A STUDY OF A MULTI-CENTURY CORAL STABLE ISOTOPE RECORD FROM RAROTONGA, SOUTHWEST SUBTROPICAL PACIFIC, FOR THE PERIOD 1726-1997

by

Lei Ren

A Dissertation

Submitted to the University at Albany, State University of New York
in Partial Fulfillment of
the Requirements for the Degree of
Doctor of Philosophy

College of Arts & Sciences

Department of Earth & Atmospheric Sciences

2002

ABSTRACT

This study presents a 271-year (1726-1997) subseasonal oxygen and carbon isotopes ($\delta^{18}O$ and $\delta^{13}C$) records from a coral colony of *porites lobata* at Rarotonga (21.5°S, 159.5°W) in the southwest subtropical Pacific. A new method is introduced whereby the effects of sea surface temperature (SST) can be separated from those of seawater $\delta^{18}O$ composition ($\delta^{18}O_{sw}$) on coral $\delta^{18}O$ by using the coupled coral Sr/Ca and $\delta^{18}O$. The reconstructed $\delta^{18}O_{sw}$ at Rarotonga using this method shows that it contributes significantly to the annual changes of coral $\delta^{18}O$ for the period 1726-1997. While changes of $\delta^{18}O_{sw}$ account for ~39% of the total coral $\delta^{18}O$ variation, changes of SST account for ~61%. The reconstructed $\delta^{18}O_{sw}$ also shows a positive linear correlation with a satellite-based estimated salinity for the period 1980-1997 (r=0.72). This linear correlation between reconstructed $\delta^{18}O_{sw}$ and salinity makes it possible to use the reconstructed $\delta^{18}O_{sw}$ to estimate the past interannual and decadal salinity changes in this region.

Applying a similar method to coral δ^{13} C, the effects of kinetic and metabolic activity on coral δ^{13} C were also quantitatively separated. The results show that the variation of coral δ^{13} C appears to be mainly caused by variation of metabolic activity rather than that of kinetic activity in both tropical and subtropical regions. For the tropical regions, δ^{13} C variation in corals is predominantly influenced by changes of metabolic activity (~90%), while for subtropical regions, approximately 70-75% of the total variation of coral δ^{13} C is due to the effects of metabolic activity.

The interannual and interdecadal variability in coral $\delta^{18}O$ at Rarotonga for the period 1726-1997 was also examined. The results suggest that although Rarotonga is located outside of the center of action of ENSO, it is generally sensitive to ENSO variability in this region. In addition, the decadal variability (~12 yr) was further differentiated from the interdecadal-scale variability (~32 yr) for the period 1726-1997 at Rarotonga. Based on the analysis of both tropical and subtropical coral data and comparisons with the Nino3.4 SST index and PDO index, it was hypothesized that the decadal and interdecadal variability might result from separate forcing mechanisms.

ACKNOWLEDGMENTS

Without the help and encouragement of many people, this work would have been impossible. I am deeply indebted to my supervisor Brad Linsley for all these years of his guidance and support, and for his consistent encouragement and inspiration throughout the course of this work. I learnt a great deal from him who not only encouraged independence but also provided direction with his thoughtful comments and challenging questions. The financial support that he provided is also gratefully acknowledged. I would also like to thank my two other committee members Andrei Lapenis and Gerard Wellington for reading this thesis and for their thoughtful comments and constructive criticisms of my early draft. I am thankful to Steve Howe for his assistance with isotopic analysis. His professional guidance and his patience in helping me using the mass spectrometer greatly improved the quality of my analysis results.

I would like to thank the Geological Sciences faculty and students of the Department of Earth & Atmospheric Sciences, who have helped me in innumerable ways through my graduate career. I'm particularly grateful to John Delano, John Arnason, and Grag Happer for their clearly-stated, thought-provoking questions and comments at the graduate students' seminars. Thanks are also to my fellow graduate students whom I could always turn to for help during the years I stayed here. In particular, I thank Steffi Dannenmann, Bin Zhu, and Adam Schoonmaker. I am especially indebted to Taohong Li, my best friend for all these years, not only for helping me in preparing my seminars of each semester and revising my thesis during the final stage of my work, but most importantly for being the one I could always count on through these important years of

my life. I am thankful for the great opportunities in the Geological Sciences Department where I really benefited a great deal over the years.

Finally, a special word of thanks goes to my parents who led me into this fantastic geological world, and my husband Andrew. I am greatly grateful to them for their encouragement, patience, and endurance with me during my adventurous university career over the past years. Without their support and understanding, the completion of this work would have been impossible.

This research was supported by National Science Foundation grants ATM-9901649 and NOAA grants NA96GP0406 to Brad Linsley, and National Science Foundation grants ATM-9619035 and NOAA grant NA96GP0470 to Gerard Wellington.

TABLE OF CONTENTS

ABSTRACT		ii
ACKNOWL	EDGMENTS	iv
TABLE OF	TABLE OF CONTENTS	
CHAPTER 1	PROBLEMS AND OBJECTIVES	
CHAPTER 2	 1.1 Introduction 1.2 Main Problems 1.3 Present Objectives 1.4 Chapter Overview DECONVOLVING THE δ¹8O _{SEAWATER} COMPONENT	2 6 9 11
	FROM SUBSEASONAL CORAL δ ¹⁸ O AND SR/CA AT RAROTONGA IN THE SOUTHWESTERN SUBTROPICAL PACIFIC FOR THE PERIOD 1726-1997	
	Abstract 2.1 Introduction 2.2 Study Area 2.3 Methods and Results 2.4 Separating the Contributions of SST and $\delta^{18}O_{sw}$ from $\delta^{18}O$	14 15 17 20
	in Corals 2.4.1 Method 2.4.2 Error Estimate 2.5 Discussion	23 30
	2.5.1 The Reconstruction of Precipitation and Salinity from $\delta^{18}O_{sw}$	32
	2.5.2 The Separate Influences of $\delta^{18}O_{sw}$ and SST on $\delta^{18}O$ in Corals	34
	2.5.3 The Decadal Variability in Rarotonga Coral δ^{18} O and δ^{18} O _{sw}	38
	$2.5.4$ The Long-term Trend in $\delta^{18}O_{sw}$ 2.6 Summary	41 42
CHAPTER 3	SEPARATING THE EFFECTS OF KINETIC AND METABOLIC FRACTIONATION ON CORAL δ^{13} C	
	Abstract 3.1 Introduction 3.2 Study Area 3.3 General Characteristics of Coral δ ¹³ C	45 46 49 50

	3.4 Method	54
	3.5 Application to Other Coral δ^{13} C Records	65
	3.6 Discussion	76
	3.7 Concluding Remarks	79
CHAPTER 4	INTERANNUAL TO INTERDECADAL VARIABILITY IN THE SUBTROPICAL PACIFIC: A RECONSTRUCTION FROM THE MULTI-CENTURY OXYGEN ISOTOPE RECORD FROM A RAROTONGA CORAL	
	Abstract	82
	4.1 Introduction	83
	4.2 Study Area and Coral Cores	85
	4.3 Methods	
	4.3.1 Sample Preparation	87
	4.3.2 Mass Spectrometer Analysis	89
	4.3.3 Chronology	90
	4.3.4 Time Series Analysis	92
	 4.4 Results and Discussion 4.4.1 General Characters and Frequency Analysis of Coral δ¹⁸O 	93
	4.4.2 Interannual Variance	95
	4.4.3 Decadal/Interdecadal Variance	93
	Characteristics of Decadal/Interdecadal Variance	100
	Potential Mechanisms of Decadal/Interdecadal Variance	103
	4.5 Summary	108
CHAPTER 5	EXAMINATION OF THE REPRODUCIBILITY OF $\delta^{18}O$ AND $\delta^{13}C$ IN CORALS	
	Abstract	111
	5.1 Introduction	112
	5.2 Sample Preparation and Chronology of Core 3R	114
	5.3 Oxygen Isotopes	116
	5.4 Carbon Isotopes	
	 5.4.1 Reproducibility of Carbon Isotope in Corals 5.4.2 Complexity of Effects of Various Factors on Coral δ¹³C 	124 129
CHAPTER 6	CONCLUSIONS AND PROBLEMS FOR FUTURE WORK	
	6.1 Summary	133
	6.2 Future Research	136

APPENDICES	
Appendix I Coral δ^{18} O and δ^{13} C in cores B and C from	
Rarotonga for the period 1726-1997	152
Appendix II Coral $\delta^{\bar{1}8}$ O and $\delta^{\bar{1}3}$ C in core 3R from Rarotonga	
for the period 1926-2000	176

139

REFERENCES