reservoir at the high water mark. 125 miles of shoreline defines the reservoir's 42 square mile surface area. Construction of the

reservoir took more than 500 men and several years to complete. It submerged 3872 graves in 22 cemeteries, which required transburial. It submerged 5 miles of the Fonda, Johnstown and Gloversville Railroad, which was never rebuilt. It submerged 67 miles of roads connecting several communities beneath the waters and required more than 50 miles of new highway construction. Five existing bridges were destroyed and 10 new bridges had to be erected. Finally, over 1250 structures were moved, burned or pulled down to make room for the water.

During the last several decades, there has been increasing dispute between the downstream interests and the recreational users of the water of the Great Sacandaga Lake. The downstream interests include 11 power generation dams, Hudson River water/Ocean water salt front concerns at Poughkeepsie and flood control at Albany. Large water-level variations to accommodate the downstream interests cause large areas of very shallow water in the Great Sacandaga, especially in midsummer, which is a situation the landowners and boating public find unacceptable.

The spillway at the Conklingville Dam is 771' above sea level. At high water the study area (Kenyon Island Group including Mead and Deer Islands) have shorelines at about 768'. The lake levels begin to drop long before the low water season approaches. Water is let out at an average rate of just less than 1 1/2" per day and by the low water season over 150 billion gallons of water have been drained from the reservoir. It takes approximately 6 months for the transformation from high water to low water level status. As the water is let out at Conklingville at the rate of a little less than 4' per month, the shallowest parts of the shallow water shoal that surrounds the Kenyon Island Group becomes more and more dangerous to recreational boaters of the lake. The situation

becomes so serious that almost nine miles of shallow warning buoys are strung around the greater shoal, warning boaters not to enter. Boaters are relegated to the marine buoyed Sacandaga River Channel for much of the boating season to get around the shallows.

In an effort to reach accommodations to all, an Upper Hudson-Sacandaga Comprehensive Settlement Negotiation Team is presently studying water level remedies (see Table below). The proposed remedy involves raising the high water mark to 770' over the next twenty years. Any remedy they agree upon will be a compromise to both upstream and downstream interests.

ADIRONDACK BOARDSAILING CLUB ADIRONDACK COUNCIL ADIRONDACK HYDRO DEVELOPMENT CORPORATION ADIRONDACK MOUNTAIN CLUB ADIRONDACK PARK AGENCY ADIRONDACK RIVER OUTFITTERS AMERICAN WHITEWATER ASSOCIATION FEDERAL ENERGY REGULATING COMMISSION FEEDER CANAL ALLIANCE FINCH, PRUYN AND COMPANY, INCORPORATED **FULTON COUNTY GOMEZ AND SULLIVAN** GREAT SACANDAGA LAKE ASSOCIATION GREAT SACANDAGA LAKE FISHERIES FEDERATION GREAT SACANDAGA LAKE MARINAS **HUDSON RIVER RAFTING** HUDSON RIVER-BLACK RIVER REGULATING DISTRICT **INDIVIDUALS** INTERNATIONAL PAPER MERCER COS NATIONAL PARK SERVICE **NEW YORK RIVERS UNITED** NEW YORK STATE CONSERVATION COUNCIL NYS DEPARTMENT OF ENVIRONMENTAL CONSERVATION NEW YORK STATE ENERGY GRID NIAGARA MOHAWK **HUDSON RIVER MANAGER** SARATOGA COUNTY TOWN OF HADLEY TROUT UNLIMITED UNITED STATES FISH AND WILDLIFE SERVICE

WILDWATERS, INCORPORATED

Appendix B

PARTICLE SIZE VOLUME STATISTICS

AND

CORING INTERPRETATIONS

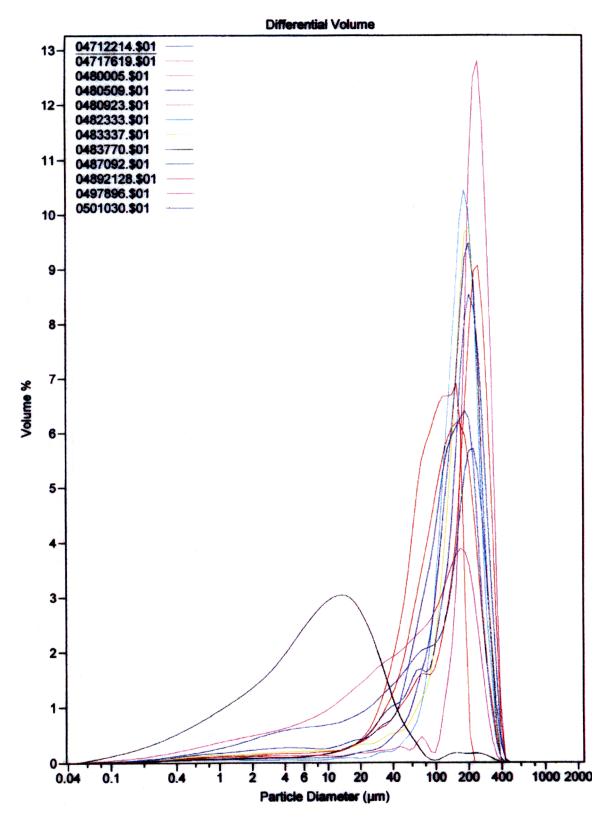


Figure 3.1a Coulter LS Particle Size Analyzer, Sediment/Core Lab, Geology Department, Union College.

LS Particle Size Analyzer

6 Jul 1999

	- SEDIMENT/CORE LAB,	GEOLOGY D	DEPT., UNION COLLEGE	6 Jul 1999
File name:	04712214.\$01	Group ID:	047122140	
Sample ID: Operator: Comments:		Run number:	2	
Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur:	Fluid Module 9:32 30 Jun 1999 100	Run length:	120 Seconds	
Fluid: Software:	Water 2.09	Firmware:	2.02- 2.02	
File name: Sample ID:	04717619.\$01	Group ID:	047176192	
Operator:		Run number:	12	
Comments: Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur:	Fraunhofer PIDS included Fluid Module 11:34 30 Jun 1999 81 3% 26%.	Run length:	119 Seconds	
Fluid: Software:	Water 2.09	Firmware:	2.02 2.02	
File name: Sample ID:	0480005.\$01	Group ID:	0480005	
Operator: Comments:		Run number:	14	
Optical model: LS 230 Start time: Pump speed: Obscuration: PiDS Obscur:	Fraunhofer PIDS included Fluid Module 11:51 30 Jun 1999 81 2% 19%	Run length:	120 Seconds	
Fluid: Software:	Water 2.09	Firmware:	2.02 2.02	

2.09

Software:

LS Particle Size Analyzer

6 Jul 1999

	- SEDIMENT/CORE LAB,	GEOLOGY D	EPT., UNION COLLEGE	O JUL. 18
File name:	0480509.\$01	Group ID:	0480509	
Sample ID: Operator: Comments:		Run number:	4	
Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur:	Fluid Module 10:27 30 Jun 1999 81 6%.	Run length:	119 Seconds	
Fluid: Software:	Water 2.09-	Firmware:	2.02 2.02	
File name: Sample ID:	0480923.\$01	Group ID:	0480923	
Operator:		Run number:	10	
Comments: Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur:	81 2%	Run length:	120 Seconds	
Fluid: Software:	Water 2.09	Firmware:	2.02 2.02	
File name: Sample ID:	0482333.\$01	Group ID:	0482333	
Operator:		Run number:	3	
LS 230 Start time: Pump speed: Obscuration: PIDS Obscur:	1 % 12%	Run length:	119 Seconds	
Fluid:	Water	F1	0.00	

Firmware:

2.02 2.02

LS Particle Size Analyzer

6 Jul 1999 SEDIMENT/CORE LAB, GEOLOGY DEPT., UNION COLLEGE-

0483337

File name:

0483337.501

Group ID:

Sample ID:

Operator:

Run number: 5

Comments:

Optical model: Fraunhofer PIDS included

LS 230

Fluid Module

Start time:

10:35 30 Jun 1999

81

Pump speed: Obscuration: PIDS Obscur:

3% 30% Water

Fluid: Software:

2.09

Firmware:

Run length:

2.02 2.02

120 Seconds

File name: Sample ID: 0483770.501

Group ID:

Run length:

0483770

119 Seconds

Operator:

Comments:

Run.number: 15

Optical model:

Fraunhofer PIDS included Fluid Modute

LS 230 Start time:

12:00 30 Jun 1999

Pump speed: Obscuration:

81 22% PIDS Obscur: 91%

Fluid: Software: Water 2.09

Firmware:

File name:

0487092.\$01 Sample ID:

Group ID:

Run number: 1

0487092

119 Seconds

2.02 2.02

Operator.

Comments: Fraunhofer PIDS included

Optical model:

LS 230 Fluid Module

81

Start time:

9:53 30 Jun 1999

Pump speed: Obscuration:

2% PIDS Obscur: 16% Fluid:

Software:

Water 2.09

Run length:

Firmware:

2.02 2.02

LS Particle Size Analyzer

6 Jul 1999

	- SEDIMENT/CORE LAB,	GEOLOGY E	DEPT., UNION COLLEGE -	0 301 1338
File name: Sample ID:	04892128.\$01	Group ID:	04892128	
Operator:		Run number:	1	
Comments: Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur: Fluid:	Fluid Module 9:23 30 Jun 1999 100 3% 23%	Run length:	120 Seconds	
Software:	Water 2.09	Firmware:	2.02 2.02	
File name: Sample ID:	0497896.\$01	Group ID:	0497896	
Operator:		Run number:	16	
LS 230 Start time: Pump speed: Obscuration: PIDS Obscur:	Fraunhofer PIDS included Fluid Module 12:10 30 Jun 1999 81 8% 60%	Run length:	119 Seconds	
Fluid: Software:	Water 2.09	Firmware:	2.02 2.02	
File name: Sample ID:	0501030.\$01	Group ID:	0501030	
Operator:		Run number:	1	
Comments: Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur: Fluid:	Fraunhofer PtDS included Fluid Module 8:41 30 Jun 1999 81 5% 43% Water	Run length:	120 Seconds	
Software:	2.09	Firmware:	2.02 2.02	

LS Particle Size Analyzer

6 Jul 1999

SEDIMENT/CORE LAB	GEOLOGY DEPT.	, UNION COLLEGE —
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04712214.\$01

Calculations from 0.040 µm to 2000 µm

1/~		-	_
V CI	u	m	ш

100.0%

Mean: Median: 168.6 um

174.2 µm

S.D.: C.V.: 89 µm 52.8%

D(3,2): 28.77 µm Mode: 203.5 µm

0.00299 Right skewed Skewness: -0.668 Platykurtic Kurtosis:

% < 90 10 50 75 39.67 283.5 Size µm 101.6 174.2 231.4

Volume Statistics (Arithmetic)

Volume Statistics (Arithmetic)

04717619.\$01

Calculations from 0.040 µm to 2000 µm

Volume

100.0%

Mean: Median: D(3,2):

128.8 µm 120.6 µm 26.19 µm S.D.: C.V.: Skewness:

74.1 µm 57.5% 0.526 Right skewed

-0.108 Platykurtic

90

230.8

Mode: 153.8 µm Kurtosis: % < 10 25 50 75

Size µm 38.32 72.66 120.6 176.9

0480005.\$01

Calculations from 0.040 µm to 2000 µm

Volume

Mode:

100.0%

Mean: Median: D(3,2):

218.8 µm 226.6 µm 25.56 µm 245.2 µm

S.D.: C.V.: Skewness: Kurtosis:

83.3 µm 38.1%

-0.758 Left skewed 0.676 Leptokurtic

% < 25 50 75 90 10 116.3 181.1 226.6 273.5 316.5 Size µm

LS Particle Size Analyzer

6 Jul 1999

	Gi	"DIMENIA		ib, GEOLOG	IT DEPT., UNION CO
		Volume Statistics (Arithmetic)			0480509.\$01
Calculations	from 0.0)40 μm to 200	0 µm		
Volume Mean: Median: D(3,2): Mode:		100.0% 124.3 µm 111.6 µm 7.889 µm 223.4 µm		S.D.: C.V.: Skewness: Kurtosis:	99.8 µm 80.3% 0.433 Right skewed -0.905 Platykurtic
% < Size µm	10 4.735	25 29.25	50 111.6	75 204.2	90 264.5
		Volume Star	tistics (A	rithmetic)	0480923.\$01
Calculations	from 0.0	40 μm to 200	0 µm		
Volume Mean: Median: D(3,2): Mode:		100.0% 179.1 µm 190.8 µm 22.32 µm 245.2 µm		S.D.: C.V.: Skewness: Kurtosis:	96.6 µm 53.9% -0.157 Left skewed -0.837 Platykurtic
% < Size µm	10 35.12	25 103.4	50 190.8	75 251.0	90 301.3
		Volume Stat	istics (Aı	rithmetic)	0482333.\$01
Calculations	from 0.0	40 µm to 2000) µm		
Volume Mean: Median: D(3,2): Mode:		100.0% 169.5 µm 168.9 µm 24.61 µm 185.3 µm	:	S.D.: C.V.: Skewness: Kurtosis:	67 µm 39.5% 0.0157 Right skewed 0.183 Leptokurtic
% < Size µm	10 88.99	25 126.3	50 168.9	7 5 212.2	90 254.5

LS Particle Size Analyzer

6 Jul 1999

	•			, ororoc	,, <u>DE</u> , 7., 0 , 110.100.		
		Volume Statistics (Arithmetic)			0483337.\$01		
Calculations	Calculations from 0.040 µm to 2000 µm						
Volume Mean: Median: D(3,2): Mode:		100.0% 162.7 μm 169.9 μm 18.45 μm 203.5 μm	C	.D.: .V.: kewness: urtosis:	80.5 µm 49.5% -0.228 Left skewed -0.389 Platykurtic		
% < Size μm	10 35.35	25 114.3	50 169.9	75. 217.4	90 262.4		
		Volume Star	tistics (Arit	hmetic)	0483770.\$01		
Calculations	from 0.0	140 μm to 200	0 µm				
Volume Mean: Median: D(3,2): Mode:		100.0% 15.55 µm 8.051 µm 1.887 µm 13.61 µm	C. SI	D.: .V.: kewness: urtosis:	30.2 µm 194% 6.37 Right skewed 50.4 Leptokurtic		
% < Size μm	10 0.831	25 2.719	50 8.051	75 17.69	90 31.42		
		Volume Stat	tistics (Ariti	hmetic)	0487092.\$01		
Calculations	from 0.0	40 μm to 2000	0 µm				
Volume Mean: Median: D(3,2): Mode:		100.0% 166.6 µm 169.0 µm 22.49 µm 203.5 µm	C. Si	D.: V.: kewness: urtosis:	75.7 µm 45.4% -0.0467 Left skewed -0.308 Platykurtic		
% < Si ze μm	10 65.51	25 115.6	50 169.0	75 218.5	90 263.8		

LS Particle Size Analyzer

6 Jul 1999

 SEDIMENT/CORE LAB.	GEOLOGY DEPT	LINION COLLEGE -
OFDINICIALIOOME FAD.	GEOLOGI DEFI.	. UNION COLLEGE -

'	Volume	Statistics	(Arithmetic)
---	--------	------------	--------------

04892128.\$01

Calculations from 0.040 µm to 2000 µm

Volume

100.0%

Mean: Median: D(3,2):

Mode:

93.80 µm 91.08 µm

16.01 µm

Skewness: 153.8 µm Kurtosis:

S.D.: 47.4 µm 50.6% C.V.:

0.0746 Right skewed -0.762 Platykurtic

% < 10 Size µm 32.64 25 58.70

91.08 130.3 90 160.2

Volume Statistics (Arithmetic)

0497896.\$01

Calculations from 0.040 µm to 2000 µm

Volume

100.0%

Mean: Median: D(3,2):

85.18 µm 60.84 µm 5.095 um

S.D.: C.V.: Skewness: 79.1 µm 92.8%

Mode: 168.8 um

Kurtosis:

0.851 Right skewed -0.2 Platykurtic

25 10 50 75 90 Size µm 3.078 16.14 60.84 141.2 204.7

Volume Statistics (Arithmetic)

0501030.\$01

Calculations from 0.040 µm to 2000 µm

Volume

100.0%

Mean: Median:

129.9 µm 126.3 µm S.D.: C.V.:

78.7 µm 60.6%

D(3,2): 13.02 µm Skewness: 0.281 Right skewed Mode: 185.3 µm Kurtosis: -0.445 Platykurtic

% < 10 25 75 50 Size µm 18.55 71.51 126.3 185.5 234.5

LS Particle Size Analyzer

6 Jul 1999 -SEDIMENT/CORE LAB, GEOLOGY DEPT., UNION COLLEGE -Particle 04712214.\$01 04717619.\$01 0480005.\$01 0480509.\$01 0480923.\$01 0482333.\$ Diameter Volume Volume Volume Volume Volume Volume % μm % % % % % 1.000 0.455 0.543 2.46 0.811 0.350 0.644 2.000 0.779 0.872 1.23 5.20 1.52 0.457 5.000 2.83 2.75 1.93 10.52 3.27 0.885 20.00 7.90 6.58 9.31 2.09 12.04 1.72 50.00 18.77 37.95 3.95 20.12 16.34 20.17 125.0 49.95 40.92 52.68 34.08 44.98 64.40 25.31 250.0 18.60 6.81 38.41 12.88 11.22 500.0 0 0 0 0 0 0 1,000 0 0 0 0 0 0 2,000 0 0 0 0 0

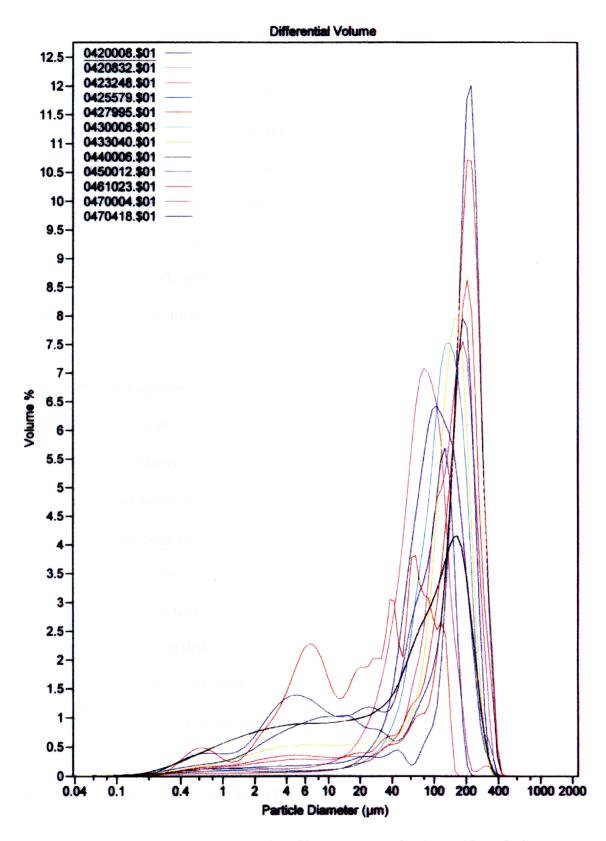


Figure 3.1b Coulter LS Particle Size Analyzer, Sediment/Core Lab, Geology Department, Union College.

LS Particle Size Analyzer

6 Jul 1999

- SEDIMENT/CORE	LAB. GEOI	OGY DEPT.	UNION COLLEGE -
-----------------	-----------	-----------	-----------------

File name:

0420008.\$01

Group ID:

Sample ID:

Operator:

Run number: 2

0420008

120 Seconds

2.02 2.02

Comments:

Optical model:

Fraunhofer PIDS included

LS 230

Fluid Module

Start time: Pump speed: 10:06. 30 Jun 1999:

81

Obscuration: 2% PIDS Obscur: 24%

Fluid: Water 2.09

Software:

Firmware:

Group ID:

Run number: 2

Run length:

File name: Sample ID:

0420832.\$01

0420832

Operator: Comments:

Fraunhofer PIDS included

Optical model: LS 230

Fluid Module Start time: 15:04 29 Jun 1999

Pump speed: 81

Obscuration: 14% PIDS Obscur: 49% Water

Fluid: Software:

2.09

Firmware:

Run length:

2.02 2.02

119 Seconds

File name:

0423248,\$01

Group ID:

Run number: 7

0423248

Sample ID:

Operator: Comments:

Fraunhofer PIDS included

Optical model: LS 230

Fluid Module 10:50 30 Jun 1999

Start time:

Pump speed: 81

Obscuration: 4% PIDS Obscur. 25% Fluid:

Software:

Water

2.09

Firmware:

Run length:

2.02 2.02

119 Seconds

Software:

2.09

LS Particle Size Analyzer

6 Jul 1999

	- SEDIMENT/CORE LAB,	GEOLOGY D	EPT., UNION COLLEGE -	0 1m taaa
File name:	0433040.\$01	Group ID:	0433040	
Sample-ID: Operator: Comments:		Run number:	13	
Optical model: LS 230	Fraunhofer PIDS included Fluid Module		•	
Start-time: Pump speed: Obscuration: PIDS Obscur: Fluid:	11:41 30:Jun 1999 81 5% 49% Water	Run length:	120 Seconds	
Software:	2.09	Firmware:	2.02 2.02	
File name: Sample ID:	0440006.\$01	Group ID:	0440006	
Operator: Comments:		Run number:	9	
Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur:	Fraunhofer PIDS included Fluid Module 11:09 30 Jun 1999 81 9% 73%	Run length:	120 Seconds	
Fluid: Software:	Water 2.09	Firmware:	2.02. 2.02	
File name: Sample ID:	0450012.\$01 a0450012	Group ID:	Sacandaga	
Operator: Comments:	Jaime	Run number:	3	
Optical model: LS 230	Fraunhofer PIDS included Fluid Module			
Start time: Pump speed: Obscuration: PIDS Obscur:	15:24 29 Jun 1999 81 19% 54%	Run length:	120 Seconds	
Fluid:	Water			

Firmware:

2.02 2.02

LS Particle Size Analyzer

6 Jul 1999

	- SEDIMENT/CORE LAB,	GEOLOGY D	DEPT., UNION COLLEGE -
File name: Sample ID:	0425579.\$01	Group ID:	0425579
Operator: Comments:	jaime	Run number:	2
Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur: Fluid:	Fluid Module. 8:15: 30 Jun 1999: 81 2% 18%:	Run length:	120 Seconds
Software:	Water 2.09	Firmware:	2.92 2.02
File name: Sample ID: Operator:	0427995.\$01	Group ID: Run number:	0427995
Comments: Optical model: LS-230	Fraunhofer PIDS included Fluid Module	•	
Start time:	11:28 30 Jun 1999	Run length:	119 Seconds

Pump speed: 81 Obscuration: 1% PIDS Obscur. 14% Fluid: Water

Software: 2.09

Group-ID: 0430006

Run number: 6

2.02 2.02

Firmware:

File name: Sample ID: Operator: Comments:

Fraunhofer PIDS included Optical model:

0430006.\$01

LS 230 Fluid Module

Start time: 10:43 30 Jun 1999 81 Pump speed:

Obscuration: 3% PIDS Obscur: 30% Fluid: Water Software: 2.09

Run length: 120 Seconds

Firmware: 2.02 2.02

Water 2.09

Fluid: Software:

LS Particle Size Analyzer

	CEDIMENTICODE : AD	0501.001		6 Jul 1999
	- SEDIMEN I/COKE LAB	GEOLOGY D	PEPT., UNION COLLEGE	, , , , , , , , , , , , , , , , , , ,
File name: Sample ID:	0461023.\$01	Group ID:	0461023	
Operator: Comments:		Run number:	8	
Optical model: LS 230. Start time: Pump speed: Obscuration: PIDS Obscur:	Fluid Module. 10:59: 30 Jun: 1999: 81: 4%- 37%	Rum length:	120 Seconds	
Fluid: Software:	Water 2.09	Firmware:	2.02 2.02	
File name: Sample ID:	0470004.\$01	Group ID:	0470004	
Operator.	jaime	Run number.	1	
Comments: Optical model: LS 230 Start time: Pump speed: Obscuration: PIDS Obscur: Fluid:	Fraunhofer PIDS included Fluid Module 8:05 30 Jun-1999 81 4%. 34% Water	Ru n length.	120 Seconds	
Software:	2.09	Firmware:	2.02 2.02	
File name: Sample ID:	0470418:\$01 0470418	Group ID:		
Operator: Comments:		Run number:	1	
Optical model: LS 230 Start time: Pump speed: Obscuration:	Fluid Module 14:44 29 Jun 1999 81 13%	Run length:	119 Seconds	
PIDS Obscur:	50%			

Firmware:

2.02 2.02

LS Particle Size Analyzer

6 Jul 1999

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	Ji		UNE EAL	o, GEULUI	ST DEPT., UNIONFOU
		Volume Sta	atistics (Ar	ithmetic)	0461023.\$01
Calculations	-from 0.0	040 μm to 200)0 µm		
Volume Mean: Median: D(3,2): Mode:		160,0% 152.0 µm 159.7 µm 14.42 µm 203.5 µm	(S.D.: C.V.: Skewness: Curtosis:	88.5 µm 58.2% -0.0137 Left skewed -0.635 Platykurtic
% < Siz e µm	10 13.77	2 5 87.32	50- 159.7	75- 214:3	90 262.0
		Volume Sta	tistics (Ari	thmetic)	0470004.\$01
Calculations	from 0.0	40 µm to 260	0 μm		
Volume Mean: Median: D(3,2): Mode:		100.0% 140.0 µm 140.5 µm 14.82 µm 185.3 µm	S	i.D.: LV.: kewness: turtosis:	78.8 µm 56.3% 0.0732 Right-skewed -0.562 Platykurtic
% < Size μm	10 20.38	25 85.75	50. 140.5	75 195.9	90 243.7
		Volume Stat	tistics (Ari	thmetic)	0470418.\$01
Calculations	from 0.0	40 µm to 200	0 µm		
Volume. Mean: Median: D(3,2): Mode:		100.0% 117.2 μm 129.7 μm 8.060 μm 185.3 μm	C	.D.: .V.: kewness: urtosis:	88.7 µm 75.7% 0.0691 Right skewed -1.3 Platykurtic
% < Size µm	10 3.974	25 17.02	50 129.7	75 190.8	90 230.9

LS Particle Size Analyzer

6 Jut 1999

SEDIMENT/CORE LAB,	GEOLOGY DEPT.	UNION COLLEGE -

		EUIMEN I/C	UKEL	B, GEOLOG	SY DEPT., UNION CO
		Volume Sta	atistics (/	0420068.\$01	
Calculations	from 0.	040 µm to 200	20 µm		
Volume		100,0%			
Mean:		186.9 µm		S.D.;	82,9 µm
Median:		197.0 µm		C.V.:	44.4%
D(3,2):		19.63 μm		Skewness:	-0.536 Left skewed
Mode:		223.4 µm		Kurtosis:	0.0867 Leptokurtic
% <	10	25	50	75	90
Size pm	42.58	147.8	197.0	240,6	282.1
		Volume Sta	itistics (A	rithmetic)	0420832.\$01
Calculations	from 0.6	740 µm to 200	10 μm		
Volume-		100.0%			
Mean:		38.39 µm		S.D.:	37.1 µm
Median:		25.28 µm		€.V.:	96.7%
D(3,2):		5.680 µm		Skewness:	0.949 Right skewed
Mode:		66:44 µm		Kurtosis:	-0:136 Platykurtic
% <	10	25	50	75	90
Size µm	2.857	6.678	25.28	63.12	96.87
		Volume Star	listics (A	rithmetic)	0423248.\$01
Calculations	from 0.0	40 μm to 200	0 µm		
Volume		100.0%			
Mean:		77.9 8 µm		S.D.:	44.6 µm
Median:		73.86 µm	•	C.V.:	5 7.2%
D(3,2):		12.10 µm	;	Skewness:	1.02 Right skewed
Mode:		80.08 µm	ı	Kurtosis:	3.24 Leptokurtic
% <	10.	25	50 :	75	90
Size µm	25.08	48.25	73.88	103.5	132.9

LS Particle Size Analyzer

6 Jul 1999

SEDIMENT/CORE LAB, GEOLOGY DEPT.	1 MINOR OCH FOR
	, Oliver GOLLEGE

		Volume Sta	atistics (Ar	rithmetic)	0425579.\$01
Calculation	s from 0.	040 µm to 200	00 µm		
Volume- Mean: Median: D(3,2): Mode:		190.0% 109.0 pm 100.1 pm 15.94 pm 105.9 pm	S.D.: C:V.:		69.2 µm 55.2% 0.686 Right skewed 0.46 Leptokurtic
% < Size µm	10 40.90	25 65, 6 7	50 100.1	75 145,6	90 191.9
		Volume Sta	tistics (Ari	thmetic)	0427995.\$01
Calculations	from 0.0	40 μm to 200	0 µm		
Volume Mean: Median: D(3,2): Mode:		100.0% 185.3 µm 191.0 µm 29.66 µm 203.5 µm	S	i.D.: i.V.: kewness: urtosis:	80.3 µm 43.4% -0.218 Left skewed -0.179 Platykurtic
% < Size µm	10 67.34	25 137.7	50 191.0	75 238.6	90 284.5
		Volume Stat	istics (Arit	hmetic)	0430006.\$01
Calculations	from 0.0	40 µm to 2000) µm		
Volume Mean: Median: D(3,2): Mode:		100.0% 125.4 µm 121.5 µm 18.62 µm 140.1 µm	C. Si	D.: V.: (ewness: utosis:	64.9 µm 51.7% 0.42 Right skewed 0.343 Leptokurtic
% < Size µm	10 45.89	25 80.49	50 121.5	7 5 166.6	90 209.4

LS Particle Size Analyzer

6 Jul 1999

SEDIMENT/CORE LAB, GEOLOGY DEPT., UNION COLLEGE	P

Volume Statistics (Arithmetic)

0433040.\$01

Calculations from 0.040 µm to 2000 µm

Volume

100.0%

Mean: Median: 119.1 µm 127.2 µm S.D.: C.V.: 74.4 µm 62.5%

D(3.2): Mode:

8.611 µm 153.8 µm

Skewness: Kurtosis:

-0.0549 Left skewed -0.814-Platykurtic

% < 10 25 50 75 90 Size µm 5.488 59.49 127.2 172.9 212.3

Volume Statistics (Arithmetic)

0440006.\$01

Calculations from 0:040 µm to 2000 µm

Volume

100.0%

Mean: Median: D(3.2):

Mode:

83.70 um 65.34 um 5.064 µm 168.8 µm

S.D.: C.V.: Skewness:

78.2 µm 93.5% 0.769 Right skewed -0.31 Platykurtic

Kurtosis: % < 10 25 50 75 90 Size um 1.948 9.899 65.34 140.5 198.8

Volume Statistics (Arithmetic)

0450012.\$01

Calculations from 0.040 µm to 2000 µm

Volume.

100.0%

Mean: Median: D(3,2):

Mode:

61.74 µm 54.09 um 5.640 µm

127.6 µm

S.D.: C.V.: Skewness: Kurtosis:

54 µm 87.4%

0.445 Right skewed -1.1 Platykurtic

%< 10 25 50 75 90 Size µm 2.509 7.543 54.09 108.1 139.9

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LS Particle Size Analyzer

6 Jul 1999 - SEDIMENT/CORE LAB, GEOLOGY DEPT., UNION COLLEGE -Particle 0420008.\$01 0420832.\$01 0423248.\$01 0425579.\$01 0427995.\$01 0430006.\$ Diameter Volume Volume Volume Volume Volume Volume μm % % % % % % 1:000 0.979 0.352 0.848 244 0.654 1.01 2.000 0.559 1.58 11.26 0.707 1.01 1.21 5,000 3.24 26.94 1.80 1.65 1.77 4.00 20.00 3.51 22.23 18.59 10.11 4.12 6.00 50.00 7.34 13.65 29.66 50.61 41.10 60.40 125.0 61.32 2.85 32.62 58.74 44.23 12.50 250.0 20.61 0 0.542 2.21 20.23 3.47 500.0 0 0 0 0-0 0 1,000 Œ 0 0 0 Q-0

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Core descriptions

In cores #1 and #4 we encountered small rocks in gravel horizons, which were about 2" in diameter and these rocks blocked the recovery of complete cores until the cutting shoe chewed them up enough to swallow them (Figure 2.1). This Figure should be referred to whenever detailed descriptions of the core horizons are being considered. None of the cores encountered any rocks of enough consequence to force us to relocate the truck for a second push.

Core #1 was located just SSW of the four corners of old downtown Osborne Bridge. Its file number is 049 and its GPS (Global Positioning Satellite) coordinates are $N(43^{\circ}-10'-801'')$ and $W(74^{\circ}-08'-998'')$. The elevation was approximately 766' above sea level. The site was also just north or on the edge of an end moraine, which was revealed further into the investigation of the study area. Every time we encountered any coarse soils/sediments, the cutting shoe plugged to a certain degree and the core recoveries were never complete, such was the case with Core #1. The probe moved down through 6" of gravel which could have been the remnants of an old road or at least the edge of one. This road could have been the road from Osborne Bridge to Northampton. The old New York State Water Commission Map #20 had not been cleaned-up at that time and no visible sign of an old road was detected at the site. We felt we were on the former lands of either J. Paul or E. Hammond (see Figure 1.5). The probe moved down through another 8" of developed soil horizon that was also gravely. Both of the above horizons were sampled together for further testing. Then we hit some very coarse gravely clay with stones as large as the cutting shoe, about 5" into it. Beyond this, the probe was unable to collect any more sediment for the next 18". The second section continued to

push aside the same sediments until it rounded and swallowed the obstructing cobble, more than 16" further down. It is believed that this horizon was a remnant braided or lateral stream flowing parallel to the adjacent end moraine. The next 13" were the same coarse gravely clay with smaller cobbles, 1-1 ½" at the largest. The last 18-19" of the first 8' push probe comprised a coated gravely clay. No attempt will be made here to predict the parent environment of this coated horizon but it did stand out from the background with further analysis and was completely saturated and assumed to be the water table. This was not hard to imagine since the site was the closest to the Sacandaga River, flowing just to the Northeast.

Core #2 was located just *NW* of Osborne Bridge on the former lands of Foster

Noyes (see Figure 1.5). It was located on the banks of the Sacandaga River Floodplain, which is pervasive north to the Northville Bridge, even at low water status. Its file number is 048 and its GPS coordinates are *N* (43°-10′-865″) and *W* (74°-09′-809″). The elevation was approximately 765′ above sea level. The top 5″ was a very heavy sand indicative of a sandbar, but may also have been involved with ice rafted clast mechanics during spring melt ice jams due to the presence of interspersed pebbles. Next was a small 3″ layer of a developed soil horizon, often referred to as the A-horizon. The probe then pushed through 14″ of light gravel with black sands and another 10″ of heavy gravel with black sands. All of the gravel horizons above could be considered glaciofluvial in nature, related to the annual melts on the front side of a retreating glacier. The bottom of the first section of probe core was approximately 1½ -2″ of fine sand. The top of the next 4′ probe had 2″ of fine gravel with black sands. The next horizon down was 33″ thick and contained alternating layers of light and dark sands with black sands. We first

determined them to be varved sands <50 µm, deposited in quiet waters within one year, adjacent to a glacier. Upon later investigation, after running a few Ostracode Studies in both light and dark layers, questions were raised as to the real interpretation. Below this very fine sand horizon lay 7" of fine sand with black sands, 4" of fine sand without black sands, 7" more of medium sand with black sands and 4" of clay and fine sand with black sands. The bottom 36" of this 12" probed core was fine sand with alternating bands of interspersed black sands. The water table (saturation zone) was encountered at 6'. The last two 4' sections are indicative of beach or shoreline deposits alternating with quiet water deposition in a lake basin or floodplain influenced by a persistent river. It implies that all of these variations in sediment deposition may have already existed before glaciation or that they arrived after a glacial retreat. Another possibility exists, which is that they could be deformed glaciofluvial sediments of a former outwash plain. This poses more questions, which are talked about throughout the study. Core site #2 revealed a wide variety of core horizons and subsequently, several horizons were examined more thoroughly in this core than in any other. Nearly the whole of the upper 3 feet of the core horizons are probably the result of postglacial, fluvial deposition. Everything below 3' was consistent with a subglacial environment, typical of a kettle lake deposition.

Core #3 was located just E of Mead Island on the former lands of L. D. Fritcher (see Figure 1.5). It was at an elevation of about 765' above sea level. Its file number is 050 and its GPS coordinates are N (43°-10'-821") and W (74°-09'-791"). The site was formerly farmed and a few small stone piles existed adjacent to Mead Island. The top 11" were a combination of a developing soil horizon and heavy beach sands at the uppermost few inches. This may be due to the nearness of the shoreline of Mead Island,

but it is clear, Mr. Fritcher farmed right up to the shore/tree line in his day. As the probe continued down the next 9" it encountered more heavy sand with black sands and small amounts of clay in the matrix. The next 3" contained gravel and black sands followed by another 3" of clayey gravel with black sands. The bottom 10" of the first section was again heavy sand and gravel with black sands. These soils were coarse enough to again clog the cutting shoe of the probe and another incomplete core was acquired. More gravel, perhaps a little lighter, met the advance of the second probe length and until the cutting shoe chugged its way through this 18" horizon, more than 15" was lost. The bottom 15" were again clayey gravel with black sands and was in the saturated zone. It's never clear where the probe core-cutting shoe gets hung up but the depth of the core bottom is still established. Core #3 is interpreted as stoss and lee topography where lodgement till slopes around an obstacle such as the Mead Island drumlin. This is likely subglacial in nature and could be another clue that everywhere within the study area the sediments could be simply made up of ice-contact deformations.

Core #4 was located at the NW corner of Mead Island adjacent to a backwater inlet that faces the Sacandaga River floodplain to the east. At first glance, the inlet appears to be a streambed, but no upstream streambed or water flowing into it exists. Another such structure exists less than ½ mile north at the Northampton Beach State Campsite, which is used as a boat launch. Both of these structures are, in fact, relict streambed outlets and probably connected to the West. There the structures drained a kettle, which today is the Tamarack Swamp, due west about 1 mile. Kettles are subglacial happenings and further evidence that the study area is probably made up of ice-contact deformations.

Core #4's file number is 047 and its GPS coordinates are $N(43^{\circ}-10'-715'')$ and $W(74^{\circ}-10'-252'')$. Its elevation was precisely at the high water mark of 768' above sea level. The site was on the former lands of Charles and Arietta Mead (see Figure 1.5). The first 4' of the first core section was medium beach sand with black sands, representative of the high water mark sites. Beneath this was approximately 14" of developed soil horizon. The next horizon down was very heavy gravel with black sands and cobbles large enough to plug the cutting sole, creating a 12" blank in the profile. It finally chewed the cobble up enough to swallow it and continued down through another 10" of light gravel with black sands. The gravel continued to lighten and contained less black sands as the second 4' section captured the top 8'. Somewhere along the push the remaining 36" of light gravel was heavy enough to compress the profile so that the sample was 6" short a complete core section. The same can be said for the next pushed section as it went on to capture 11" of clayey light gravel with black sands, 8" of light gravel with black sands and 2" of red clay light gravel with black sands before hitting the water table at 10'. Somewhere along this section push, probably while capturing the next 18" of medium gravel with black sand the collected sediments were less than complete. The remaining 4' in this third section was heavy sand with heavy black sands.

We had prepared for at least one deeply pushed core section and considered the water table at such a site. After we hit the water table at site #4 we installed a pointed cutting sole on the last 4' before lowering it down the core hole. It was designed to be ridged until the 12' mark was hit and then released to capture the sediments by the core liner in the same way that the sediments are collected above them. It worked well and we

collected the next 14" of heavy sand with black sands. The remaining 16" down to the 16" mark were heavy gravel with up to 1" stones, which gave us an 18" blank in this core capture. This horizon was in a matrix of fine sands and black sands. This core site had horizons that are consistent with streambed sediments and reflect a wandering high water flow subglacial drainage system of the kind mentioned earlier, perhaps such as in draining a kettle.

Core #5 was located WSW of Mead Island on the edge of the largest eroded micro-island of a hummock field in the local greater Mead Islands group. The microisland is usually submerged until late summer. The site was picked just on the eastern edge of the old Osborne Bridge Road and was 765' above sea level. Its file number is 046 and its GPS coordinates are $N(43^{\circ}-10'-527'')$ and $W(74^{\circ}-10'-230'')$. It is unclear by the 1908 New York State Water Commission Maps who the former owner was but it could have been Horace Kenyon. In a straight line, the site was about 900' from core site #4 (see Figure 1.5). The first section behaved like the previous pushes, with gravel hangups at the cutting sole but we collected 80% of the sediments leaving a 10" blank at the top. The top 13" recovered was typical soil developed light gravely loam with black sands. We pushed down through 18" of coated (oxidized) medium gravel with black sands and then another 8" of coated light gravel with black sands. The second section saw the same gravel compression and recovery was absent 8" at the top. The top 7" of this section was the same-coated light gravels with black sands that preceded it. Then we ran into 18" of heavy sand with black sands and some banded fine sand/clay matrix segments. The bottom 18" continued to be heavy sand with small bands of clay matrix

and heavier black sand layers. The hummocky topography of this site is consistent with kame and kettle microenvironments associated with subglacial ice-contact deformations.

Core #6 was located about 800' further SW along the old Osborne Bridge gravel road from site #5 (see Figure 1.5). This site is approximately 760' above sea level and it is unclear who the former owner was, but was likely to be Horace Kenyon. The site was on the east side of the road and appears to be a former dirt road heading northwest to an anonymous unimproved road connected to the Bunker Hill road. Its file number is 045 and its GPS coordinates are $N(43^{\circ}-10'-431'')$ and $W(74^{\circ}-10'-421'')$. Both this site and site #7 are similar regarding their sediment profiles. The top 12" of the first pushed section was developed soil horizon made up of loamy clay. Beneath this was 36" of densely packed fine sand and clay with what appears to be varved banding. The top 18" of the second section was more of the same densely packed fine sand/clay with banding. The next deeper horizon was 24" of fine sand and clay with black sand banding and was completely saturated (water table). The final 6" was fine sand and clay with black sands (the emphasis on the first component in abundance). This sediment profile is consistent with kettle lake environments, which are subglacial and represent ice-soil contact seasonal melting depositions. In retrospective, we should have marked a direction onto the side of the liners at both this site and core #7. That would have given us a better slump profile of these two sites, which were the only sites whose core horizons could be correlated in the study area.

Core #7 was located at a bend in the old Osborne Bridge Road, SW of core #6 by some 1200' (see Figure 1.5). It's not perfectly clear from the maps but this property could have been formerly the same lands of Ed Gifford. This site was also at the 760'

elevation above sea level and on the west side of the road. Its file number was 044 and its GPS coordinates were N (43°-10′-332″) and W (74°-10′-487″). The top 6″ of the probe were probably the gravels of the old roadbed. The next 6″ was a typical clayey loam developed soil horizon. The next 36″ of the first section was densely packed fine sand and clay with varved banding. The top 12″ of the second section was the same densely packed fine clay and varved banding. The remaining 36″ of this 8′ deep core was fine sand and clay with black sand banding and was completely saturated (water table). Again, this sediment profile is consistent with a subglacial kettle lake depositional environment.

Core #8 was sited approximately 800' further SW along the old Osborne Bridge Road. It is on the east side of the road at 764' above sea level and was probably the former lands of Ammon Blowers (see Figure 1.5). Its file number is 043 and its GPS coordinates are $N(43^{\circ}-10'-262'')$ and $W(74^{\circ}-10'-603'')$. The site was chosen because it was an apparent sandbar, which was of interest. The top 16" was deposited in the last 70 years and consists of heavy coated sand with black sands and a little fine sand and clay in the upper few inches. The next 24" in the first section consisted of a well-defined soil horizon of sand and light gravel with black sands. The last 8" was heavy sand and clay with varved banding. The next section of probe compressed somewhere along the push perhaps because of the first 12" of heavy sand and clay, or more likely because of the dark heavy coated sand and alternating clay bands and black sand banding which made up the last 31" of sediments. The last 31" of heavy coated sand horizon was completely saturated (water table). The sediment profile in this 8' core below the top newly formed sandbar represent alternating quiet water and fast water episodes of deposition. The

horizons are interpreted as stoss and lee topography where lodgement till slopes around large obstacles, such as the drumlin just to the *NW* of the site. This drumlin has never been identified, so it is named the *Ambrosino Drumlin*. It has an elevation of 880' above sea level, just southeast of Sweets Crossing. It is not clear as of this writing, but this drumlin appears to be of continental glaciation, rather than of the outlet or valley glacier type.

Core #9 is located at the high water mark (768') on the shoreline of the old Green's Beach Campsites. It would be just south of the old Osborne Bridge Road and on the former lands of Ammon Blowers (see Figure 1.5). Its file number is 042 and its GPS coordinates are $N(43^{\circ}-10'-182'')$ and $W(74^{\circ}-10'-798'')$. The top 8" of the upper 4' section was coarse heavy beach sand with black sand. The next 6" down was coated gravel with black sands. Below that was 18" of loamy heavy coated sand with black sands all showing typical soil development. The last 16" of the upper section were fine sand and clay mixtures, which lightened downward and contained black sands. This section slipped while in the liner or during splitting in half lengthwise. The bottom 4' section compressed at the lowest coarse gravel horizon and resulted in about a 7" blank at the top of the core. In (Figure 2.1) the last core in the rack, to the right of the circled #9, a 4" plug of dark gravel separated and at the top of the core section is evidenced. This is because during handling at the core site, the whole of the contents of the core slipped toward the bottom and protruded by 4". To avoid destroying the integrity of the protruded sediment, it was cut-off and slid down into the blank space at the top of the liner. The upper 6" of this core section consisted of fine sand and clay mixtures and black sands. The next 18" down was made up of medium sand with varved bands of

heavy black sands, especially in the upper half portion of the horizon. The bottom 16" was progressively coarser gravel going downward with black sands. The last 4" was darker due to the zone of saturation (water table). Core #9 was approximately 1200' further SW along the old Osborne Bridge Road and although it has the same profile of quiet water alternating fast water depositional character, the sediments seem to have undergone greater than usual secondary weathering. This is evident in the shades of color in the horizons, indicating groundwater's of totally different charge and suspended solids loads. The horizons are interpreted as stoss and lee topography where lodgement till slopes around large obstacles, perhaps such as an Ambrosino drumlin type.

Appendix C

USGS WATER WELL RECORDS

Date: Mon, 21 Jun 1999 09:38:28 -0400

From: "Ronald V Allen, Hydrologic Technician, Troy, NY " <rallenGusgs.gov>

To: aa7148@cnsvax.albany.edu

Cc: "Ronald V Allen, Hydrologic Technician, Troy, NY " <rallen@usgs.gov>

Subject: GWSI database retrieval, well records.

Arthur Ambrosino:

re: well records within 2-miles of: 431000, 0741000.

The 2-mile retrieval located one record (local # FU 137). I re-searched using a 5-mile radius and picked-up 26 records. They are sorted by ascending latitude.

footnotes:

county, 035 = Fulton

lat-long accuracy, F = within 5-seconds

type analysis, B = common ions

Primary use of water, H = domestic

S = stock

N = industrial

C = commercial

aquifer code, 112TILL = pleistocene till

112SAND = sand

112SDGV = " gravel

371LLFL = Little Falls Dolomite

371PSDM = Potsdam Sandstone

364CNJR = Canajoharie Shale

400PCMB = Precambrian Erathem

lith code, TILL = till

SHLE = shale

DLMT = dolomite

SAND = sand

ROCK = bedrock, undifferentiated

GRVL + gravel

Method constructed, D = dug

C = cable tool

type of finish, W = walled

Please contact me if you have questions.

Ron Allen

U.S. Geological Survey, Water Resources Division

U.S. Geological Survey WRD Pay: (518)285-5601

-+- U.S. Geological Survey, WRD Fax: (518)285-5601
G|S 425 Jordan Road USGS, NY web site,
Troy, NY 12180-8349 http://ny.usgs.gov

1DATE: 06/16/99 GWSI retrieval, well records within 2 miles of, 43 10 00; 074 10 00. PAGE la

LAT ALTITUDE

LAT/	LONG L	ONG	OF LAND	ALTIT	UDE 1	:OPO-
LONGITUE	E DATU	IM a	CCURACY	SURFACE	DAT	LATITUDE TUM GRAPHIC
	L WELL NUM			SORFACE E-ID	COUNT	
(DEGREES			CODE	(FEET)		DE) SETTING
(5201020	, (002	,	CODE	(FEEL)	(00	DE) SETTING
FU 135			4305500	7410590	1 035	430550
0741100	NAD27	F	82		NGVD29	-
FU 148		-	4306040	-		430605
0741000	NAD27	F			NGVD29	-
FU 136		_	4306120			430610
0741030	NAD27	F			NGVD29	-
		•	, 0	0.00	1101023	
FU 147			4306260	7409440	1 035	430625
0740945	NAD27	F			NGVD29	-
		•	,,			
FU 146			4306410	7409220	1 035	430640
0740920	NAD27	F			NGVD29	420040
	.u.u.	•	00	0.00	WG V D Z J	
FU 118			4307280	7414050	1 035	430730
0741405	NAD27	F			NGVD29	430730
0147400	HED2	F	, ,	0.00	NGVD23	_
FU 117			4307400	7414000	1 035	430740
0741410	NAD27	F			NGVD29	430740
0.47410	TADE I	£	04	0.00	NGVD29	-
FU 115			4308150	7414200	1 035	430815
0741420	NAD27	F			NGVD29	420012
0/41420	NAD21	r	/9	0.00	NGVD29	_
FU 116			4300170	7414000	1 035	430815
0741410	NAD27		4308170			430013
FU 114	NAD27	F			NGVD29	420025
0741420	NAD27		4308240			430825
FU 145	NADZ /	F			NGVD29	420045
0740755	M DOT	_	4308430			430845
	NAD27	F			NGVD29	420050
FU 144 0740725	N3 D27	-	4308480			430850
0/40/25	NAD27	F	/91	0.00	NGVD29	-
FU 113			4200100	7442050		420010
	222007	_	4309100			430910
0741325	NAD27	F	791	0.00 1	NGVD29	-
DEL 120						424440
FU 138	*** ***	_	4311110			431110
0741200	NAD27	F			NGVD29	-
FU 108		_	4311330			431135
0741325	NAD27	F			NGVD29	
FU 107		_	4311410			431140
0741315	NAD27	F		_	NGVD29	-
FU 143			4312030		-	431205
0740805	NAD27	F	890	0.00	NGVD29	
				_		
FU 142			43122907		1 035	431230
0740835	NAD27	F			NGVD29	-
FU 141			43125707	7409020:	1 035	431255
0740900	NAD27	F	950	0.00	NGVD29	-
FU 139			43135407	7411220	035	431355
0741120	NAD27	F	750	0.00	NGVD29	-
FU 112			43092507	74131001	L 035	430925
0741310	NAD27	F			NGVD29	-
				-		

FU 111			431026074134601	035	431025
0741345	NAD27	F	900.00 NGVD	29	-
FU 110			431035074134001	035	431035
0741340	NAD27	F	890.00 NGVD	29	-
FU 109			431048074133901	035	431050
0741340	NAD27	F	880.00 NGVD	29	_
FU 137		•	431053074115901	035	431055
0741200	NAD27	F	820.00 NGVD	29	-
FU 140		4:	31355074111101 0	35	431355
	NAD27	F	750.00 NGVD	29	-
1DATE: 06				l, well re	cords within 2
miles of,	43 10 00;	074 10	00.	PAGE	1b

DEPTH	DEPTH		DDD##1	PRIMARY BOTTOM	DIAMETER		T 7/00V
TO	OT	METHOD	DEPTH TYPE	USE OF	TYPE OF		LITH-
TOP	BOTTOM	CONST-	OF WELL	OF CASING	OF CASING	AQUIFER	
(FEET)	WELL NUI (FEET)	RUCTED	(FEET) FINISH	WATER (FEET)	ANALYSES (IN)	CODE	CODE
FU 135	26	D	26	Н		112TILL	TILL
FU 148	20	_	₩ 67.0	54 H		364CNJR	SHLE
FU 136		С	- 65.0	н	6.00		
		-	-			364CNJR	SHLE
18.0 FU 14 7		С	- 53.0		6.00		
		С	-		6.00		
FU 146	~~	_	117	н			
		С			B 6.00		
FU 118			134	н			
		-	-		В	371LLFL	DLMT
114 FU 117		С	130	s	6.00	371LLFL	DLMT
13.0		С	-		6.00		
FU 115		_	77.0 -	H 			
		С	_		 6.00	112SAND	SAND
FU 116		c	99.0	н	6.00	112SAND	SAND
FU 114			109	н			
			-		В	400PCMB	ROCK
82.0		С	~		6.00		
FU 145		D	10.0	H 	24.0	112SDGV	GRVL
FU 144		-	42.0	Н			
		С	-		 6.00	371LLFL	D LMT
FU 113			73.0	н			
		-	-			112SAND	SAND
		С	-		6.00		

FU 112		_	41.0	Н		112TILL	TILL
FU 111		С	40.0		6.00	112TILL	TILL
	-	а	40.0		48.0	11211111	TILL
FU 110			65.0		40.0	371LLFL	DLMT
		С	-	"	6.00	3,10010	DIMI
FU 109		_	19.0	н		112TILL	TILL
		D	-		20.0		
FU 137	•		186	N			
	-	-	-				
		С	-		8.00		
FU 138		_	140	C		371LLFL	DLMT
15.0		С			6.00		
FU 108		_	38.0	H		112TILL	TILL
FU 107		D	-				
FU 107		D	12.0	Н	36.0	112TILL	TILL
FU 143		D	-		36.0		
		_	14.0	H			
		_	-				
						112SAND	SAND
		Ď	_			11201212	21112
FU 142		_	18.0	н		112SAND	SAND
		D	-		24.0		
FU 141			62.0	S			
		-	-				
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3.00		С	-		3.00		
FU 139			128	C			
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					_	1100000	~~
		С			B 6.00	112SDGV	GRVL
FU 140		L	125		6.00	4000000	DOCT
4.00		С	125	H	6.00	400PCMB	ROCK
1DATE: 06	/16/99	C		retries		records with	nin 2
miles of,		0: 074 10		Terries	PAGE	lecords with	1111 6
		-, o, - 10			FAGE	10	

	DISCHARGE
LOCAL WELL NUMBER	REMARK
OWNER	(GPM)
	(0222)
FU 135	100 GPH, HARD WATER, TEMP 51
NYSZCZY, NICK	77
FU 148	FROM BULL.GW-24.; 30= AVG.ANN.WL.
HUGHES JOHN	12.0
FU 136	PUMPED AT 6 GPM
	TOMPED AT 0 GPM
	FROM BULL.GW-24. ; 30= AVG.ANN.WL.
MINKLER HARRY	6.00
FU 147	
PELCHER JOE	FROM BULL.GW-24.; 30= AVG.ANN.WL.
PELCHER JUE	
FU 146	Formerly site 430643074092201.
	
	FROM BULL.GW-24.; 30= AVG.ANN.WL.
FYFE JAMES	~ ~
FU 118	Formerly site 430723074140601.
No Marri	
	FROM BULL.GW-24. ; 30= AVG.ANN.WL.
SART LAUREN	5.00

FU 117 FROM BULL.GW-24.; 30= AVG.ANN.WL. WARNER V. J. 6.00 FU 115 Formerly site 430816074142001. FROM BULL.GW-24.; 30= AVG.ANN.WL. VAN NOSTRAND DONALD FU 116 FROM BULL.GW-24.; 30= AVG.ANN.WL. WOJSKI A. 4.00 FU 114 Formerly site 430826074141001. FROM BULL.GW-24.; 30= AVG.ANN.WL. ZEA EARL 2.50 FROM BULL.GW-24. ; 30= AVG.ANN.WL. FU 145 BOICE KATHERN 3.00 FU 144 Formerly site 430849074073701. FROM BULL.GW-24.; 30= AVG.ANN.WL. HAYDEN RALPH FU 113 Formerly site 430912074131901. FROM BULL.GW-24. FERENCY ESTHER FU 112 FROM BULL.GW-24.; 30= AVG.ANN.WL. WARNER BURDETTE 7.50 FU 111 FROM BULL.GW-24.; 30= AVG.ANN.WL. WARNER CYRUS 3.00 FU 109 FROM BULL.GW-24. ; 30= AVG.ANN.WL. STEWART EDITH FU 137 NOT USED AT PRESENT FROM BULL.GW-24.; 30= AVG.ANN.WL. U. S. FELDSPAR CORP. 40.0 FU 138 FROM BULL.GW-24.; 30= AVG.ANN.WL. FERGUSON AND MITTLSTAEDT 5.00 FU 108 FROM BULL.GW-24.; 30= AVG.ANN.WL. SHAW ED FU 107 FROM BULL.GW-24.; 30= AVG.ANN.WL. EDWARDS PALMER FU 143 Formerly site 431203074081001. FROM BULL.GW-24.; 30= AVG.ANN.WL. DARLING M. FU 142 FROM BULL.GW-24.; 30= AVG.ANN.WL. MADISON ANDY FU 141 Formerly site 431255074090301. FROM BULL.GW-24.; 30= AVG.ANN.WL. LINDSEY M. H. 3.00 FU 139 Formerly site 431353074112501. FROM BULL.GW-24.; 30= AVG.ANN.WL. HOBBS WILLIAM E. 13.0 FU 140 FROM BULL.GW-24.; 30= AVG.ANN.WL. STANDHARDT FRED 1.00

Appendix D

ALS CHEMEX ANALYSES AND MINERALOGY



Chemex Labs, Inc. Analytical Chemists ' Geochemists ' Registered Assayers 994 Glendale Ave., Unit 3, Sparks Nevada, U.S.A. Sparks PHONE: 775-336-5395 FAX: 775-355-0179

EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE. ALBANY, NEW YORK į.

Comments: ATTN. ARTHUR M. AMBROSINO

CERTIFICATE

A9919903

CO) - EARTH & ATMOSPHERIC SCIENCES 00 10 umples submitted to our lab in Mississauga, ON. ifs report was printed on 22-JUN-1999.

Pulp; prepped on other workorder DESCRIPTION SAMPLE PREPARATION NUMBER SAMPLES -SOE X 299

		ANALYTICAL	ANALYTICAL PROCEDURES		
CHEMEX	NUMBER SAMPLES	DESCRIPTION	МЕТНОВ	DETECTION	UPPER
9301	H	% ICP + IC	ICP	0.03	25.0
9341		ppm: ICP + ICP-MB	ICP-MB	0.1	1000
7076	rı .	Promi ICP + ICP-MS	ICP	27	10000
5056	٦,	ICP + ICP-MS	ICP-MS/ICP	9.05	1000
9306	٠,	pper ICP + ICP-M8	ICP-MS/ICP	0.01	10000
2010	4 -	Co pper ICF + ICF-MB package	ICB-NB/ICP	0.05	200
2000		DE CONTRACT OF THE CONTRACT OF	ICE	0.01	75.0
200		ppm tcr + icr-Mg	ICP-MS	0.01	200
6086	-	TOTAL	ICP-MG	0.02	200
9310	-	TOTAL TOTAL STATE	ICP	(10000
9311		100	TCL-MS/ICE	7. 0	10000
9313	1 ~	tobar ICP + ICP-Ma	101	·	10000
9313	-	DDG: ICP + ICP-MS			000
9315	-1	CP + ICP-M3 Da	a		000
9316	-	ppe ICP	TOP-MS	1 4	7
9317		ppa ICP +	ICP-MB/ICP	9 0	10000
9318	-	ppa: ICP	ICP-MB	0.2	500
9319	н	LI ICP + IC	ICP	0.01	15.00
9320	н,	Ppm ICP + ICP-MS	ICP	ın	10000
222	н,	Pres ICP + ICP-NS	ICP	0.0	10000
22.0	٦,	ppm: ICP +	ICP-MS/ICP	0.7	10000
100	٠,	Dom: ICE	ICP-MB	6 .5	200
9326	1.	r ppm: 1CF + 1CF-KS package	104	2	10000
9326	٠.	THE TOTAL TO	HOP	0.01	10.00
9327	٠,		TCL-MR		200
9328	1	No Sec Ind a Transaction	ICE-MB/ICE	0.05	100.0
9329	-		101	0.01	10.00
9330	-	Dra. ICB - TCB-Me	101/201 101 101	N :	10000
9331	-	TOP + TOP - NO	201 AUT	60.0	100.0
9332	-	TOP - TOP - TOP - NO	107-101	0.0 0.0	200
0113		The tot to the	TCK-MB	0.03	200
71.0	• •		10F-10	0.7	200
200	4,-		ICP	0.01	10.00
200	•	ppm icr + icr-Ms	ICP-NS/ICP	0.1	10000
200	٠.	Ding Tole	ICP-MB	٥.۶	200
	٠,	Dom: Tole	ICP	-	10000
	٠.	Month ICP + I	ICP-MS	٥.	200
8278	•	zn ppm: ICP + ICP-MS package	ICP	~	10000
			-		
-					

CFRTIFICATION

3 699			
1-B 1 122-JUN 1991990 :45855 :RCO		Na % (ICP)	2 . 0.7
Page Number 1:B Total Pages 1 Cerditate Date: 22-JUN 1999 Invoice No. 19919903 P.O Number 145855 Account RCO	9	Ag ppm (ICP)	2.3
A P F S C C C C C C C C C C C C C C C C C C	A9919903	Rb ppm (ICP)	• •
:	A	K % (ICP)	3.6
Q	YLYSIS	P ppm (ICP)	C C C C C C C C C C C C C C C C C C C
SCIENCES	OF AN	Nb ppm (ICP)	v. 65
EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE. ALBANY, NEW YORK 12222 SI GSL SI GSL REGISSINO	CERTIFICATE OF ANALYSIS	Ni ppm (ICP)	æ vi
	CERTIF	No ppm (ICP)	1.2
To: EARTH SUNY, 1400 W ALBAN 1222 Project: Comments:		MD ppm (ICP)	730
		Mg & (ICP)	0.53
Ssayers sayers sarks sarks 1179		Li ppm (ICP)	34.3
Pegistered As Registered As Sp Sp Sp Sp Sp Sp Sp Sp Sp Sp Sp Sp Sp		Pb ppm (ICP)	18.0
ochemists Unit 3.		La ppm (ICP)	81.0
Chemex Labs, Inc. Analytical Chemists 'Geothemists' Registered Assayers 994 Glendale Ave., Unit 3. Sparks Nevada, U.S.A. 89431 PHONE: 775 356-5395 FAX: 775-355-0179		Fe % (ICP)	2.39
Analytical of 994 GI Nevad PHON		PREP CODE	0668
		SAMPLE	0101500
	_		

3				
1.C 22-JUN 1991990 45855 RCO				
Page Number 1.C Total Pages 1 Certificate Date: 22.JUN 1999 Invoice No. 19919903 P.O. Number 45655 Account RCO				
Page Total Certifi Invok P.O.	A9919903			
	A 96	Zn ppm (ICP)	3	
:	ANALYSIS	Y ppm Zn (ICP)	37 30.7	
CIENCES	OF ANA!	V ppm Y (1CP)	37 EEBT	
EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE. 12222 t: G.S.L nemis: ATTN: ARTHUR M. AMBROSINO			0 .	
~~>~	CERTIFICATE	W ppm U	0.7	
To. EARTH SUNY 1400 V 1400 V ALBAN 12222 Project : Comments:		Ti % H	0.33	
		Th ppm 1	0.00	
DC.		Tl ppm 7	99	
lbS, legistored Asi Spa Spa Spa Spa Spa Spa		Te ppm (ICP)	50.0	
Cochemists - F. Unit 3.		Ta ppm (ICP)	05.0	
Chemex Labs, Inc. Analytical Chemists' Geochemists' Registered Assayers 994 Glendalle Ave. Unit 3. Nevada, U.S.A. 89431 PHONE: 775-356-5395 FAX: 775-355-0179		Sr ppm (ICP)	366	
Che Analytical C 994 Gb Nevad PHONI		PREP		
<u>(</u>		877		
4		SAMPLE	0101500	



Chemex Labs, Inc.
Analytical Chemists Geochemists Registered Assayers
994 Glendale Ave., Unit 3.
Sparks
Nevada, U.S.A.
89431
PHONE: 775.386 5395 FAX: 775.355.0179

EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE. ALBANY, NEW YORK ပ္

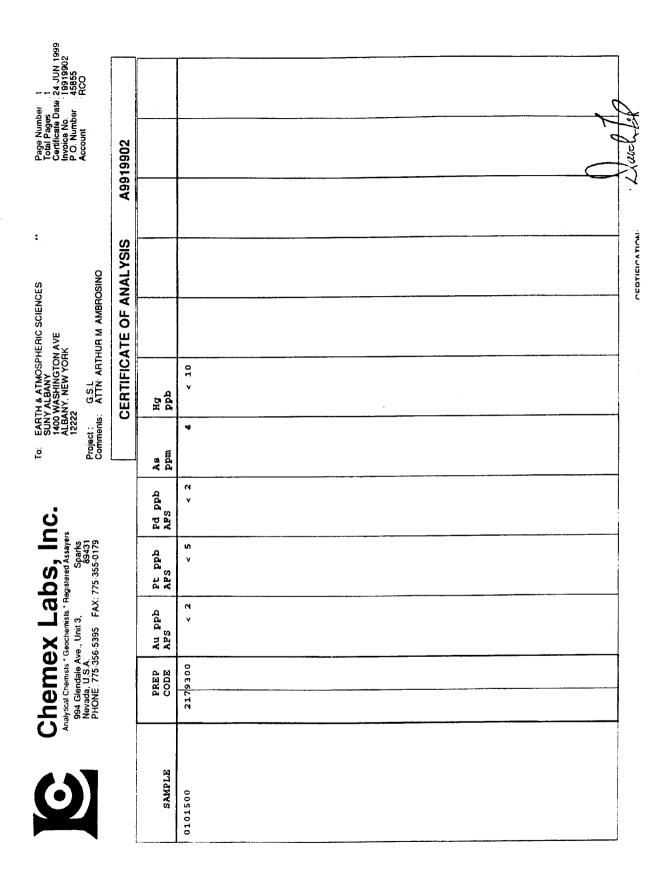
Comments: ATTN: ARTHUR M AMBROSINO

A9919902 CERTIFICATE

(CO) - EARTH & ATMOSPHERIC SCIENCES G.S.L 45855 olect Olect umples submitted to our lab in Mississauga, ON. if report was printed on 24-JUN-1999.

Geochem ring entire sample T127 special code DESCRIPTION SAMPLE PREPARATION NUMBER SAMPLES -- **--**NEX NO DE 217

	UPPER	10000 10000 10000 10000 10000	o invest ased on or by veilable
	DETECTION LIMIT	<i>ଜଳ</i> ଜଳ ପ୍ର	d. Any decision to be defermined by appetitive finwator grada State in the first
OCEDURES	METHOD	PA-ICP-AFS PA-ICP-AFS PA-ICP-AFS AS-HYDRIDE/EDL AAS-PLAMELESS	nt of the semple submitte the claim or depositives tals collected by the pro
ANALYTICAL PROCEDURES	DESCRIPTION	ICP-fluorescence package ICP-fluorescence package ICP-fluorescence package HNO3-aqua regia digest HNO3-HCl digestion	The results of this assay were based solely upon the content of the sample submitted. Any decision to invest should be made only after the potential investment value of the claim or deposit has been determined based on the results of assays of multiple samples or geodotic materials collected by the prospective investor or by a qualified person selected by hinther and based on an evaluation of all engineering data which is available. Statement required by Navada State Law NRS 519
		Au ppb: 1	s assay we will stee the same of mu
;	NUMBER SAMPLES	палан	of the part of the
	CHEMEX	977 8 77 8 77 8 77 8 77 8 77 8 77 8 77	The result to the result the result to the result to the result the result the result to t





Chemex Labs, Inc.

Analytical Chemists * Geochemists * Registered Assayers 994 Glendala Ave.. Unit 3, Sparks Nevada, U.S.A. 89431 PHONE: 775-356-5395 FAX: 775-355-0179

EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE. ALBANY, NEW YORK ë

Comments: ATTN: ARTHUR M. AMBROSINO

A9919911 CERTIFICATE

(CO) - EARTH & ATMOSPHERIC SCIENCES G.S.L 45856 00 # C

umples submitted to our lab in Mississauga, ON. its report was printed on 22-JUN-1999.

SAMPLE PREPARATION	DESCRIPTION	Pulp; prepped on other workorder	
SAM	HEMEX NUMBER SAMPLES	60	
	JEMEX	299	

The results of this easay were based solely upon the content of the sample submitted. Any decision to invest should be made only after the potential investment value of the claim or deposit has been determined based on the results of assays of mustiple samples of geologic materials collected by the prospective investor to by a qualified person selected by him her and based on an evaluation of all engineering data which is available concerning any proposed project.

Chemex Labs, Inc. Analytical Chemists Geochemists Registered Assayers 994 Glendale Ave., Unit 3, Sparks Nevada, U.S.A 89431 PHONE: 775-356-5395 FAX: 775-355-0179

EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE ALBANY, NEW YORK ë

Project: G.S.L Comments: ATTN: ARTHUR M. AMBROSINO

Page Number: 1.A
Tolal Pages: 1
Certificate Date: 22-JUN 1999
Invoice No. 19919911
P.O. Number: 45856
Account RCO

Ge ppm (ICP)	HOHHH NOUGH	ं ज ज ल ल ल
Ga ppm (ICP)	18.9 9.8 17.3 16.9	13.9
	लि चच लाल ४	пни
Co pps (ICP)	8.23 8.88 8.90 9.60	6 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4
Cr ppm (ICP)	294 273 273 260 166	269 259
Cs ppm (ICP)	0.90 0.50 0.60 0.55 0.85	000 9.00 000
Ce ppm (ICP)	52.7 36.2 45.5 56.8 62.0	м д д п с д д ц д
Ca % (ICP)	22.40 22.40 22.60 24.60	4.4.4.00000000000000000000000000000000
cd ppm (ICP)	0.50 0.14 0.36 0.38	00.00 00.00 00.00
Bi ppm (ICP)	0.00	0.00 0.00 0.03
Be ppm (ICP)	1.55 1.40 1.55 1.55	11.30 1.30 . 55
Ba ppm (ICP)	490 400 510 490 590	6 6 8 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
Sb ppm (ICP)	0.0 0.1 0.1 1.0 0.1	0 · · · · · · · · · · · · · · · · · · ·
A1 % (ICP)	6.50 5.99 6.82 7.50	5 5 6 6 5 5 6 6 6 6 6 6 6 6 6 6 6 6 6 6
PREP	2999	-
SAMPLE	0491832 0427995 0425579 042333	0487092 0480509 0480505 0480505
	PREP A1 % Sb ppm Ba ppm Bi ppm Cd ppm Ca % Ce ppm Cr ppm Cr ppm Cr ppm Ga ppm Ga ppm Cr ppm (ICP)	PREP A1 % Sb ppm Ba ppm Ba ppm (ICP)

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CERTIFICATION.

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Page Number 1.C Total Pages 1. Certificate Date: 22.UN 1999 Invoice No 19919911 P.O. Number 48856 Account RCO				
ge Number tal Pages rtificate Dat roice No. O. Number count				70
20528	A9919911			
;		Zn ppm (ICP)	96 60 50 50 50 50 50 50 50 50 50 50 50 50 50	A & L
, C	ALYSIS	Y ppm (ICP)	32.9 26.6 33.9 32.6	8 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2
SCIENCES	OF ANALYSIS	V ppm (ICP)	105 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8	9 5 4 9 5 5 1
EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY ALBANY, ALGANY, NEW YORK 12222 G.S.L. G.S.L. ATTN ARTHIRM AMBROSING	CERTIFICATE	U ppm (ICP)	100011	0 1 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
T ~ 2 ~		W ppm (ICP)	40440	0 1 0
To: EARTI SUNY 1400 V ALBAN 12222 Project: Comments		ri & (ICP)	10000 0.50 0.50 0.50 0.50 0.50 0.50	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
•		Th ppm (ICP)	7.44NN	4 N 4
savers arks 9431		Tl ppm (ICP)	0 0 0 0 0 4 0 0 0 0 0 0 0 0 0 0 0 0 0 0	000
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Chemex Labs, Inc. Analytical Chemists · Goochomists · Registered Assayers 994 Glendale Ave. Unit 3. 994 Glendale Ave. Unit 3. PHONE: 775-356-5395 FAX: 775-356-0178		Ta ppm (ICP)	1.05 0.50 0.70 11.25 0.75	0.50 0.70 0.75
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Chemex Labs, Inc.

Analytical Chemists * Geochemists * Reg stered Assayers 994 Glendale Ave. Unit 3, Sparks Newda U.S.A. 89431 PHONE: 775-366-5395 FAX: 776-365-0179

To EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE ALBANY. NEW YORK 12222

Comments: ATTN ARTHUR M AMBROSINO

ANALYTICAL PROCEDURES

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mples submitted to our lab in Mississauga, ON. is report was printed on 24-JUN-1999.

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SAMPLE PREPARATION	DESCRIPTION
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Geochem ring entire sample T127 special code

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The results of this assay were based solely upon the content of the sample submitted. Any decision to invest should be made only after the potential investment value of the claim or deposit has been determined based on the results of assays of mutiple samples of geologic materials collected. By the prospective investor or by a qualified person selected by him/her and based on an evaluation of all engineering data which is available concerning say proposed project.

Page Number 1 Total Pages 1 Certificate Date 24-UNN-1999 Invoice No. 19919909 P.O. Number 45856 Account RCO	A9919909			
& ATMOSPHERIC SCIENCES "BANY ASHINGTON AVE Y. NEW YORK G.S.L ATTN: ARTHUR M AMBROSINO	TE OF ANALYSIS			
To. EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE. 12222 Project: G.S.L Comments: ATTN: ARTHUR M AMBROSIN	CERTIFICATE	As Hg ppm ppb	, 4, 4, 10, 10, 10, 10, 10, 10, 10, 10, 10, 10	A A B D D D D D D D D D D D D D D D D D
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Chemex Labs, Inc. Analytical Chemists 'Geochemists' Registerd Assayers 994 Glendale Ave., Unit 3. Sparks Nevada, U.S.A. PHONE: 775-355-0179		PREP CODE	2179300 2179300 2179300 2179300 2179300	2179300 2179300 2179300 9300
		SAMPLE	0491832 0427995 0425579 048233	0487092 0480509 04800005



Chemex Labs, Inc.

Analytical Chemists* Geochemists* Registered Åssayers 994 Glendale Ave., Unit 3, Sparks Nevada, U.S.A. 89431 PHONE: 775-356-5395 FAX: 775-355-0179

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Comments: ATTN. ARTHUR M. AMBROSINO

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The results of this seasy were based solely upon the content of the sample submitted. Any decision to invest should be made only after the potential investment value of the claim or deposit has been determined based on the results of seasys of multiple samples of geologic matchals collected by the prospective investor or by a qualified person selected by him her and based on an evaluation of all engineering data which is evaluation concerning any proposed project

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. AMBROSINO	
G.S.L ATTN: ARTHUR M	
Project: Comments:	

Ge ppm (ICP)	44460	. 4.0.4.4.	566.44	च च च - : । तिल्ल
Ga ppm (ICP)	20.5 17.5 16.9 17.2	21.0 18.7 17.9 19.5	19.8 28.3 15.7 20.3	100.01
(doi)	40054	4 C L 9 9	15 77	∞ →
Co ppm (ICP)	9.2 10.0 8.2 7.8 10.0	10.0 11.2 17.2 10.8	19.4 7.4 11.8 8.6	ता. व. च.
Cr ppm (ICP)	2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2	214 271 355 223 275	247 56 366 219 172	4222 4221 4331
Cs ppm (ICP)	1.25 0.65 0.90 1.30	1.05 0.75 0.55 0.80	0.70 3.65 0.85 0.90	0111
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Ca % (ICP)	1.61 2.08 1.26 1.56	2.5.4 2.0.4 2.0.4 2.0.5 3.0.6 3.0.6	1.92 1.35 1.65 1.65 1.65	8.1.1 1.9.3 1.9.3
cd ppm	0.32 0.22 0.24 0.14 0.34	00.00	00.30	0.38
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Be ppm (ICP)	2.65 1.75 2.10 1.65 2.70	1.90 1.30 2.25 0.05	1.75 3.70 2.00 2.50	1.90
Ba ppm (ICP)	570 500 4 90 500 580	540 340 570 550	630 670 570 570	0 4 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8 8
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Chemex Labs, Inc. Analytical Chemists 'Geochemists' Registered Assayers 894 Gelendale Ave. Unit 3. Sparks Nevada, U.S.A. 88431 PHONE: 775-356-5395 FAX: 775-355-0179

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Project: G.S.L Comments: ATTN: ARTHUR M. AMBROSINO

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	Mg % (ICP)	0.56 0.52 0.59 0.59	0.59 0.84 1.18 0.76	0.71 0.94 0.71 0.60	0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
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	La ppm (ICP)	20.5 21.0 19.0 26.0 32.5	22.0 3.0.0 19.0	29.5 57.5 16.0 24.0	4444 3446 3.0.8
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	ri k (ICP)	0.52 0.60 0.41 0.30	0.00 1	0.57	000 000 000 000
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	Sr ppm (ICP)	226 249 167.0 203 175.0	230 153.5 226.5 226.5 206 226.5		251 251 251 251 251 251 251
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CERTIFICATION:



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Comments: ATTN: ARTHUR M. AMBROSINO

CERTIFICATE

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OVERBURDEN DRULING MANAGEMENT LIMITED

September 07, 2000

Mr. Arthur M. Ambrosino Trackside I, Apt. 3C2 Johnstown, NY U.S.A. 12095

Fax: (514) 442-5825

Dear Mr. Ambrosino:

e: Heavy Mineral Suite of Estuary Sand Sample AA-01, New York

We have completed our heavy mineral processing of the above composited drill core sample and are pleased to report the results.

We extracted the -2.0 mm heavy minerals using a shaking table followed by a heavy liquid separation at S.G. 3.2 (Table 1), performed a ferromagnetic separation to remove magnetite, and sieved off two nonferromagnetic size fractions — medium (0.25-0.5 mm) and fine (0.125-0.25 mm) sand — for mineral logging. To ensure that we would be able to log all of the minerals accurately, we checked their paramagnetic properties beforehand and analyzed representative grains by energy dispersive x-ray spectrometry (EDS; Table 2) using our scanning electron microscope (SEM). These analyses were done on unpolished grains and, although reliable, were not sufficiently precise to be recorded for publication.

The medium and fine sand sized heavy minerals were found to contain 7.1 and 7.6 weight percent magnetite respectively (Table 1). The percentages of nonferromagnetic minerals were determined by counting grains and are therefore expressed in grain percent of the nonferromagnetic fraction, not weight percent of the total concentrate. The minerals were scattered thinly over a lined plate and 200 contiguous grains were counted along the lines. The minimum count of one grain therefore equals 0.5 percent (Table 3).

The major minerals present are almandine, hornblende, two clinopyroxenes (brown to dark green augite and pale green to colourless diopside), orthopyroxene (dark to pale brown hypersthene, bronzite and enstatite) and ilmenite. Minor minerals occurring at the 0.5-2 percent level are titanite, leucoxene, apatite, zircon and goethite. Only a low trace of epidote was observed and hematite, rutile, kyanite and staurolite are conspicuously absent. The general abundance of pyroxene and ilmenite suggests major inputs from a mafic plutonic terrane. If this is the case, the unexamined 1-2 mm fraction of our heavy mineral concentrate may contain gabbro grains.

September 07, 2000

Page 2 Mr. Arthur M. Ambrosino Heavy Mineral Suite of Estuary Sand Sample AA-01

The fine sand fraction contains significantly more orthopyroxene (16.5 versus 7 percent) and correspondingly less clinopyroxene (17.5 versus 24.5 percent) than the medium sand fraction (Table 3), possibly because the orthopyroxene occurs in less equant, splintery grains, but the proportions of the other major and minor nonferromagnetic minerals and of magnetite (Table 1) are similar. Therefore it may be assumed that the minerals of the -0.125 mm suite, which are too fine-grained to log, are also similar.

I hope these observations are helpful. Please call me if you have any questions.

Sincerely,

Stuart Averill, President

Bulk Sample Tabling Weights and Descriptions	Sample Description	Madrix <2.0 mm	Percentage Colour	GR 15 OT SO SD ST CY SAND CLAY R CLASS	O O U + Y - B B SILIY SAND
Description	Sample Descr			OX.	
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		Clasts >2.0 mm	Porce	ğ	84 82 02 8.0 P 10 90
anomple J		2		V/S	2
Bulk S				5	Δ.
				Table 1	0.8
	-	Weight (kg)		Commercial	0.2
		Weigh		Toble Split	8.2
				Buk Rec'd	4.8
				Anyle Number)O-VV

HMC Processing Weights

							2	< 2.0 mm Table Concentrate	Concell							1
							H	vy Liquie	Separate	OR S.G.3	2					
					Herrette	rremagnetic Praction (Com (g)					Nonferron	Hypertic Fr	retion (g)		
Sample Number	Total	Herry	E S	40.125 alle	0.125 to 0.25 com	W1% af 0.125 to 0.25 tim HMC	0.25 to	West of 0.25 to 0.5 to 19MC	0.5-2.0 nom	Top]	<0.125 seen (what)	40.106 BBB	0.25 to	0.25 to 0.5 tum	0.5 to 1.0 area	1.0-2.0 com
10-44	AA-01 1276.4 696.6	9969	4.4	17.7	152	7.5%	23	7.1%	3.2	535.4	38.0	168.1	187.8	108.6	28.3	4.6

Table 1 - Sample weights and descriptions.

No. of Grains	Description	Binocular Microscope I.D.	SEM I.D.
1	Dark green	Augite or diopside	Augite
i	Brown	Augite or orthopyroxene	Augite
1	Beige-white	Clinopyroxene	Augite
3	Colourless, elongate	Kyanite or diopside	Diopside
1	Very pale green	Clinopyroxene	Diopside
5	Brown-orange to brown	Orthopyroxene	Hypersthene-bronzite
2	Pale brown to colourless	Enstatite or epidote	Bronzite-enstatite
1	Orange-brown	Leucoxene or goethite	Goethite
1	Colourless to white	Apatite	Apatite

Table 2 - SEM analyses of representative medium sand grains by EDS.

	0.25-0.	5 mm	0.125-0.	25 mm
Mineral	No. of Grains	% of Grains	No. of Grains	% of Grains
Almandine	69	34.5	65	32.5
Homblende	50	25.0	42	21.0
Augite	28	14.0	18	9.0
Diopside	21	10.5	17	8.5
Orthopyroxene	14	7.0	33	16.5
Ilmenite	13	6.5	21	10.5
Titanite	1	0.5	0	0.0
Leucoxene	0	0.0	1	0.5
Apatite	2	1.0	2	1.0
Zircon	1	0.5	0	0.0
Goethite	1	0.5	1	0.5
	200	100	200	100

Table 3 - 200-grain counts of medium and fine sand size fractions.

umples submitted to our lab in Sparks, its report was printed on 29-SRP-2000.

ICO - EARTH & ATMOSPHERIC SCIENCES

G.S.L.

CERTIFICATE

SAMPLE PREPARATION

NUMBER

HEMEX

235

ALS Chemex Chemex Lates are Analysized Chemen 1 Gas strongers 1 Repres area Assumers 994 Generals Ave. Unit 3 Newdon Class Assumers PHONE 775:356 5395 FAX 775:355 0179

EARTH & ATMOSPHERIC SCIENCES SUNY ALBANY 1400 WASHINGTON AVE ALBANY NEW YORK 12222 USA <u>,</u>

A0025777

Comments: ATTN: ARTHUR M. AMBROSINO CC: ARTHUR M. AMBROSINO

ICATE A0025777			ANALYTICAL PROCEDURES	ROCEDURES		
ATMOSPHERIC SCIENCES	CHEMEX	NUMBER SAMPLES	DESCRIPTION	METHOD	DETECTION	UPPER
ed to our lab in Sparka, NV. printed on 29-SEP-2000.	382, 332, 873, 372, 451,		Tio2 %: Percoxide fush-Fe precip Tho2 %: Assay Ce %: Assay La %: Assay Beo %: Acid decomposition Fe %: HClO4-HN03-HP digestion	AAS NAA NAA NAA TITRATION AAS	0.001 0.001 0.001 0.001 0.001	100.0 100.0 100.0 100.0 100.0 100.0
PLE PREPARATION						
DESCRIPTION						
Pan con ring to approx 150 mesh Pulp; prepped on other workcrder						
	a seriasing					
	The res should should the res qualifier concert	ulls of the be made utts of as a person ning any p	The results of this assay were based solely upon the content of the sample submitted. Any decision to invest should be made only after the potential investment value of the claim or deposit has been determined based on the results of assays of multiple samples of geologic materials collected by the prospective investor or by a qualified person selected by him/her and based on an evaluation of all engineering data which is available concerning any proposed project	ent of the sample subrathe claim or deposit he claim or deposit herials collected by the valuation of all enginees Statement required b	nitted. Any decisions been determined processions processions which is need to be not some any Nevada State Lavy Nevada State Lav	n to invest d based on for or by a s available

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A CA		SAMPLE	TIGSL01 THGSL02 YGSL03 CEGSL04 LAGSL05	e Casa

Test T871 - Sacandaga Lake, New York

SAMPLE:	T871A 20101	:	T871B 20102		T871C 20103	
Slimes(%)	18.5		7.7		11.1	
THM (%)	8.5		4.1		4.3	
THM d ₅₀ (μm)	193		190		301	
THM MINERALOGY	Calc.	SEM	Calc	SEM	Calc.	SEM
Titanomag	7.2	1.2	5.9	1.6	5.7	-
Altered Ilmenite		1.2		-		0.8
Ilmenite	16.8	20.8	13.7	14.8	9.4	9.2
Ilm./Silicate/Fe oxide		2.8		1.2		2.0
Rutile	0.3	ا الانتقال المساورة	0.4	-	0.3	0.4
Other Ti Minerals	4.7		4.4		3.7	
Zircon	0.6	0.8	0.6	1.2	0.3	0.4
Monazite	0.2	70.0	0.3	81.2	0.3 80.2	87.2
Other Heavy Minerals	70.1	73.2	74.7	01.2	00.2	07.2
ILMENITE PRODUCT CHEM	<u>IISTRY</u>					
Best M	lag Fractions	MicroProbe	BMF	MP	BMF	MP :
TiO ₂	44.11	47.53	42.00	48.00	37.22	47.28
FeT	37.12	39.04	34.94	38.50	33.90	39.26
Al ₂ 0 ₃	0.68		1.41	7.33 4.7	2.16	
CaO	0.54	0.03	1.26	0.01	1.65	0.09
MgO	0.74	0.51	0.93	0.44	1.58	0.42
P ₂ O ₅	0.02		0.03		0.08	
MnO	0.79	0.47	0.94	0.58	0.81	0.56
SiO ₂	2.81	0.19	6.17	0.28	9.80	0.23
Cr ₂ O ₃	0.04	0.03	0.03	0.03	0.05	0.03
V_2O_5	0.26		0.21		0.24	
ZrO ₂	0.02		0.01		0.01	
ILMENITE PRODUCT (Cond	luctor)					
As % of conductor fraction	54.8		35.7		31.4	
As % of THM	14.6		10.5		7.5	
As % of sample	1.2		0.4		0.3	

Note: MicroProbe average oxide values for 83, 88 and 66 grains respectively.

Appendix E RADIOLOGICAL SAFETY DATA

EG&G ORTEC G V - I (191) WAN32 I4W02.74 07-MAR-00 14:30:38 Page U.C.L.A. Radiation Safety Spectrum name: AN1E202.AN1

Sample description

Monsanite Sand Samples; 03/06/2000

Spectrum Filename: An1E202.An1

Acquisition information

Start time 06-Mar-00 10:34:46

Live time 3600 Real time 3608 Dead time .23%

Detector/Geometry IDs **3** O

Detector system

CHS-GAMMA SPEC1 MCB 9

Calibration

Filename: An1E202.An1

Created: 10-Feb-00 15:58:22 & 15-Feb-00 16:20:38

I-129 Calibration; 10/20/99; charcoal cartridge- 2" from the det

.797 keV; Gain .301 keV/channel Zero offset

Library Files

Main analysis library: C:\USER\WYCHCTER.LIB

Analysis parameters

Start channel 10 for an energy of 3.81keV

Stop channel 8144 for an energy of 2455.19keV

Peak rejection level 50.000%

Activity scaling factor 1.0000E+00/(3.7000E+04* 1.0000E+00)

2.7027

E-05

Detection limit method:

MDA - EG&G ORTEC method

Additional random error: 1.0000000E+00 Additional systematic error: 1.0000000E+00

Background width: average of five points.

Corrections

Status Comments

15-Feb-00 12:00:00 Decay correct to date YES

Pbc
02:21

Energy calibration normalized difference: .1313

* * * * * * * * * * * * * * * * * * *	*** U N	IDENT	IFIED	PEAK	SUMI	MAR	Y ****
PEAK CE CTED	NTROID	BACKGROUND	NET AREA	EFFICIENCY	UNCERT	FWHM	SUSPE
CHANNEL IDE	ENERGY	COUNTS	COUNTS	* AREA	1 SIGMA %	keV	NUCL
+-+-+-+	-+-+-+	-+-+-+-+-	-+-+-+-+	-+-+-+-+-+	-+-+-+-+	+-+-4	+-+-
+-+-+-+							
53.02 D	16.76	1810.	203	5. 6.944E+04	3.70	1.461	
64.34 D	20.17	1702.	. 179	9. 7.238 E+04	4.01	1.467	-
2451.09 D	739.02	184.	198	5. 0.000E+00	2.44	2.394	-
2578.47 D	777.39	210.	89	9. 0.000E+00	4.04	2.419	-

s Peak fails shape tests.

mt.:

D Peak area deconvoluted.

This section based on library: C:\USER\WYCHCTER.LIB

ZN-65 .000D	24.36	8.13	0.	0.	.000	.00	
BI-207	32.56	10.60	1.	0.	.000	.00	
.000D PB-214	33.22	10.80	1025.	0.	.000	.00	1
.449D PB-210	33.22	10.80	1.	0.	.000	123.29	
.000D TH-234	41.53	13.30	912.	2881.	.800	2.38	1
.454D PA-234	42.52	13.60	3713.	278.	.077	31.60	1
.455D Y-88 .456D	44.35	14.15	4157.	379.	.105	24.62	1
CO-57 .288D	45.21	14.41	1343.	0.	.000	.00	
Y-88 .459D	49.83	15.80	4392.	501.	.139	19.25	1
SN-113	77.06	24.00	2544.	1073.	.298	7.32	1
SN-113 .475D	77.76	24.21	3097.	514.	.143	15.95	1
SN-113	88.02	27.30	2767.	799.	.222	9.95	1
SN-113	90.01	27.90	2868.	672.	.187	11.92	1
CS-137	103.03	31.82	2542.	845.	. 235	9.11	1
CS-137	104.26	32.19	3062.	308.	.086	26.05	1
CS-137	118.24	36.40	2467.	836.	. 232	9.09	1
EU-152 .504D	128.61	39.52	2950.	3 59 .	.100	22.05	1
BI-212 .504D	129.73	39.86	3314.	0.	.000	.00	1
EU-152 .505D	130.60	40.12	3065.	319.	.089	25.17	1
EU-155 .778D	137.87	42.31	1726.	0.	.000	.00	
EU-154 .417D	137.87	42.31	674.	0.	.000	162.12	
EU-155	140.16	43.00	1105.	0.	.000	.00	
EU-154 .513D	140.16	43.00	1105.	0.	.000	318.02	
EU-155	147.80	45.30	1994.	0.	.000	.00	
EU-152 .514D	148.13	45.40	3326.	283.	.079	29.38	1
PB-210 .517D	151.85	46.52	3724.	0.	.000	.00	1

EU-152 152.12 46.60 3469. 255. .071 33.21 1 .517D

Nuclide M	Channel	Energy	Background	Net area	Cnts/sec	Uncert	FWH
EU-155	159.09	48.70	3085.	0.	.000	.00	
EU-154	159.09	48.70	2635.	0.	.000	372.03	
EU-154	163.41	50.00	3695.	149.	.041	58.34	1
PB-214	174.04	53.20	3270.	19.	.005	422.44	1
EU-155 .145D	196.65	60.01	5078.	0.	.000	.00	1
PA-234 .547	207.25	63.20	4528.	0.	.000	.00	1
TH-234 .692s	208.35	63.53	2819.	115.	.032	62.53	
TL-208 .565D	239.13	72.80	6538.	725.	.201	16.20	1
BI-207 .478D	239.13	72.80	1854.	0.	.000	44.84	
PB-214 .568D	245.81	74.81	8074.	148.	.041	86.30	1
PB-212 .727D	245.81	74.81	5239.	0.	.000	10736.27	
TL-208 .569D	246.34	74.97	6811.	1373.	.381	8.92	1
BI-207 .447D	246.34	74.97	5199.	0.	.000	10790.98	
PB-214 .572D	253.44	77.11	6128.	158.	.044	70.32	1
PB-212 .000D	253.44	77.11	2558.	0.	.000	491.34	
TL-208 .586D	278.98	84.80	5501.	573.	.159	18.78	1
BI-207 .314D	278.98	84.80	1767.	0.	.000	68.70	
EU-155 .589D	284.46	86.45	6921.	17.	.005	674.48	1
PB-214 .591D	286.95	87.20	6148.	287.	.080	39.02	1
PB-212 .247D	286.95	87.20	4190.	0.	.000	299.61	

BI-207 .217D	287.28	87.30	4549.	0.	.000	.00	
PB-214 .595D	295.59	89.80	6197.	129.	.036	86.67	1
PB-212 .156D	295.59	89.80	1091.	0.	.000	49.66	
U-235 .245D	296.12	89.96	2113.	0.	.000	.00	
TH-234 .600D	304.16	92.38	6127.	106.	.029	104.74	1
TH-234 .601D	305.55	92.80	6013.	231.	.064	47.91	1
U-235 .497D	307.38	93.35	3993.	0.	.000	.00	
PA-234 .604D	311.76	94.67	6158.	64.	.018	173.00	1
PA-234 .611D	324.28	98.44	5718.	58.	.016	184.25	1
AC-228 .612D	327.63	99.45	5720.	0.	.000	.00	1
PA-234 .613D	328.46	99.70	5679.	0.	.000	.00	1
U-235 .000D	346.07	105.00	1977.	0.	.000	.00	
EU-155 .623D	347.10	105.31	5495.	122.	.034	86.63	1
U-235 .000D	359.81	109.14	1255.	0.	.000	.00	
PA-234 .633D	365.99	111.00	5600.	101.	.028	104.87	1
PA-234 .639D	377.61	114.50	5198.	98.	.027	104.95	1
EU-152 .651D	401.79	121.78	4001.	9.	.003	962.13	1
CO-57 .652D	402.75	122.07	4010.	0.	.000	.00	1
EU-154 .654D	406.17	123.10	5062.	39.	.011	256.91	1
PA-234 .657D	413.48	125.30	4688.	0.	.000	.00	1
AC-228 .664D	426.10	129.10	4801.	0.	.000	.00	1
PA-234 .668D	433.34	131.28	4393.	60.	.017	156.81	1
CO-57 .676D	450.44	136.43	4205.	0.	.000	.00	1
FE-59 .687D	471.10	142.65	4212.	48.	.013	192.96	1
U-235	474.78	143.76	3170.	0.	.000	.00	
PA-234	496.00	150.15	2874.	98.	.027	92,28	

U-235 .372s	545.11	164.94	3414.	231.	.064	51.01	
U-235	614.12	185.72	1269.	13.	.004	375.20	1

Nuclide M	Channel	Energy	Background	Net area	Cnts/sec	Uncert	FWH
RA-226 .759D	615.02	185.99	1408.	0.	.000	.00	1
PA-234 .000D	615.05	186.00	1116.	0.	.000	223.58	
FE-59 .769D	636.11	192.34	2019.	106.	.029	60.68	1
PA-234 .782D	664.20	200.80	1342.	0.	.000	.00	1
U-235 .785D	668.58	202.12	1783.	24.	.007	250.65	1
PA-234 .786D	671.17	202.90	2108.	53.	.015	122.21	1
U-235 .790D	679.18	205.31	2249.	10.	.003	650.38	1
AC-228 .796D	692.76	209.40	2126.	123.	.034	53.82	1
PA-234 .290D	750.77	226.87	1306.	26.	.007	235.23	
PA-234 .000D	750.77	226.87	1320.	0.	.000	342.39	
PB-212 .841D	789.82	238.63	1718.	422.	.117	14.72	1
RA-224 .845D	797.69	241.00	2030.	30.	.008	210.26	1
PB-214 .846D	800.74	241.92	1728.	113.	.031	52.96	1
EU-152 .851D	809.88	244.67	1713.	6.	.002	928.67	1
EU-154 .856D	821.07	248.04	1632.	14.	.004	403.50	1
PA-234 .857D	823.92	248.90	1615.	40.	.011	143.11	1
SN-113 .866D	844.31	255.04	1210.	34.	.009	146.92	1
AC-228 .889D	894.98	270.30	1357.	50.	.014	104.80	1
PA-234 .892D	900.96	272.10	1355.	26.	.007	199.97	1

TL-208 .899D	918.42	277.36	1114.	43.	.012	111.40	1
PA-234	972.68	293.70	759.	25.	.007	159.80	1
.923D PB-214	977.73	295.22	960.	86.	.024	52.07	1
.925D PB-212	993.90	300.09	878.	52.	.015	81.16	1
.932D CR-51	1055.46	318.63	1034.	93.	.026	73.27	1
.108s AC-228 .000s	1088.00	328.43	128.	2.	.001	403.70	
AC-228 .985D	1121.10	338.40	774.	134.	.037	30.66	1
EU-152 .993D	1140.69	344.30	907.	35.	.010	124.24	1
PB-214 .613	1165.03	351.63	952.	308.	.086	23.51	1
PA-234 .027D	1224.70	369.60	760.	61.	.017	65.46	2
PA-234 .030D	1233.33	372.20	744.	25.	.007	156.39	2
SN-113 .652s	1301.51	392.74	553.	70.	.019	63.51	
AC-228 .078D	1356.84	409.40	718.	55.	.015	70.23	2
EU-152 .080D	1362.45	411.09	659.	25.	.007	148.44	2
EU-152 .550s	1468.55	443.05	272.	37.	.010	65.15	
PA-234 .137D	1520.19	458.60	368.	20.	.005	139.96	2
AC-228 .142D	1534.80	463.00	639.	51.	.014	70.91	2
CS-134 .253s	1570.95	473.89	267.	0.	.000	2310.84	
PA-234 .190D	1680.22	506.80	707.	51.	.014	75.30	2
TL-208	1693.23	510.72	665.	212.	.059	18.52	2
PA-234 .198D	1702.79	513.60	726.	50.	.014	77.83	2
PA-234 .205D	1726.03	520.60	617.	0.	.000	.00	2
PA-234 .205D	1727.36	521.00	524.	59.	.016	56.45	2
CS-134 .248D	1867.65	563.26	427.	10.	.003	286.50	2
PA-234	1878.41	566.50	501.	22.	.006	145.02	2
.252D PA-234 .254D	1885.71	568.70	518.	19.	. 005	167.38	2

PA-234 .254D	1887.57	569.26	549.	0.	.000	.00	2
PA-234	1887.57	569.26	64.	0.	.000	661.21	
CS-134 .254D	1887.67	569.29	549.	0.	.000	.00	2

Nuclide M	Channel	Energy	Background	Net area	Cnts/sec	Uncert	FWH
BI-207 .255D	1888.93	569.67	580.	0.	.000	.00	2
PA-234 .259D	1903.31	574.00	323.	13.	.004	196.94	2
TL-208	1933.65	583.14	462.	188.	.052	17.76	2
.268D EU-154	1962.07	591.70	273.	16.	.004	150.60	2
.276D PA-234 .286D	1998.92	602.80	268.	7.	.002	329.98	2
CS-134	2005.09	604.66	407.	9.	.003	302.35	2
.287D BI-214	2020.56	609.32	392.	198.	.055	15.79	2
.292D PA-234 .348s	2065.25	622.78	196.	39.	.011	62.17	
PA-234 .330D	2166.23	653.20	380.	20.	.006	137.15	2
CS-137	2194.18	661.62	359.	19.	.005	140.61	2
PA-234	2204.07	664.60	395.	2.	.000	1596.98	2
.339D BI-214	2206.89	665.45	380.	0.	.000	.00	2
.340D PA-234 .341D	2211.04	666.70	380.	20.	.006	138.66	2
PA-234	2221.33	669.80	223.	16.	.005	130.34	2
.343D PA-234	2296.68	692.50	305.	21.	.006	119.13	2
.361D EU-154	2296.68	692.50	112.	0.	.000	127.81	
.550D PA-234	2318.59	699.10	343.	40.	.011	67.68	2
.366D NB-94 .368D	2329.88	702.50	319.	38.	.011	68.70	2

PA-234	2717.90	819.40	176.	13.	.004	151.59	2
.444D							
	2735.49	824.70	182.	20.	.006	98.93	2
446D							
PA-234	2740.80	826.30	183.	0.	.000	.00	2
.447D							
PA-234	2756.74	831.10	138.	10.	.003	168.08	2
.450D							

Nuclide M	Channel	Energy	Background	Net area	Cnts/sec	Uncert	FWH
MN-54 .452D	2769.05	834.81	123.	0.	.000	.00	2
AC-228 .452D	2771.67	835.60	93.	9.	.002	157.88	2
MN-56	2808.18	846.60	105.	0.	.000	.00	2
CO-56 .000s	2805.00	845.64	27.	6.	.002	102.85	
TL-208 .465D	2854.22	860.47	126.	23.	.006	71.73	2
EU-152 .468D	2877.19	867.39	149.	15.	.004	117.11	2
NB-94 .470D	2889.50	871.10	144.	0.	.000	.00	2
EU-154 .471D	2896.47	873.20	145.	0.	.000	.00	2
PA-234 .472D	2908.09	876.70	140.	11.	.003	158.70	2
PA-234 .474D	2920.73	880.51	159.	0.	.000	.00	2
PA-234 .000D	2920.73	880.51	35.	0.	.000	357.07	
PA-234 .475D	2929.79	883.24	135.	21.	.006	81.00	2
SC-46 .407D	2949.77	889.26	56.	0.	.000	.00	
Y-88 .481D	2978.85	898.02	128.	13.	.004	123.37	2
PA-234 .481D	2980.77	898.60	125.	0.	.000	.00	2
AC-228 .051	3020.85	910.68	109.	165.	.046	16.42	2
PA-234 .492D	3067.06	924.60	94.	10.	.003	146.25	2

PA-234 .492D	3074.03	926.70	131.	7.	.002	240.84	2
BI-214 .495D	3098.43	934.05	124.	35.	.010	47.36	2
PA-234 .502s	3136.67	945.57	76.	29.	.008	61.47	
PA-234 .500	3148.04	949.00	90.	0.	.000	.00	2
EU-152 .505D	3197.83	964.00	113.	0.	.000	.00	2
AC-228 .505D	3199.82	964.60	127.	18.	.005	91.98	2
AC-228 .506D	3214.09	968.90	127.	92.	.025	20.28	2
CO-56	3242.50	977.46	86.	0.	.000	.00	
PA-234 .509D	3246.94	978.80	122.	7.	.002	238.01	2
PA-234 .509D	3252.58	980.50	129.	2.	.001	868.99	2
PA-234 .510D	3262.21	983.40	104.	12.	.003	125.87	2
EU-154 .514D	3305.02	996.30	62.	7.	.002	167.90	2
EU-154 .516D	3333.23	1004.80	119.	21.	.006	76.02	2
CO-56 .645D	3442.84	1037.83	29.	0.	.000	.00	
CS-134 .316s	3443.50	1038.03	8.	0.	.000	.00	
BI-207 .404s	3523.81	1062.23	41.	6.	.002	164.58	
RB-86 .528D	3571.60	1076.63	82.	14.	.004	98.29	2
EU-152 .529D	3602.04	1085.80	99.	24.	.007	62.72	2
EU-152 .529D	3615.38	1089.82	110.	26.	.007	60.23	2
FE-59 .530D	3646.57	1099.22	63.	11.	.003	110.19	2
EU-152 .531D	3689.21	1112.07	89.	20.	.005	71.36	2
ZN-65 .531D	3700.66	1115.52	133.	0.	.000	.00	2
BI-214 .531D	3716.45	1120.28	97.	77.	.021	21.44	2
SC-46 .531D	3717.25	1120.52	167.	0.	.000	.00	2
BI-214 .262s	3835.05	1156.02	47.	17.	.005	68.52	
CS-134 .514s	3871.65	1167.05	76.	45.	.013	44.03	
•	•						

CO-60 .301D	3892.16	1173.23	86.	0.	.000	.00	
CO-56	3898.46	1175.13	70.	0.	.000	.00	
EU-152	4025.92	1213.54	76.	14.	.004	124.29	
BI-214 .523D	4107.44	1238.11	121.	40.	.011	42.23	2
CO-56 .252D	4108.01	1238.28	76.	0.	.000	.00	

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Nuclide	Channel	Energy	Background	Net area	Cnts/sec	Uncert	FWH
М	9.14111101	n.c. gy	Buckground	wee area	01103/360	OfficeTc	T. AATT
NA-22 .516D	4228.32	1274.54	43.	0.	.000	.00	2
EU-154 .516D	4229.18	1274.80	41.	9.	.002	109.19	2
BI-214 .515D	4249.62	1280.96	74.	13.	.003	101.32	2
FE-59 .512D	4284.79	1291.56	58.	14.	.004	83.95	2
EU-152 .510D	4310.11	1299.19	53.	15.	.004	73.37	2
CO-60	4415.33	1330.90	54.	39.	.011	48.85	1
PA-234 .493D	4488.64	1353.00	45.	19.	.005	54.11	2
CO-56	4512.60	1360.22	6.	0.	.000	.00	
CS-134 .489D	4528.89	1365.13	33.	10.	.003	84.75	2
NA-24 .488D	4540.24	1368.55	45.	9.	.003	107.68	2
BI-214 .484D	4570.43	1377.65	34.	27.	.008	35.84	2
PA-234 .477D	4625.00	1394.10	25.	7.	.002	116.30	2
EU-154 .476D	4637.28	1397.80	38.	6.	.002	144.67	2
BI-214 .474D	4649.55	1401.50	32.	19.	.005	46.62	2
BI-214 .471D	4671.05	1407.98	22.	26.	.007	31.78	2
EU-152 .471D	4671.38	1408.08	47.	0.	.000	.00	2

PA-234 .449D	4819.09	1452.60	21.	3.	.001	264.09	2
AC-228	4840.98	1459.20	1036.	127.	.035	37.01	2
.446D K-40	4846.12	1460.75	210.	973.	.270	3.83	2
.445D AC-228	4954.90	1493.54	7.	40.	.011	24.05	
.362s BI-214	5016.44	1512.09	0.	34.	.009	17.15	
.452s AC-228	5283.55	1592.61	21.	40.	.011	34.19	
.442s	F200 70	1507.50	_	_			
EU-154 .407s	5299.79	1597.50	3.	6.	.002	56.68	
BI-212 .435s	5377.64	1620.97	4.	30.	.008	25.07	
AC-228 .420s	5408.19	1630.18	4.	13.	.004	38.31	
BI-214 .303D	5511.34	1661.28	4.	2.	.001	136.31	2
PA-234 .296D	5535.29	1668.50	8.	9.	.003	53.56	2
PA-234 .377s	5617.29	1693.22	0.	14.	.004	26.73	
BI-214 .271s	5735.94	1728.99	2.	23.	.006	25.41	1
BI-214 .006s	5852.31	1764.08	7.	66.	.018	14.71	1
BI-207 .362D	5872.68	1770.22	3.	0.	.000	.00	
CO-56	5876.90	1771.49	3.	0.	.000	.00	
MN-56 .302s	6008.60	1811.20	1.	0.	.000	.00	
Y-88	6091.00	1836.04	1.	0.	.000	87.18	
.241s BI-214	6126.93	1846.88	8.	9.	.002	65.50	
.480s CO-56	6684.00	2014.85	0.	1.	.000	95.92	
.129s CO-56	6748.00	2034.15	0.	2.	.001	70.71	
.302s MN-56	7008.13	2112.60	2.	0.	.000	.00	
.559D BI-214	7023.00	2117.08	1.	1.	.000	141.42	
.257s BI-214	7314.32	2204.94	10.	2.	.001	232.70	
.309s BI-214 .000s	8118.00	2447.35	12.	2.	.001	281.07	

s Peak fails shape tests.

D Peak area deconvoluted.

* * * * * * * *	SUMMARY	OF L	I B R A R Y P	E A K U S	A G E **
		ENERGY keV	ACTIVITY CODE microCi	MDA VALUE	COMMENTS
+-+-+-+		-+-+-+-	+-+-+-+-+-+-		-+-+-+-+ -+-
NA-22 9.99E+01	0.00000E+00	1,274.54	0.000E+00 !(0.0000E+00	9.50E+02
NA-24 1.00E+02	0.00000E+00	1368.55	0.000E+00 %	0.0000E+00	6.26E-01
9.98E+01		2754.10	0.000E+00 =	0.0000E+00	
K-40 1.07E+01	0.00000E+00	1460.75	0.000E+00 (0.0000E+00	4.68E+11
SC-46 1.00E+02	0.00000E+00	1120.52	0.000E+00 !(0.0000E+00	8.38E+01
1.00E+02		889.26	0.000E+00 &(0.0000E+00	
CR-51 9.83E+00	0.00000E+00	320.07	0.000E+00 &	0.0000E+00	2.77 E+01
MN-54 1.00E+02	0.00000E+00	834.81	0.000E+00 !(0.0000E+00	3.12E+02
MN-56 9.90E+01	0.00000E+00	846.60	0.000E+00 (0.0000E+00	1.07E-01
2.72E+01		1811.20	0.000E+00 %	0.0000E+00	
1.43E+01		2112.60	0.000E+00 ?(0.0000E+00	
FE-59 5.65E+01	0.00000E+00	1099.22	0.000E+00 %	0.0000E+00	4.51E+01
4.32E+01		1291.56	0.000E+00 %	0.0000E+00	
7.326701		192.34	0.000E+00 %	0.0000E+00	

```
3.11E+00
                           142.65
                                                    0.0000E+00
                                    0.000E+00 %
1.03E+00
 CO-56
          0.00000E+00
                           846.75
                                    0.000E+00 %
                                                    0.0000E+00 7.73E+01
1.00E+02
                          1238.28
                                    0.000E+00 ?
                                                    0.0000E+00
6.76E+01
                          2598.58
                                    0.000E+00 =
                                                    0.0000E+00
1.69E+01
                          1771.49
                                    0.000E+00 ?
                                                    0.0000E+00
1.57E+01
                          1037.83
                                    0.000E+00 ?
                                                    0.0000E+00
1.40E+01
                          2034.92
                                    0.000E+00 %
                                                    0.0000E+00
7.89E+00
                          3253.64
                                    0.000E+00 =
                                                    0.0000E+00
7.41E+00
                          1360.22
                                    0.000E+00 ?
                                                    0.0000E+00
4.33E+00
                          2015.36
                                    0.000E+00 %
                                                    0.0000E+00
3.08E+00
                          3202.30
                                    0.000E+00 =
                                                    0.0000E+00
3.04E+00
                                    0.000E+00 ?
                          1175.13
                                                    0.0000E+00
2,28E+00
                          3273.19
                                    0.000E+00 =
                                                    0.0000E+00
1.75E+00
                          977.46
                                    0.000E+00 ?
                                                   0.0000E+00
1.44E+00
                          3010.20
                                    0.000E+00 =
                                                   0.0000E+00
1.00E+00
          0.00000E+00
CO-57
                          122.07
                                                   0.0000E+00 2.70E+02
                                    0.000E+00 !(
8.56E+01
                          136.43
                                    0.000E+00 !(
                                                   0.0000E+00
1.06E+01
                           14.41
                                    0.000E+00 ?(
                                                   2.5799E-04
9.50E+00
EG&G ORTEC G V - I ( 191) WAN32 I4W02.74 07-MAR-00 14:30:38
U.C.L.A. Radiation Safety
                                 Spectrum name: AN1E202.AN1
Nuclide Ave activity
                         Energy
                                   Activity Code Peak MDA Comments
CO-58
          0.00000E+00
                          810.75
                                   0.000E+00 % 0.0000E+00 7.08E+01
9.94E+01
```

CO-60 1.00E+02	0.00000E+00	1332.51	0.000E+00	& (0.0000E+00	1.93E+03
9.99E+01		1173.23	0.000E+00	? (0.0000E+00	
ZN-65 5.08E+01	0.00000E+00	1115.52	0.000E+00	! (0.0000E+00	2.44E+02
3.87E+01		8.13	0.000E+00	? (1.9134E-06	
	0.000007.00	1076 62	0.0007.00	•		1.067.00
RB-86 8.76E+00	0.00000E+00	1076.63	0.000E+00 9	፟	0.0000E+00	1.86E+01
Y-88 9.94E+01	0.00000E+00	1836.01	0.000E+00	o,o	0.0000E+00	1.07E+02
9.40E+01		898.02	0.000E+00 9	oic .	0.0000E+00	
5.25E+01		14.15	1.656E-04		8.4177E-05	
8.70E+00		15.80	1.414E-03		5.6452E-04	
	0.0000=.00	074 10				7 415 06
NB-94 1.00E+02	0.00000E+00	871.10	0.000E+00 !	•	0.0000E+00	7.41E+06
1.00E+02		702.50	0.000E+00 %	00	0.0000E+00	
NB-95 9.90E+01	0.00000E+00	765.82	0.000E+00	(0.0000E+00	3.52E+01
SN-113 6.42E+01	0.00000E+00	391.71	0.000E+00 %	o o	0.0000E+00	1.15E+02
3.90E+01		24.21	5.050E-04		1.6907E-04	
		24.00	1.959E-03	;	3.1440E-04	
2.07E+01		27.30	3.508E-03		7.5013E-04	
1.10E+01		27.90	1.626E-02	4	4.1225E-03	
2.10E+00		255.04	0.000E+00 %	ŝ (0.000 0E+0 0	
2.07E+00						
CS-134 9.76E+01	0.00000E+00	604.66	0.000E+00 %	š (0.0000E+00	7.53E+02
8.54E+01		795.76	0.000E+00 %	; (0.0000E+00	
1.54E+01		569.29	0.000E+00 !	! (0.0000E+00	
		801.84	0.000E+00 !	! (0.0000E+00	
8.73E+00						

```
563.26
                                    0.000E+00 %
                                                    0.0000E+00
8.38E+00
                          1365.13
                                    0.000E+00 %
                                                   0.0000E+00
3.04E+00
                          1167.86
                                    0.000E+00 ?
                                                   0.0000E+00
1.80E+00
                          475.35
                                    0.000E+00 &
                                                   0.0000E+00
1.46E+00
                          1038.50
                                    0.000E+00 %
                                                   0.0000E+00
1.00E+00
CS-137
          0.00000E+00
                           661.62
                                    0.000E+00 %
                                                   0.0000E+00
                                                               1.10E+04
8.46E+01
                           32.19
                                    6.810E-03
                                                   3.6768E-03
3.70E+00
                           31.82
                                    3.423E-02
                                                   6.7480E-03
1.92E+00
                           36.40
                                    6.601E-02
                                                   1.2997E-02
1.04E+00
EU-152
          7.38047E-04
                          40.12
                                    7.380E-04 (
                                                   3.6687E-04
                                                                4.64E+03
3.00E+01
                          121.78
                                    0.000E+00 %
                                                   0.0000E+00
2.92E+01
                          344.30
                                   0.000E+00 %
                                                   0.0000E+00
2.70E+01
                         1408.08
                                   0.000E+00 -
                                                   0.0000E+00
2.12E+01
                           39.52
                                   1.596E-03 +
                                                   7.3273E-04
1.60E+01
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Nuclide	Ave activity	Energy	Activity	Code	Peak MDA Comments
1.46E+01		964.00	0.000E+00	-	0.0000E+00
1.36E+01		1112.07	0.000E+00	clo	0.0000E+00
1.30E+01		778.90	0.000E+00	cio	0.0000E+00
1.03E+01		1085.80	0.000E+00	olo	0.0000E+00
9.00E+00		45.40	1.791E-03	+	1.0865E-03
7.62E+00		244.67	0.000E+00	qio	0.0000E+00
7.021.00					

4.18E+00		867.39	0.000E+00 %	0.0000E+00
2.84E+00		444.00	0.000E+00 %	0.0000E+00
2.26E+00		411.09	0.000E+00 %	0.0000E+00
2.10E+00		46.60	6.644E-03 +	4.5427E-03
1.72E+00		1089.82	0.000E+00 %	0.0000E+00
		1299.19	0.000E+00 %	0.0000E+00
1.71E+00		1212.89	0.000E+00 %	0.0000E+00
1.45E+00				
EU-154 4.05E+01	0.00000E+00	123.10	0.000E+00 %	0.0000E+00 3.10E+03
3.55E+01		1274.80	0.000E+00 %	0.0000E+00
1.97E+01		723.30	0.000E+00 %	0.0000E+00
1.76E+01		1004.80	0.000E+00 %	0.0000E+00
lication		43.00	0.000E+00 %	4.5412E-04 Energy dup
1.13E+01		873.20	0.000E+00 !	0.0000E+00
1.07E+01		996.30	0.000E+00 %	0.0000E+00
lication		42.31	0.000E+00 %	6.5748E-04 Energy dup
6.60E+00		248.04	0.000E+00 %	0.0000E+00
		591.70	0.000E+00 %	0.0000E+00
4.60E+00		48.70	0.000E+00 %	1.7911E-03 Energy dup
lication		756.70	0.000E+00 %	0.0000E+00
4.10E+00		1597.30	0.000E+00 %	
1.70E+00				0.0000E+00
lication			0.000E+00 %	0.0000E+00 Energy dup
1.60E+00		1397.80	0.000E+00 %	0.0000E+00
1.00E+00		50.00	0.000E+00 %	8.7175E-03
EU-155 3.27E+01	0.00000E+00	86.45	0.000E+00 %	1.7177E-04 1.81E+03
2.18E+01		105.31	0.000E+00 %	0.0000E+00

lication		43.00	0.000E+00 ?	4.6009E-04 Energy dup
lication		42.31	0.000E+00 ?	1.1056E-03 Energy dup
		48.70	0.000E+00 ?	2.0873E-03 Energy dup
lication		45.30	0.000E+00 ?	5.4037E-03
1.36E+00		60.01	0.000E+00 !	6.3765E-03
1.21E+00	•			
TL-208 8.60E+01	0.00000E+00	583.14	0.000E+00 (0.0000E+00 2.15E-03
2.25E+01		510.72	0.000E+00 (0.0000E+00
1.20E+01		860.47	0.000E+00 %	0.0000E+00
		277.36	0.000E+00 %	0.0000E+00
6.50E+00		74.97	1.110E-02 +	2.0843E-03 Energy dup
lication		72.80	1.065E-02 +	3.5695E-03 Energy dup
lication		763.30	0.000E+00 !(0.0000E+00
1.70E+00		84.80	1.157E-02 +	
lication		04.00	1.13/2-02 +	4.4931E-03 Energy dup

Nuclide	Ave activity	Energy	Activity	Code	Peak MDA Comments
PB-210 2.43E+01	0.00000E+00	10.80	0.000E+00	olo	4.7401E-06 7.45E+03
4.00E+00		46.52	0.000E+00	!	2.3912E-03
PB-212 4.31E+01	0.00000E+00	238.63	0.000E+00	(0.0000E+00 4.43E-01
lication		77.11	0.000E+00	g _o	2.2626E-04 Energy dup
lication		74.81	0.000E+00	olo	6.1086E-04 Energy dup
lication		87.20	0.000E+00	oio	6.8807E-04 Energy dup
TICALION		300.09	0.000E+00	0,0	0.0000E+00

3.27E+00		22.00		
lication		89.80	0.000E+00 ?(1.2324E-03 Energy dup
PB-214 3.71E+01	0.00000E+00	351.99	0.000E+00 ?(0.0000E+00 1.86E-02
1.92E+01		295.22	0.000E+00 %	0.0000E+00
lication		10.80	0.000E+00 !(1.3785E-04 Energy dup
lication		77.11	0.000E+00 %	5.7718E-04 Energy dup
7.47E+00		241.92	0.000E+00 %	0.0000E+00
lication		74.81	0.000E+00 %	1.1583E-03 Energy dup
lication		87.20	1.819E-03 +	1.4489E-03 Energy dup
1.10E+00		53.20	0.000E+00 %	6.7038E-03
1.09E+00		785.95	0.000E+00 !(0.0000E+00
		89.80	0.000E+00 %	4.9786E-03 Energy dup
lication				
BI-207 9.80E+01	0.00000E+00	569.67	0.000E+00 !(0.0000E+00 1.39E+04
7.70E+01		1063.62	0.000E+00 &	0.0000E+00
lication		74.97	0.000E+00 %	1.5093E-04 Energy dup
3.50E+01		10.60	0.000E+00 ?(3.1394E-06
lication		72.80	0.000E+00 ?(1.5808E-04 Energy dup
lication		84.80	0.000E+00 %	2.0935 E -04 Energy dup
7.00E+00		1770.22	0.000E+00 ?(0.0000E+00
4.00E+00		87.30	0.000E+00 ?(1.1271E-03
BI-212	0.00000E+00	727.17	0.000E+00 %	0.0000E+00 4.20E-02
1.18E+01		1620.56	0.000E+00 ?	0.0000E+00
2.75E+00		785.42	0.000E+00 !	0.0000E+00
2.00E+00		39.86	0.000E+00 !	1.0513E-02
1.10E+00				
BI-214	0.00000E+00	609.32	0.000E+00 (0.0000E+00 1.38E-02

4.61E+01				
1.59E+01	1764.51	0.000E+00	(0.0000E+00
	1120.28	0.000E+00	(0.0000E+00
1.50E+01	1238.11	0.000E+00	(0.0000E+00
5.92E+00	2204.12	0.000E+00	&	0.0000E+00
4.99E+00	768.36	0.000E+00	(0.0000E+00
4.89E+00		0.000E+00	,	
4.02E+00	1377.65		(0.0000E+00
3.16E+00	934.05	0.000E+00	(0.0000E+00
3.05E+00	1729.60	0.000E+00	? (0.0000E+00
2.48E+00	1407.98	0.000E+00	(0.0000E+00
	1509.19	0.000E+00	& (0.0000E+00
2.19E+00	1847.44	0.000E+00	ઇ	0.0000E+00
2.12E+00				

Nuclide	Ave activity	Energy	Activity	Code	Peak MDA	Comments
1 600.00		1155.19	0.000E+00	oło	0.0000E+00	
1.69E+00		665.45	0.000E+00	! (0.0000E+00	
1.56E+00		2447.71	0.000E+00	olo	0.0000E+00	ı
1.55E+00		1280.96	0.000E+00	oj.	0.0000E+00	I
1.47E+00		1401.50	0.000E+00	(0.0000E+00	ı
1.39E+00		806.17	0.000E+00	! (0.0000E+00)
1.23E+00		2118.54	0.000E+00	,	0.0000E+00	
1.21E+00		1661.28			0.0000E+00	
1.15E+00		1001.20	0.000E+00	6	0.00002+00	
RA-224 3.90E+00	0.00000E+00	241.00	0.000E+00	ું	0.0000E+00	3.66 E +00

RA-226 3.28E+00	0.00000E+00	185.99	0.000E+00 !(0.0000E+00	5.84E+05
AC-228 2.90E+01	0.00000E+00	.911.07	0.000E+00 (0.0000E+00	2.55E-01
1.75E+01		968.90	0.000E+00 (0.0000E+00	
1.20E+01	2	338.40	0.000E+00 (0.0000E+00	
		964.60	0.000E+00 %	0.0000E+00	
5.45E+00		794.80	0.000E+00 !(0.0000E+00	
4.84E+00		463.00	0.000E+00 %	0.0000E+00	
4.64E+00		209.40	0.000E+00 %	0.0000E+00	
4.55E+00		270.30	0.000E+00 %	0.0000E+00	
3.77E+00			0.000E+00 &(
3.71E+00		1587.90		0.0000E+00	
3.36E+00		328.00	0.000E+00 %	0.0000E+00	
2.93E+00		129.10	0.000E+00 !(0.0000E+00	
2.23E+00		409.40	0.000E+00 %	0.0000E+00	
1.95E+00		1630.40	0.000E+00 ?(0.0000E+00	
1.82E+00		835.60	0.000E+00 %	0.0000E+00	
1.62E+00		772.10	0.000E+00 %	0.0000E+00	
		99.45	0.000E+00 !(3.1646E-03	
1.36E+00		755.20	0.000E+00 !(0.0000E+00	
1.10E+00		1495.80	0.000E+00 &(0.0000E+00	
1.05E+00		1459.20	0.000E+00 (0.0000E+00	
1.04E+00			,		
TH-234 9.80E+00	6.52686E-03	13.30	6.527E-03 (1.9815E-04	2.41E+01
3.90E+00		63.29	0.000E+00 %	1.4047E-03	
3.00E+00		92.80	1.672E-03 -	1.6318E-03	
2.57E+00		92.38	0.000E+00 %	1.9137E-03	

PA-234 1.14E+02	5.47395E-05	13.60	5.474E-05 (3.4359E-05 2.79E-01
2.51E+01		98.44	0.000E+00 %	1.7480E-04
		946.00	0.000E+00 %	0.0000E+00
2.00E+01		131.28	0.000E+00 %	0.0000E+00
2.00E+01		94.67	0.000E+00 %	3.0778E-04
1.55E+01	•	883.24	0.000E+00 %	0.0000E+00
1.20E+01		003.24	0.000E+00 %	0.00002+00
1.10E+01		926.70	0.000E+00 %	0.0000E+00
		569.26	0.000E+00 %	0.0000E+00 Energy dup
lication		111.00	0.000E+00 %	0.0000E+00
8.55 E +00		733.00	0 00000 00 0	0.00007.00
8.50E+00		133.00	0.000E+00 %	0.0000E+00

Nuclide	Ave activity	Energy	Activity	Code	Peak MDA	Comments
7.80E+00		949.00	0.000E+00	_	0.0000E+00)
7.20E+00		152.70	0.000E+00	&	0.0000E+00)
lication		880.51	0.000E+00	oto	0.0000E+00	Energy dup
		226.87	0.000E+00	&	0.0000E+00	Energy dup
lication		831.10	0.000E+00	olo	0.0000E+00)
5.60E+00		808.10	0.000E+00	-	0.0000E+00)
4.90E+00		99.70	0.000E+00		9.1173E-04	
4.70E+00		699.10	0.000E+00	ભ	0.0000E+00)
4.60E+00		880.51	0.000E+00		0.000 0E +00	Energy dup
lication		898.60	0.000E+00	-	0.0000E+00	
4.00E+00		1394.10	0.000E+00	ojo	0.0000E+00	
3.90E+00						

lication	226.87	0.000E+00 &	0.0000E+00 Energy dup
3.60E+00	824.70	0.000E+00 %	0.0000E+00
3.40E+00	796.20	0.000E+00 -	0.0000E+00
3.30E+00	805.50	0.000E+00 %	0.0000E+00
3.20E+00	826.30	0.000E+00 -	0.0000E+00
3.20E+00	706.00	0.000E+00 %	0.0000E+00
3.20E+00	293.70	0.000E+00 %	0.0000E+00
3.15E+00	114.50	0.000E+00 %	0.0000E+00
lication	569.26	0.000E+00 -	0.0000E+00 Energy dup
3.10E+00	248.90	0.000E+00 %	0.0000E+00
3.10E+00	63.20	0.000E+00 -	2.1941E-03
3.00E+00	742.81	0.000E+00 %	0.0000E+00
3.00E+00	568.70	0.000E+00 %	0.0000E+00
2.80E+00	924.60	0.000E+00 %	0.0000E+00
2.80E+00	369.60	0.000E+00 %	0.0000E+00
2.70E+00	876.70	0.000E+00 %	0.0000E+00
2.40E+00	983.40	0.000E+00 %	0.0000E+00
2.20E+00	980.50	0.000E+00 %	0.0000E+00
2.20E+00	819.40	0.000E+00 %	0.0000E+00
2.00E+00	574.00	0.000E+00 %	0.0000E+00
lication	186.00	0.000E+00 &	0.0000E+00 Energy dup
1.70E+00	1353.00	0.000E+00 %	0.0000E+00
1.70E+00	566.50	0.000E+00 %	0.0000E+00
1.60E+00	786.27	0.000E+00 %	0.0000E+00
1.60E+00	780.80	0.000E+00 -	0.0000E+00
1.50E+00	793.60	0.000E+00 %	0.0000E+00

1.50E+00	669.80	0.000E+00 %	0.0000E+00
1.305+00	666.70	0.000E+00 %	0.0000E+00
1.50E+00	664 60	0.0007.00.0	0.0000=.00
1.50E+00	664.60	0.000E+00 %	0.0000E+00
1.50E+00	458.60	0.000E+00 %	0.0000E+00
	372.20	0.000E+00 %	0.0000E+00
1.50E+00	978.80	0.000E+00 %	0.0000E+00
1.40E+00			
1.40E+00	754.80	0.000E+00 %	0.0000E+00
1 400+00	506.80	0.000E+00 %	0.0000E+00
1.40E+00	1694.00	0.000E+00 -	0.0000E+00
1.30E+00	600 50	0.000	
lication	692.50	0.000E+00 %	0.0000E+00 Energy dup
1.30E+00	653.20	0.000E+00 %	0.0000E+00
1.300100			

Nuclide	Ave activity	Energy	Activity	Code	Peak MDA Comments
1.30E+00		602.80	0.000E+00	용	0.0000E+00
		1452.60	0.000E+00	olo	0.0000E+00
1.20E+00		521.00	0.000E+00	ojo	0.0000E+00
1.20E+00		520.60	0.000E+00	_	0.0000E+00
1.20E+00		513.60	0.000E+00	o _i o	0.0000E+00
1.20E+00		272.10	0.000E+00	g g	0.0000E+00
1.20E+00		202.90	0.000E+00	o,o	0.0000E+00
1.10E+00		125.30	0.000E+00	_	0.0000E+00
1.10E+00		200.80	0.000E+00		0.0000E+00
1.08E+00		1668.50	0.000E+00		0.0000E+00
1.06E+00		1000.30	0.0006+00	ō	0.00002+00

1.00E+00		623.50	0.000E+00 %	0.0000E+00	
U-235 5.40E+01	0.00000E+00	185.72	0.000E+00 %	0.0000E+00	1.39 E+09
1.05E+01		143.76	0.000E+00 &	0.0000E+00	
		205.31	0.000E+00 %	0.0000E+00	
4.70E+00		163.35	0.000E+00 &	0.0000E+00	
4.70E+00		93.35	0.000E+00 ?	1.5583E-03	
2.50E+00		109.14	0.000E+00 ?	0.0000E+00	
1.50E+00					
1.50E+00		89.96	0.000E+00 ?	1.9844E-03	
1.00E+00		202.12	0.000E+00 %	0.0000E+00	
1.00E+00		105.00	0.000E+00 &	0.0000E+00	
/ mb		-112			

- (This peak used in the nuclide activity average.
 - * Peak is too wide, but only one peak in library.
 - ! Peak is part of a multiplet and this area went negative during deconvolution.
 - ? Peak is too narrow.
 - @ Peak is too wide at FW25M, but ok at FWHM.
 - % Peak fails sensitivity test.
 - \$ Peak identified, but first peak of this nuclide failed one or more qualification tests.
 - + Peak activity higher than counting uncertainty range.
 - - Peak activity lower than counting uncertainty range.
 - = Peak outside analysis energy range.
 - & Calculated peak centroid is not close enough to the library energy centroid for positive identification.
 - P Peakbackground subtraction

***** SUMMARY OF NUCLIDES IN SAMPLE **

TIME OF COUNT TIME CORRECTED UNCERTAINTY 1 SIGMA
NUCLIDE ACTIVITY ACTIVITY COUNTING TOTAL

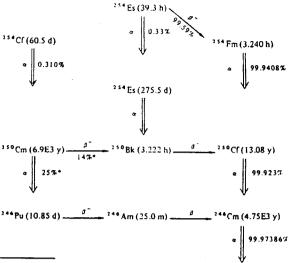
uCi/grams

+-+-+-+
NA-22 < 0.00E+00 0.00E+00
NA-24 < 0.00E+00 >12 HALFLIVES
K-40 < 0.00E+00 0.00E+00

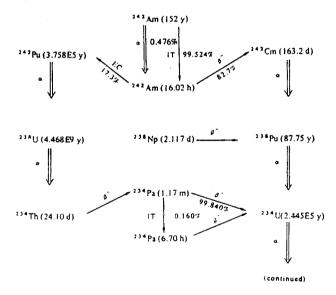
uCi/grams

```
0.00E+00
0.00E+00
 SC-46 < 0.00E+00
CR-51 < 0.00E+00
MN-54 < 0.00E+00
 EG&G ORTEC G V - I ( 191) WAN32 I4W02.74 07-MAR-00 14:30:38 Page
 U.C.L.A. Radiation Safety Spectrum name: AN1E202.AN1
MN-56
       <
            0.00E+00 >12 HALFLIVES
                                       25.17%
                                                  25.30%
                                       2.38%
                                                 ქ.40.
31.70%
         1.4794E-09 >12 HALFLIVES
0.00E+00 0.00E+00
                                      31.60%
U-235 <
----- S U M M A R Y -----
TOTAL ACTIVITY ( 3.8 to 2455.2 keV) 3.3303310E-07 uCi/grams
The library has energies which are not separable.
Analyzed by:
                 MFH
```

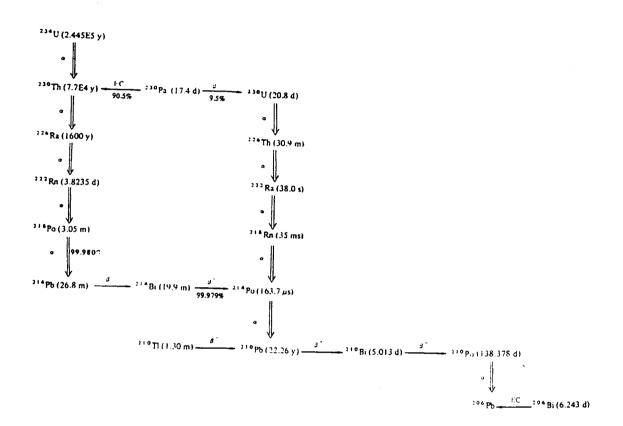
Uranium Series



*Branching ratio based on systematics; decay has not been deserved



RADIOACTIVE DECAY DATA TABLES



DIAGRAMS OF RADIOACTIVE DECAY CHAINS