Linkages Between Tropopause Polar Vortices and the Great Arctic Cyclone of August 2012

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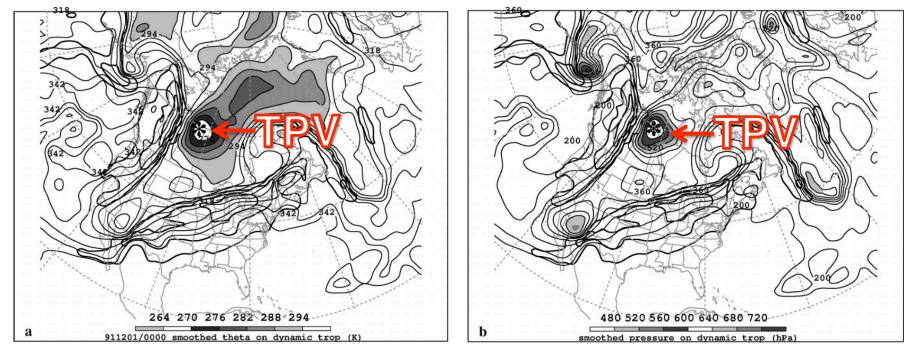
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What are Tropopause Polar Vortices (TPVs)

 TPVs are defined as tropopause-based vortices of highlatitude origin and are material features (Pyle et al. 2004; Cavallo and Hakim 2009, 2010, 2012, 2013)



(left) Dynamic tropopause (DT) wind speed (every 15 m s⁻¹ starting at 50 m s⁻¹, thick contours) and DT potential temperature (K, thin contours and shading) on 1.5-PVU surface valid 0000 UTC 1 Dec 1991; (right) same as left except DT pressure (hPa, thin contours and shading). Adapted from Fig. 11 in Pyle et al. (2004).

Motivation

- TPVs may interact with and strengthen jet streams, and act as precursors to the development of intense Arctic cyclones (e.g., Simmonds and Rudeva 2012, 2014)
- Arctic cyclones may be associated with strong surface winds and poleward advection of warm, moist air, contributing to reductions in Arctic sea-ice extent (e.g., Zhang et al. 2013)
- Heavy precipitation, strong surface winds, and large waves due to Arctic cyclones may pose hazards to ships moving through open passageways in the Arctic Ocean

The Great Arctic Cyclone of August 2012 (AC12)

- AC12 formed over Siberia on 2 August 2012 and tracked northeastward into the Arctic, reaching a minimum central sea level pressure (SLP) of 966.4 hPa at 1800 UTC 6 August in the CFSR (Simmonds and Rudeva 2012)
- AC12 led to reductions in Arctic sea-ice extent during a time in which sea ice was thin and sea-ice volume was well below normal (Zhang et al. 2013)
- Strong surface winds associated with AC12 helped to break up the thin sea ice (e.g., Parkinson and Comiso 2013)

The Great Arctic Cyclone of August 2012 (AC12)

- According to Zhang et al. (2013) sea-ice volume decreased twice as fast as normal during AC12 due to melting of the bottom and perimeter of ice floes
- Simmonds and Rudeva (2012) and Yamazaki et al. (2015) found that a TPV played an important role in the lifecycle of AC12

Outline

- Identification and synoptic examination of three TPVs, a predecessor surface cyclone, and AC12
- Impact of AC12 on Arctic sea-ice extent

• Data:

- 0.3° ERA5 (Hersbach and Dee 2016)

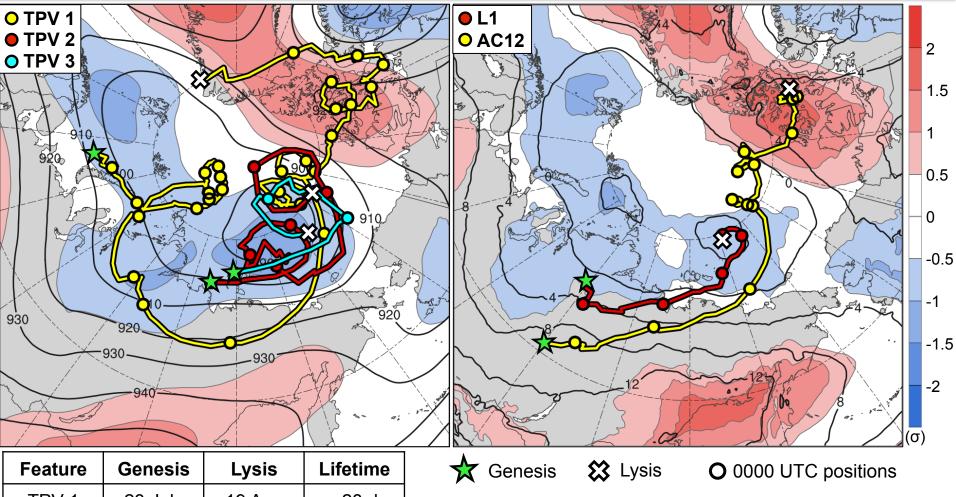
- Utilized TPV tracking algorithm developed by Nicholas Szapiro and Steven Cavallo to identify and track TPVs of interest for AC12
 - Input variables: potential temperature, relative vorticity, and wind on 2-PVU surface
 - Potential temperature minima on 2-PVU surface tracked spatially and temporally to create TPV tracks

Link for Tracking Algorithm: https://github.com/nickszap/tpvTrack

Data and Methods

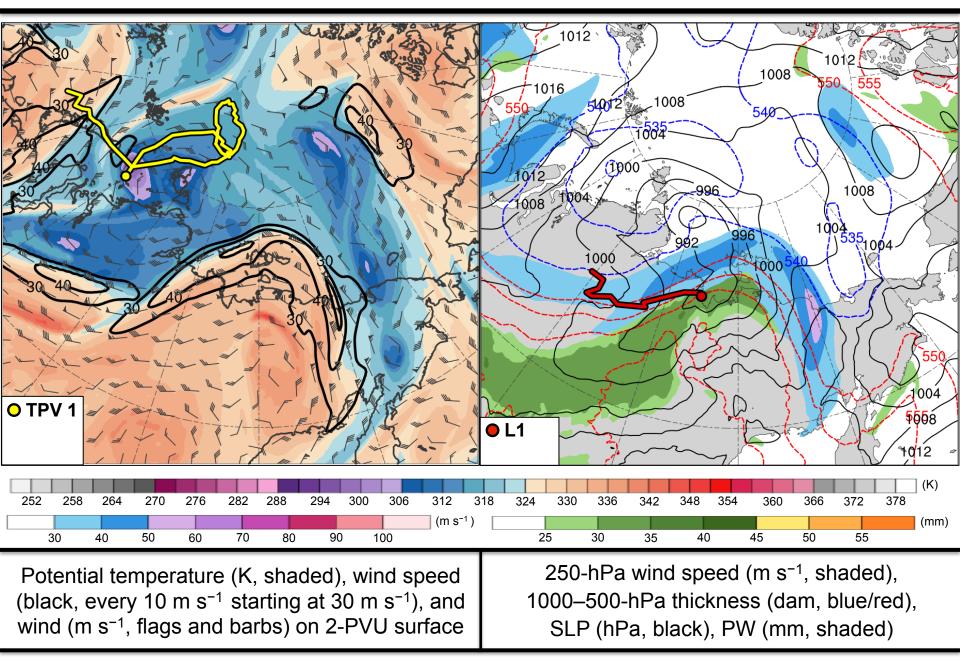
 Manually tracked the predecessor surface cyclone and AC12 by following the locations of minimum SLP

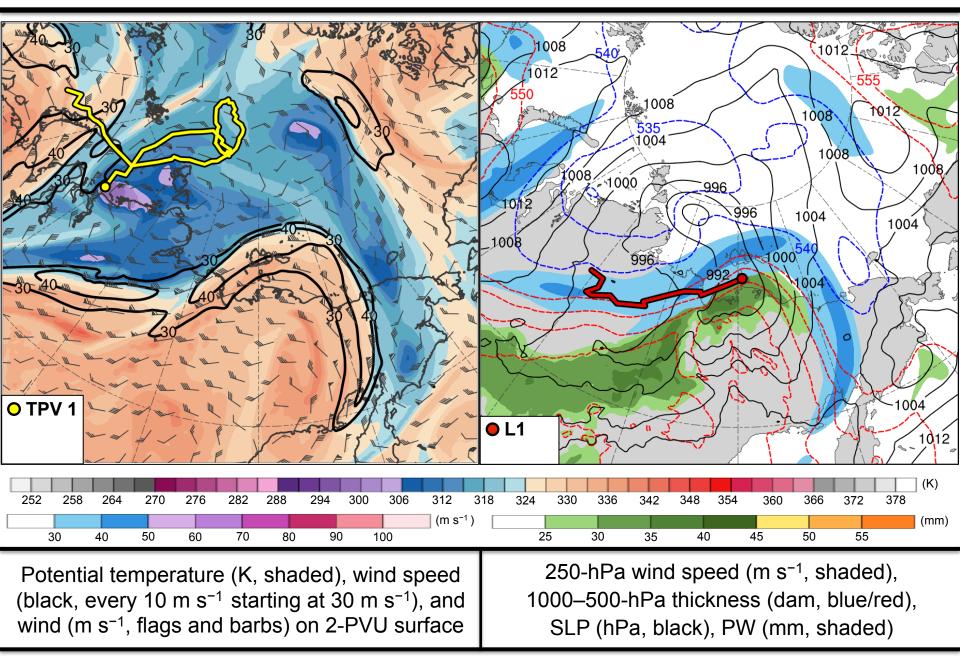
TPV and Surface Cyclone Tracks

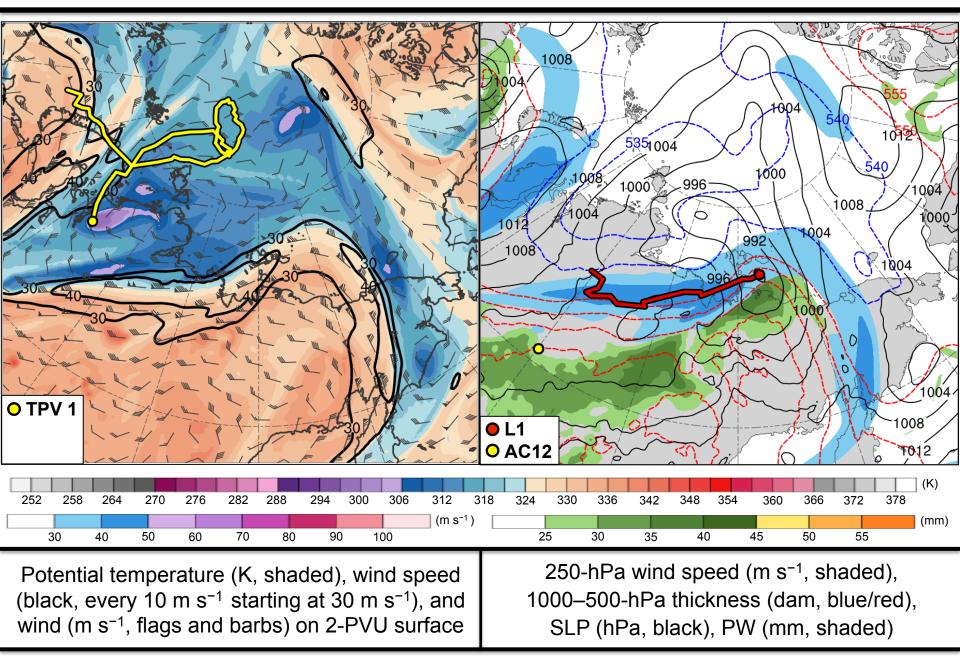


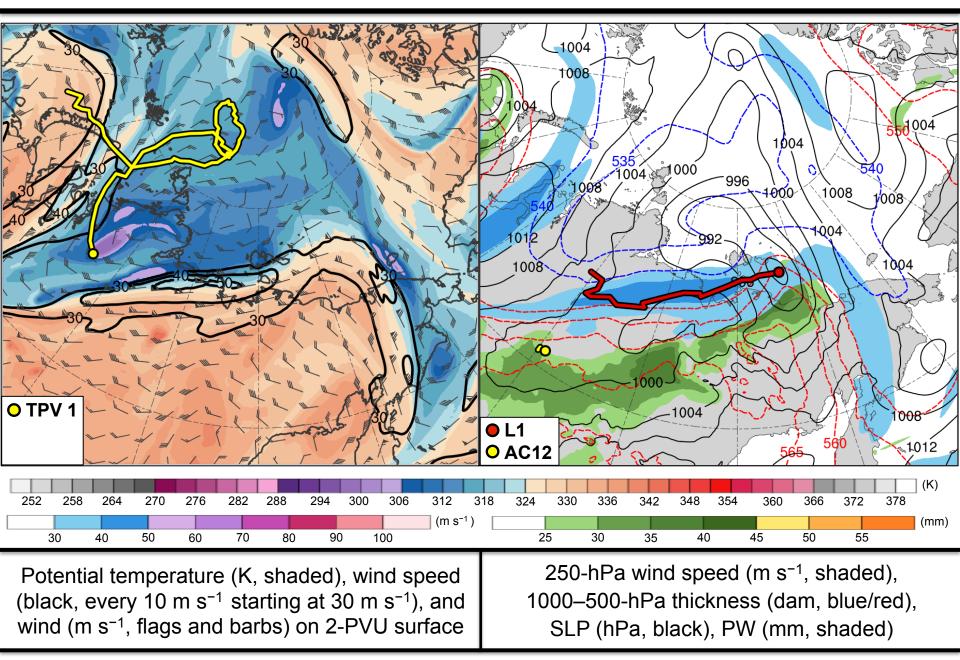
 1–7 Aug 2012 time-mean (left) 300-hPa geopotential height (dam, black) and standardized anomaly of 300-hPa geopotential height (σ, shaded); (right) 850-hPa temperature (°C, black) and standardized anomaly of 850hPa temperature (σ, shaded)

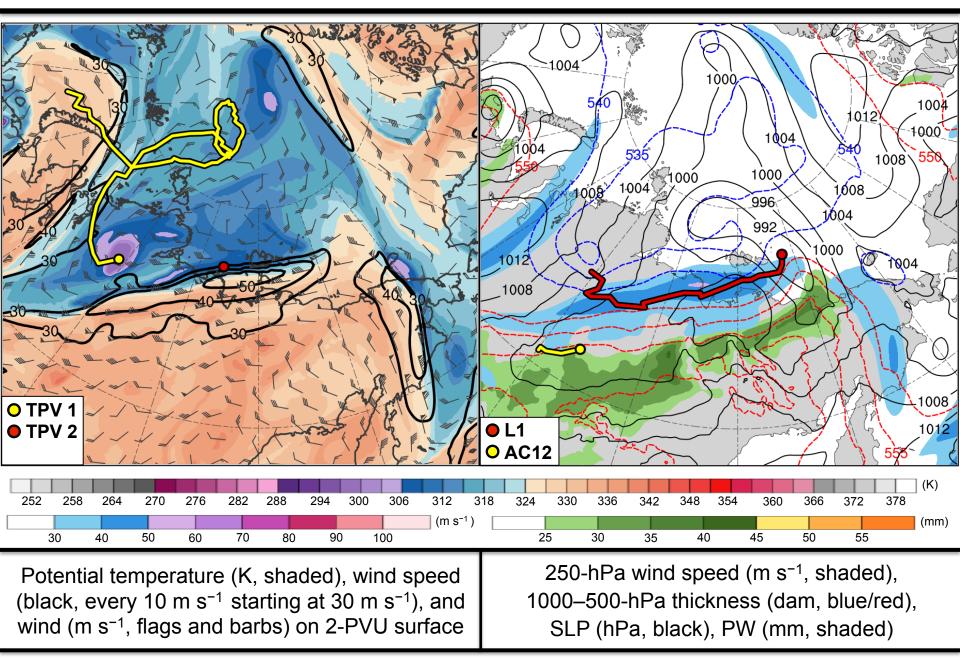
Feature	Genesis	Lysis	Lifetime
TPV 1	23 July	19 Aug	~28 d
TPV 2	3 Aug	9 Aug	~6 d
TPV 3	4 Aug	6 Aug	~3 d
L1	31 July	5 Aug	~5 d
AC12	2 Aug	15 Aug	~13 d

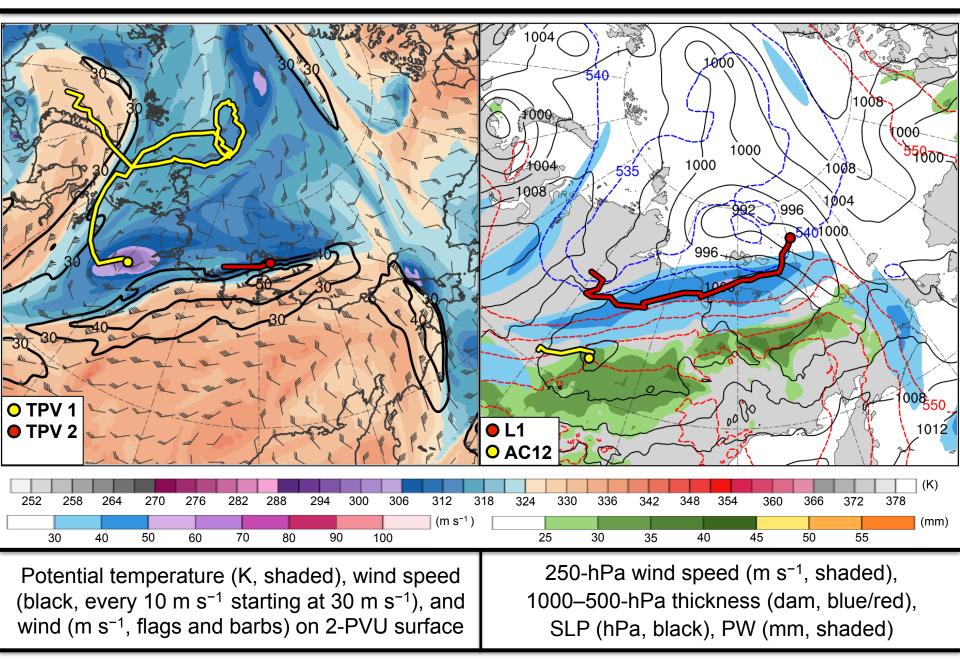


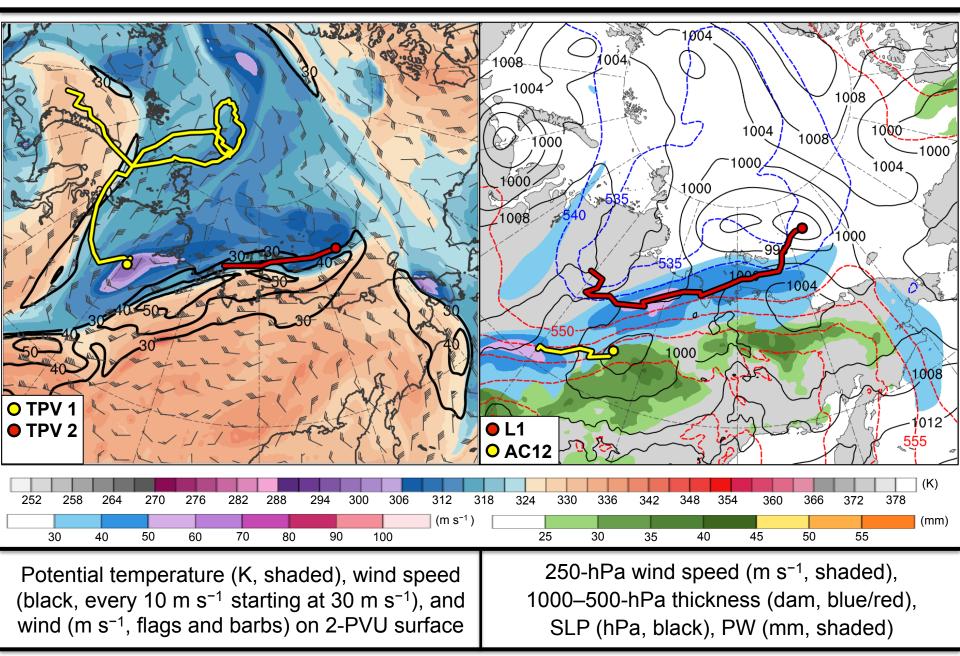


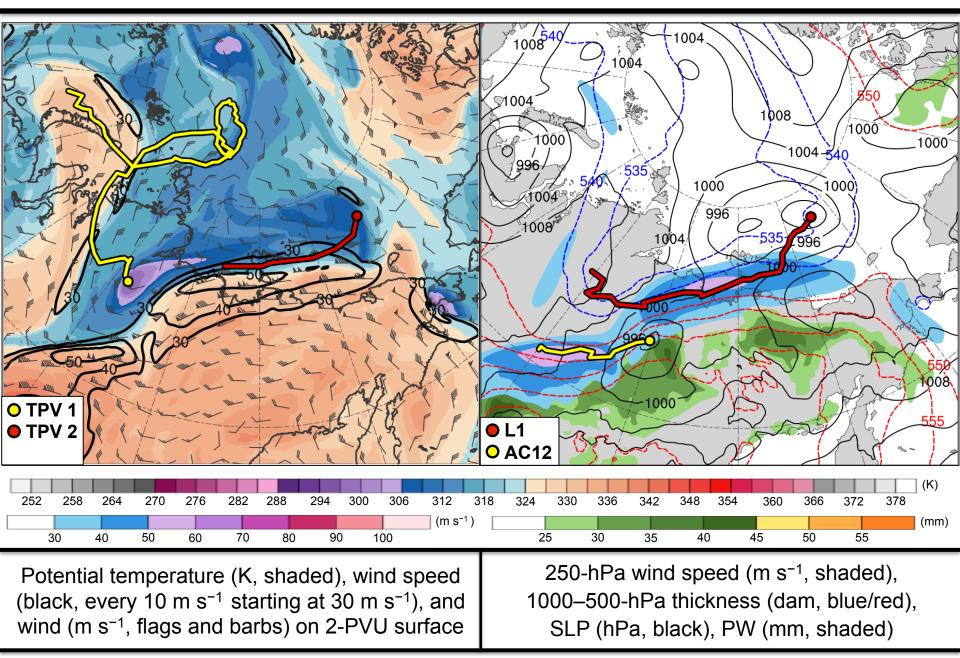


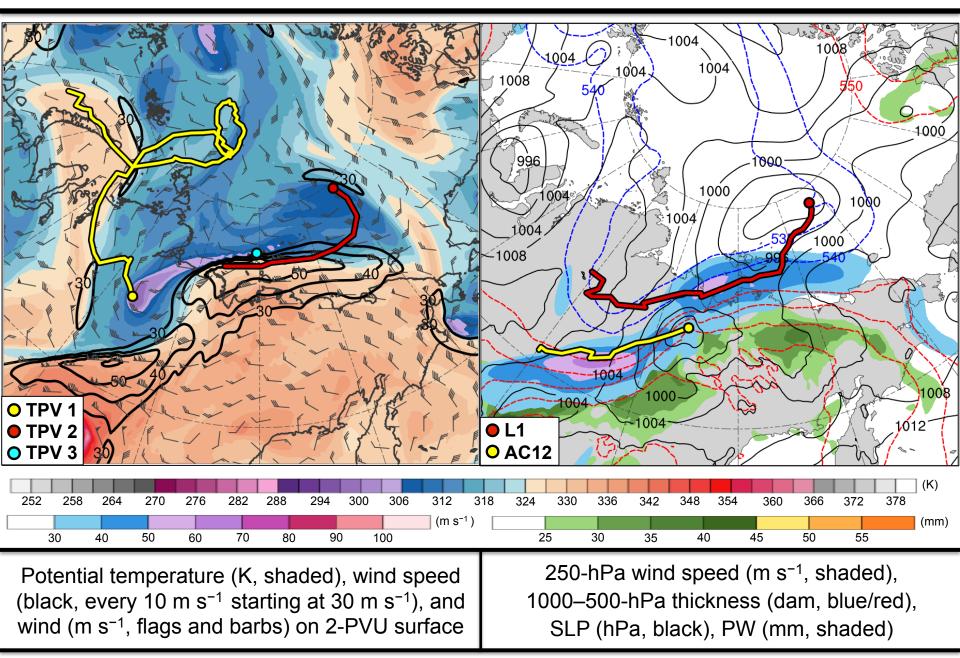


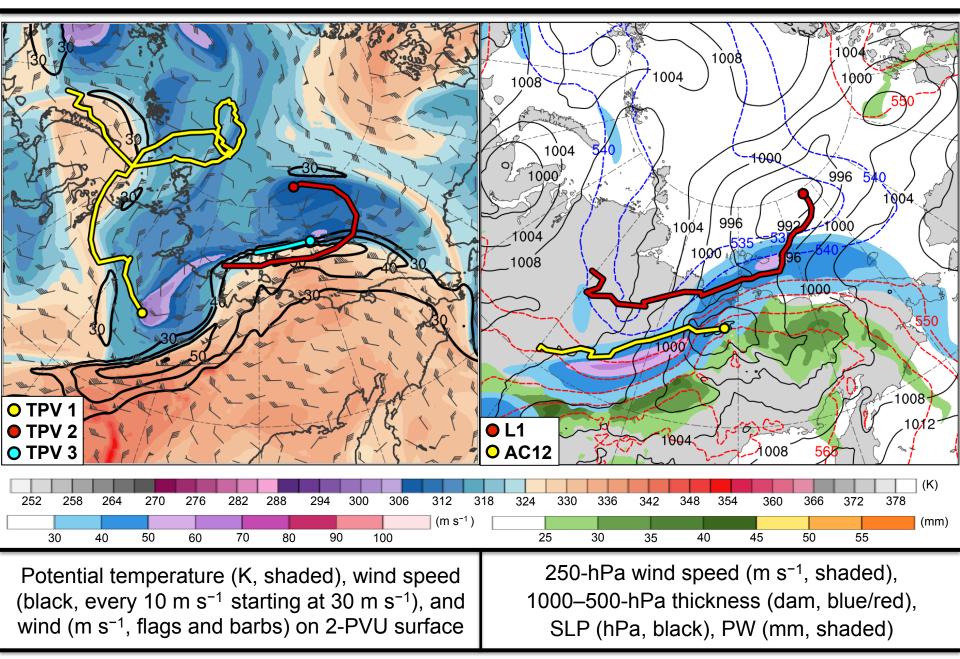


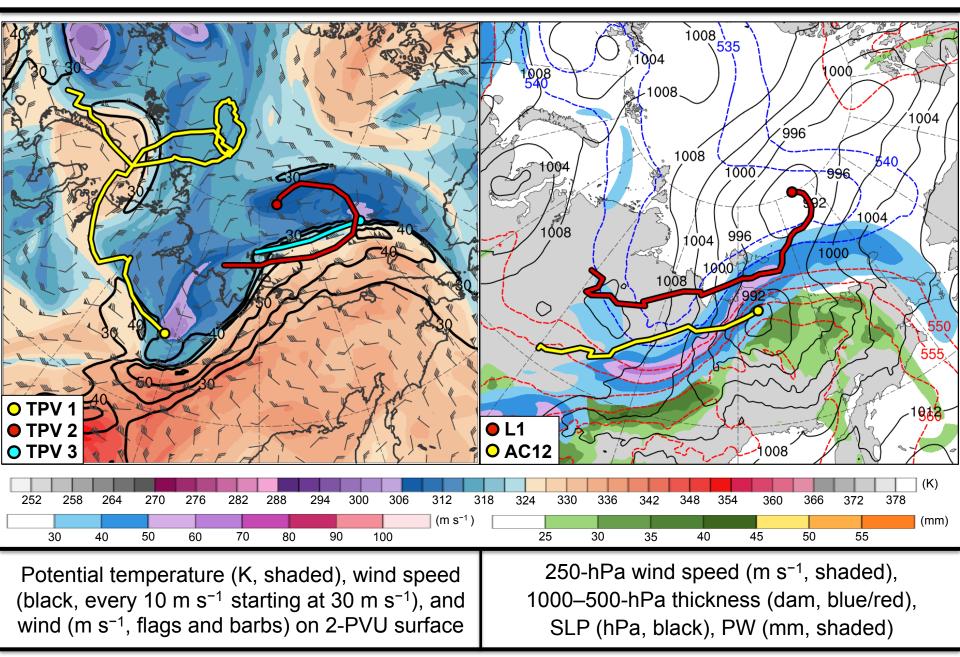


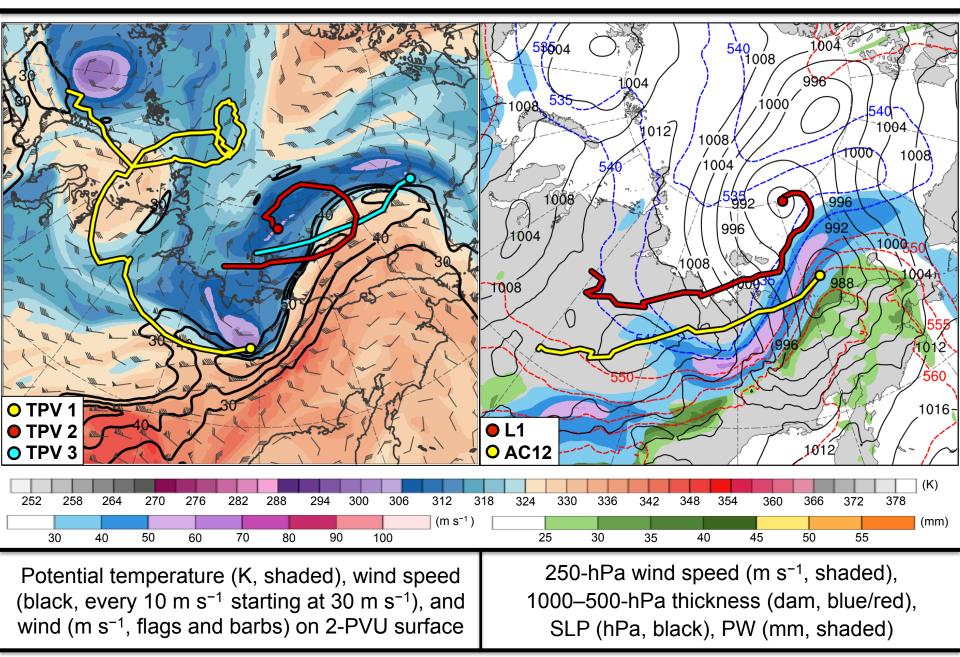


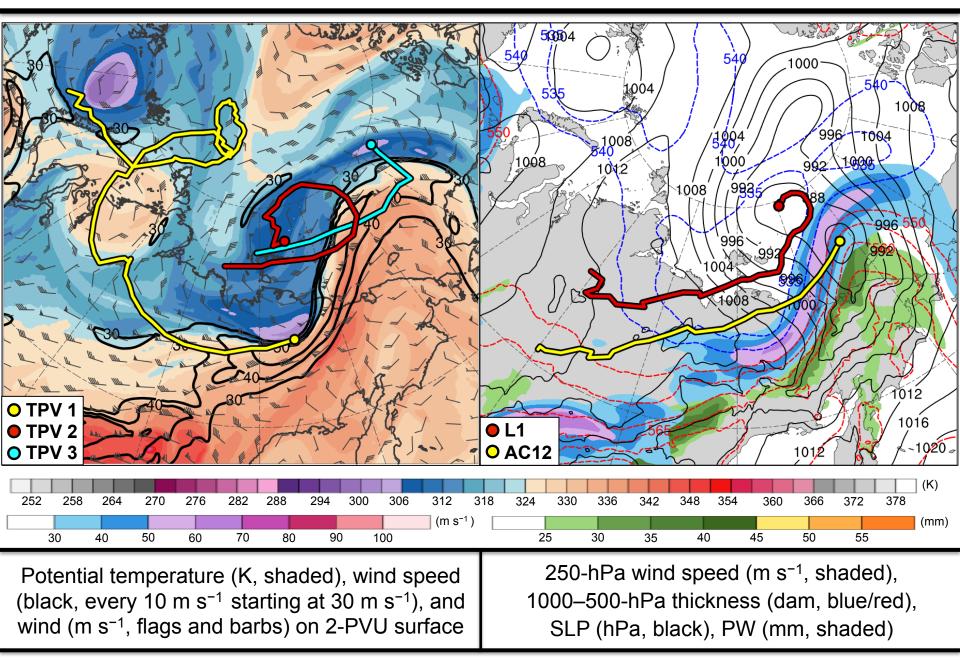


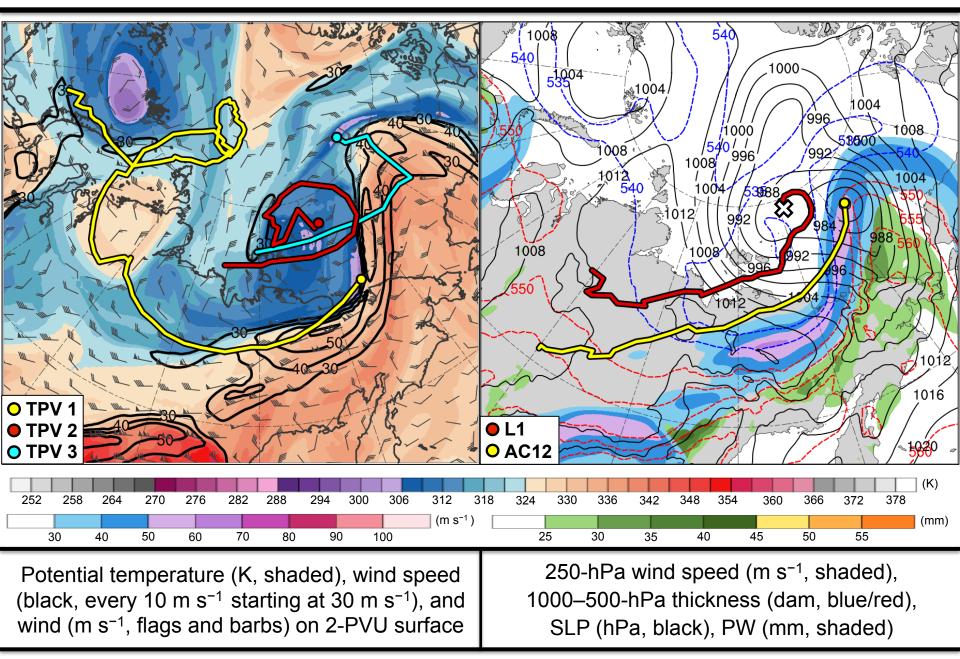


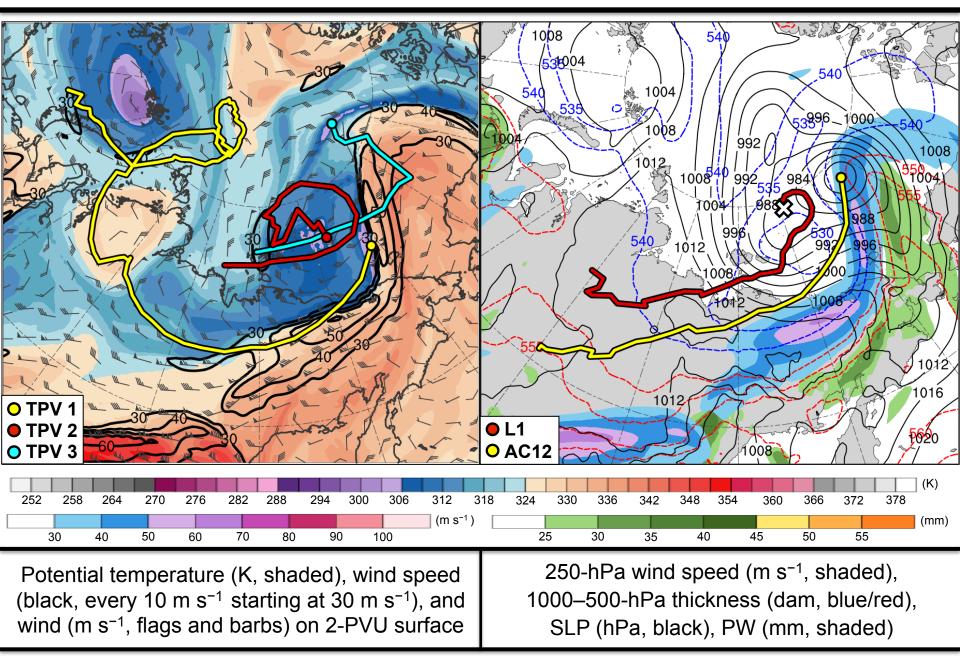


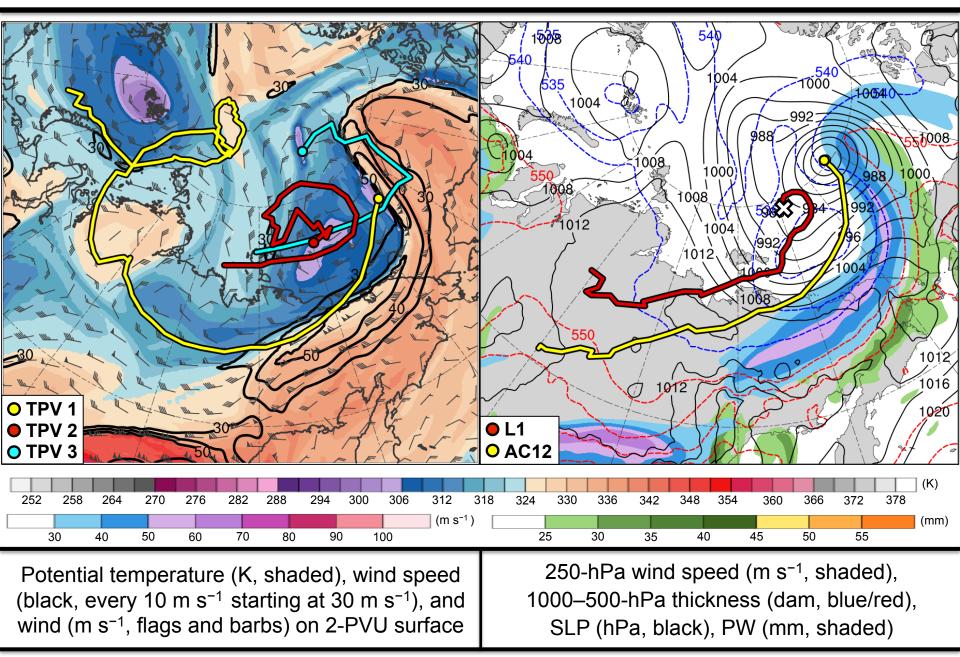


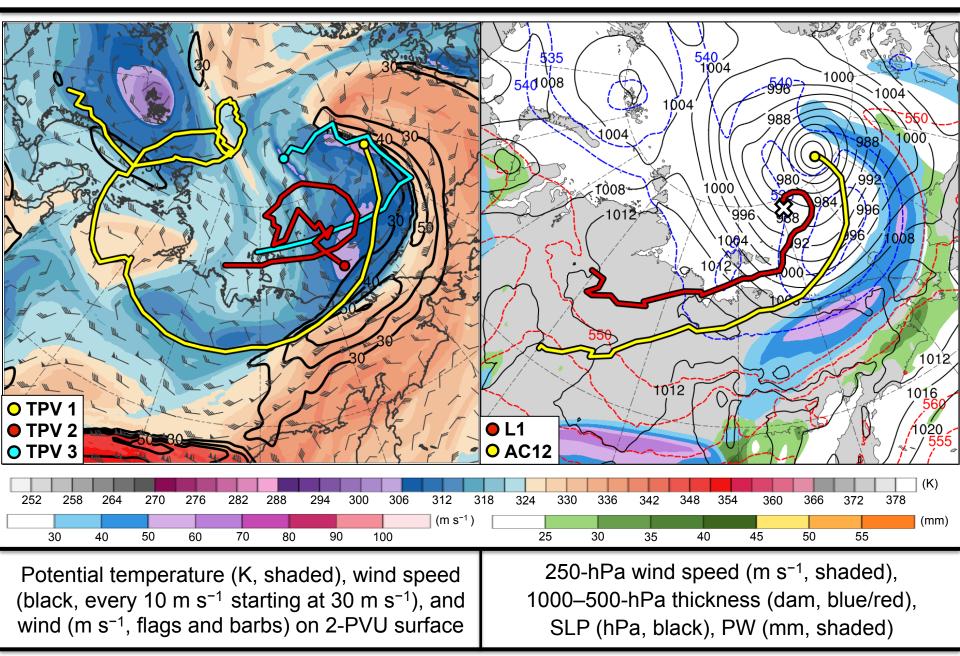


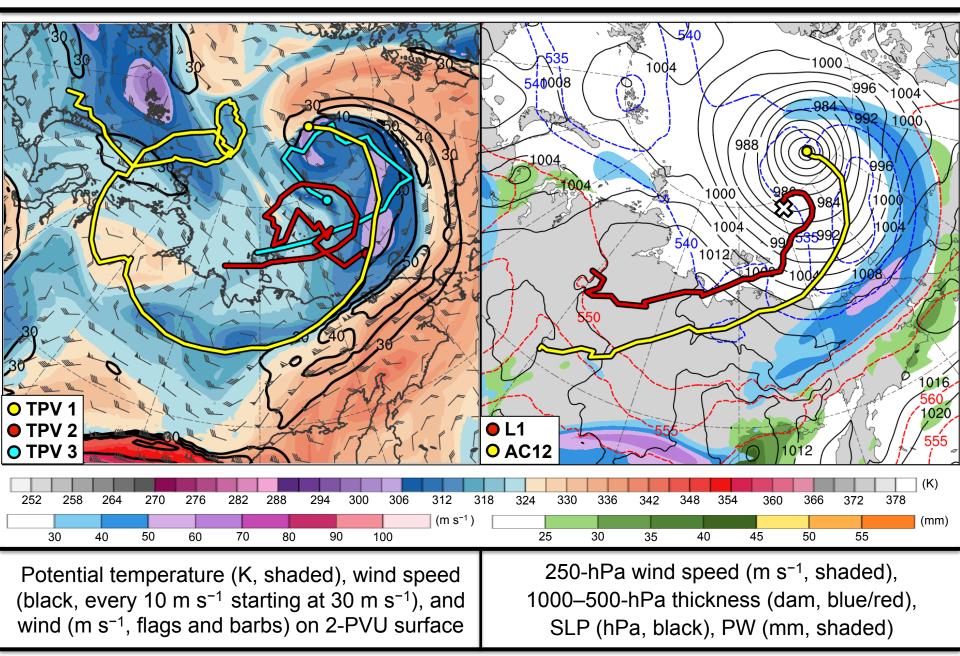


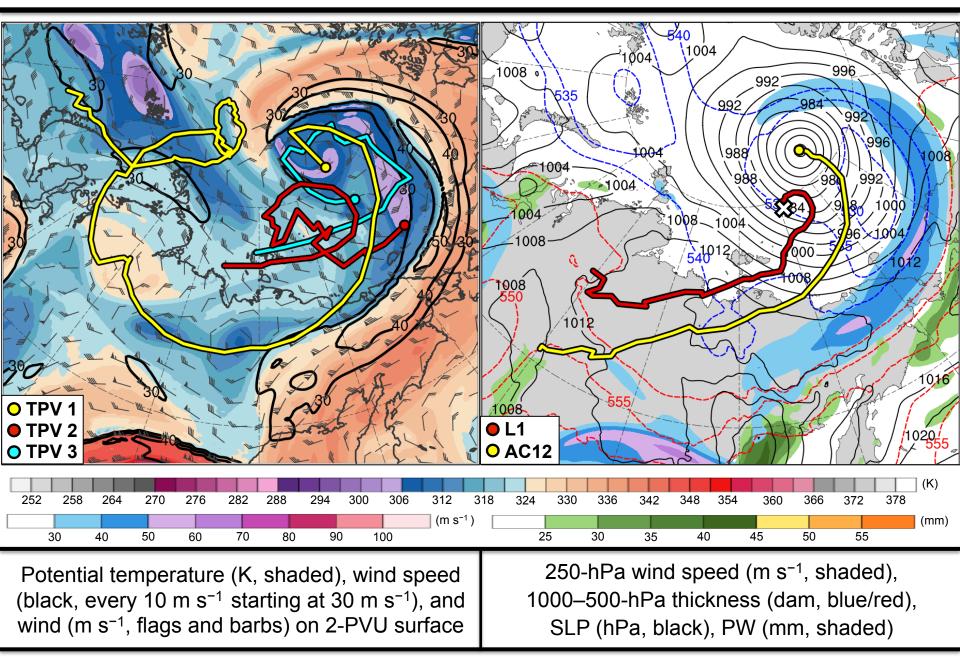


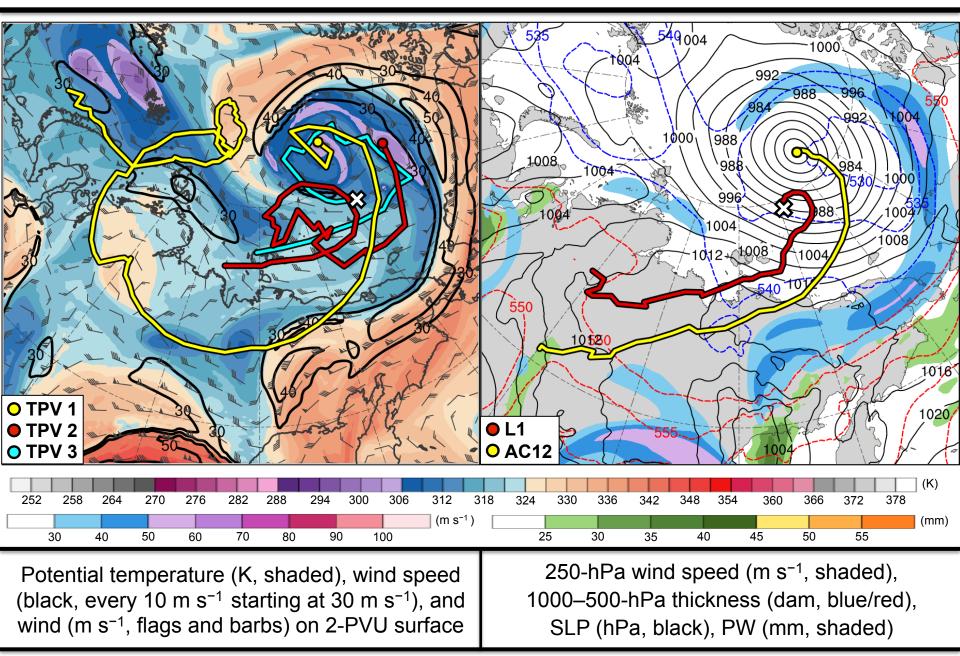


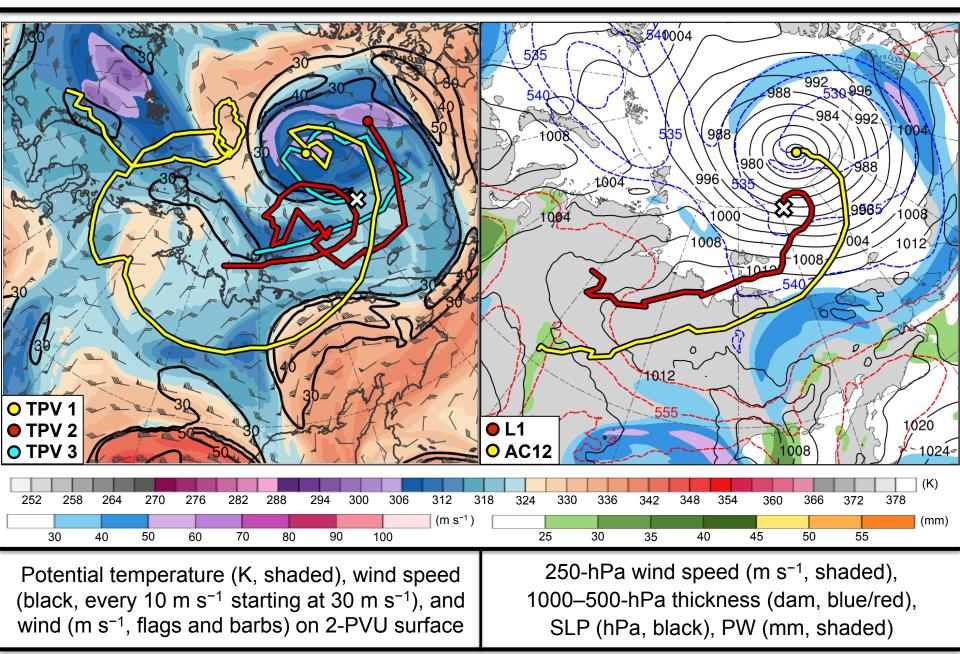


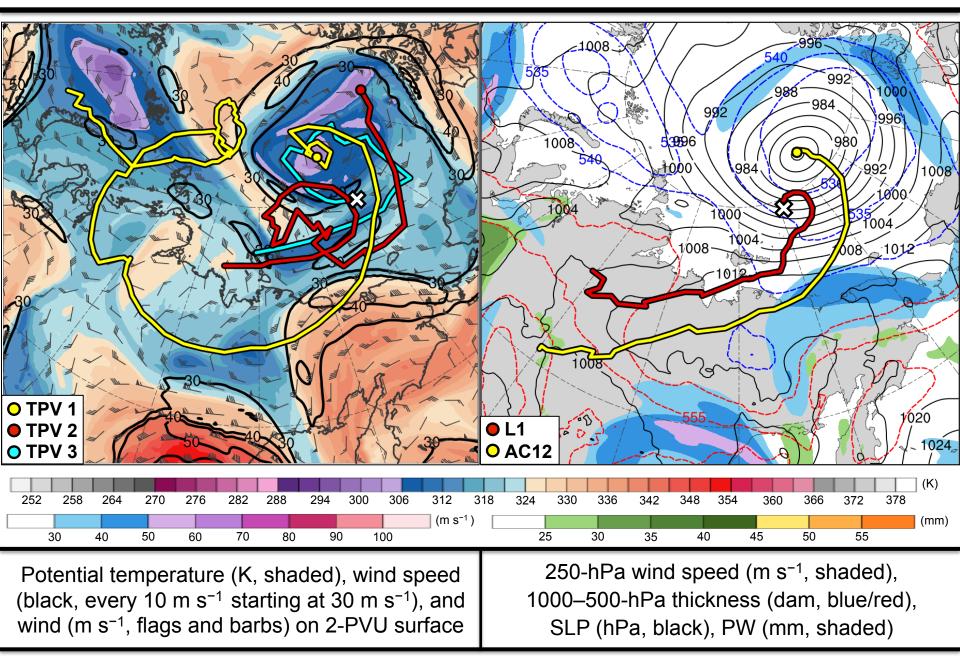


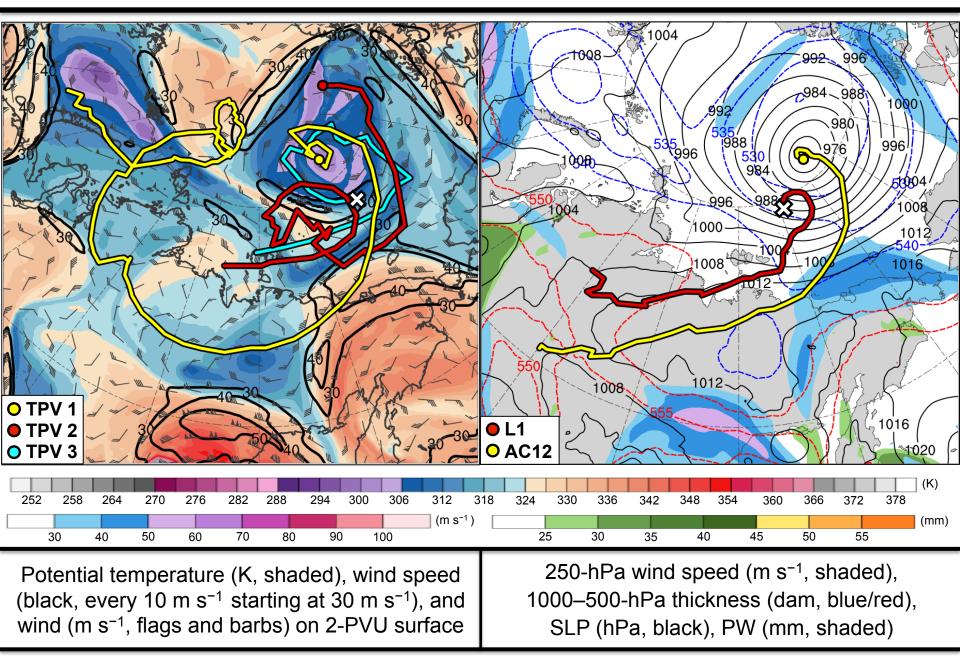




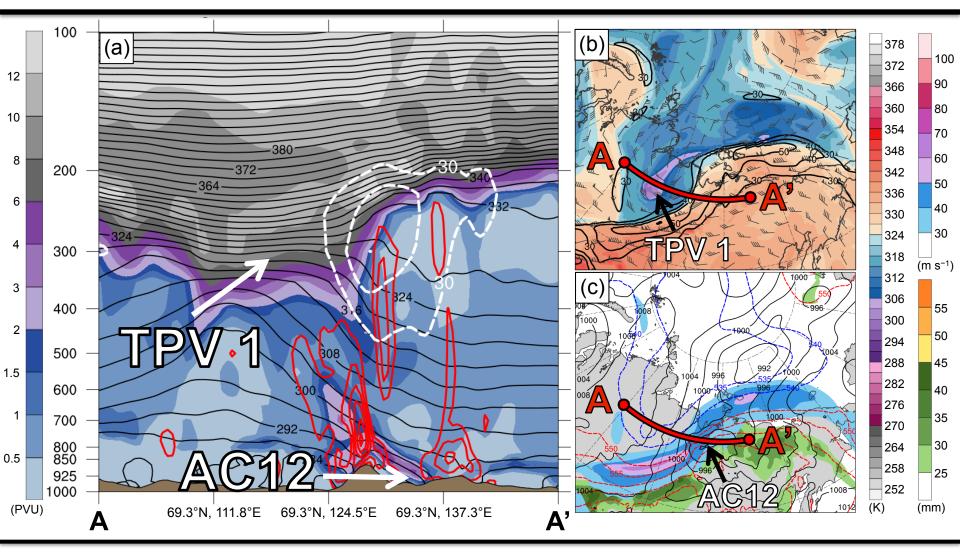








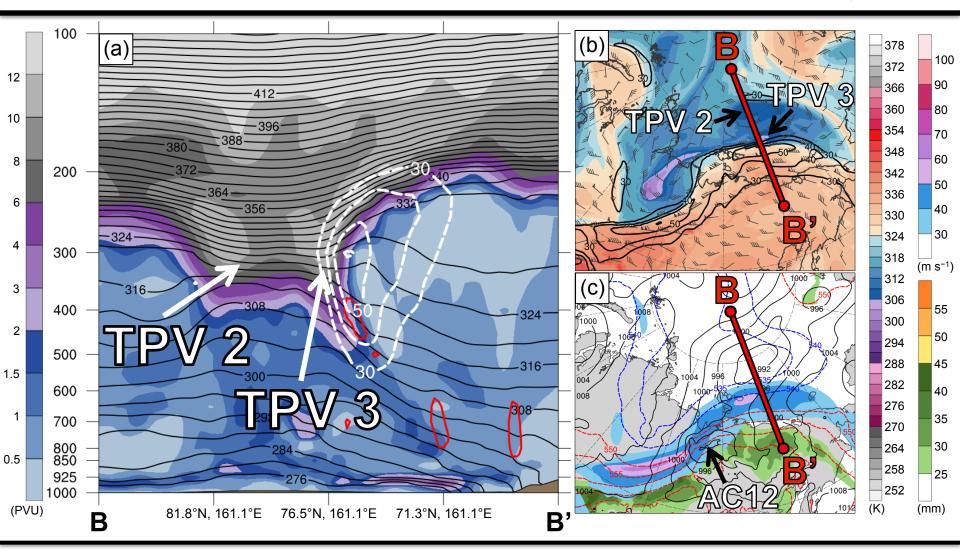
Cross Sections



(a) PV (PVU, shaded), θ (K, black), ascent (red, every 3.0 × 10⁻³ hPa s⁻¹), and wind speed (dashed white, m s⁻¹); (b) DT (2-PVU surface) θ (K, shaded), wind speed (black, m s⁻¹), and wind (m s⁻¹, flags and barbs); (c) 250-hPa wind speed (m s⁻¹, shaded),1000–500-hPa thickness (dam, blue/red), SLP (hPa, black), PW (mm, shaded)

Cross Sections

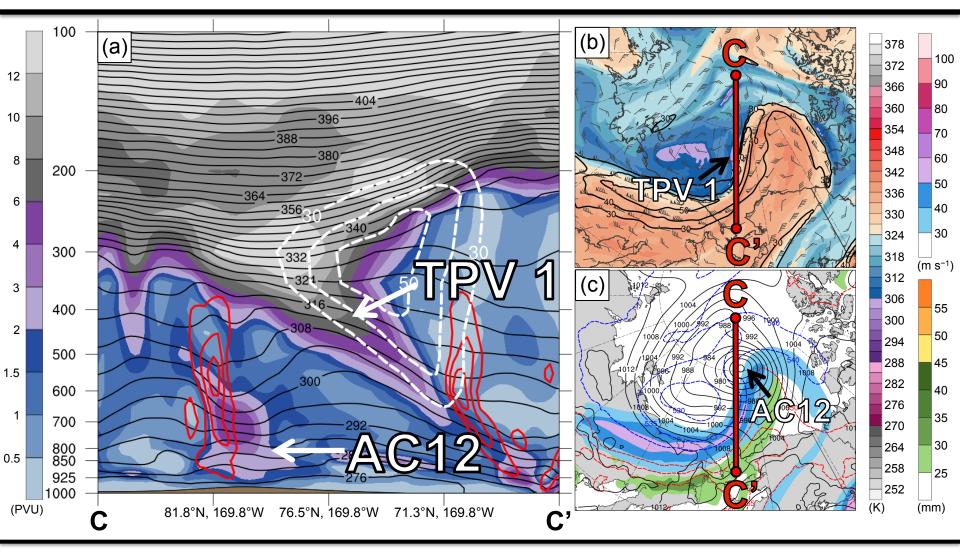
0600 UTC 4 Aug 2012



(a) PV (PVU, shaded), θ (K, black), ascent (red, every 3.0 × 10⁻³ hPa s⁻¹), and wind speed (dashed white, m s⁻¹); (b) DT (2-PVU surface) θ (K, shaded), wind speed (black, m s⁻¹), and wind (m s⁻¹, flags and barbs); (c) 250-hPa wind speed (m s⁻¹, shaded),1000–500-hPa thickness (dam, blue/red), SLP (hPa, black), PW (mm, shaded)

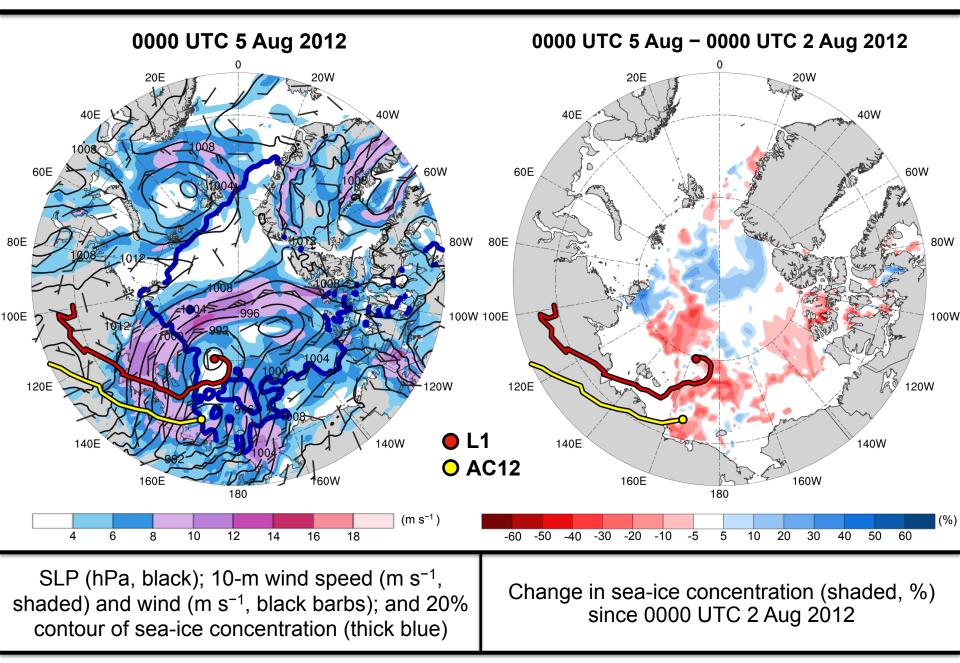
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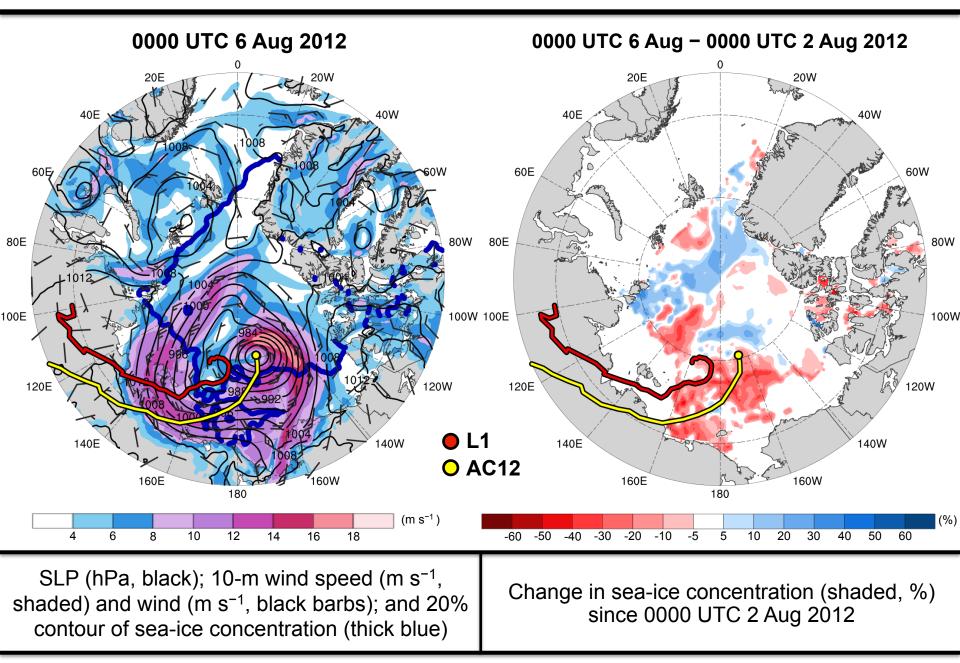
0000 UTC 6 Aug 2012

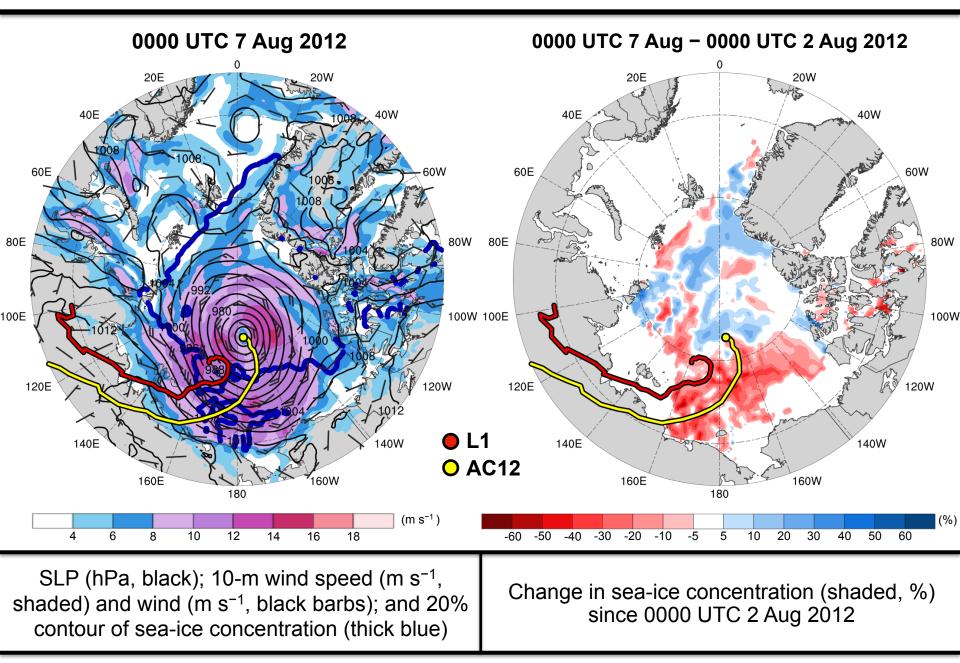


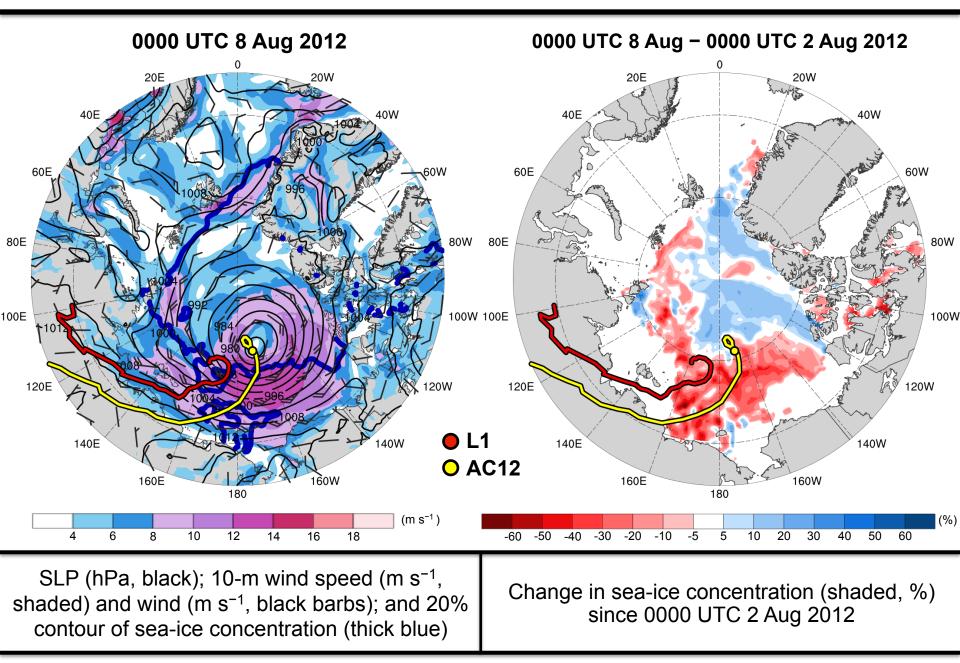
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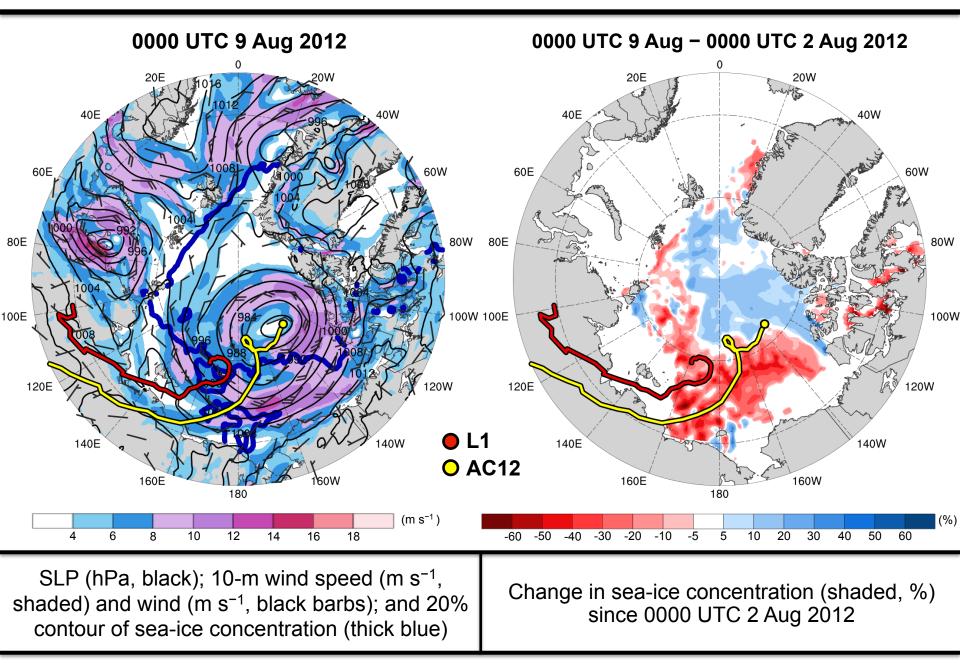
Reduction in Arctic Sea Ice

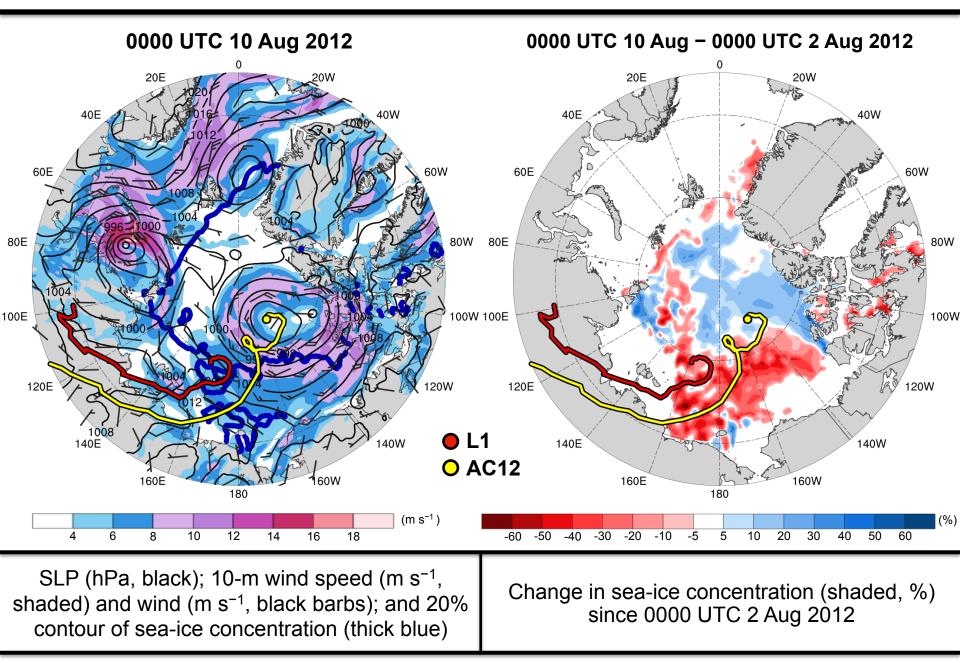


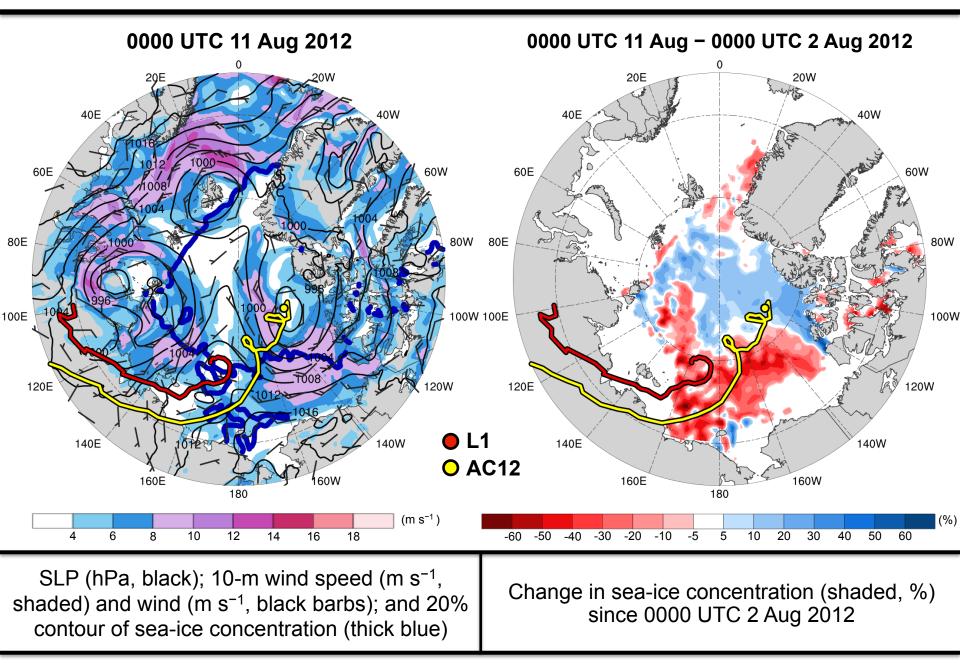


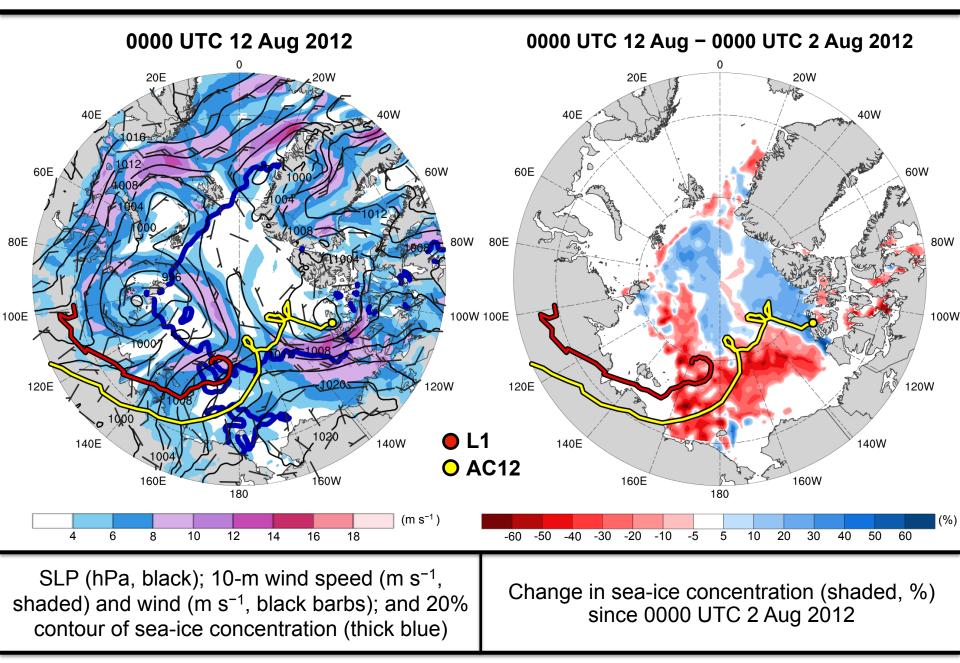


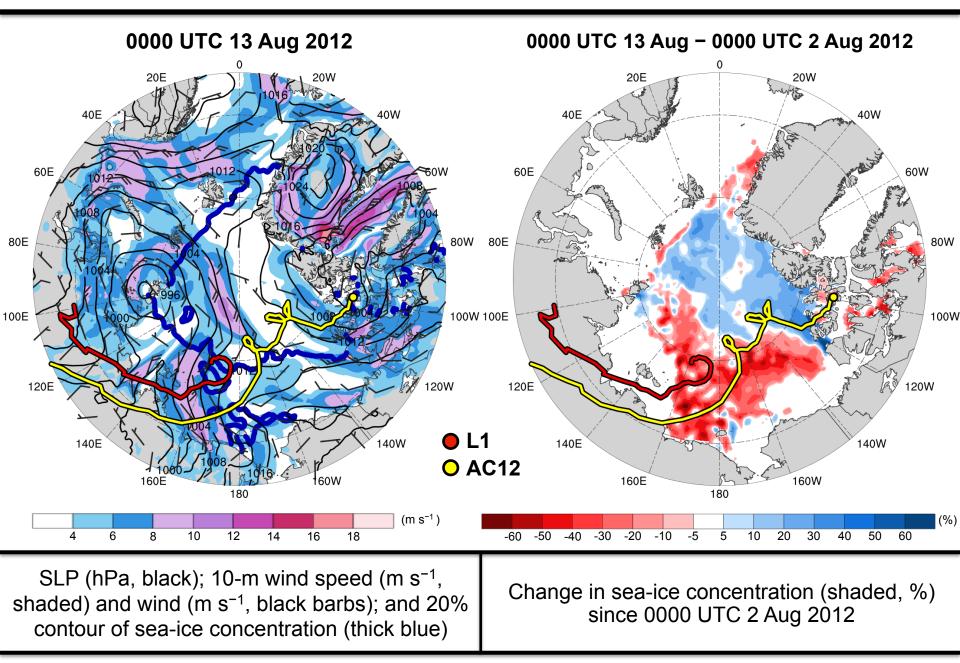


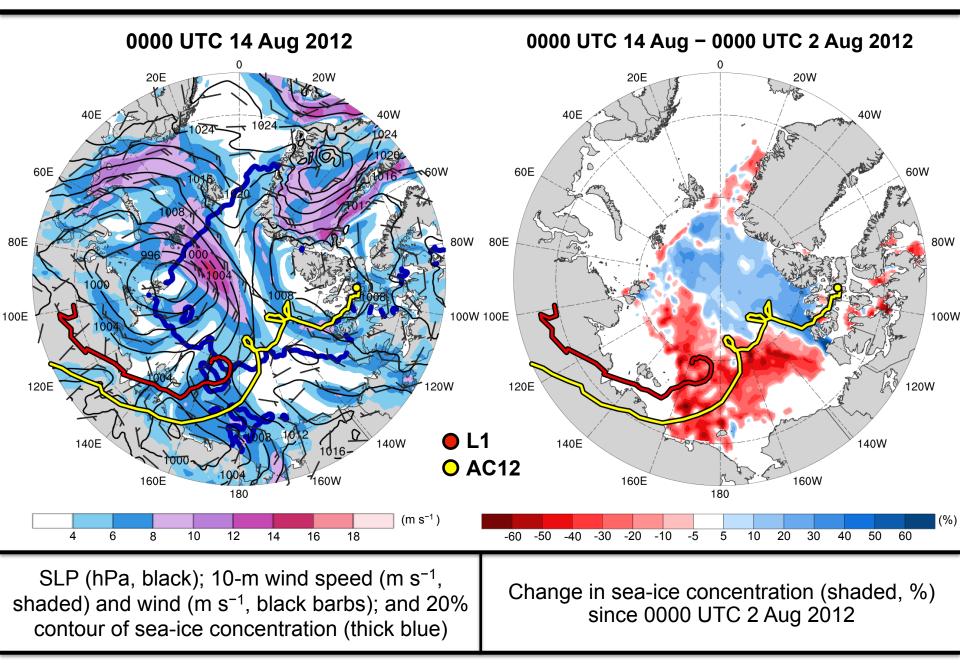












Conclusions

- TPV 1 approaches and interacts with AC12 in a region of strong baroclinicity, likely supporting the development of AC12 via baroclinic processes
- TPV-jet interactions involving TPV 1, TPV 2, and TPV 3 likely contribute to the formation of a dual-jet configuration and jet coupling over AC12
- Presence of warm, moist air and relatively strong lowertropospheric ascent in region of jet coupling likely contribute to formation of a potential vorticity (PV) tower associated with AC12

Conclusions

- Interaction between TPV 1 and the PV tower associated with AC12 likely supports the intensification of AC12
- L1 interacts and merges with AC12, which may further support the intensification of AC12
- After attaining a minimum SLP of 962 hPa, AC12 moves slowly over Arctic, where its expansive surface wind field contributes to reductions in Arctic sea-ice extent

Questions? *Email:* kbiernat@albany.edu

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Acknowledgments

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