

Intensity Variations of Subtropical Potential Vorticity Streamers:

Impact on the Environment of the Subtropical Atlantic
and Tropical Cyclone Activity

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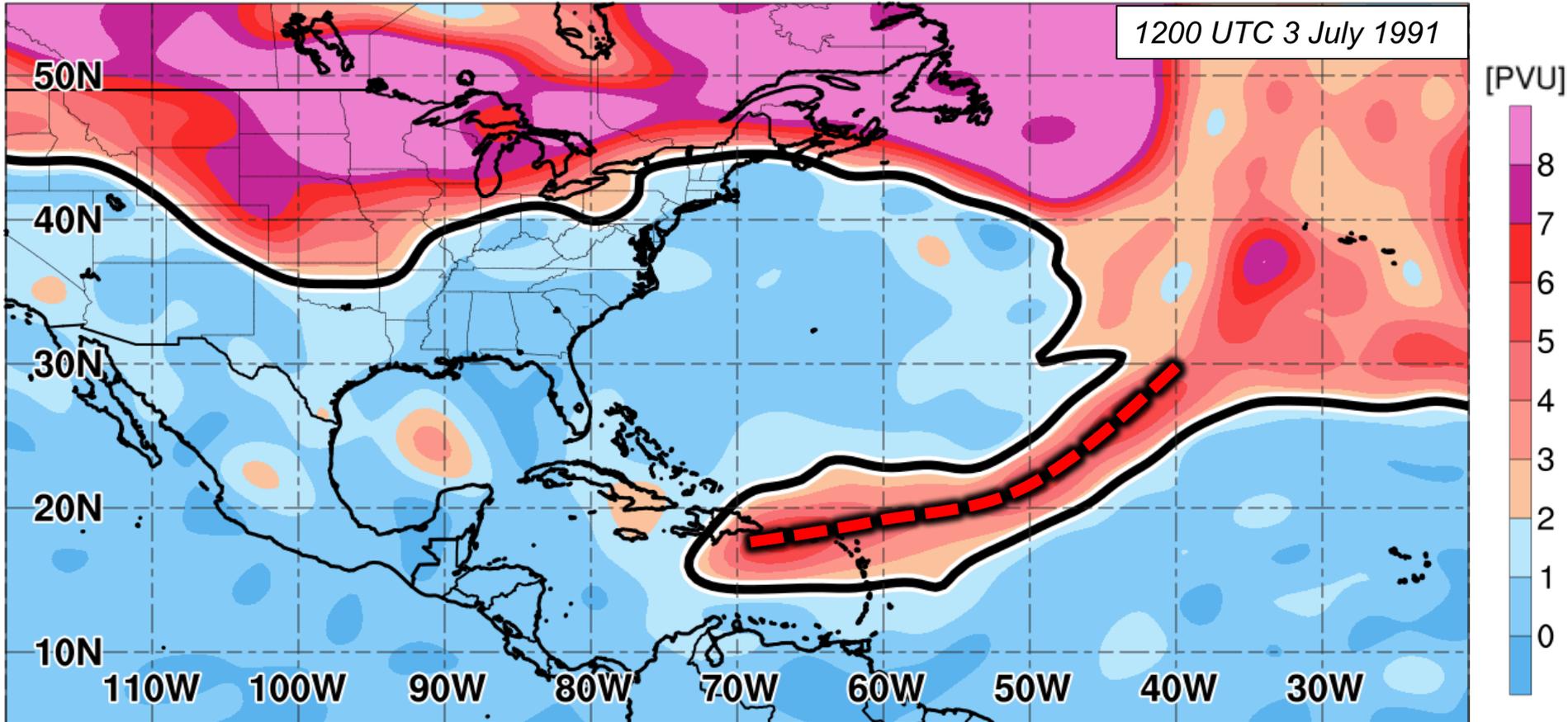
University at Albany, SUNY

NHC Seminar – 20 March 2017

What are PV Streamers

- Potential vorticity (PV) streamers are elongated filaments of high PV air
 - Correspond to positively tilted upper-tropospheric troughs

350-K PV (shaded, PVU), 2-PVU contour (black line)

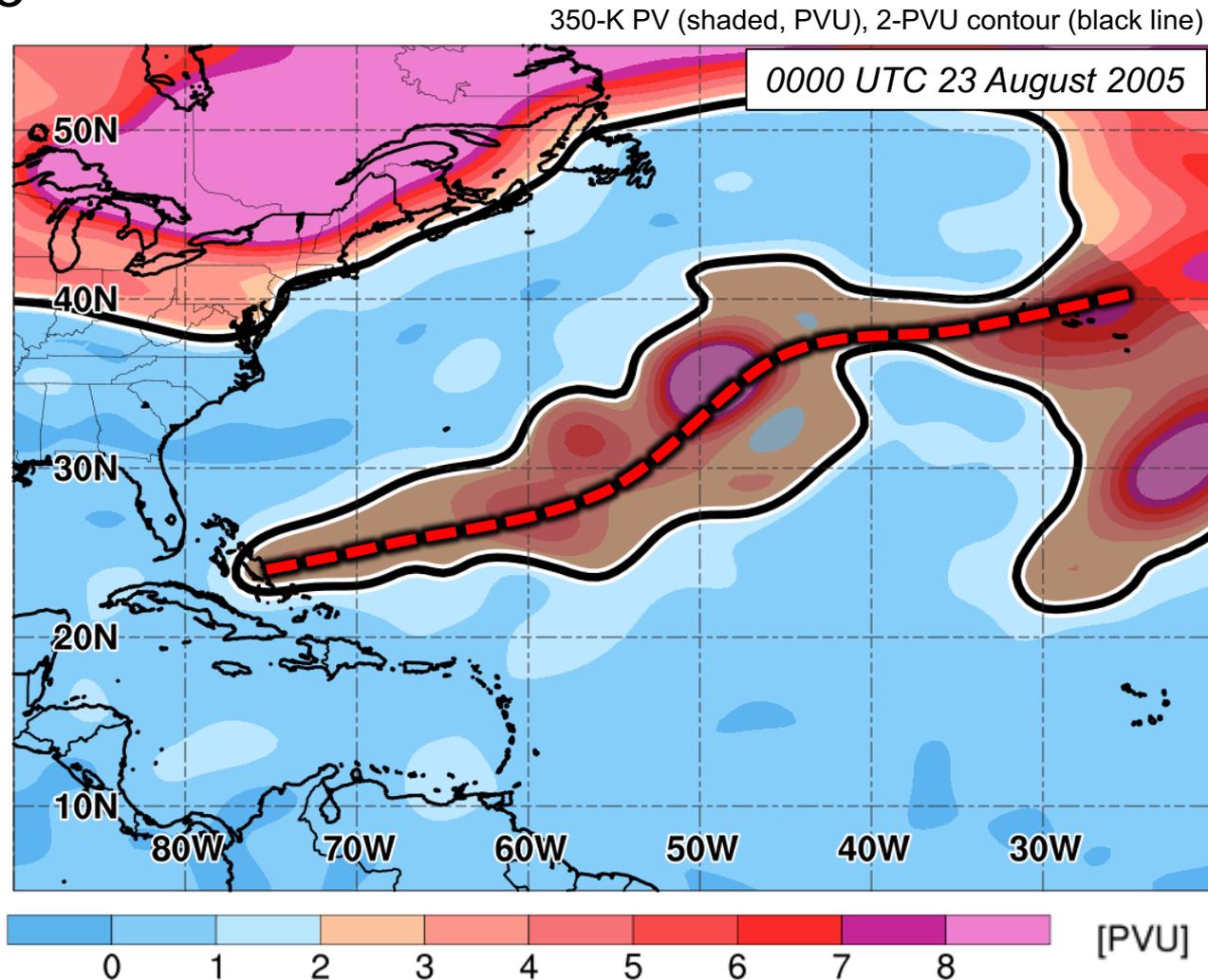


Motivation

A tale of two PV streamers

- Tale 1: August 2005

- ❖ PV streamer draped across subtropical Atlantic basin
- ❖ Small width and weak intensity in Bahamas
- ❖ A developing system in the Bahamas (Katrina) easily overcomes PV streamer induced shear

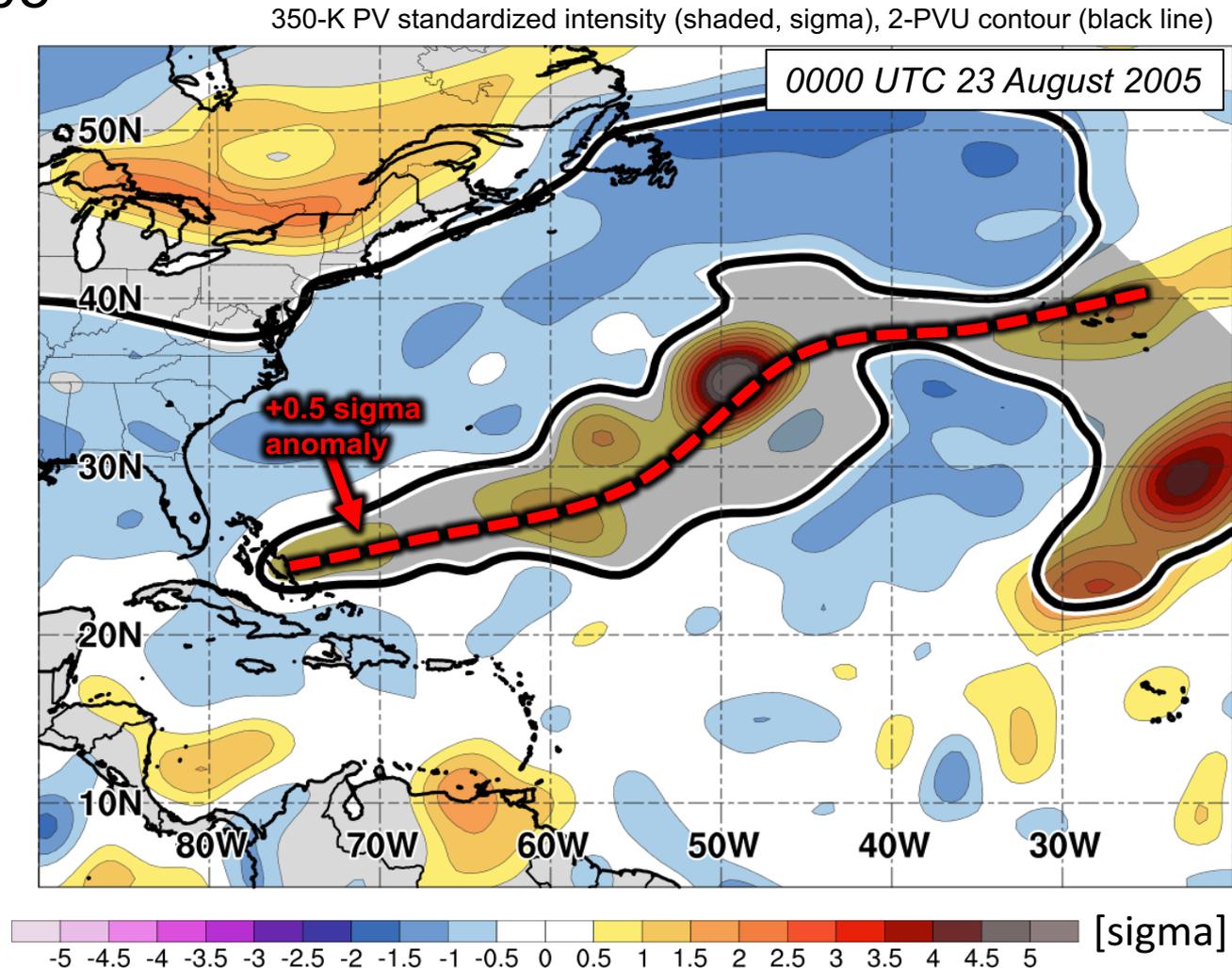


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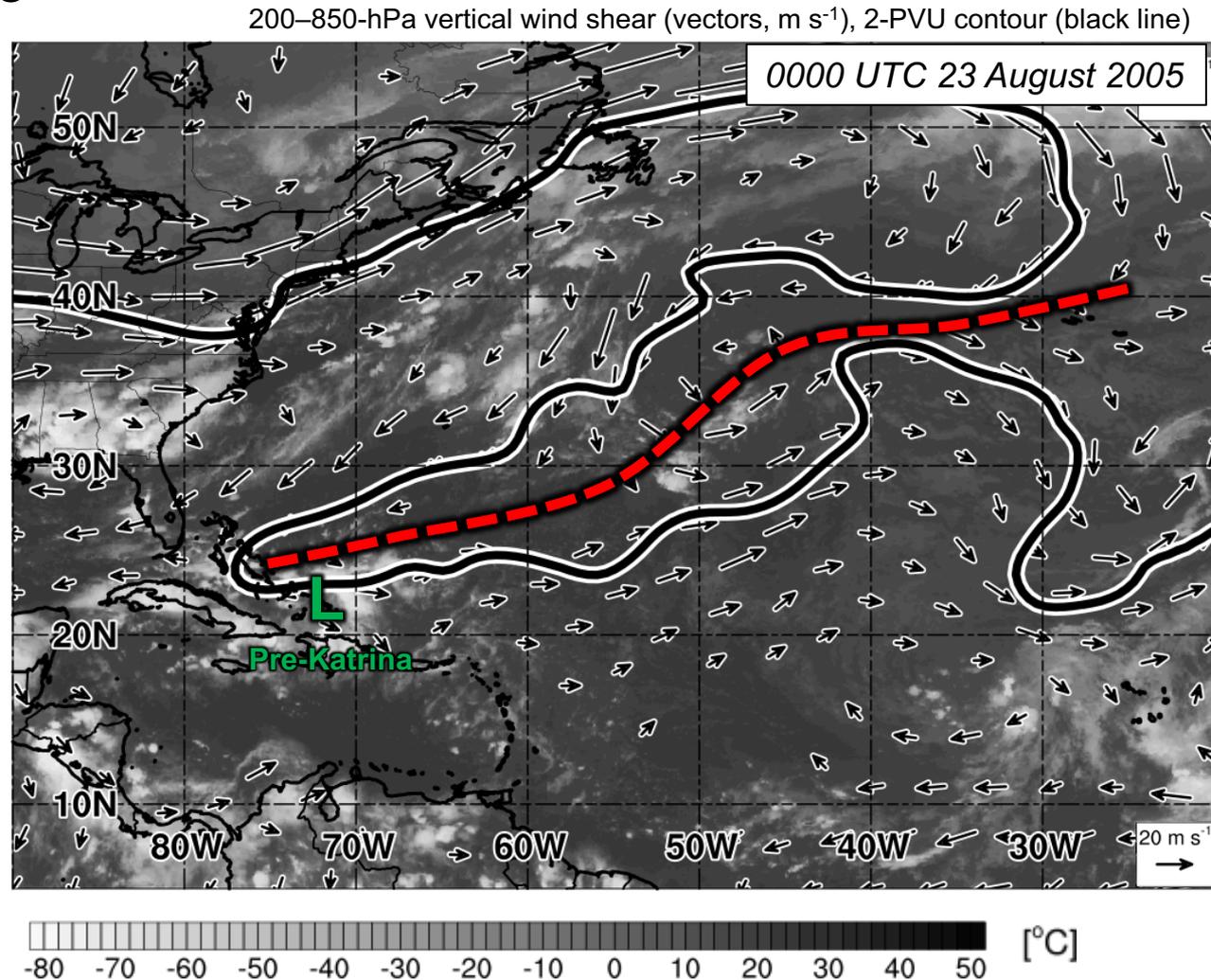


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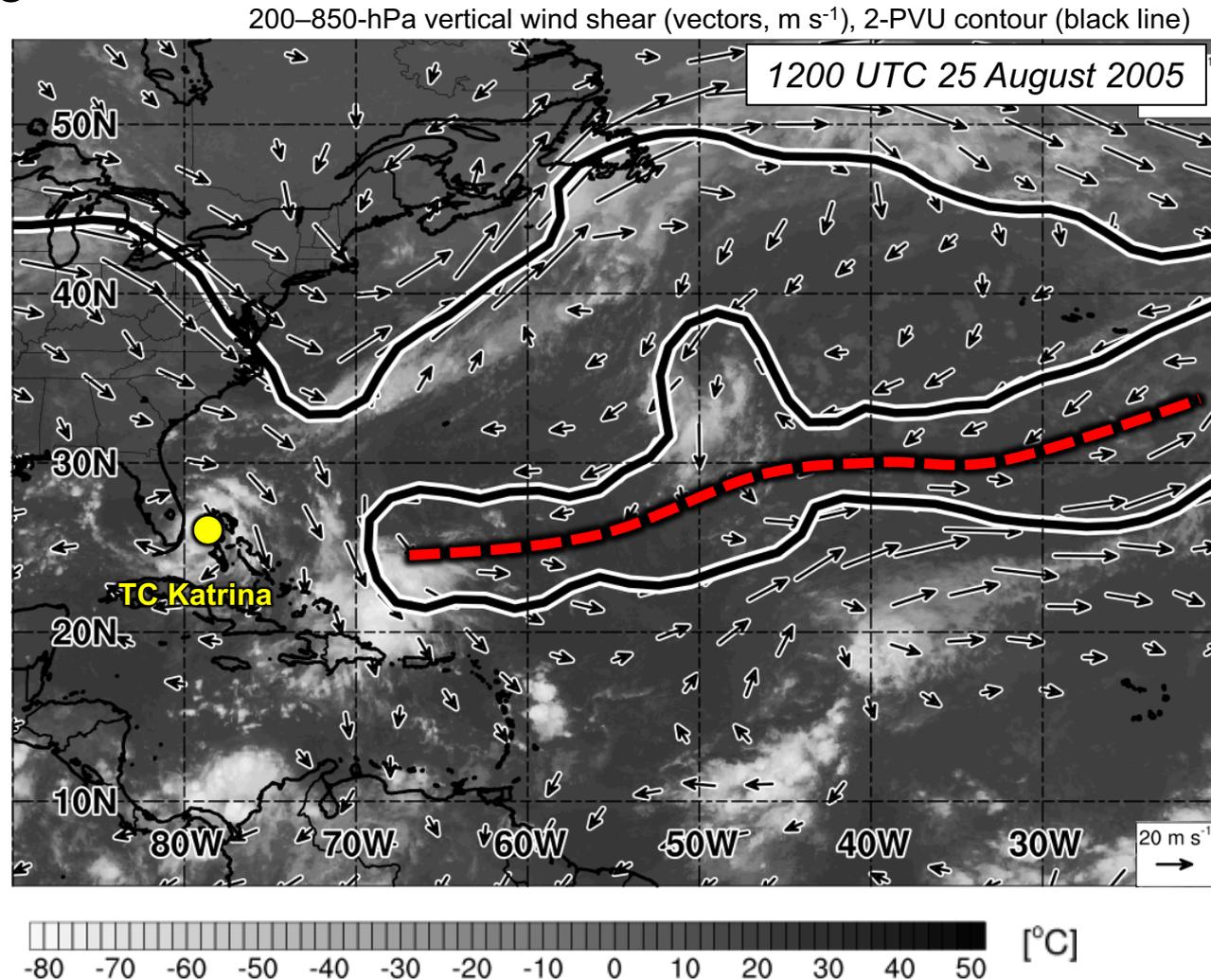


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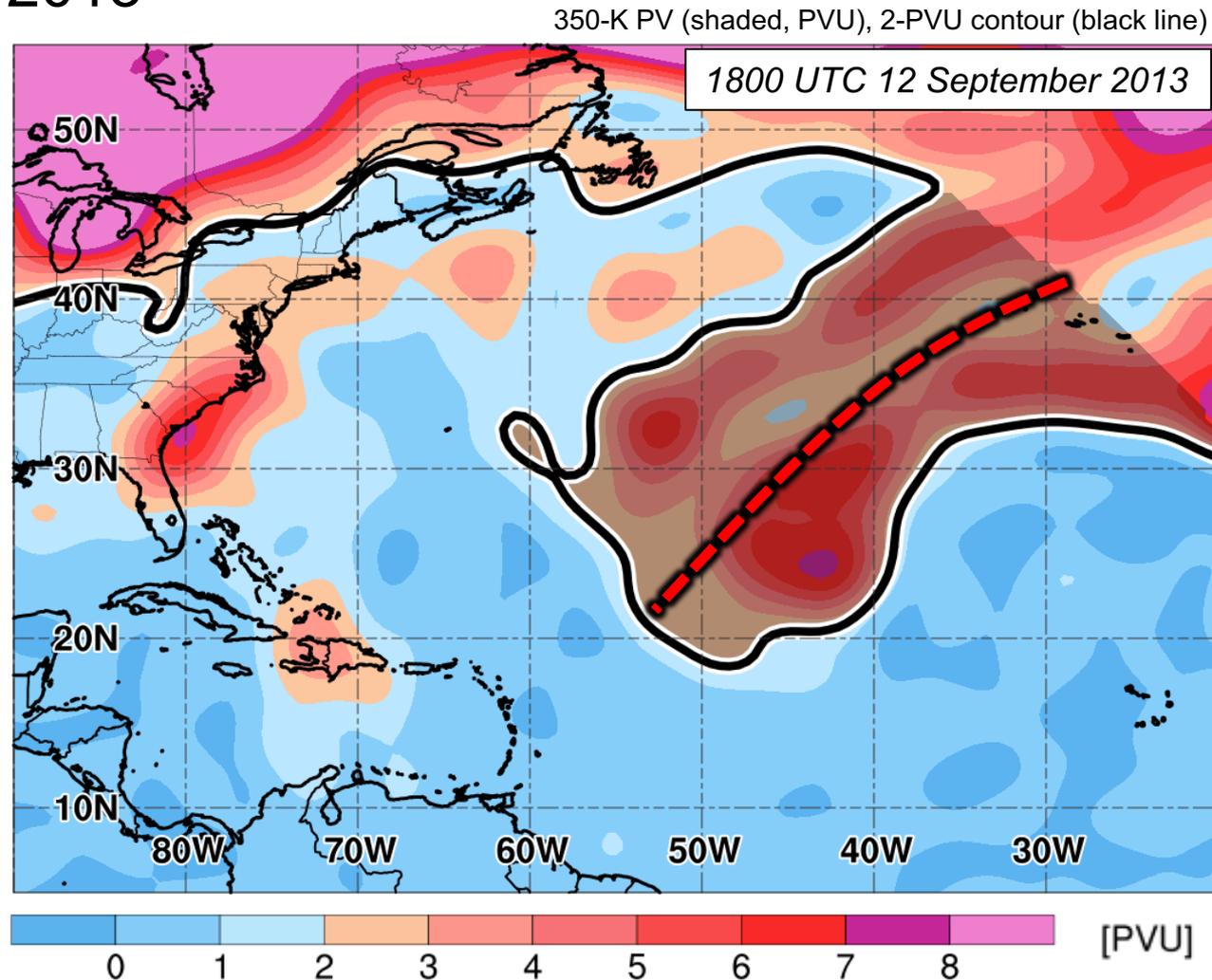


Motivation

A tale of two PV streamers

- Tale 2: September 2013

- ❖ PV streamer over the eastern Atlantic
- ❖ Large width and strong intensity in central Atlantic
- ❖ The PV streamer negatively interacted with TC Humberto due to high vertical wind shear and dry mid-latitude air

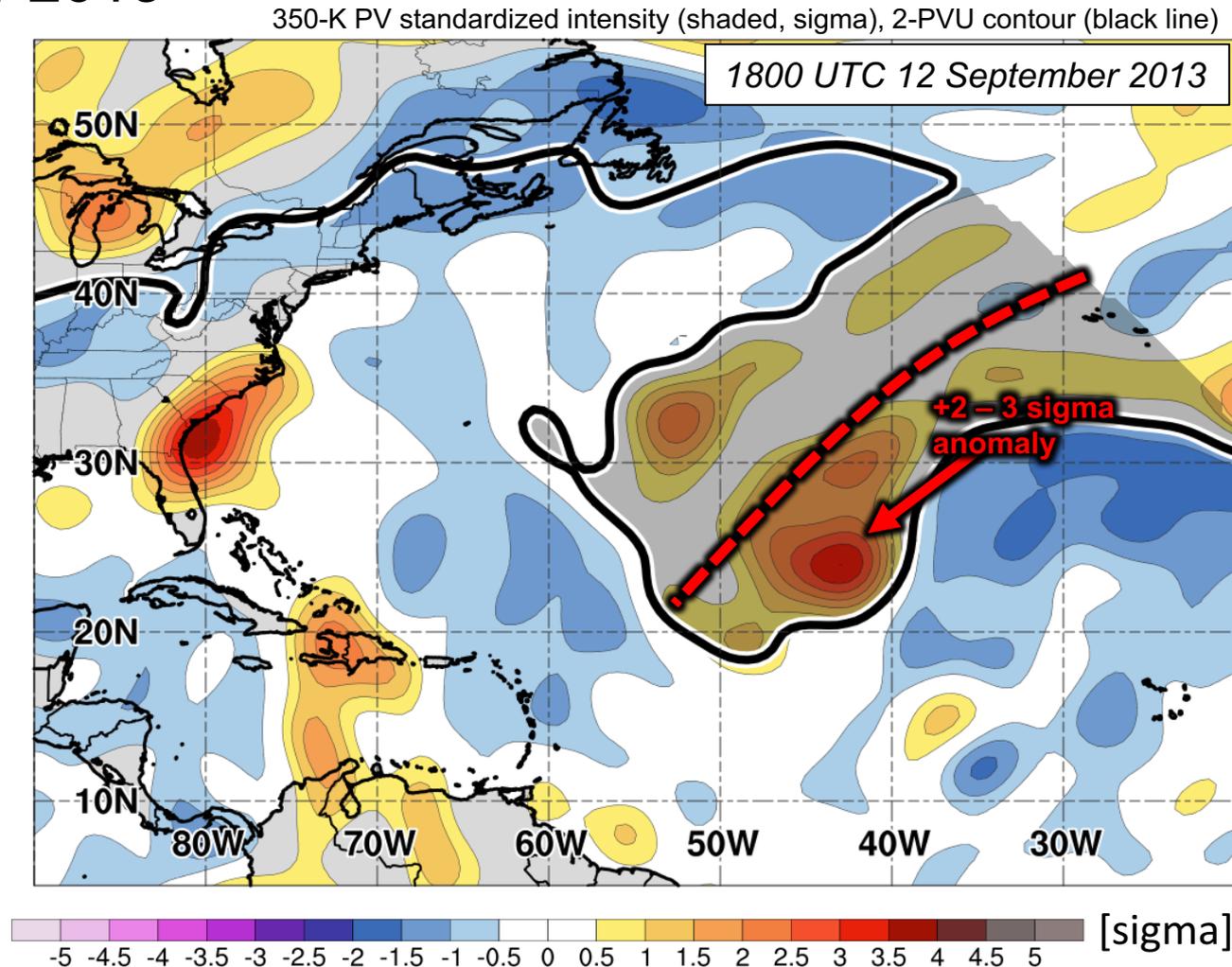


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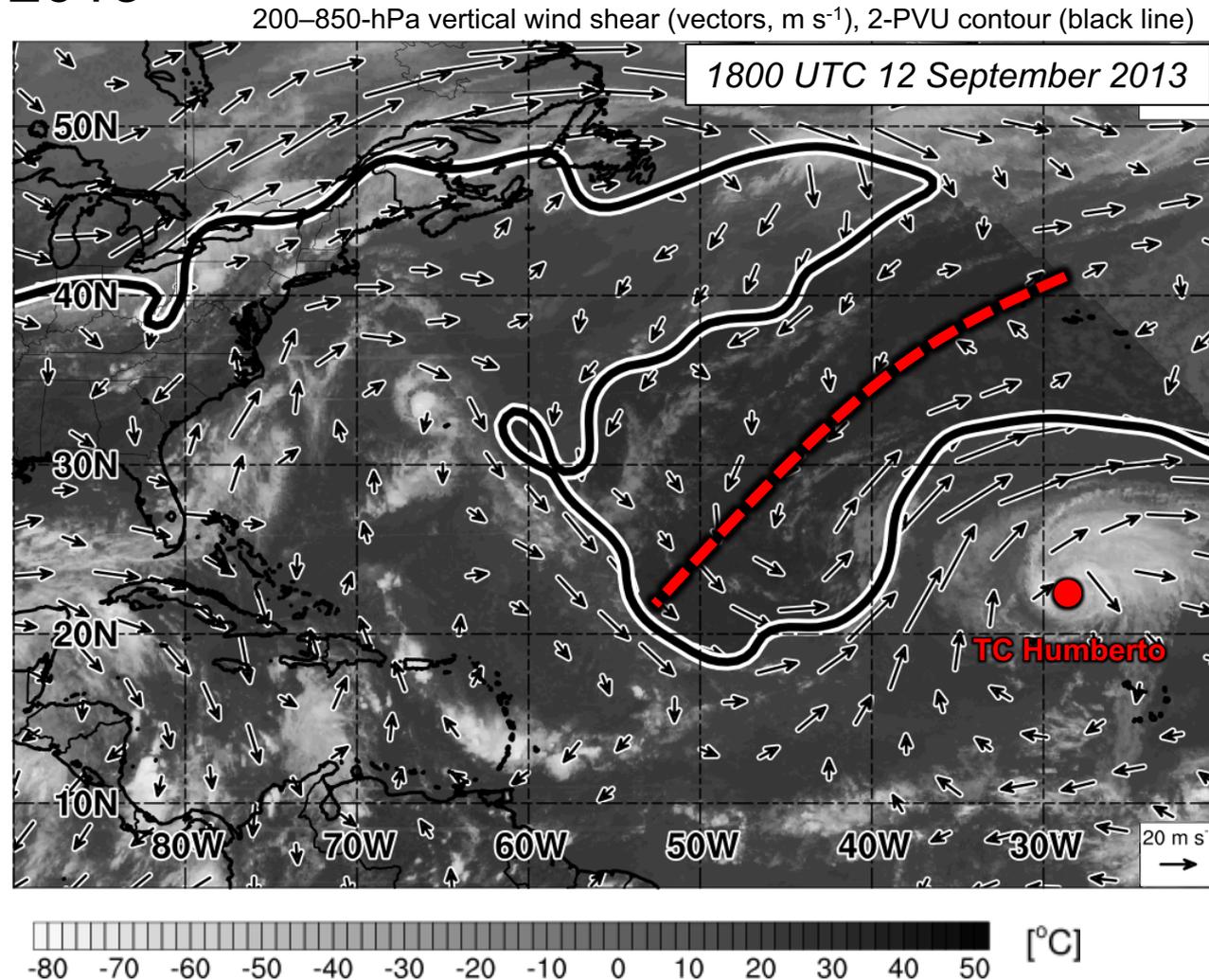


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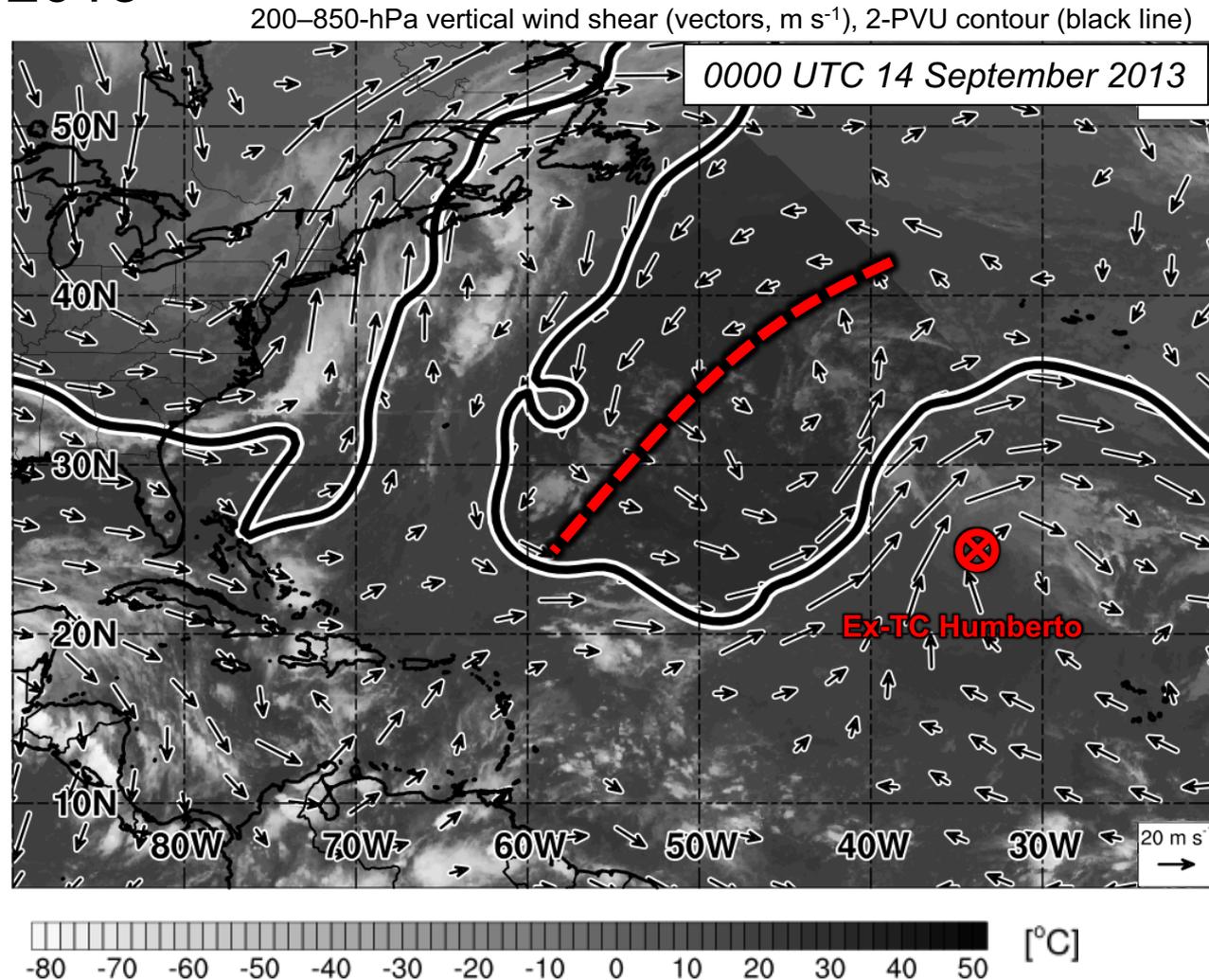
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These cases illustrate how different PV streamers alter the environment of TCs in their vicinity



Outline

- **Introduction**

 - **Motivation**

 - **Literature review of PV Streamers**

- **Methodology**

 - Identification of PV Streamers

 - Unique characteristics associated with each PV streamer

- **Results**

 - Climatological characteristics of PV streamers

 - Relationship with TC activity

 - Comparing composite PV streamers of different intensity

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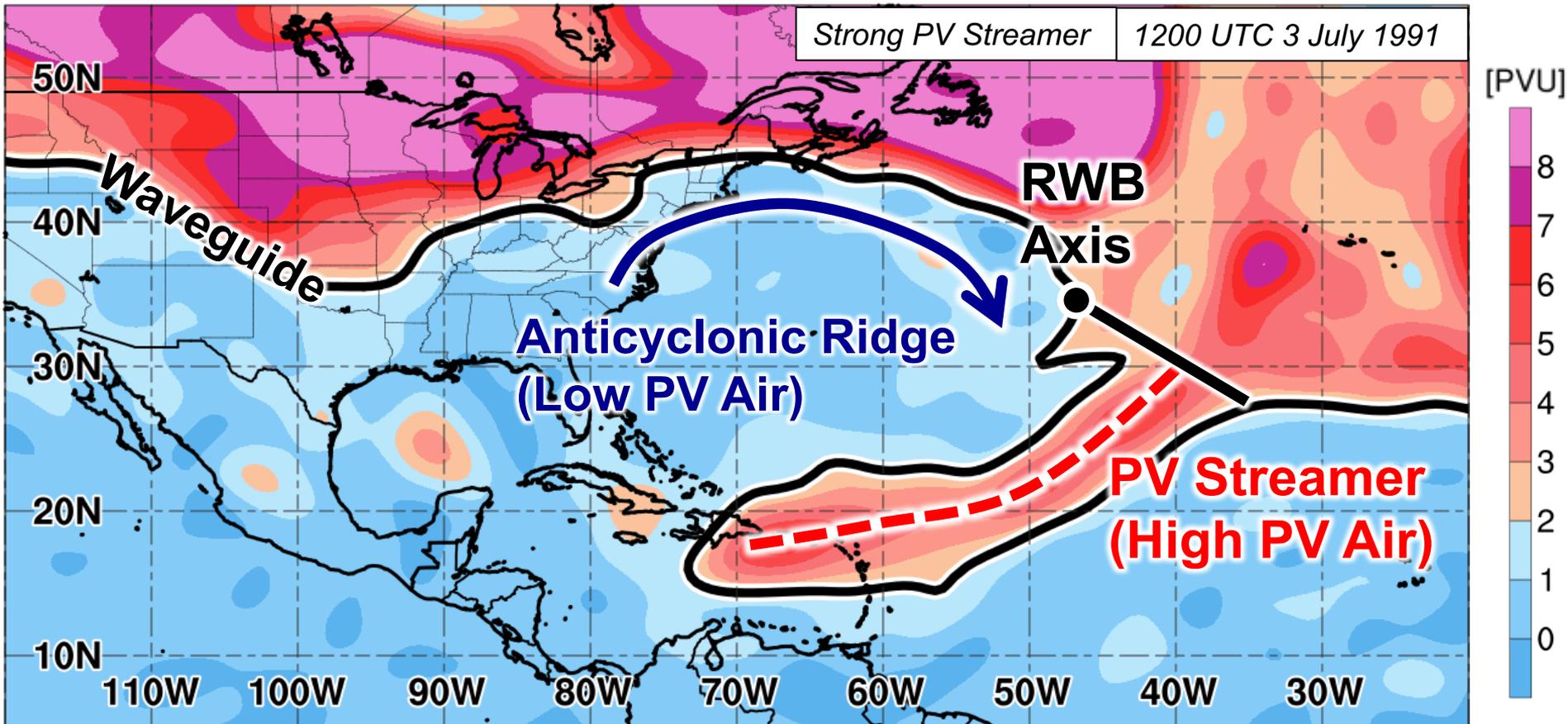
- **Relationship with TC activity**

- **Comparing composite PV streamers of different intensity**

Literature Review

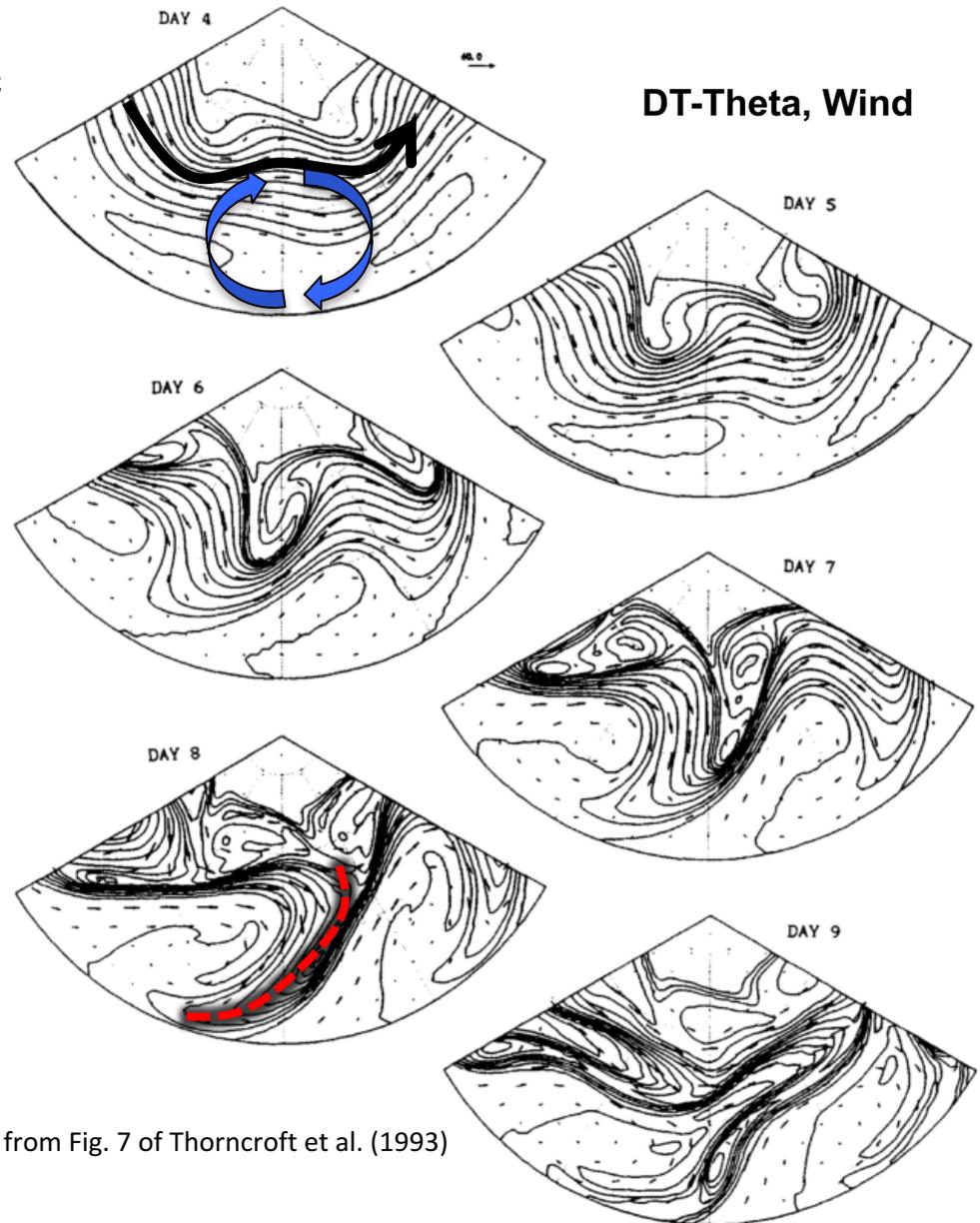
Literature Review

- Potential vorticity (PV) streamers are elongated filaments of high PV air
 - Occur in tandem with Rossby wave breaking (RWB)
 - Occur in the subtropical Atlantic with anticyclonic RWB (AWB)
 - Often Impact tropical cyclone (TC) activity in the Atlantic basin (McTaggart-Cowan et al. 2013, Galarneau et al. 2015)



Literature Review: Rossby Wave Breaking

- RWB Manifests as two characteristic baroclinic wave lifecycles
 - ✧ Anticyclonic Wave Breaking (LC1, AWB)
 - ✧ Cyclonic Wave Breaking (LC2, CWB)
- Anticyclonic meridional shear found equatorward of the waveguide
 - ✧ Thin positively tilted PV streamer
- Cyclonic meridional shear found poleward of the waveguide
 - ✧ Thick negatively tilted PV streamer

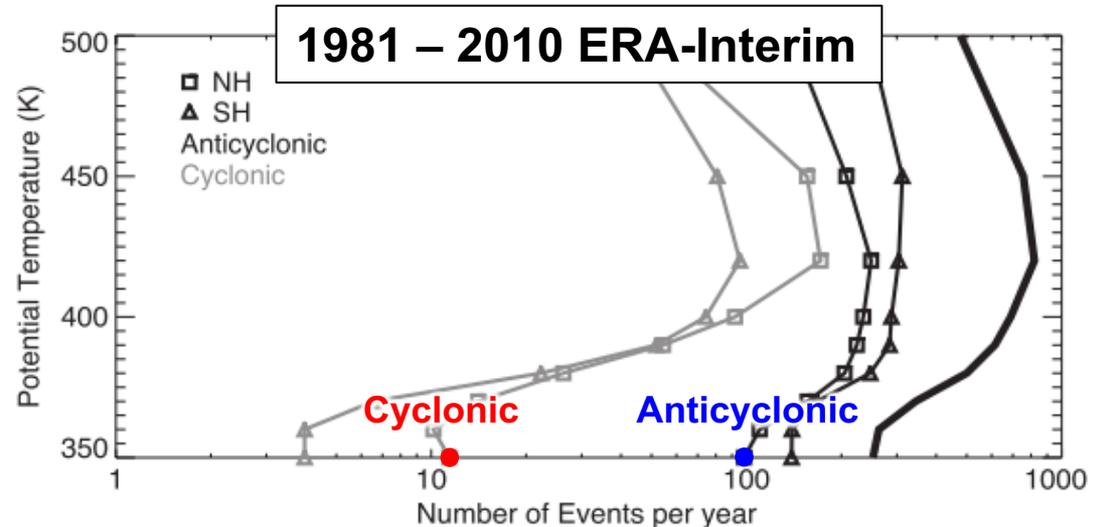
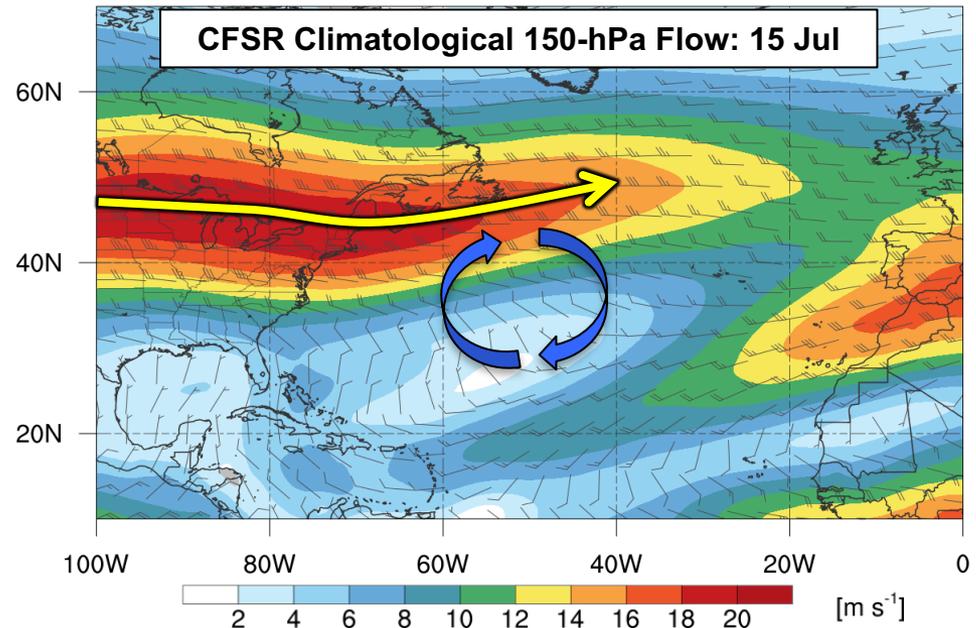


This study emphasizes the AWB pathway of occurrence

Adapted from Fig. 7 of Thorncroft et al. (1993)

Literature Review: RWB Frequency

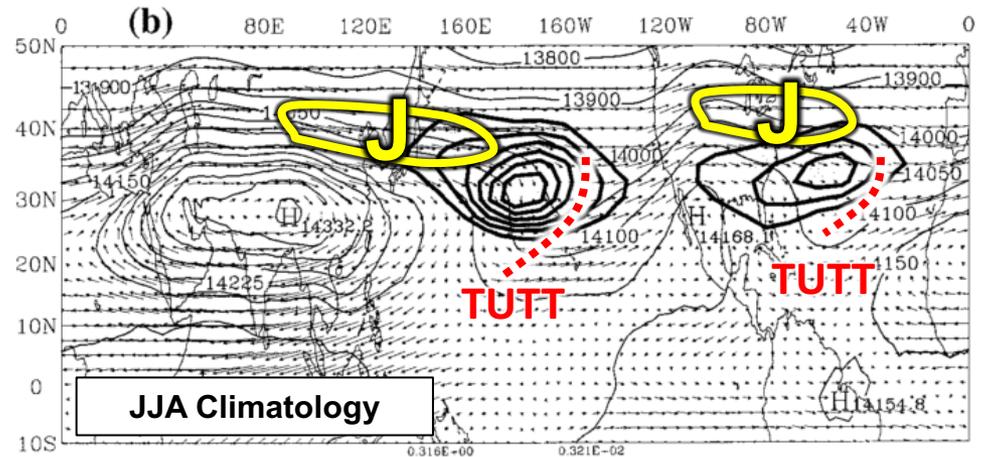
- Anticyclonic RWB is much more common in subtropical latitudes
 - ✧ Equatorward of waveguide, background barotropic meridional shear is anticyclonic
- Results in primarily positively oriented PV streamers in low latitudes



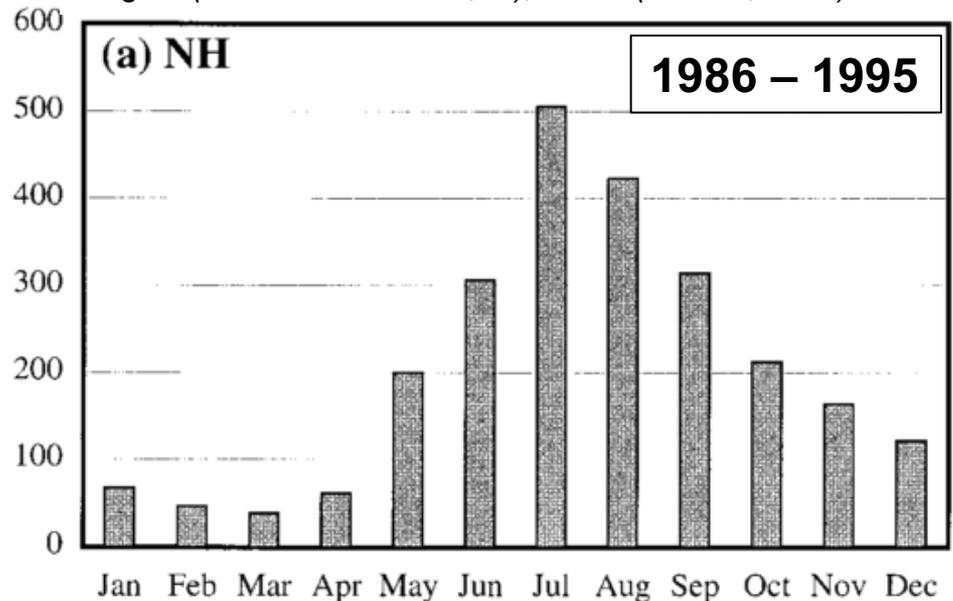
Adapted from Figs. 5 of Homeyer and Bowman (2013)

Literature Review: RWB Frequency

- RWB is favored over oceanic basins near jet exit regions
- RWB frequency peaks in the warm season when the time-mean flow along the waveguide is the weakest
- A weaker waveguide allows more perturbations in flow to become significant relative to the time mean flow (Holton 2004).



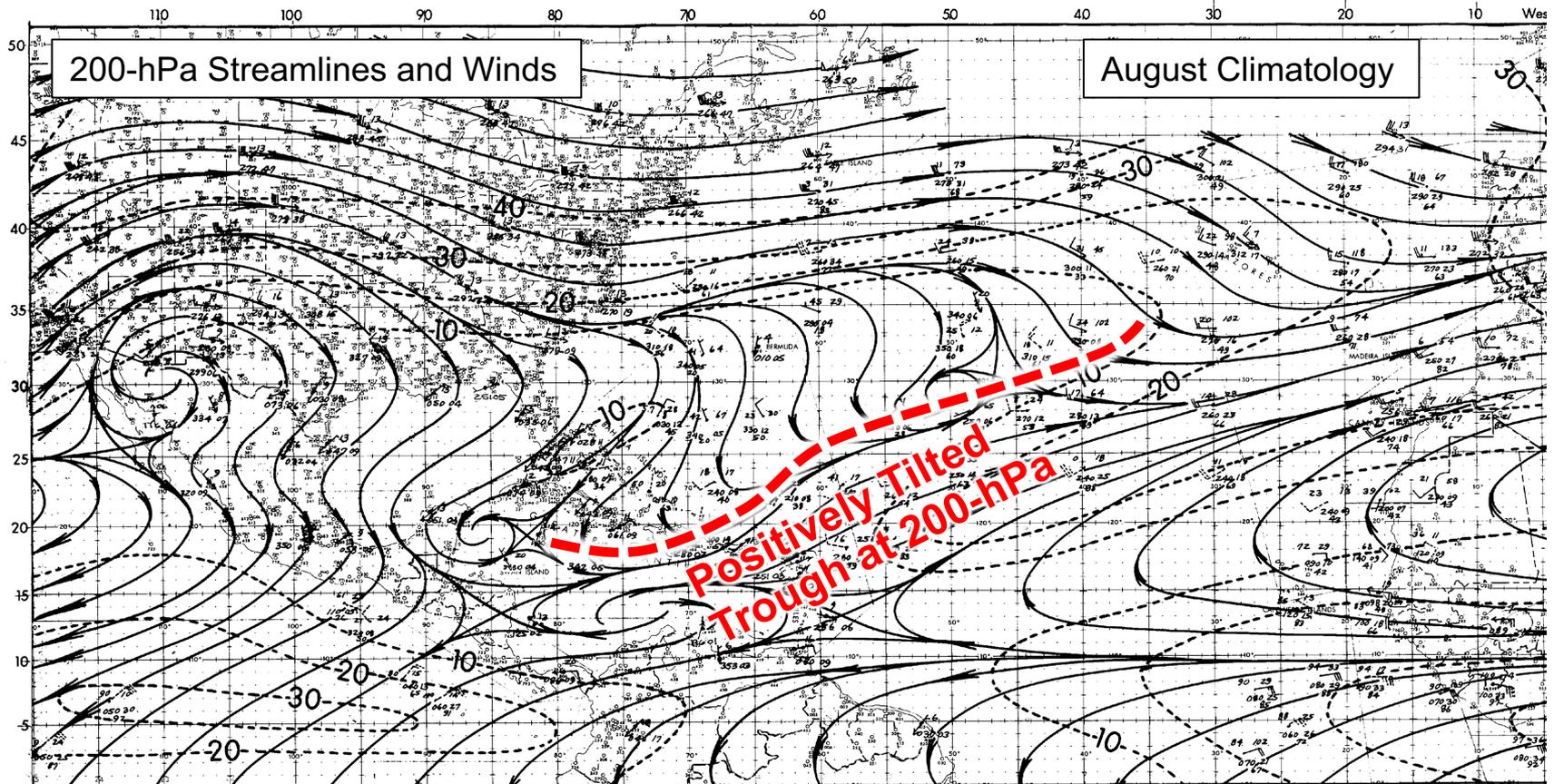
RWB frequency (thick black contours), 150-hPa geopotential heights (thin black contours, m), winds (vectors, $m\ s^{-1}$)



Adapted from Figs. 3a and 4b of Postel and Hitchman (1999)

Literature Review: RWB Linkage to TUTTs

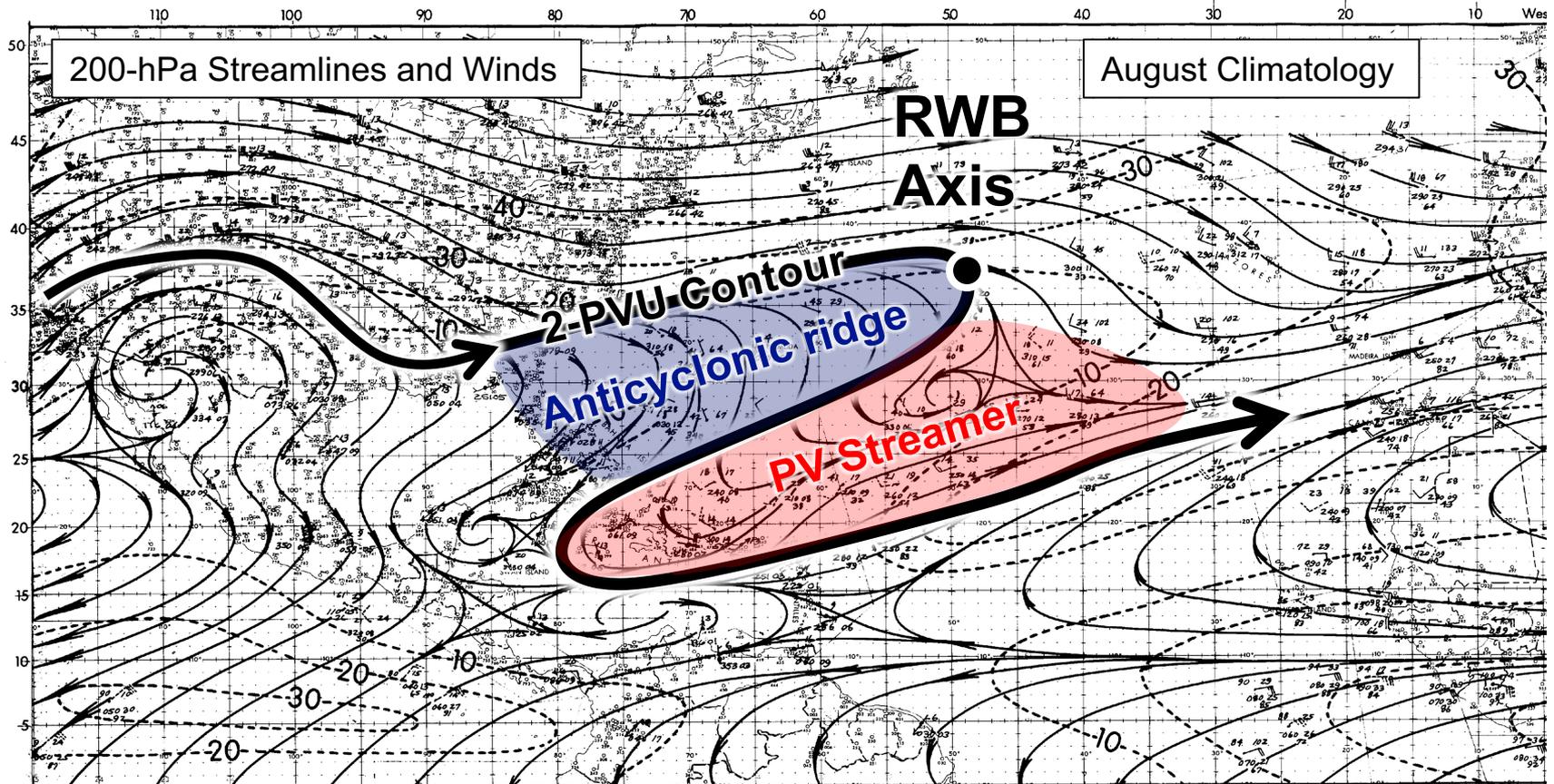
- Tropical Upper-Tropospheric Trough (TUTT) are located in both the subtropical NPAC and NATL basins (Sadler 1975, 1976)
 - Also called Mid Ocean Troughs (MOTs) in the time mean
- The dynamical component of the TUTT/MOT has been described in literature as PV streamers (Postel and Hitchman 1999; McTaggart Cowan et al. 2013)



Adapted from Sadler (1975)

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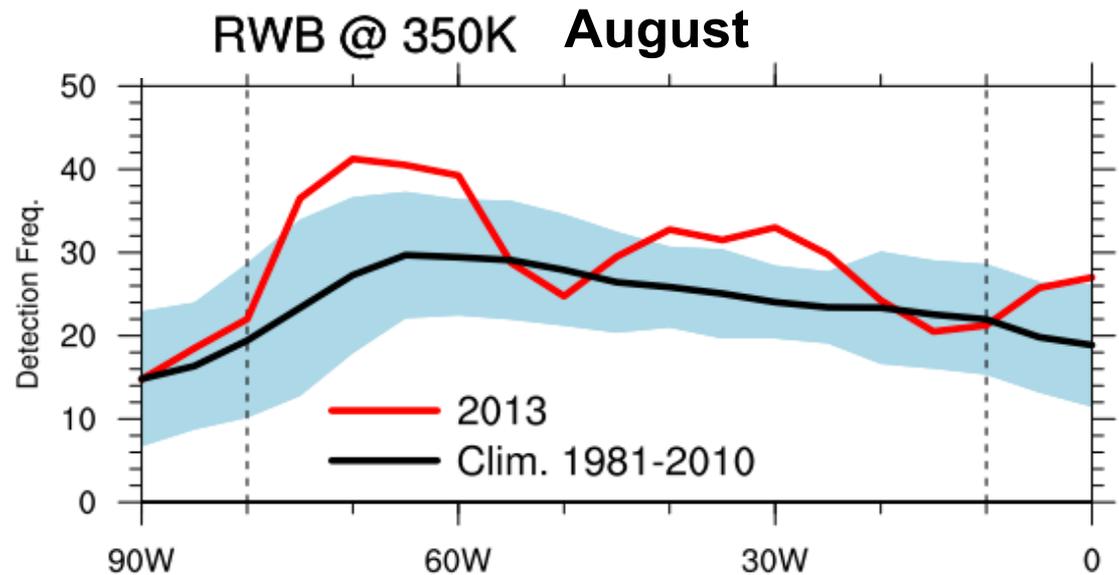


Adapted from Sadler (1975)

Literature Review: RWB Impact on TC Activity

- Recent research by Zhang et al. (2016) on the frequency of RWB in the Atlantic basin has revealed a significant negative correlation with tropical cyclone (TC) activity in August.

	HurrN	TSN	ACE
RWBFreq	-0.47	-0.39	-0.49
MDR	0.43	0.54	0.47
RSST	0.50	0.57	0.50
Niño-3.4	-0.35	-0.26	-0.34
SahelR	0.34	0.30	0.30
HCI	-0.45	-0.36	-0.41
NAO/CPC	-0.19	-0.23	-0.02
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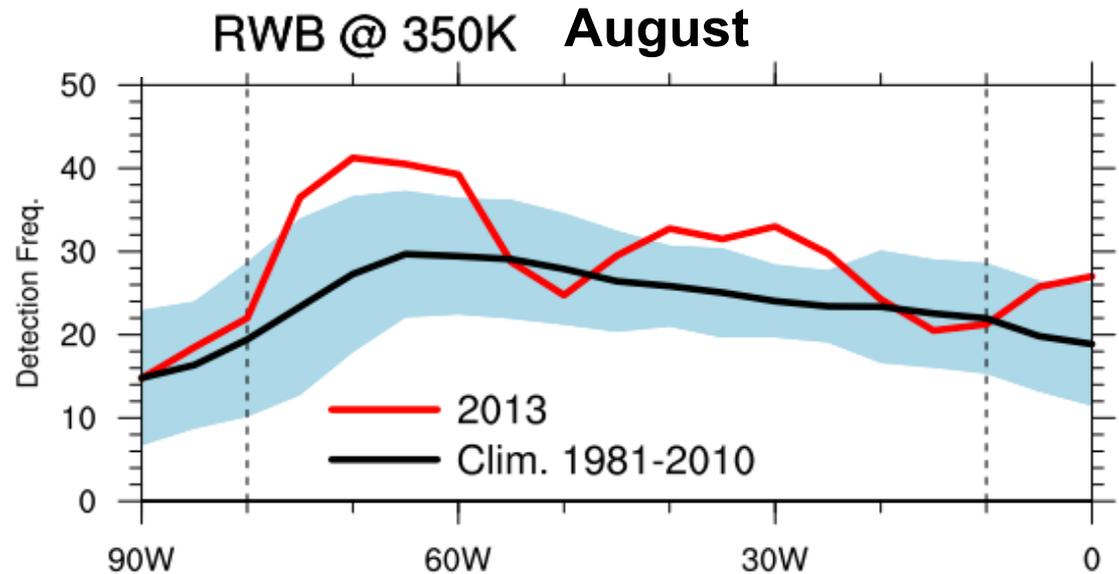


Adapted from Table 1 and Fig. 5a in Zhang et al. (2016)

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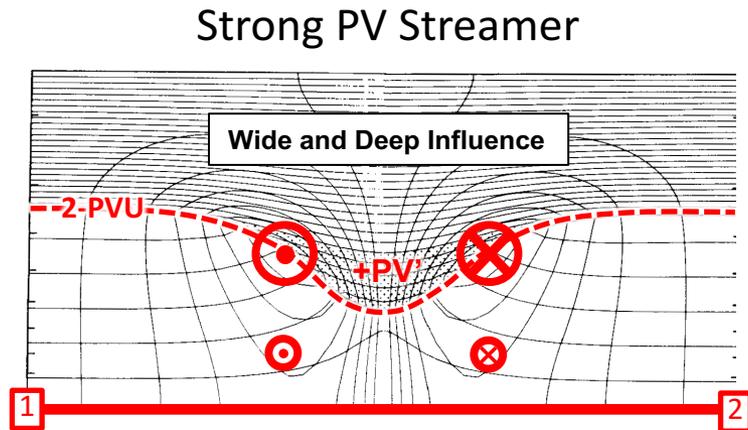


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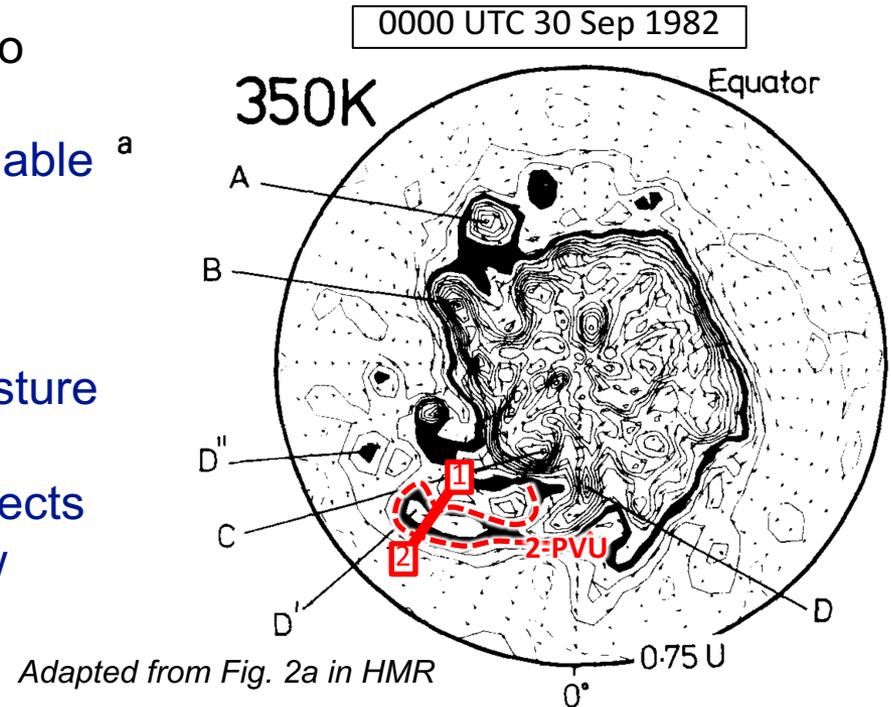
However, not all RWB events are created equal!

Literature Review: PV Streamer Intensity

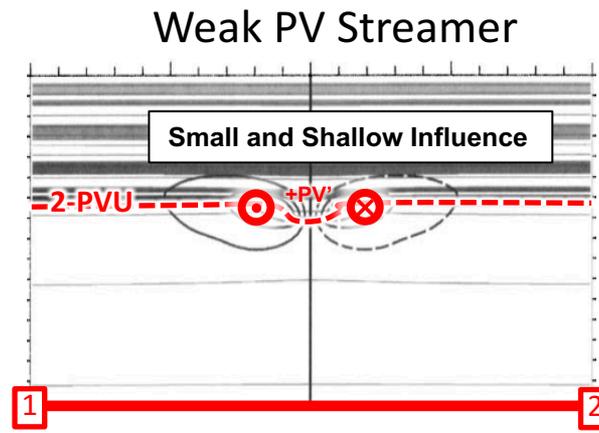
- PV streamer intensity impacts its ability to affect the troposphere below
 - ✧ Stronger intensity and larger size enable deeper and wider perturbation flow associated with the PV streamer
 - ✧ Impacts vertical wind shear, nearby tropospheric static stability, and moisture anomalies
 - ✧ Static stability of the troposphere affects size and intensity of perturbation flow



Adapted from Fig. 15a of HMR



Adapted from Fig. 2a in HMR

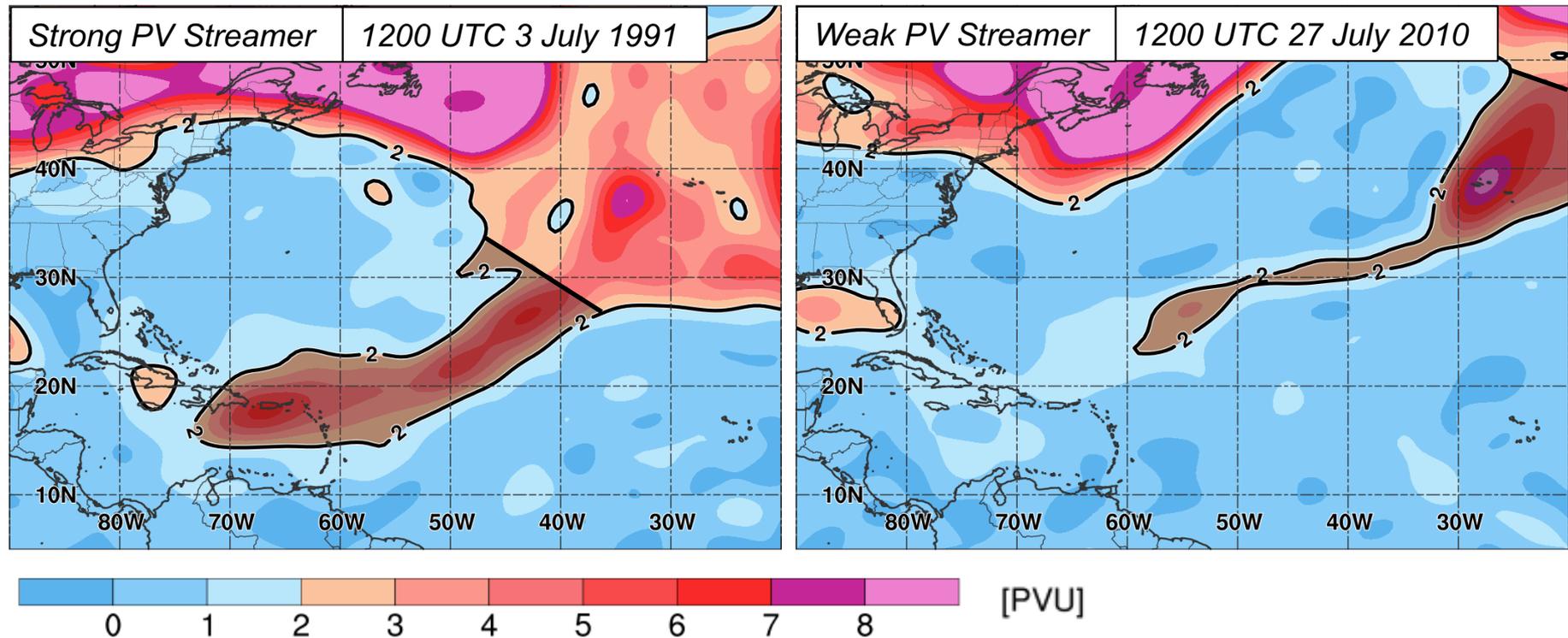


Adapted from Fig. 1 of Jukes and Smith (2000)

Literature Review: PV Streamer Intensity

- We hypothesize that in addition to RWB frequency, the size and intensity of the PV streamers they produce may significantly alter seasonal TC Activity
 - Different sizes and intensities may modify important variables for TC intensity
 - Vertical Wind Shear (VWS)
 - Moisture (PW)

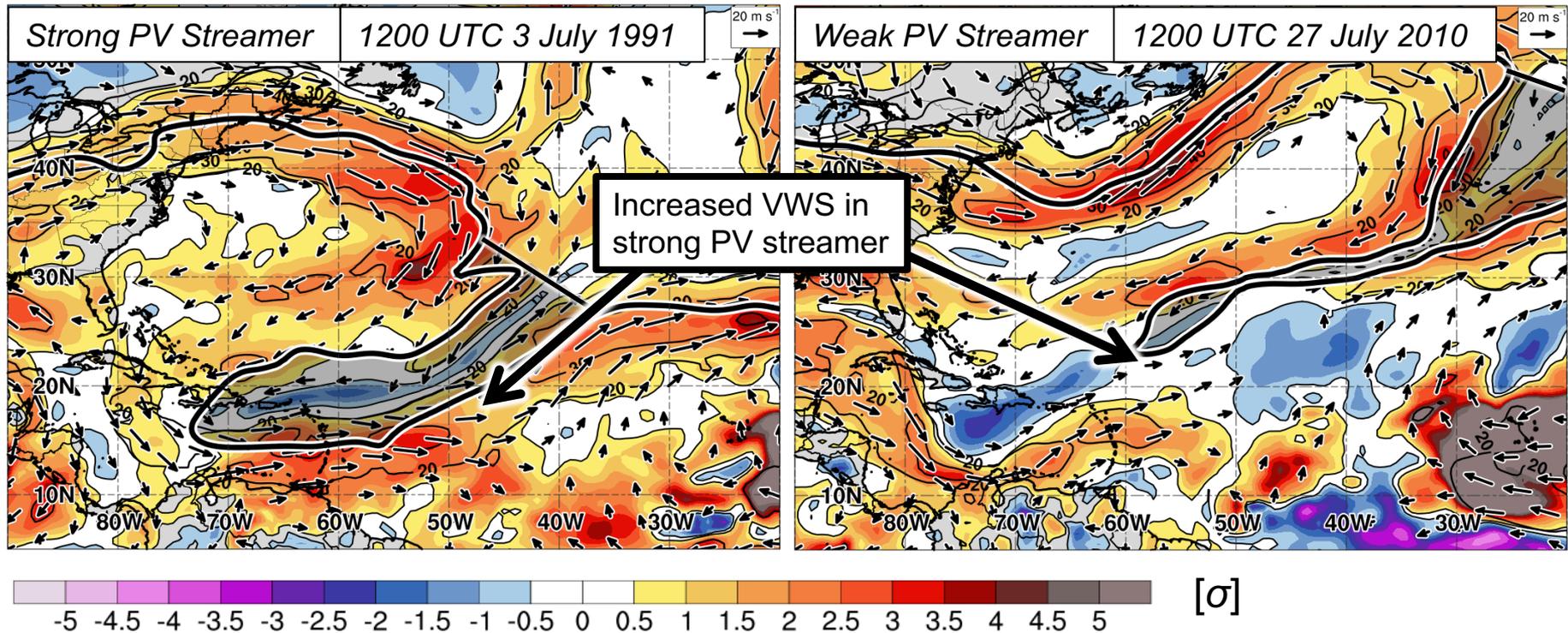
350-K PV (shaded, PVU), 2-PVU contour (black line)



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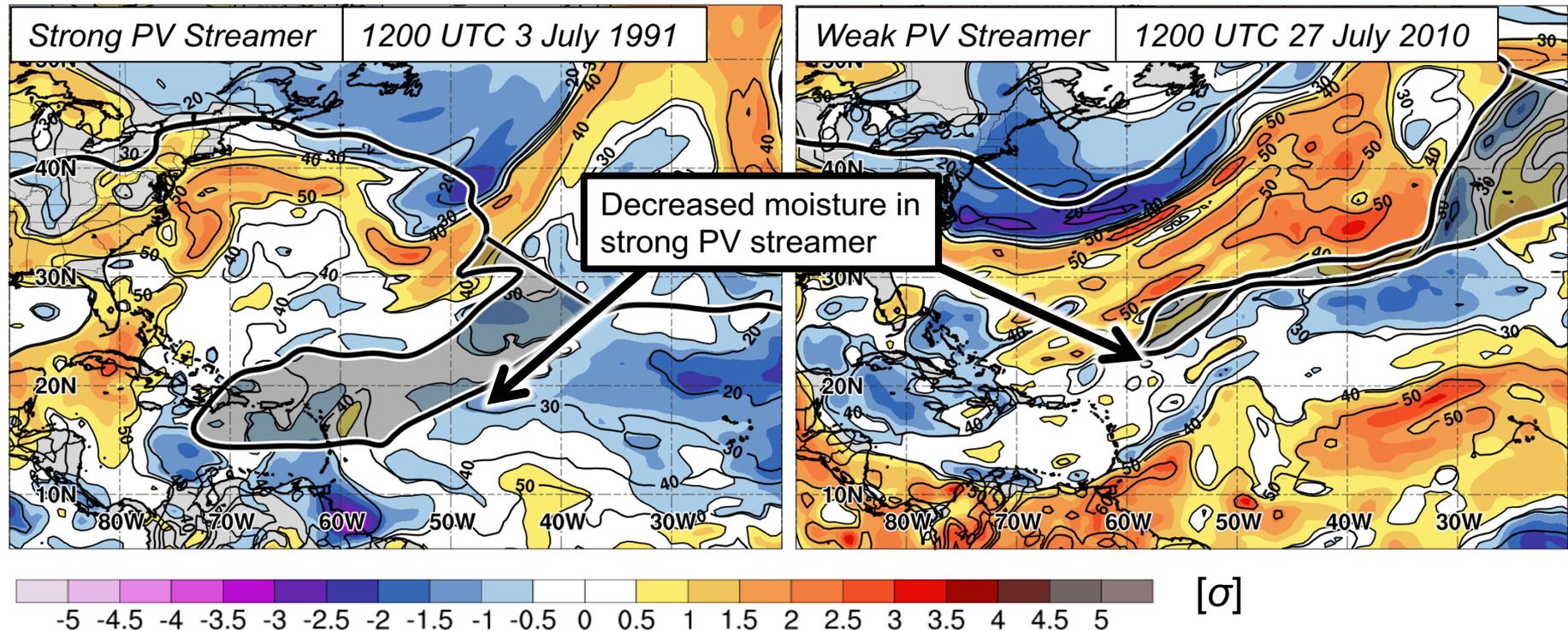
200–850-hPa Vertical Wind Shear Magnitude (black contours, m s^{-1}), Direction (vectors), Normalized Anomaly (shaded, σ)



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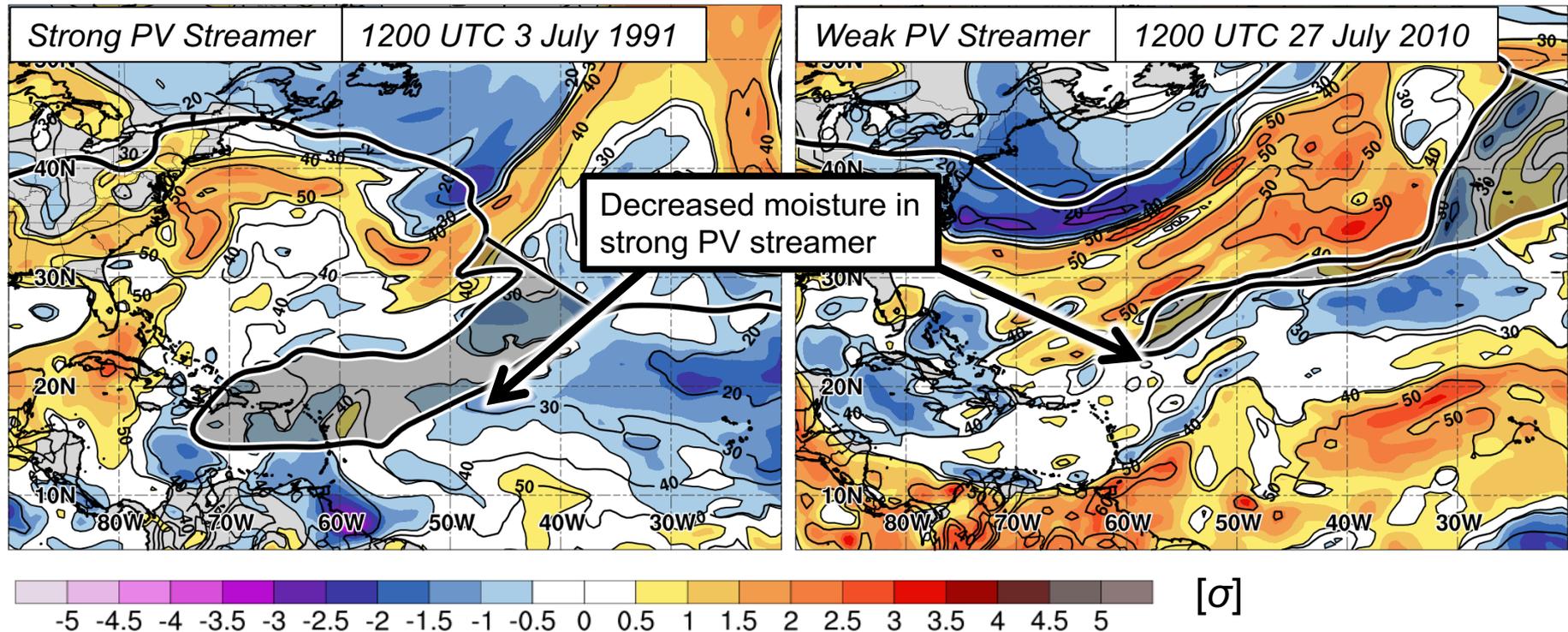
Precipitable Water (black contours, mm), and Normalized Anomaly (shaded, σ)



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We will investigate this hypothesis by composing a climatology that will quantify the size and intensity variations of PV streamers

Methodology: PV Streamer Identification

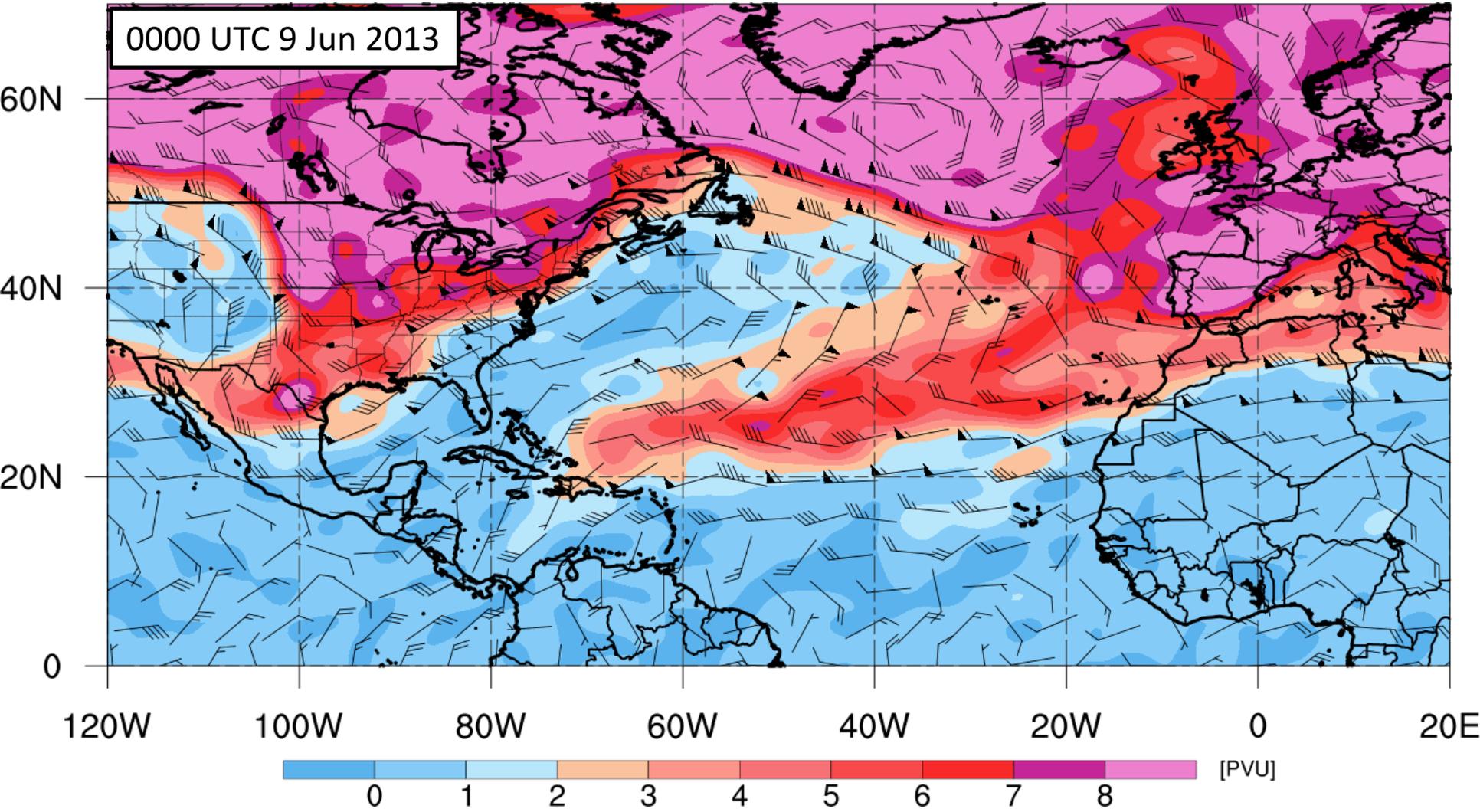
Methodology: PV Streamer Identification

- PV streamers are identified from June–November 1979–2015 using the 0.5° NCEP CFSR (Saha et al. 2010).
- A new PV streamer algorithm is created that combines previous methodologies
 - Postel and Hitchman (1999)
 - ✧ Identifies locations where RWB occurs (meridional gradient reversal in PV)
 - Wernli and Sprenger (2007)
 - ✧ Identifies elongated filaments of high PV air using width and perimeter of PV streamer
- Identification of PV streamers occurs on a isentropic surface that approximates the location of the subtropical tropopause
 - 350-K surface using the 2-PVU contour as the dynamical tropopause

Methodology: PV Streamer Identification

- Identify 2-PVU contour on 350-K surface

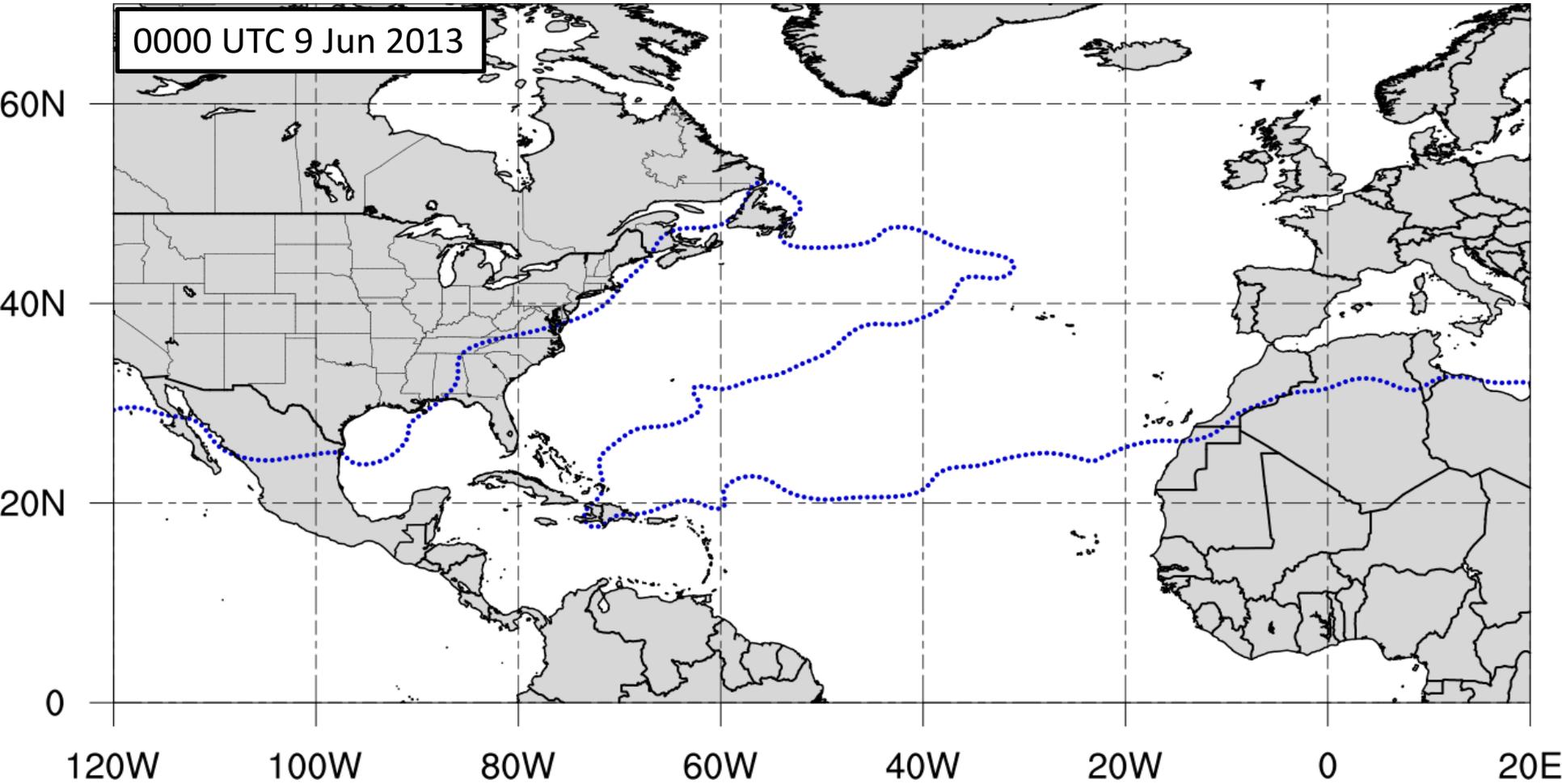
350-K PV (shaded, PVU), and winds (barbs, kt)



Methodology: PV Streamer Identification

- Identify 2-PVU contour on 350-K surface

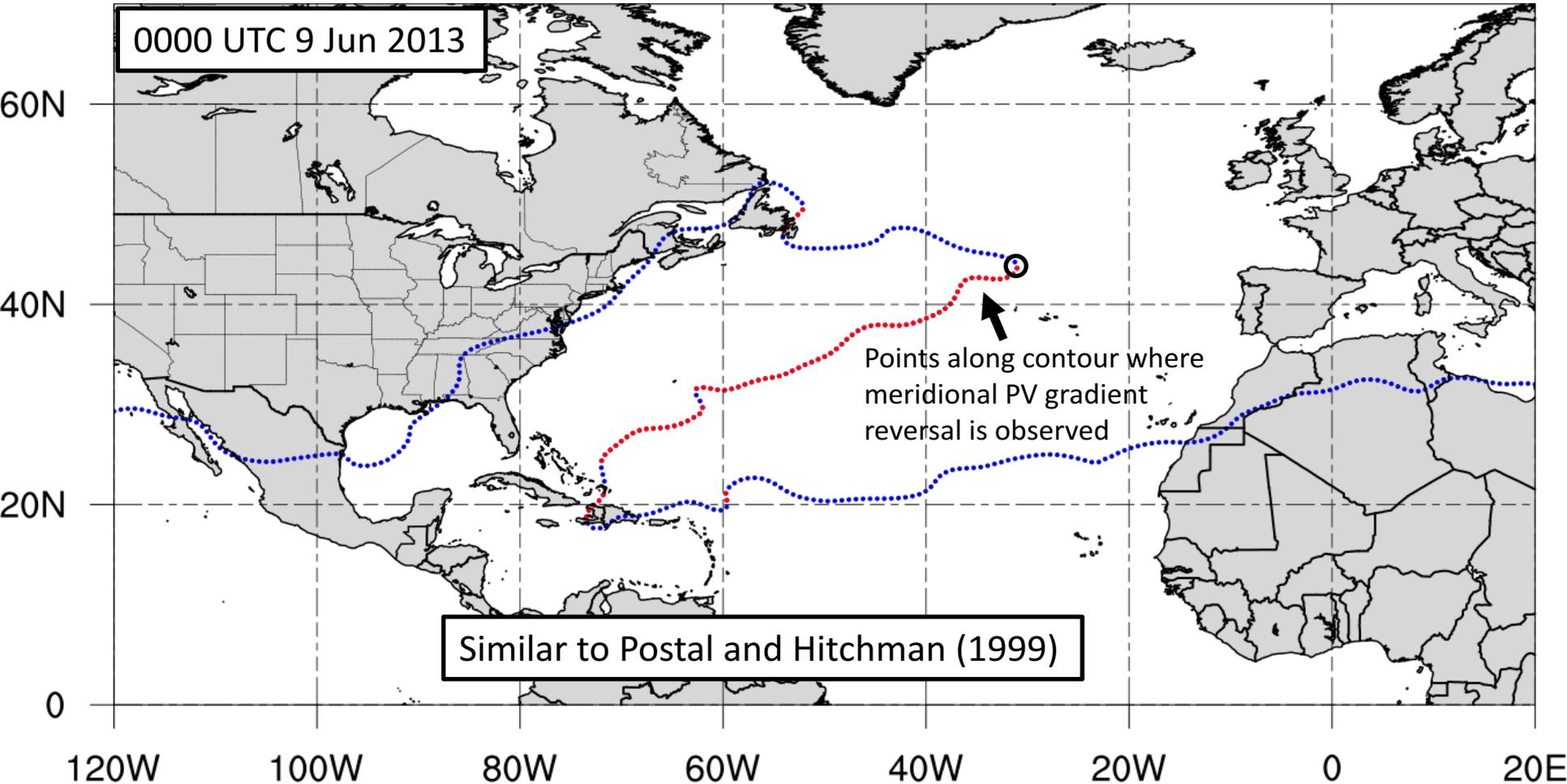
2-PVU contour on 350-K surface (blue contour)



Methodology: PV Streamer Identification

- Identify all points along contour where meridional PV gradient reversal is observed
 - First point along meridional reversal chosen as starting point of PV streamer

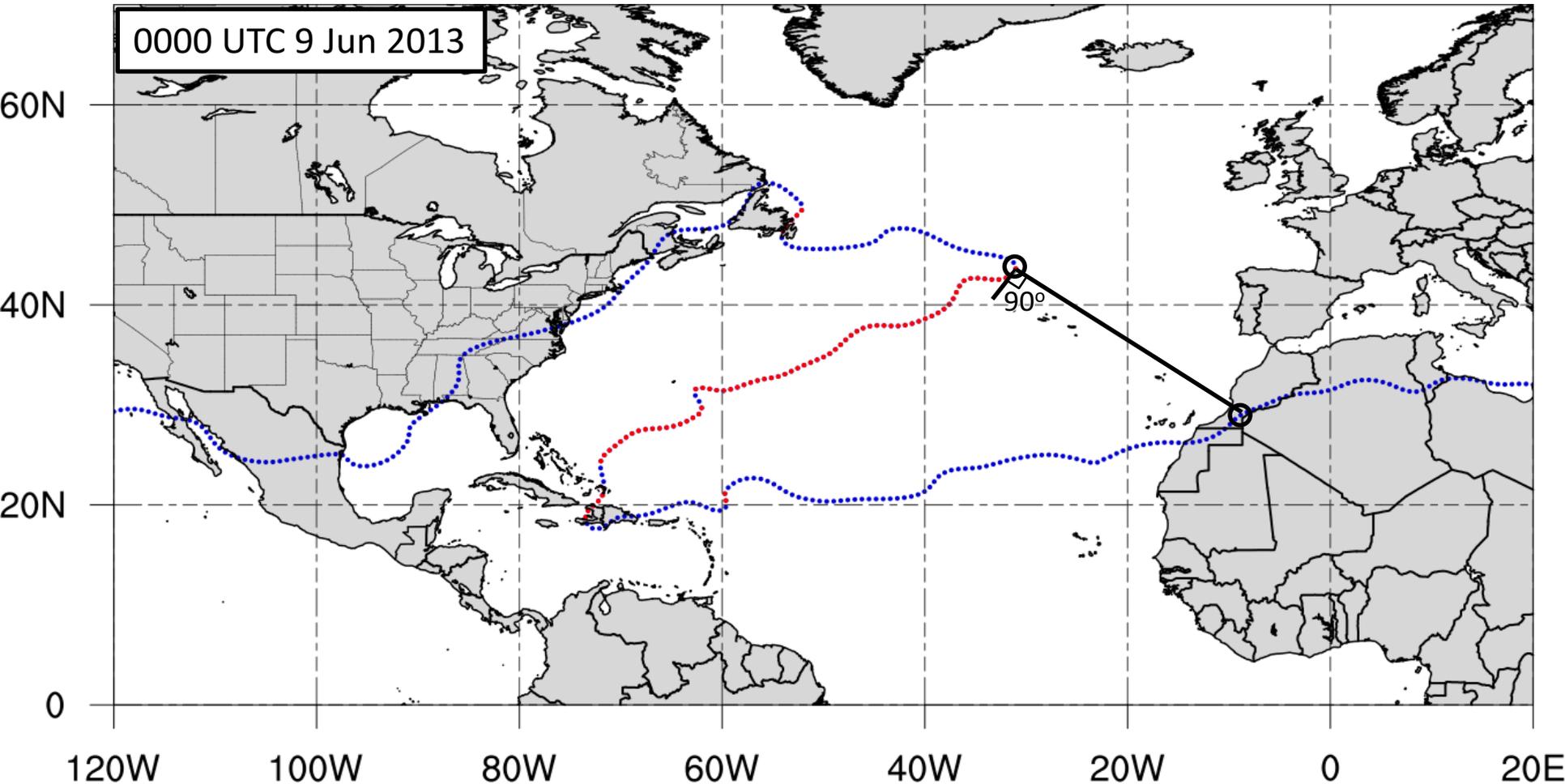
2-PVU contour on 350-K surface (blue contour), regions with meridional PV gradient reversal (red contour)



Methodology: PV Streamer Identification

- Line drawn orthogonal to line made by first few points of PV reversal
- Line ends when line crosses 2-PVU contour downstream

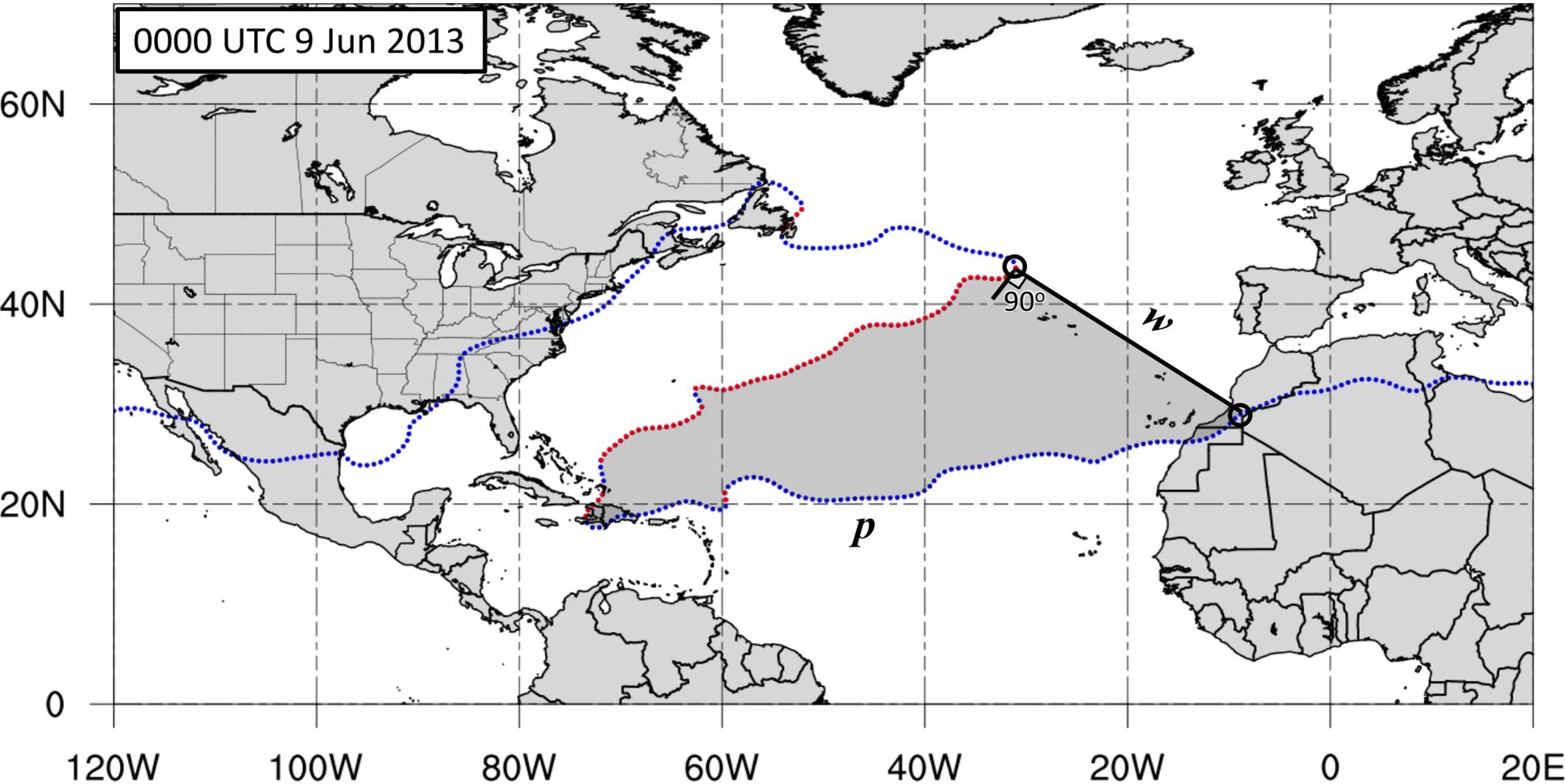
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PV streamer area (black shading), w (width between two points), p (along contour perimeter between two points)

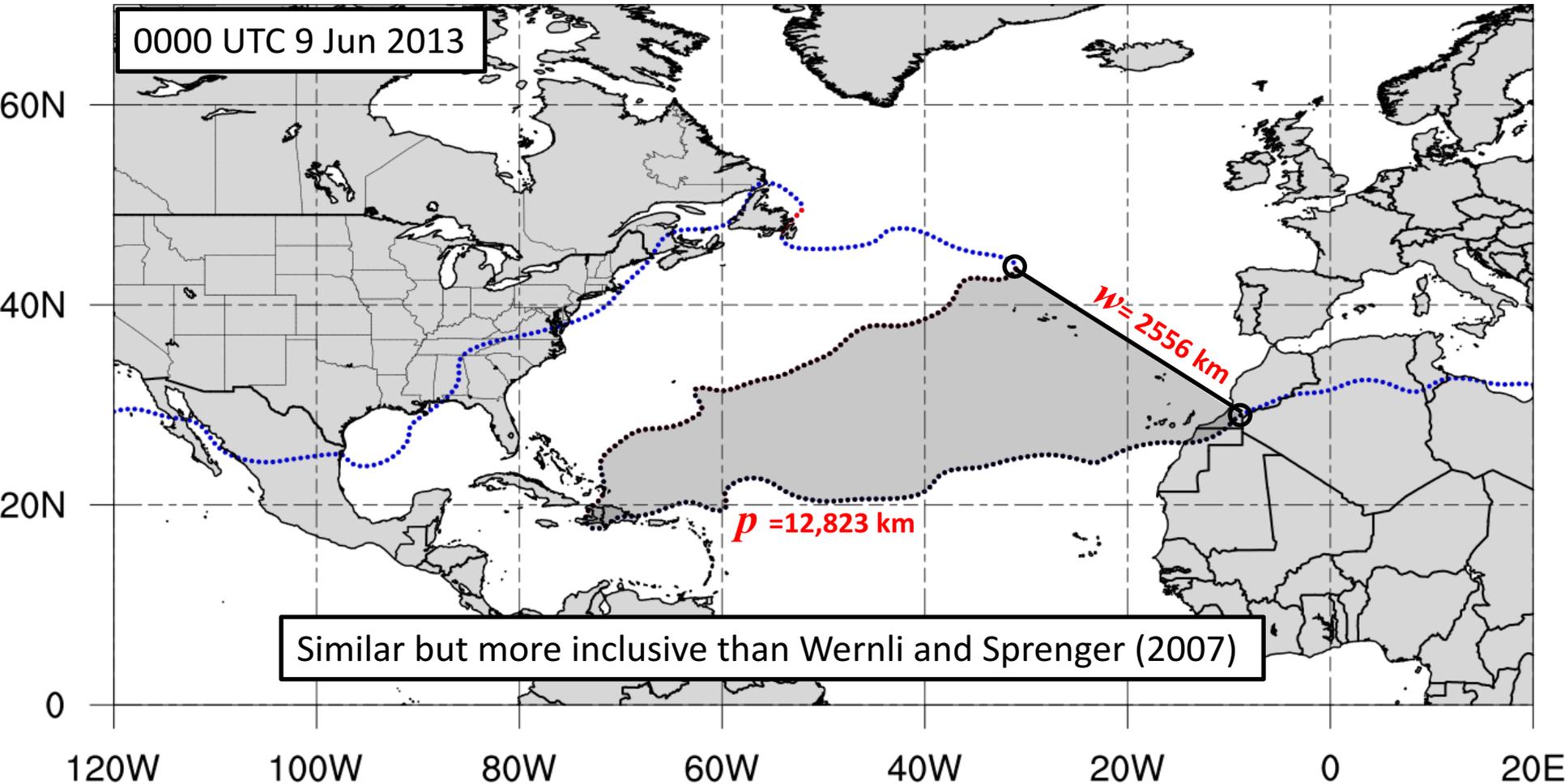


Methodology: PV Streamer Identification

- Check if PV streamer candidate is large and elongated enough

Threshold Values: p must be 3 times $>$ than w and $p > 3000$ km

PV streamer area (black shading), w (width between two points), p (along contour perimeter between two points)

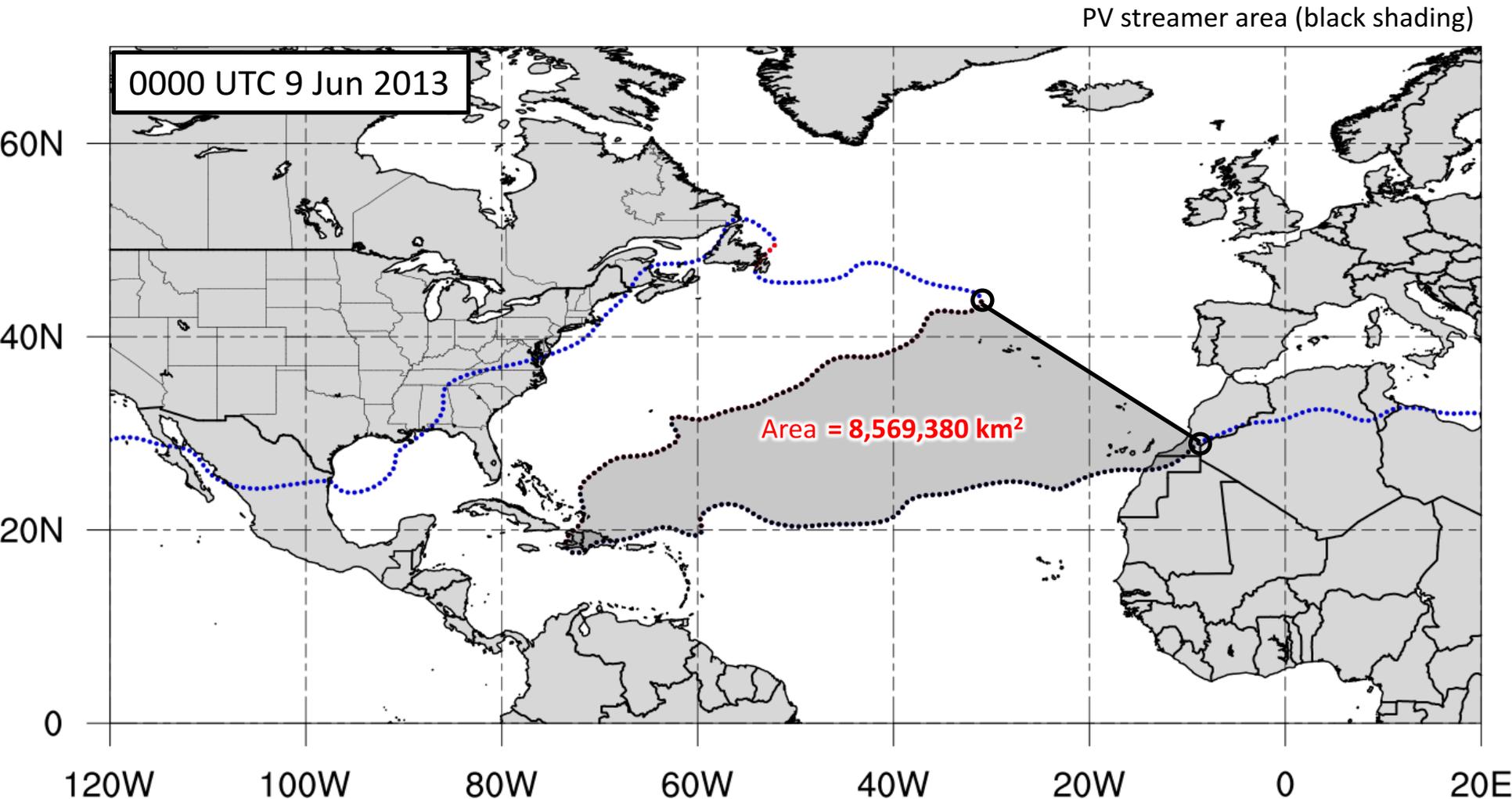


Methodology:

Unique PV Streamer Variables

Methodology: PV Streamer Variables

- PV streamer Area
 - An closed polygon allows us to calculate the area of the PV streamer

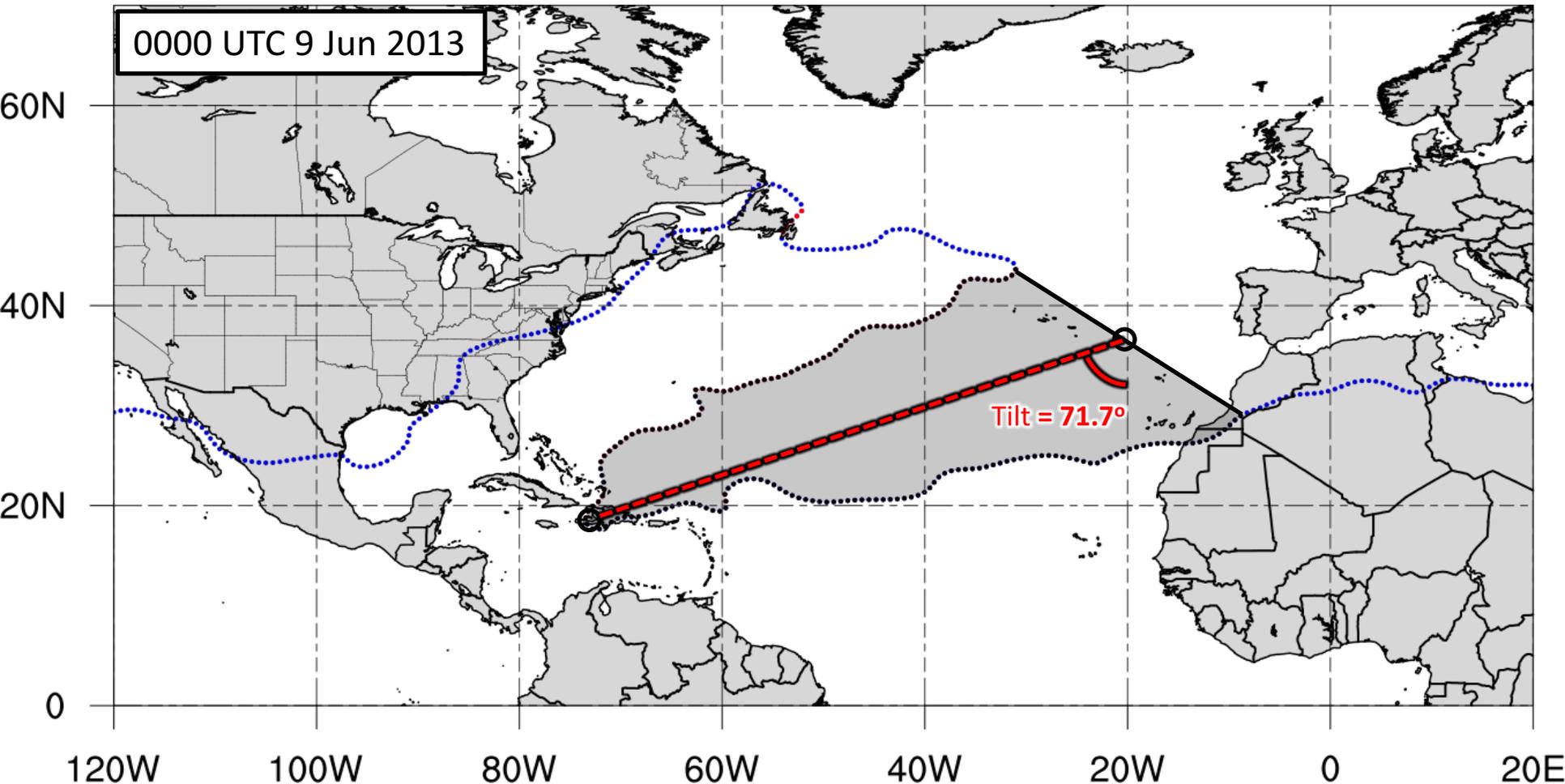


Methodology: PV Streamer Variables

- **PV streamer Tilt**

- First obtain midpoint between the start & end of PV streamer
- Find furthest location from midpoint along PV streamer perimeter
- Determine angle of line relative to N–S meridian

PV streamer area (black shading)



Methodology: PV Streamer Variables

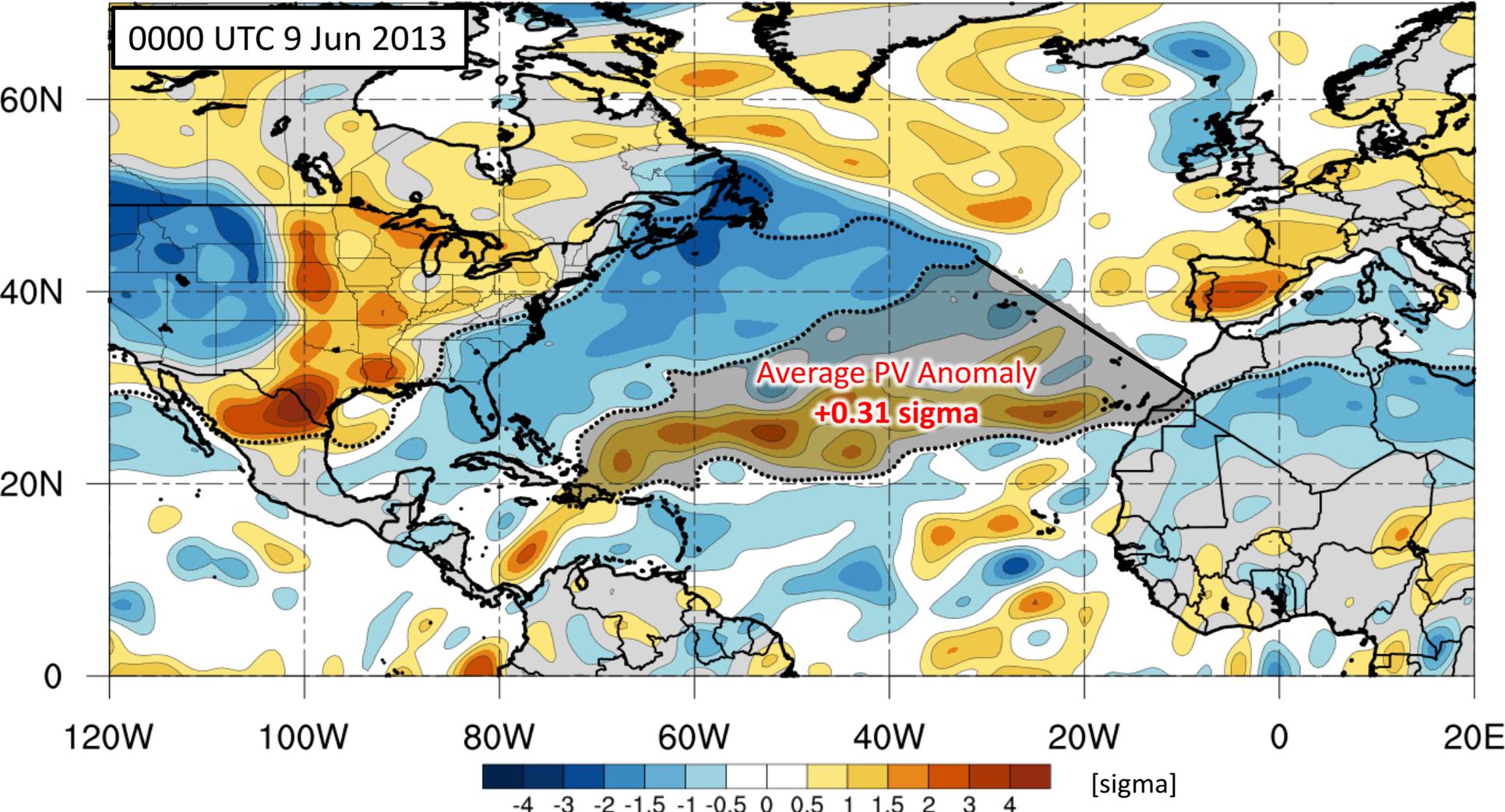
- PV streamer Intensity

- Calculated as a standardized anomaly

$$PV_{std_anom} = (PV - PV_{mean}) / PV_{sd}$$

Mean and Standard Deviations are derived from a 1979-2009 CFSR climatology

350-K Standardized PV Anomaly (shaded, Sigma), and 2-PVU contour (black contour)



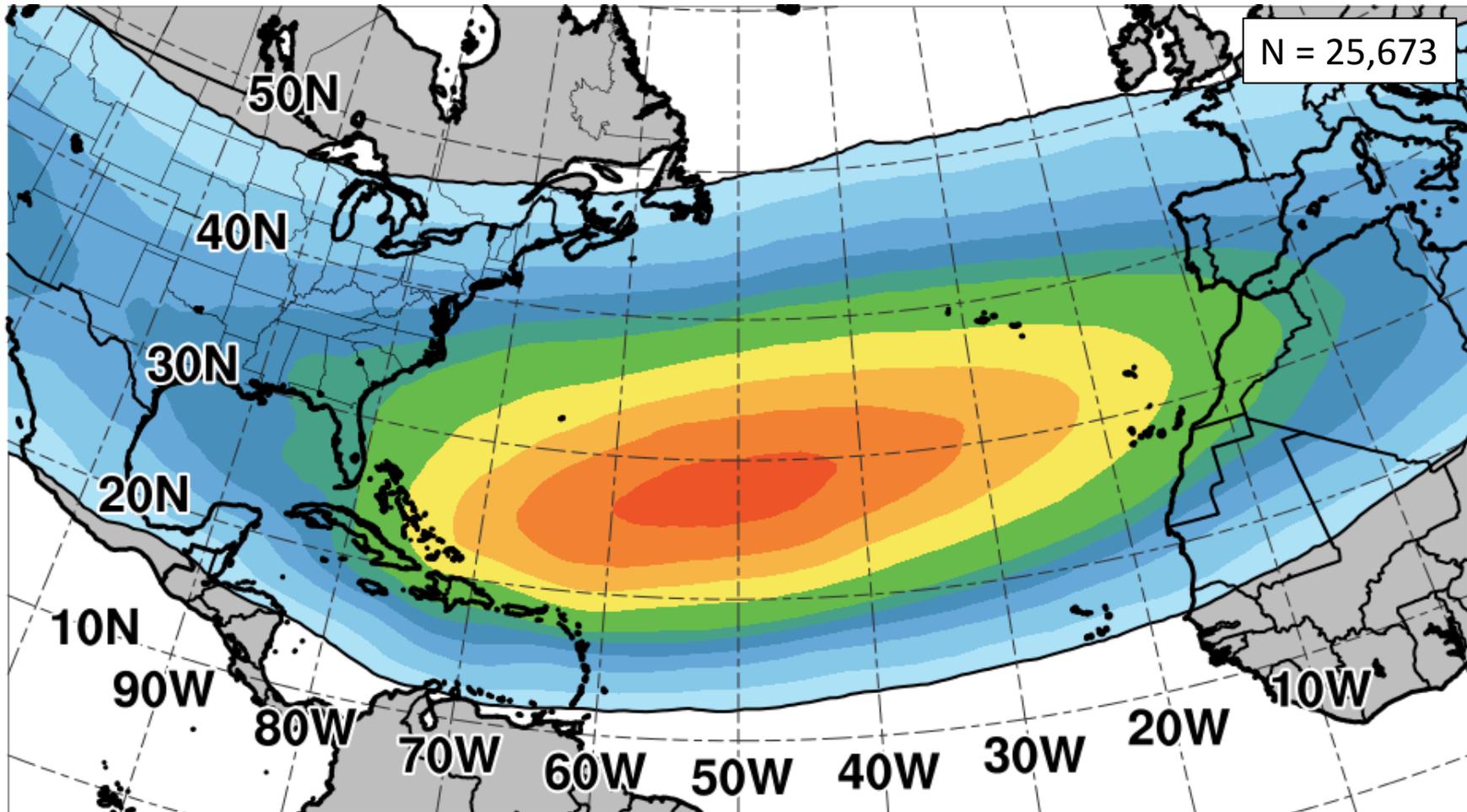
Results: PV Streamer Climatology

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- PV streamer frequency in the North Atlantic

- 1 Jun–30 Nov
 - ◇ Occurrence maximized on equatorward side of jet
 - ◇ Corresponds to climatological position of Mid-Ocean trough (i.e., TUTT)

Probability PV streamer is observed on any particular day (shading, %)

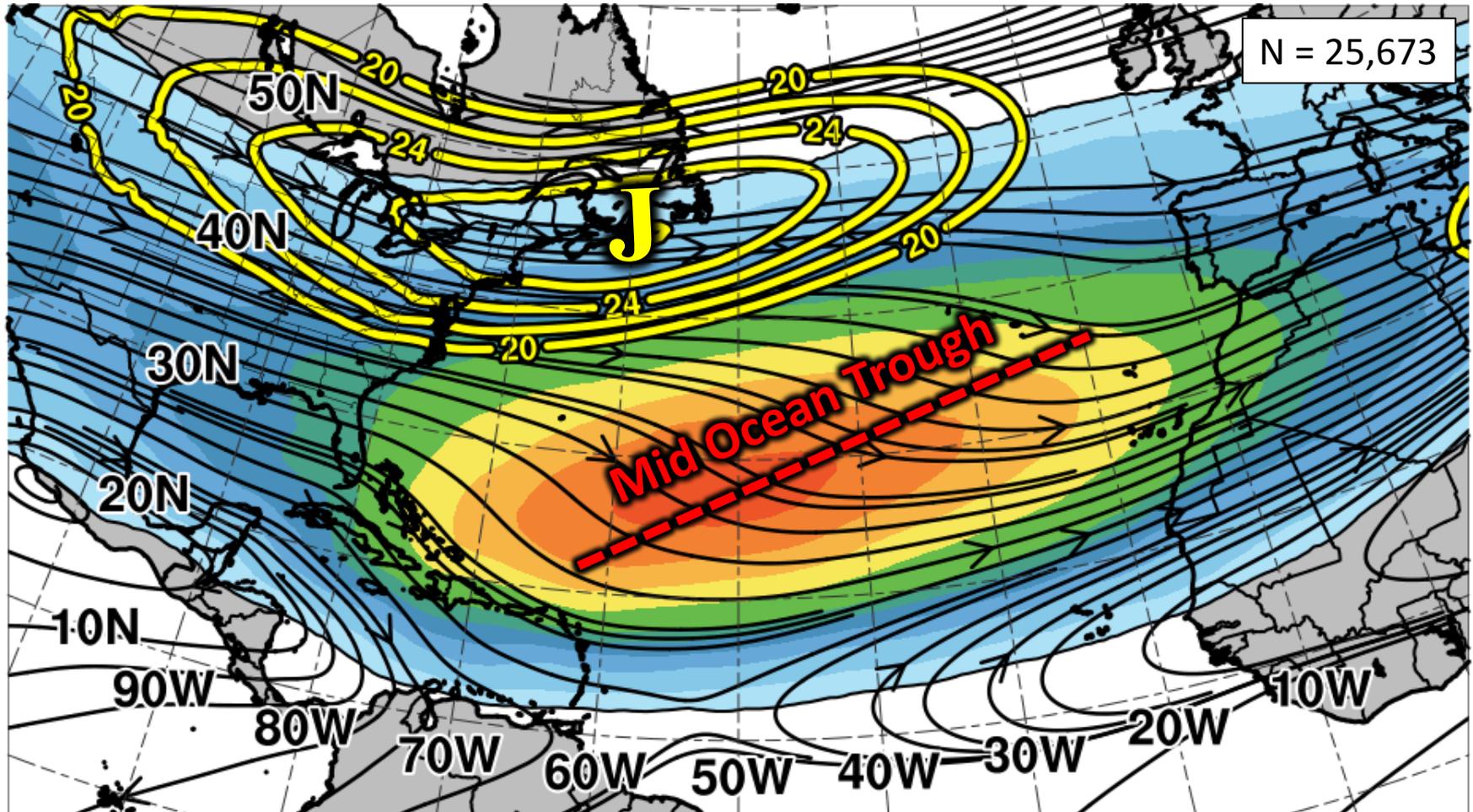


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Probability PV streamer is observed on any particular day (shading, %), 200-hPa winds (yellow contours, m s^{-1}) and streamlines (black lines)

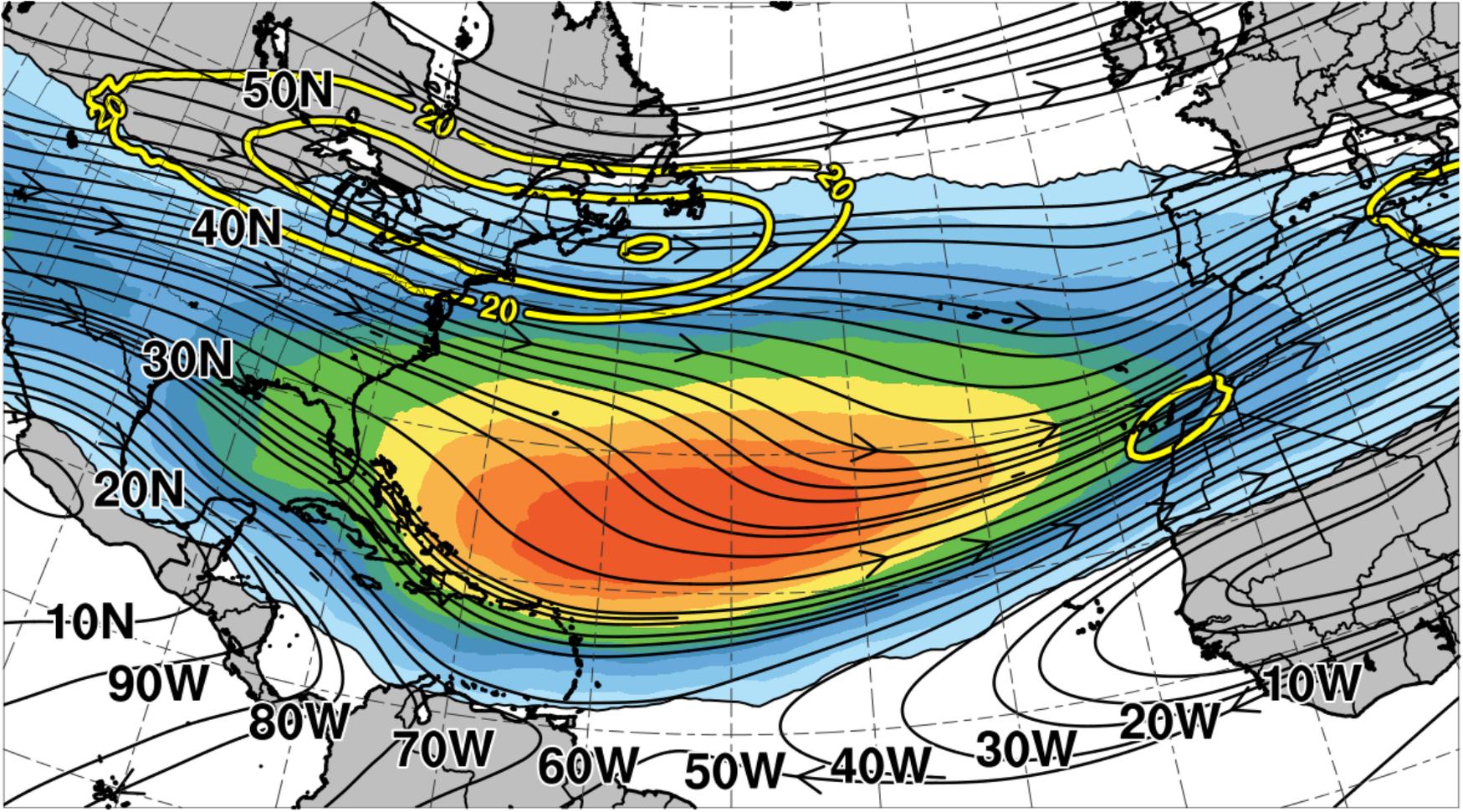


Results: Month by Month

Results: PV Streamer Climatology – June

- PV Streamer occurrence shifts over TC season
 - Westerlies dominate Atlantic basin at beginning of TC season

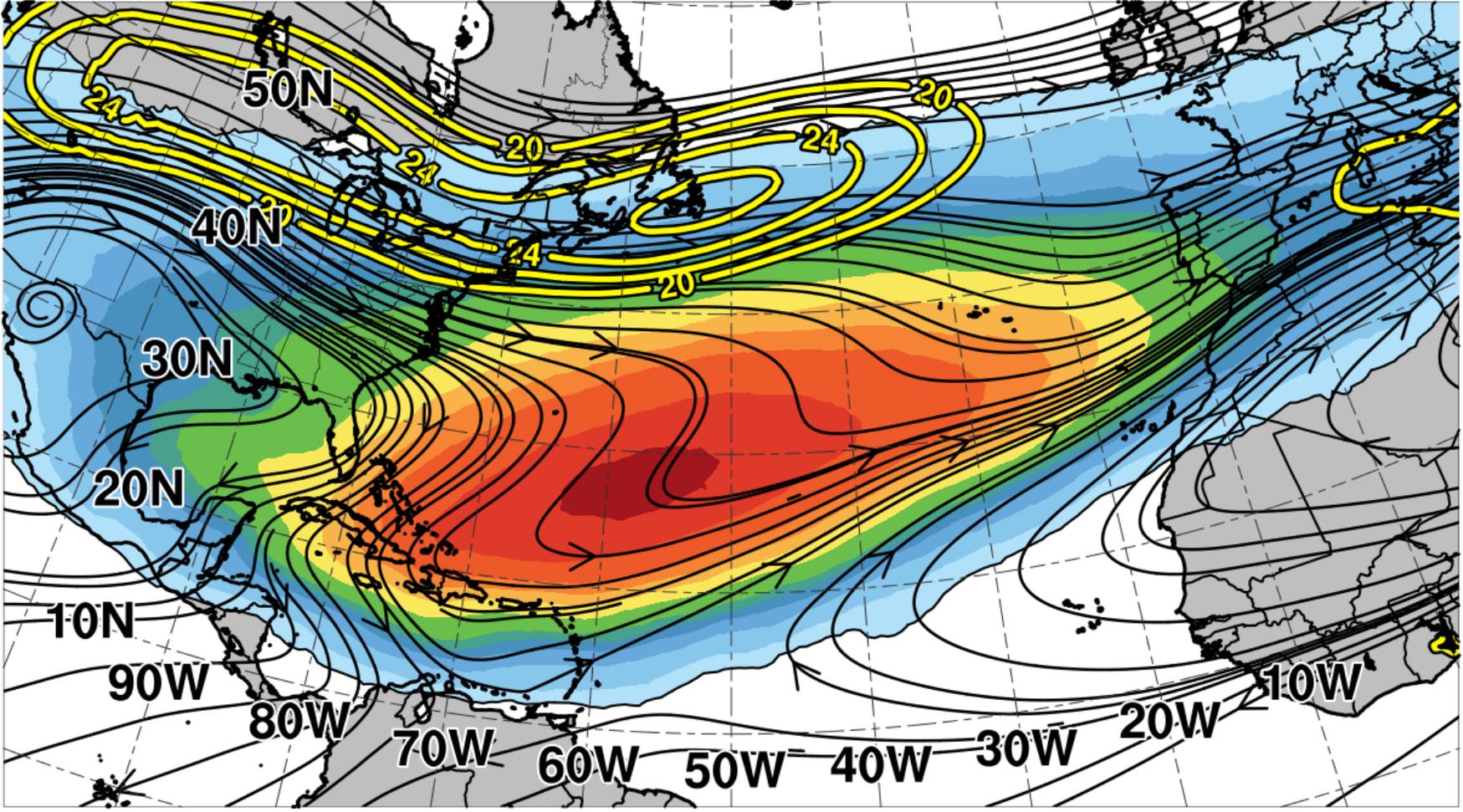
Probability PV streamer is observed on any particular day (shading, %), 200-hPa winds (yellow contours, $m s^{-1}$) and streamlines (black lines)



Results: PV Streamer Climatology – July

- PV Streamer occurrence shifts over course of season
 - Max frequency at 60°W with increased frequency over Caribbean

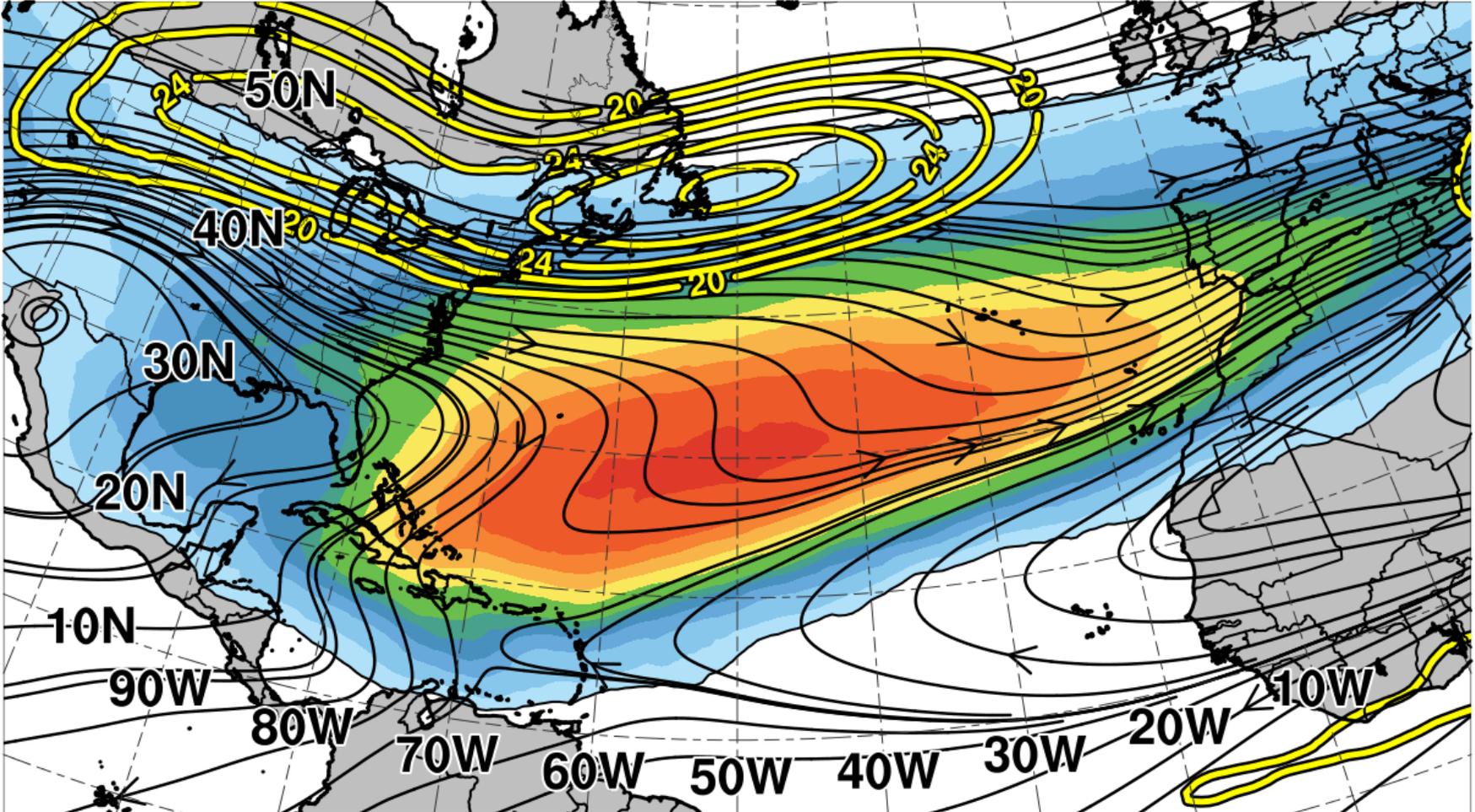
Probability PV streamer is observed on any particular day (shading, %), 200-hPa winds (yellow contours, $m s^{-1}$) and streamlines (black lines)



Results: PV Streamer Climatology – August

- PV Streamer occurrence shifts over course of season
 - Subtle shift north and west away from Caribbean

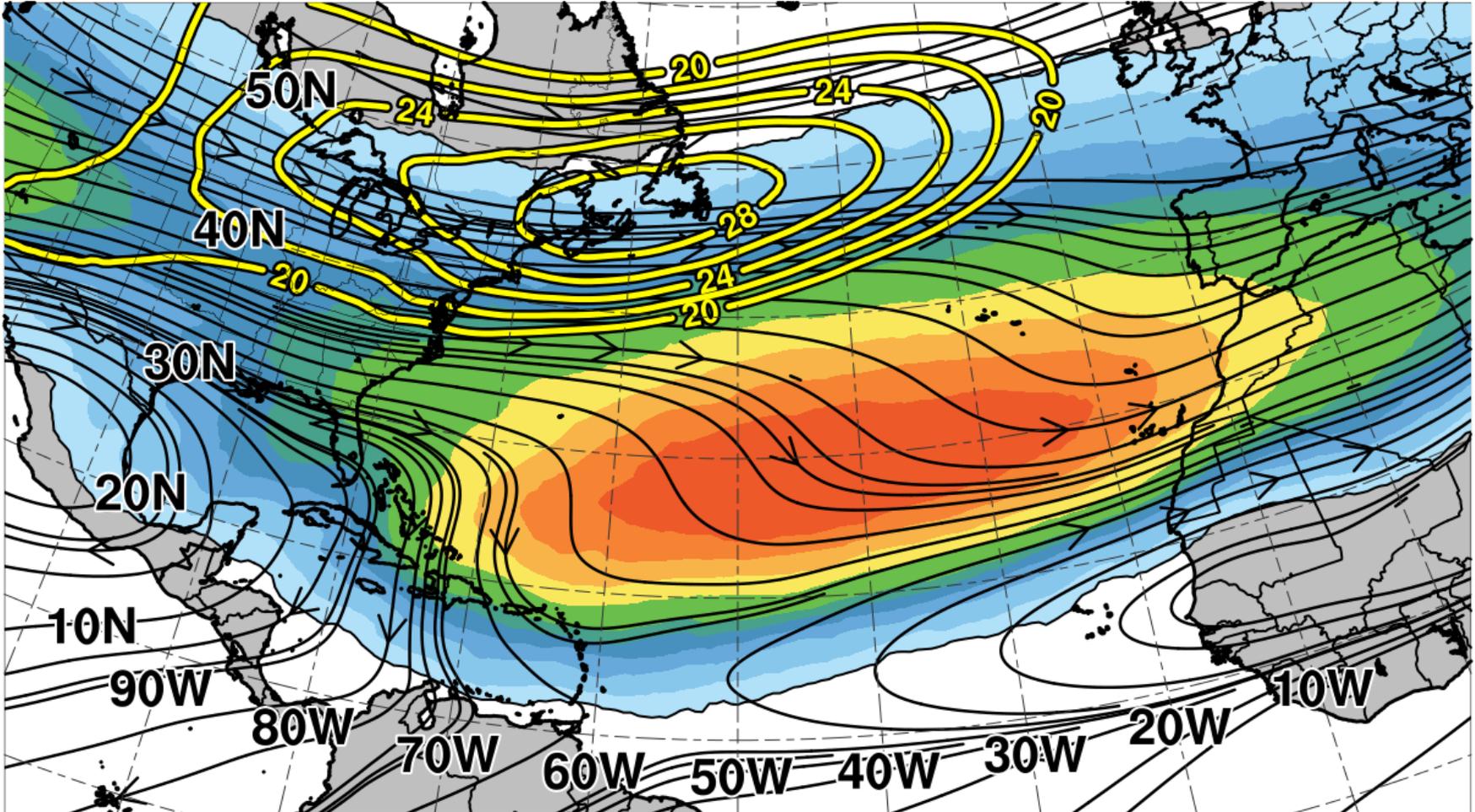
Probability PV streamer is observed on any particular day (shading, %), 200-hPa winds (yellow contours, $m s^{-1}$) and streamlines (black lines)



Results: PV Streamer Climatology – September

- PV Streamer occurrence shifts over course of season
 - More distinct shift in maxima towards eastern subtropical Atlantic

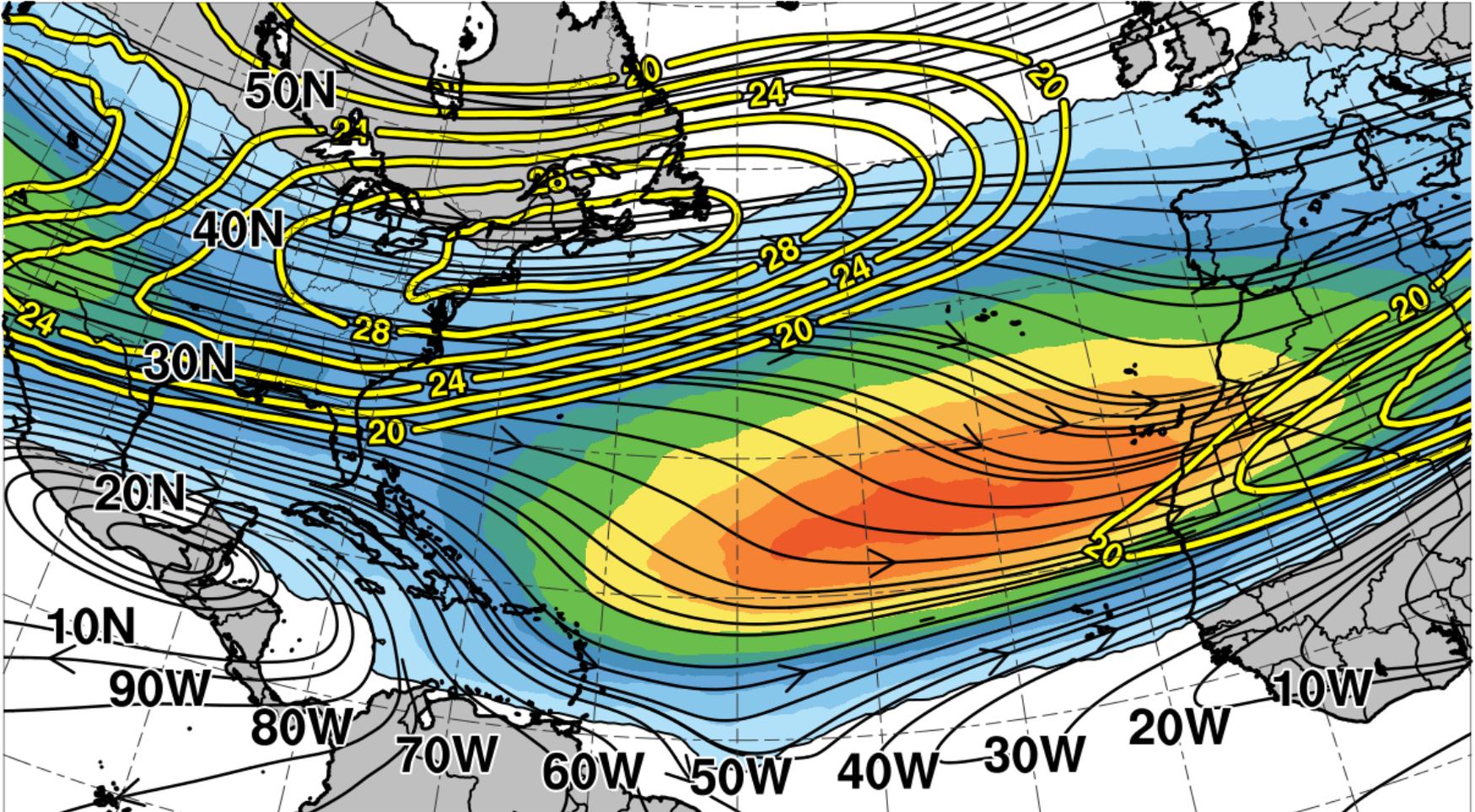
Probability PV streamer is observed on any particular day (shading, %), 200-hPa winds (yellow contours, $m s^{-1}$) and streamlines (black lines)



Results: PV Streamer Climatology – October

- PV Streamer occurrence shifts over course of season
 - Continued shift increases 200-hPa westerlies in eastern Atlantic

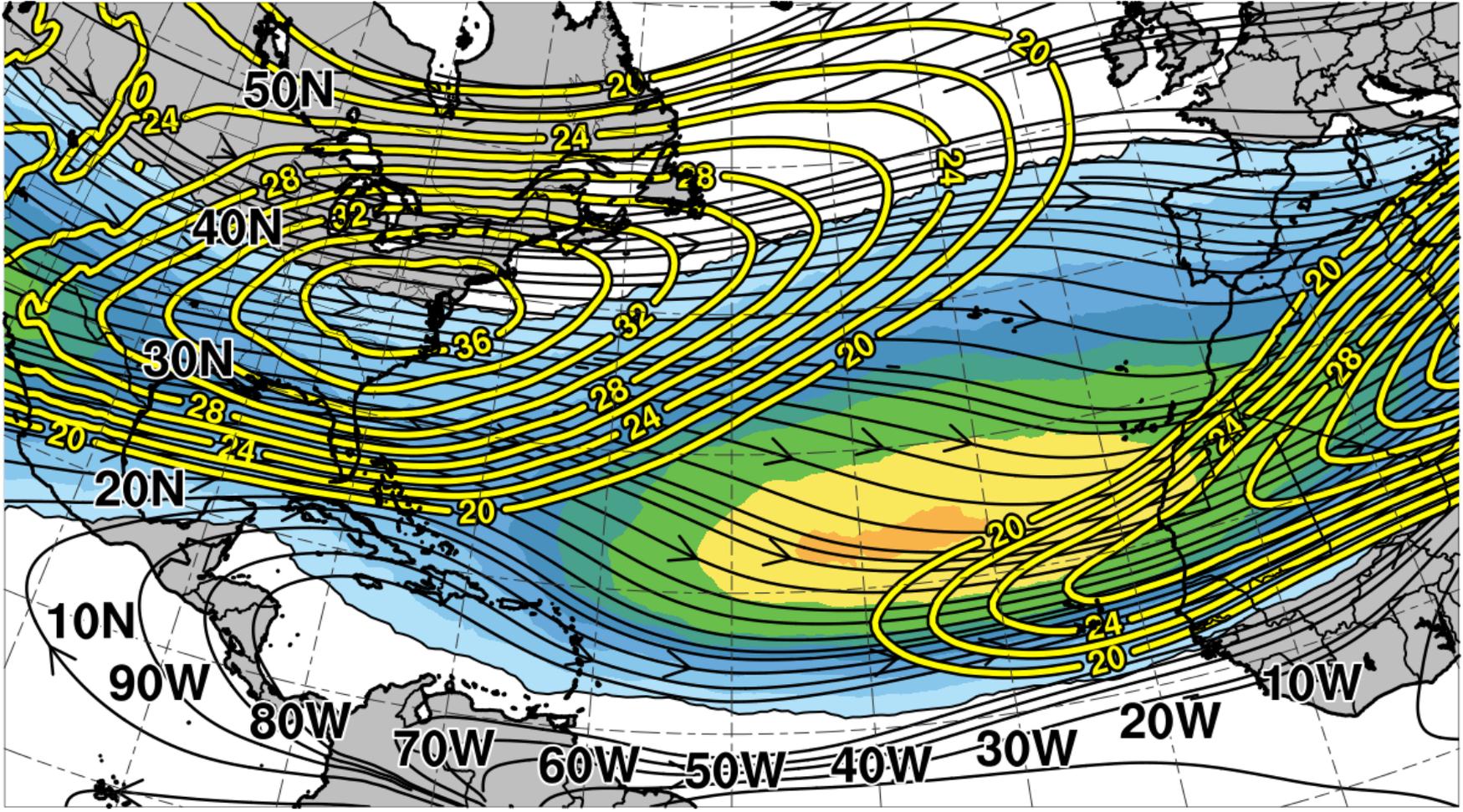
Probability PV streamer is observed on any particular day (shading, %), 200-hPa winds (yellow contours, $m\ s^{-1}$) and streamlines (black lines)



Results: PV Streamer Climatology – November

- PV Streamer occurrence shifts over course of season
 - Westerlies dominate Atlantic basin at end of TC season again

Probability PV streamer is observed on any particular day (shading, %), 200-hPa winds (yellow contours, $m s^{-1}$) and streamlines (black lines)



Results: Year to Year Variability

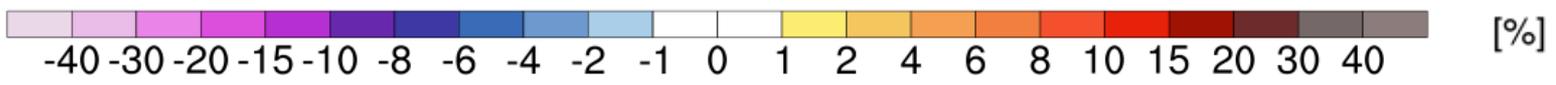
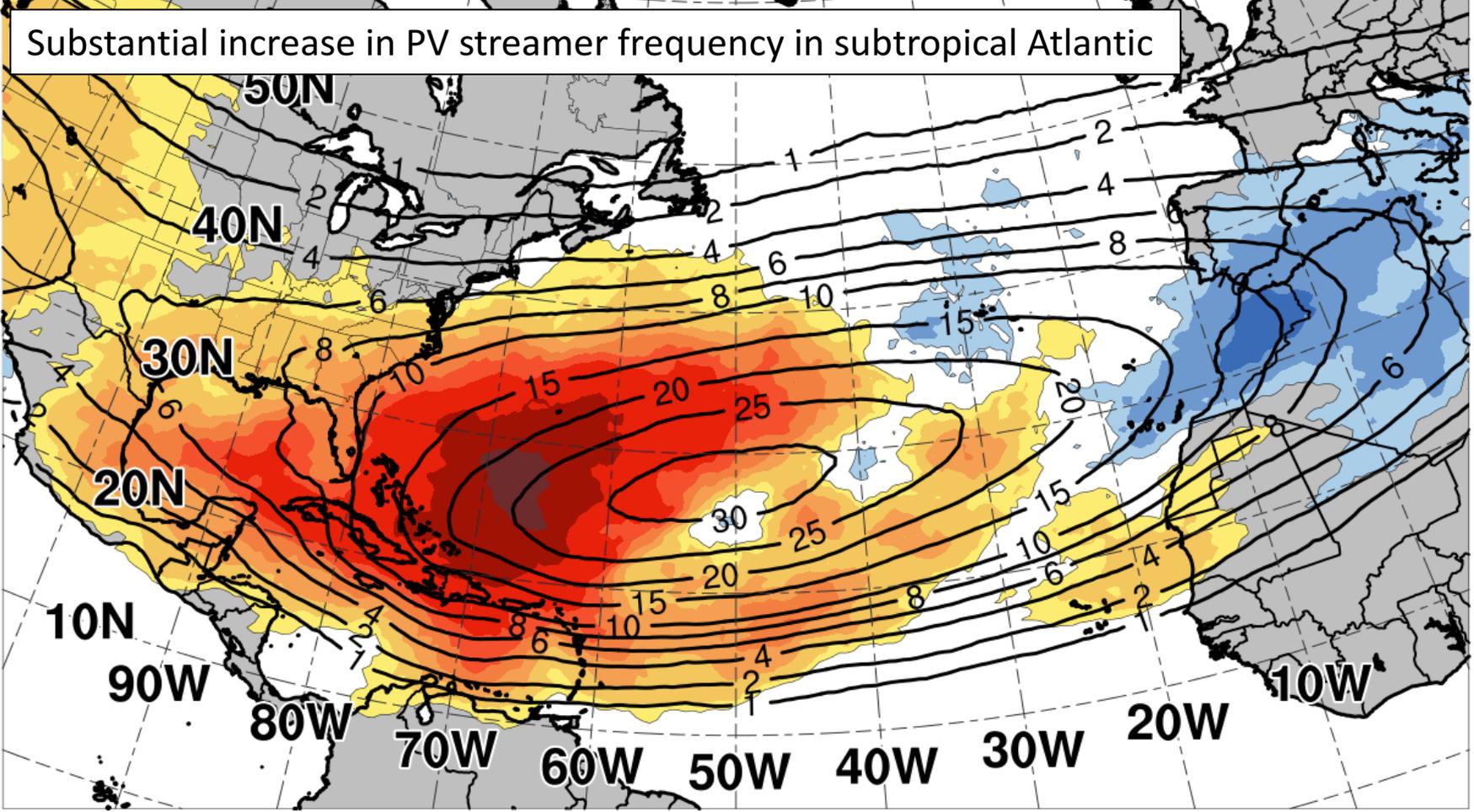
Results: PV Streamer Variability – 1994

- PV streamer activity can vary greatly from season to season

Accumulated Cyclone Energy (ACE): $32.0 \times 10^4 \text{ kt}^2$

Inactive Hurricane Season

PV streamer frequency anomaly for given season (shaded, %), Climatological frequency (black contours, %)



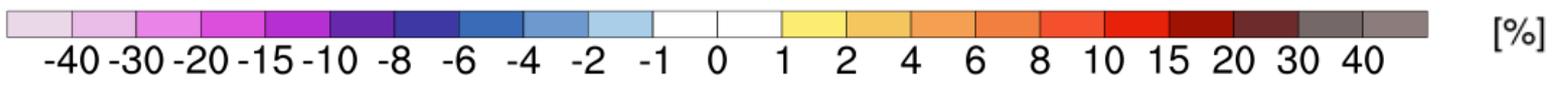
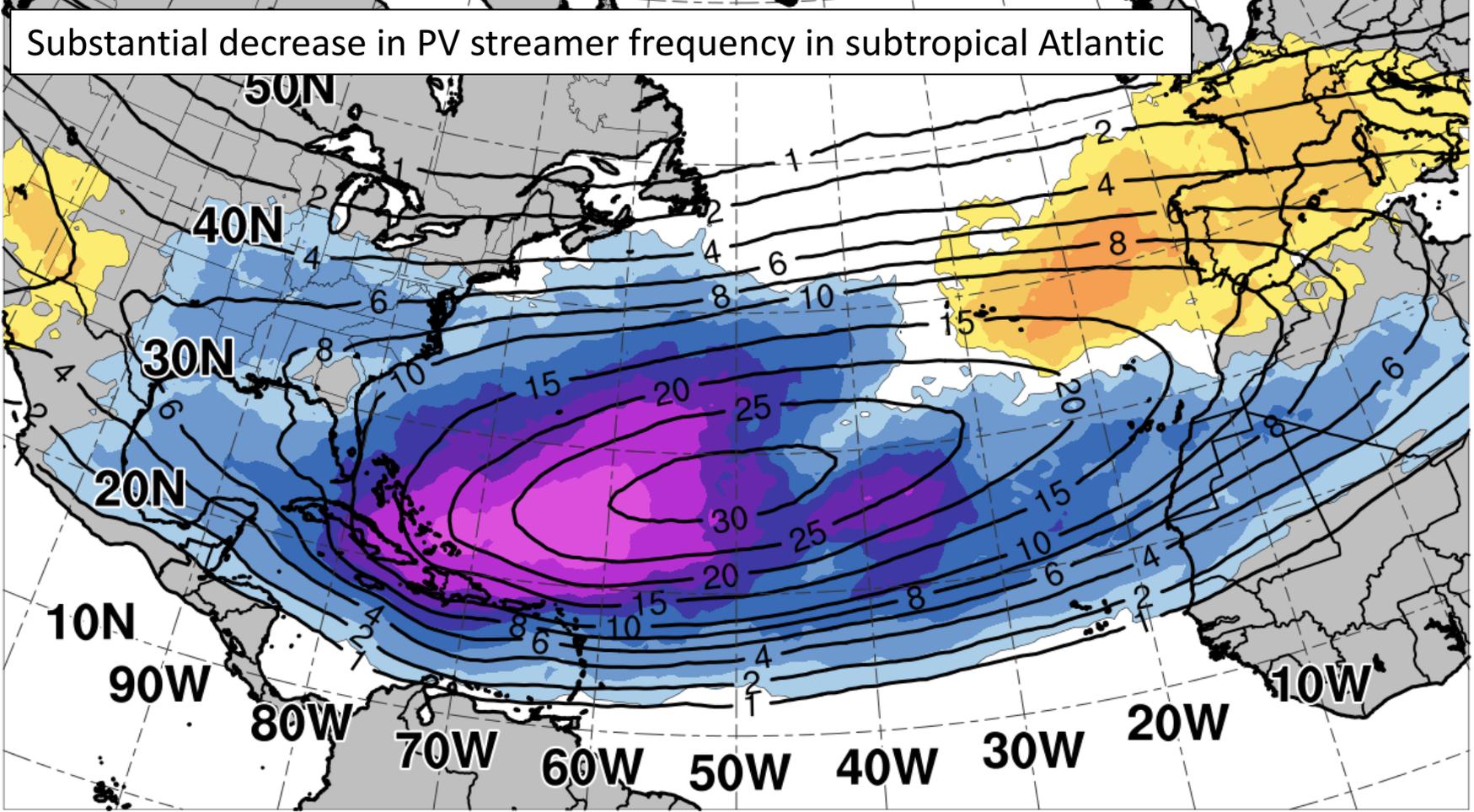
Results: PV Streamer Variability – 1995

- PV streamer activity can vary greatly from season to season

Accumulated Cyclone Energy (ACE): $227.1 \times 10^4 \text{ kt}^2$

Hyperactive Hurricane Season

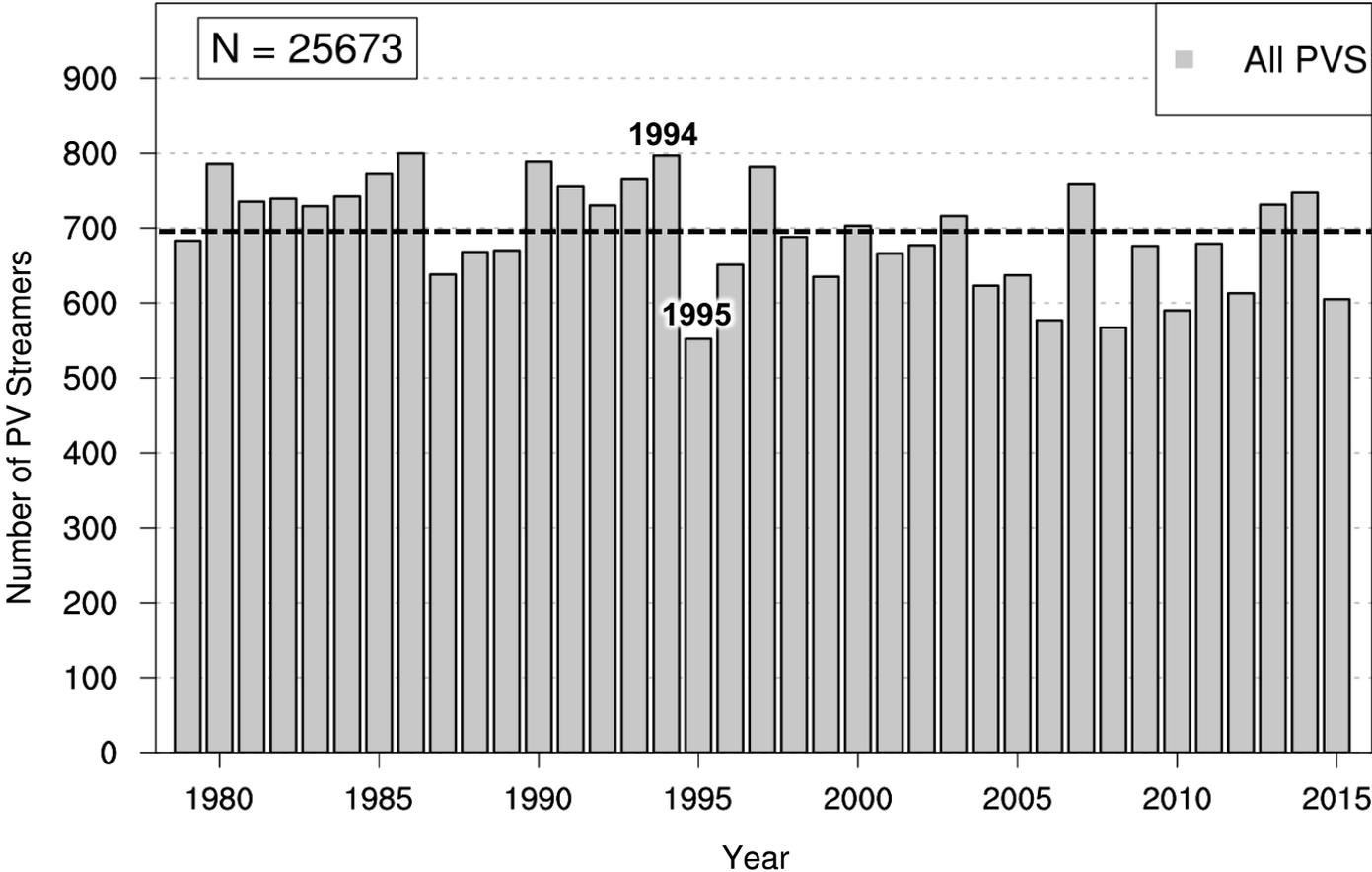
PV streamer frequency anomaly for given season (shaded, %), Climatological frequency (black contours, %)



Results: PV Streamer Variability 1979–2015

- Interseasonal variability in the number of PV streamers
 - 1 Jun–30 Nov between 100–10°W
 - Mean: 694
 - Stdev: +/- 70

Other variables important other than just number of of PV streamers

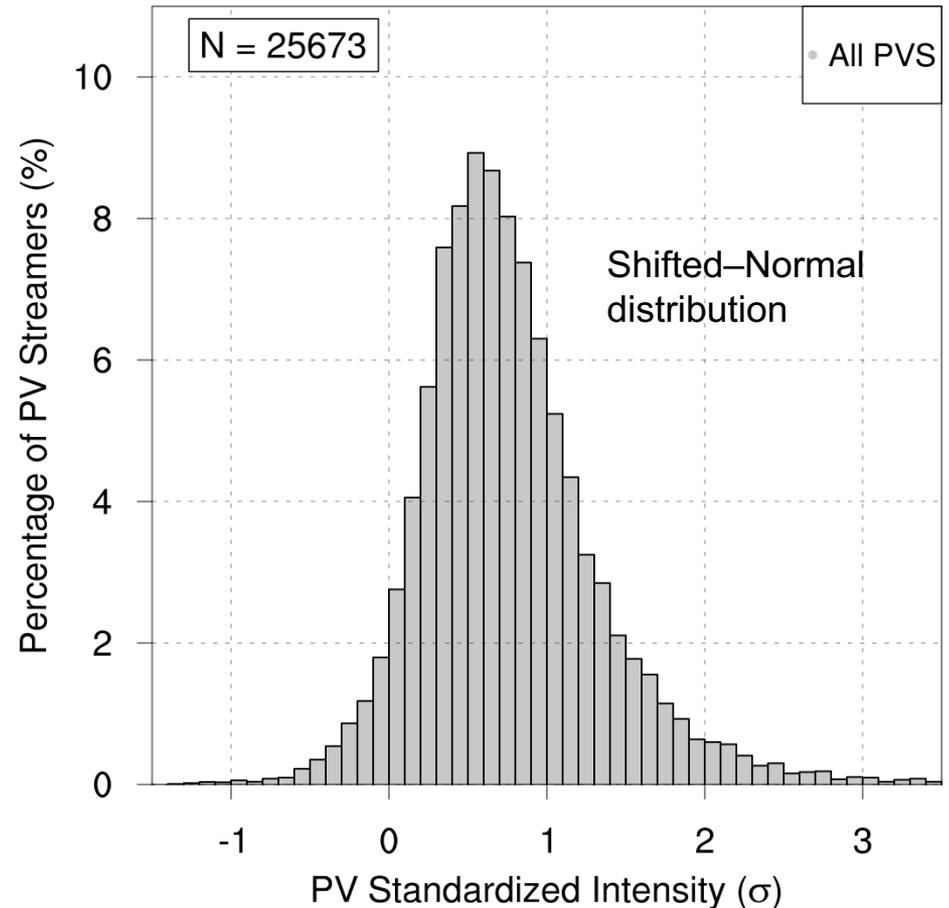


Yearly Distribution of PV Streamer Count

Results: Variability of Intensity and Area

Results: Variability of PV Streamer Intensity

- Intensity of PV streamers (standardized anomaly)
 - 1 Jun–30 Nov between 100–10°W
 - Mean: 0.76 sigma
 - Stdev: +/- 0.57 sigma

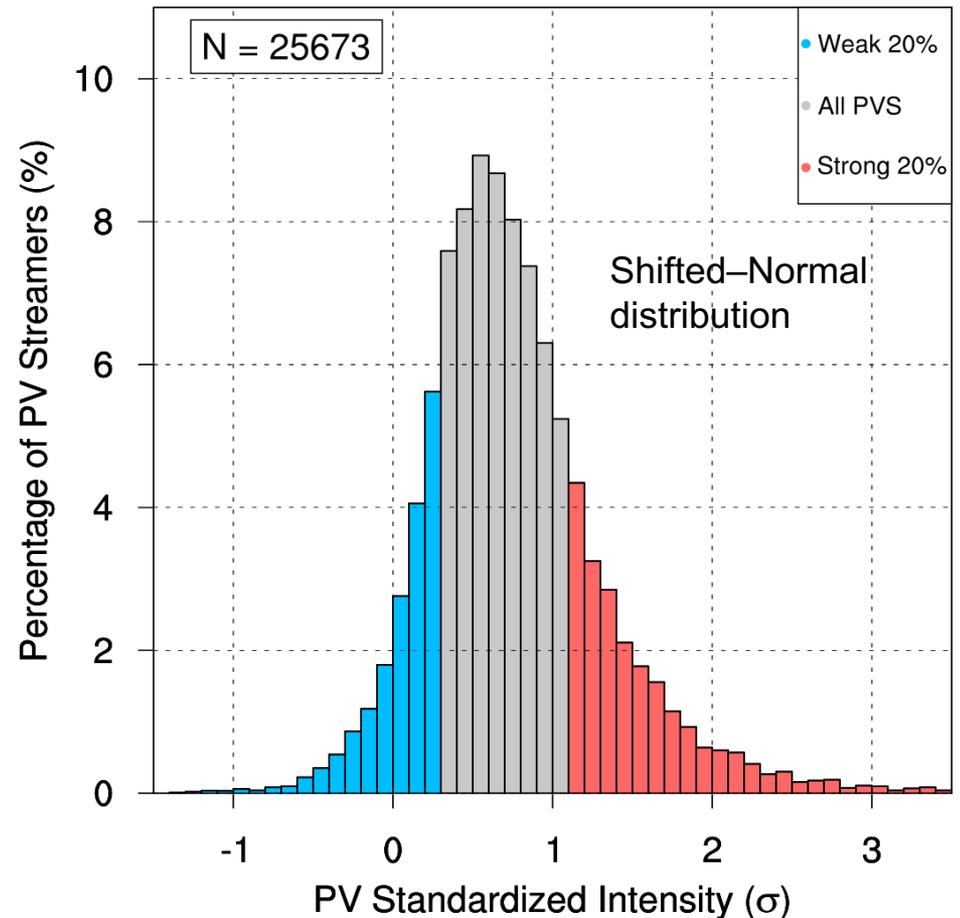


PDF of PV Streamer Intensity

Results: Variability of PV Streamer Intensity

- Intensity of PV streamers (standardized anomaly)
 - 1 Jun–30 Nov between 100–10°W
 - Mean: 0.76 sigma
 - Stdev: +/- 0.57 sigma

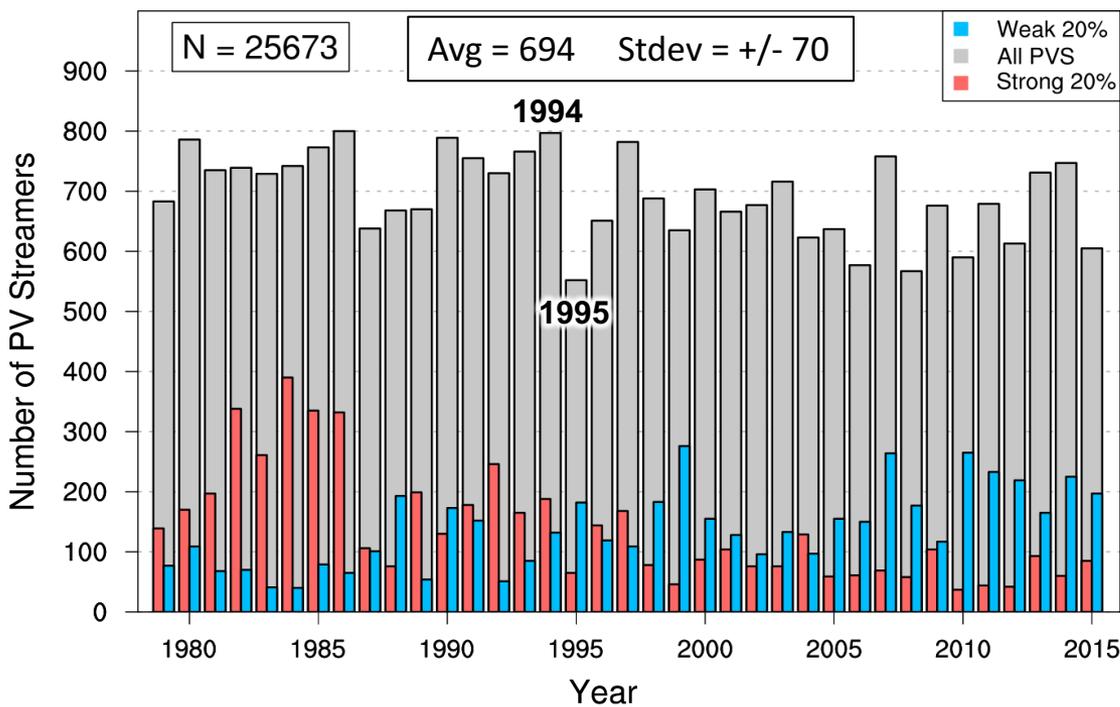
For this study, we emphasize the strongest and weakest 20 percentile of PV streamers



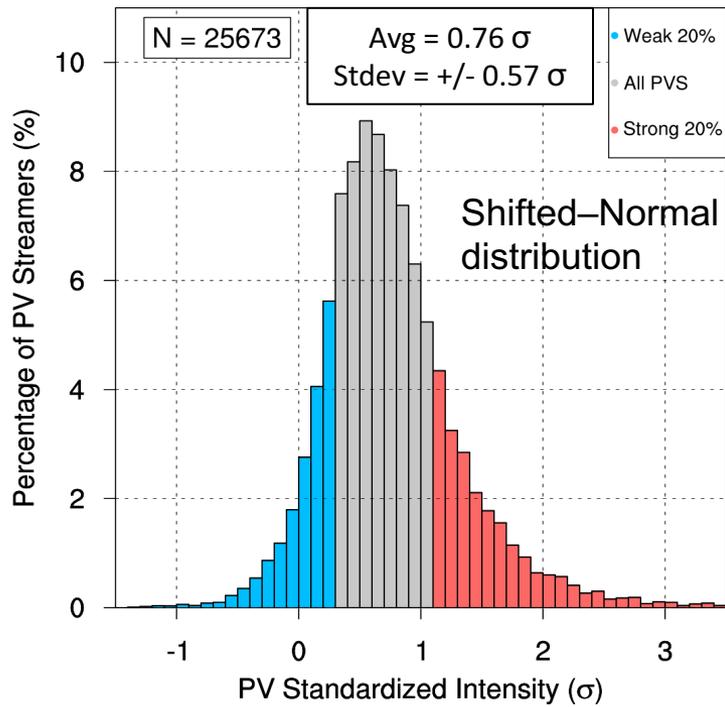
PDF of PV Streamer Intensity

Results: Variability of PV Streamer Intensity

- Interseasonal variability in the number of PV streamers
 - 1 Jun–30 Nov between 100–10°W



Yearly Distribution of PV Streamer Count

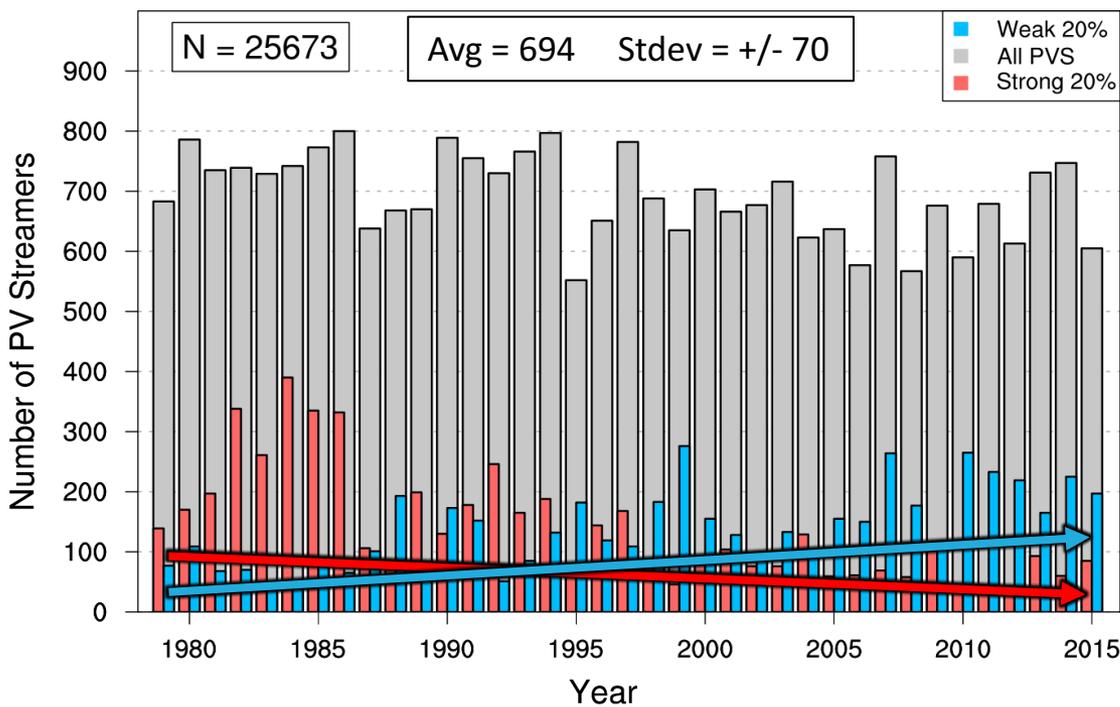


PDF of PV Streamer Intensity

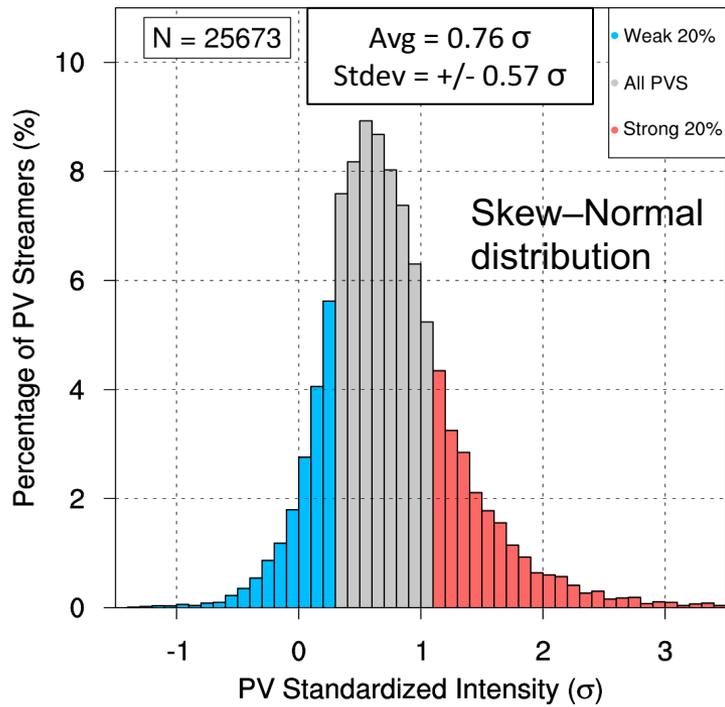
- ✧ CFSR depicts a trend towards a decrease in the number of strong PV streamers and increase in the number of weak PV streamers with time
- ✧ Questions remain if this is a real trend or is dataset dependent

Results: Variability of PV Streamer Intensity

- Interseasonal variability in the number of PV streamers
 - 1 Jun–30 Nov between 100–10°W



Yearly Distribution of PV Streamer Count



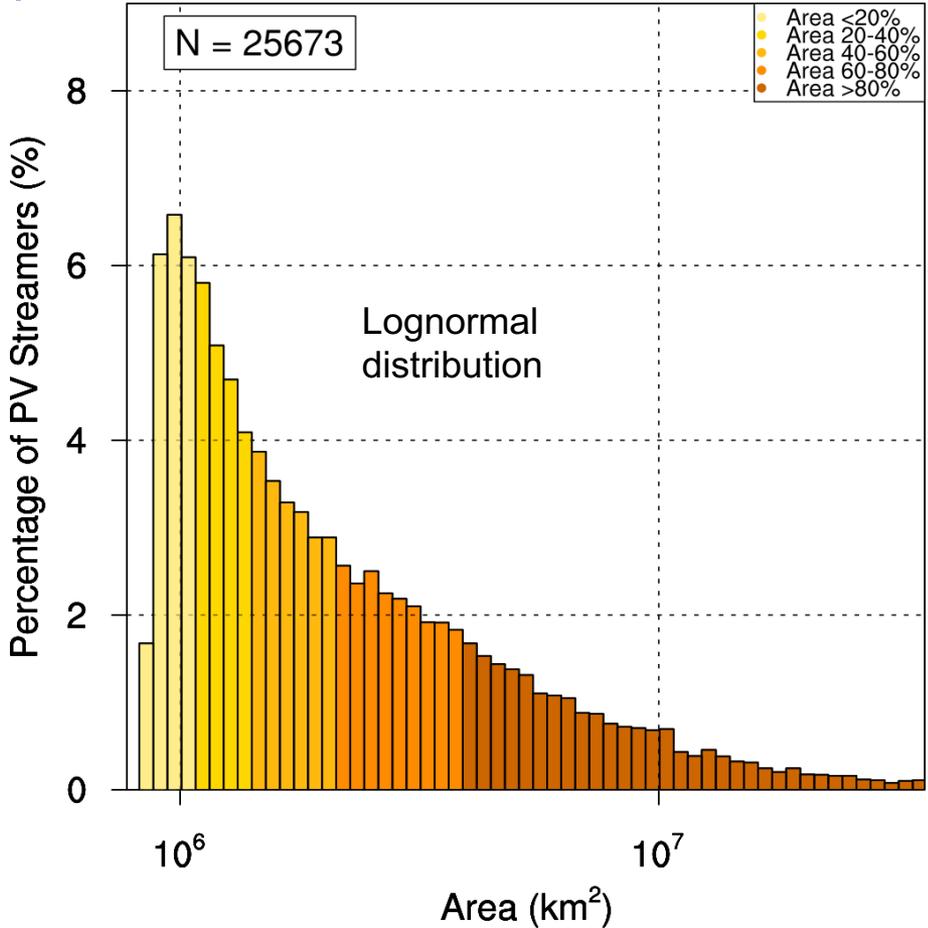
PDF of PV Streamer Intensity

- ✧ CFSR depicts a trend towards a decrease in the number of strong PV streamers and increase in the number of weak PV streamers with time
 - ✧ Questions remain if this is a real trend or is dataset dependent

Results: Variability of PV Streamer Area

- Area of PV streamers (raw value)
 - 1 Jun–30 Nov between 100–10°W
 - Mean: $4.05 \times 10^6 \text{ km}^2$
 - Stdev: $\pm 5.24 \times 10^6 \text{ km}^2$

Because of the highly skewed distribution, it is useful to use percentiles when describing area values.

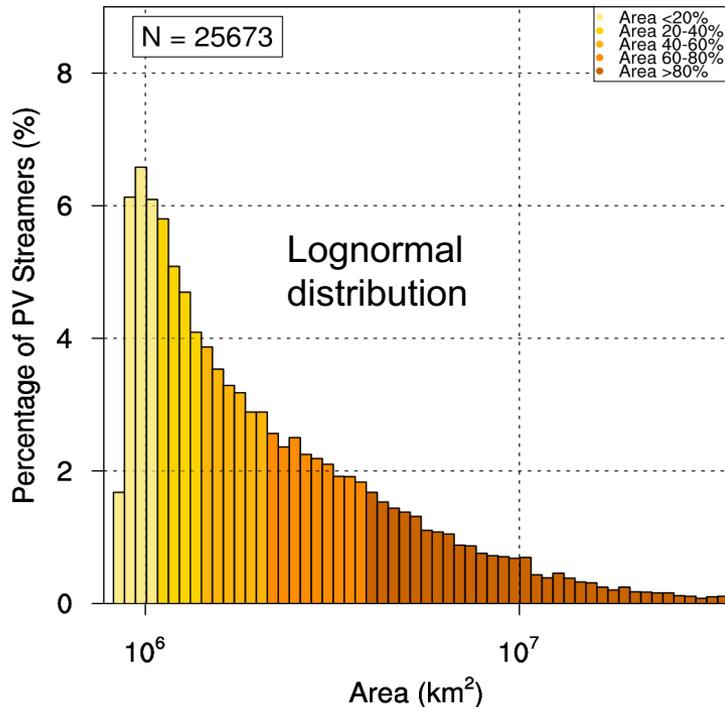
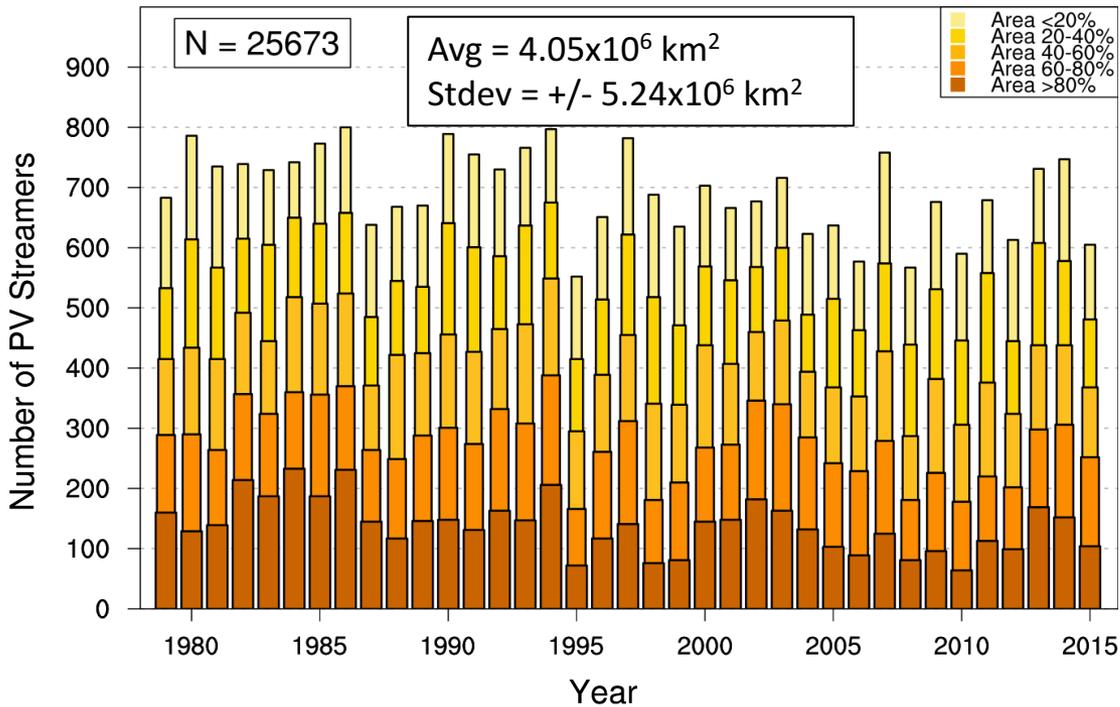


PDF of PV Streamer Area

Results: Variability of PV Streamer Area

- Interseasonal variability in PV streamer area

➤ 1 Jun–30 Nov between 100–10°W



Yearly Distribution of PV Streamer Count by Area

PDF of PV Streamer Area

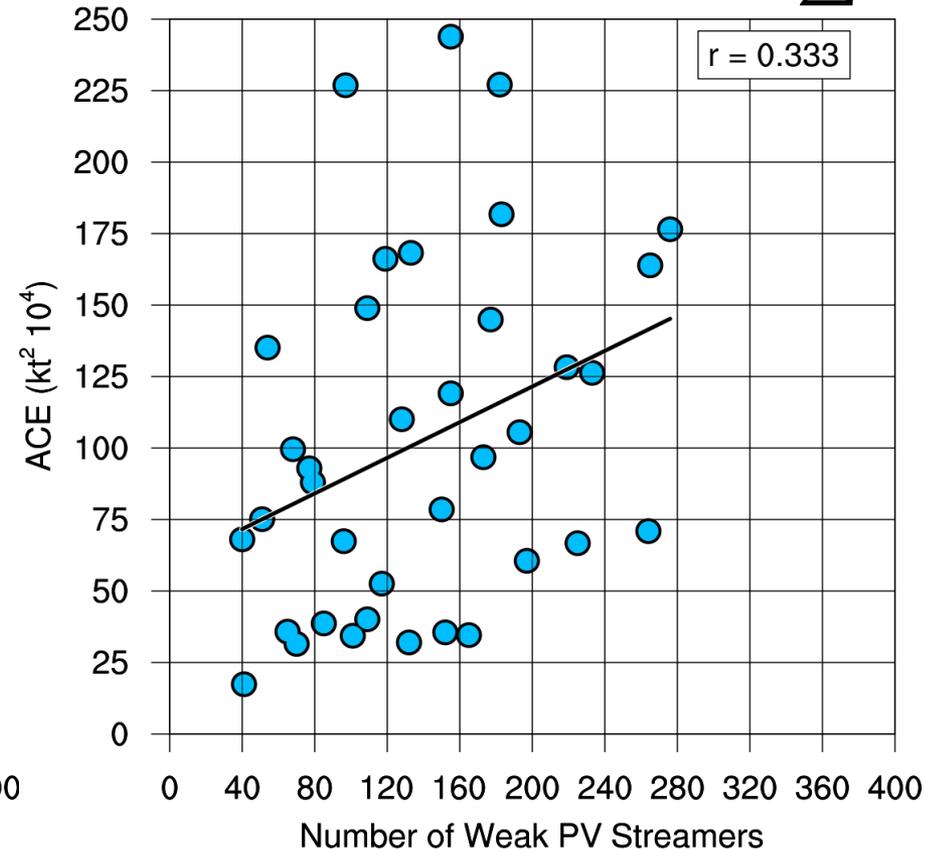
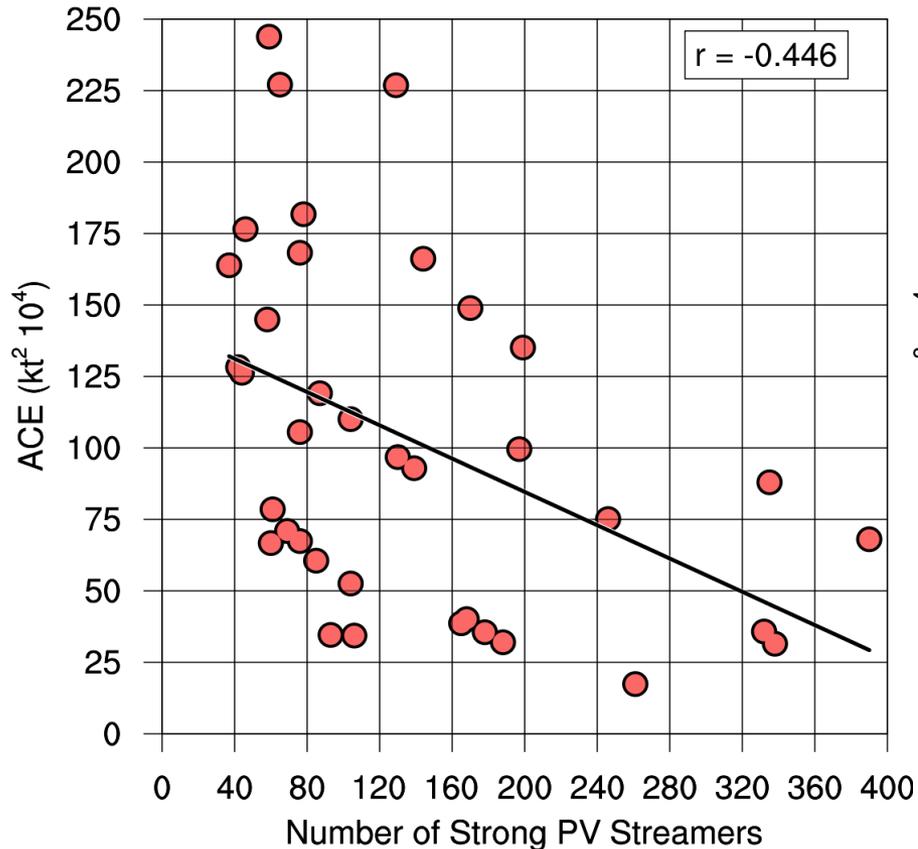
✧ Subtle decreasing trend in the number of large PV streamers with little noticeable trend in the number of small PV streamers

Results: Relationship to TC Activity

Results: Relationship to TC Activity – Intensity

- How is TC activity related to differences in PV streamer intensity

➤ Seasonal TC activity measured by accumulated cyclone energy (ACE) $ACE = 10^{-4} \sum v_{max}^2$



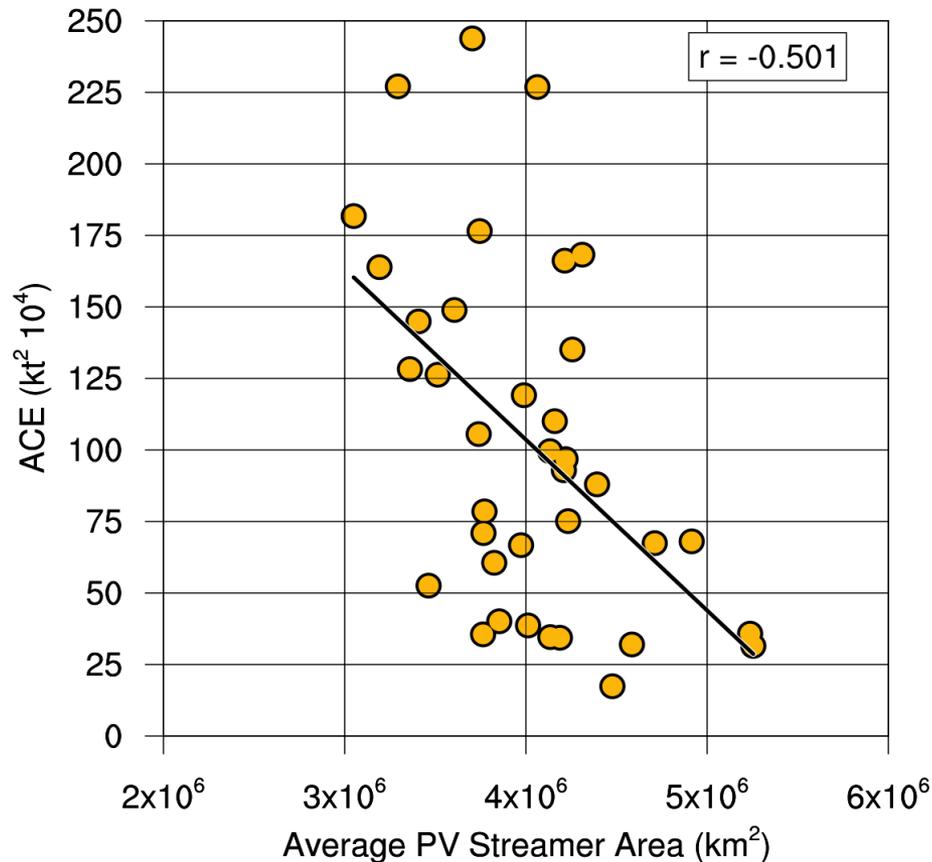
- ◇ Strong PV streamers are negatively correlated with ACE
- ◇ Weak PV streamers are positively correlated with ACE

Results: Relationship to TC Activity – Area

- How is TC activity related to differences in PV streamer area

➤ Seasonal TC activity measured by accumulated cyclone energy (ACE)

$$ACE = 10^{-4} \sum v_{max}^2$$



To further illustrate TC activity relationship to PV streamer intensity and area lets compare the most and least active TC seasons (using ACE).

✧ PV streamer area is negatively correlated with ACE

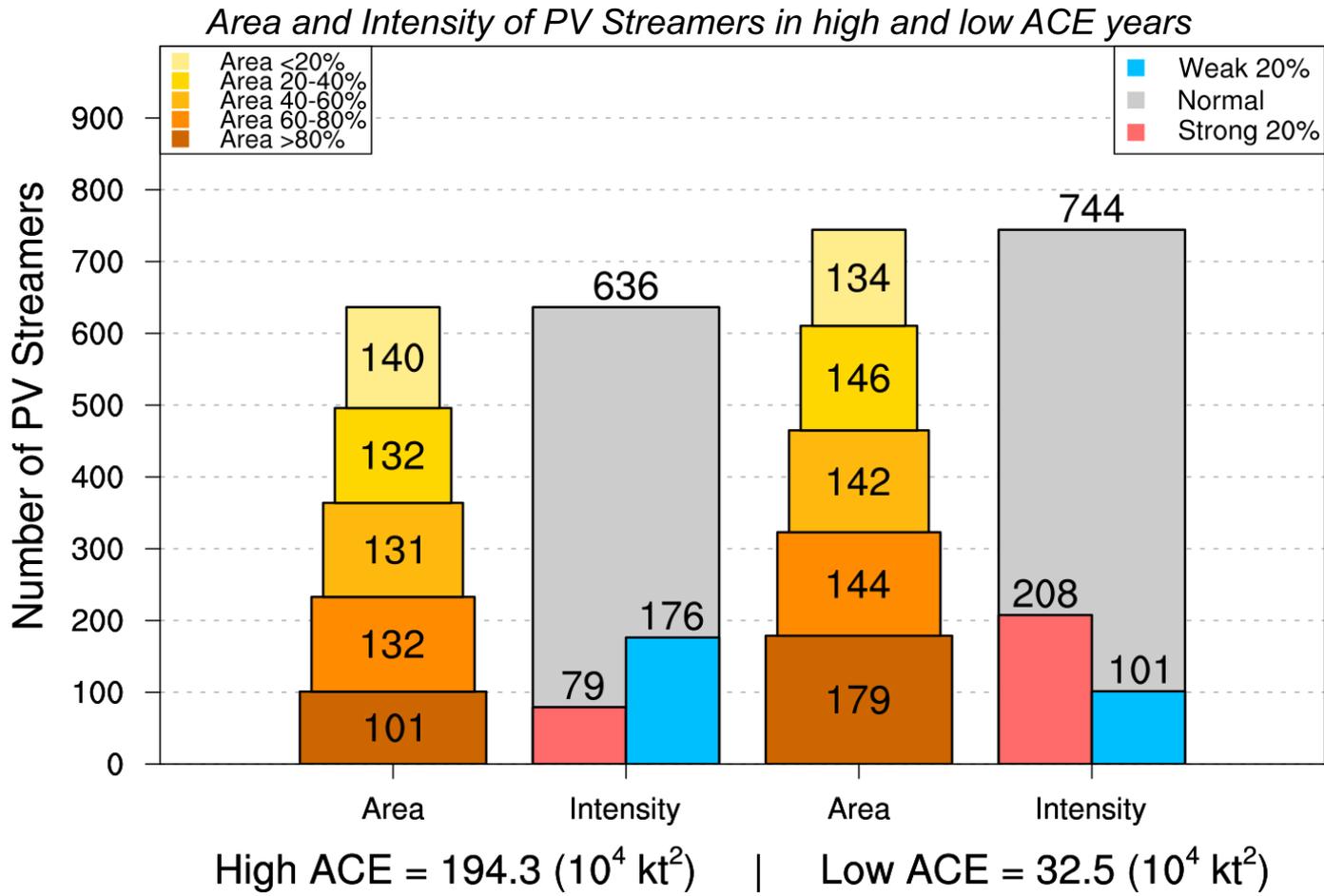
Results: TC Activity – High vs. Low ACE

- Active TC seasons (High ACE) feature:

- ✧ Fewer PV streamers (-14.5%)
- ✧ Much Fewer Strong PV streamers (-62.0%)
- ✧ Many More Weak PV streamers (+74.2%)
- ✧ Less Large PV Streamers (-56.4%)

Combine these variables into a seasonal index

Top 8 High ACE Years	Bottom 8 Low ACE Years
2005	1983
1995	1982
2004	1994
1998	1987
1999	2013
2003	1991
1996	1986
2010	1993



Results: PV Streamer Intensity Metric vs. ACE

- Putting these variables together over the course of a season

- PV Streamer Occurrence
- PV Streamer Area
- PV Streamer Intensity

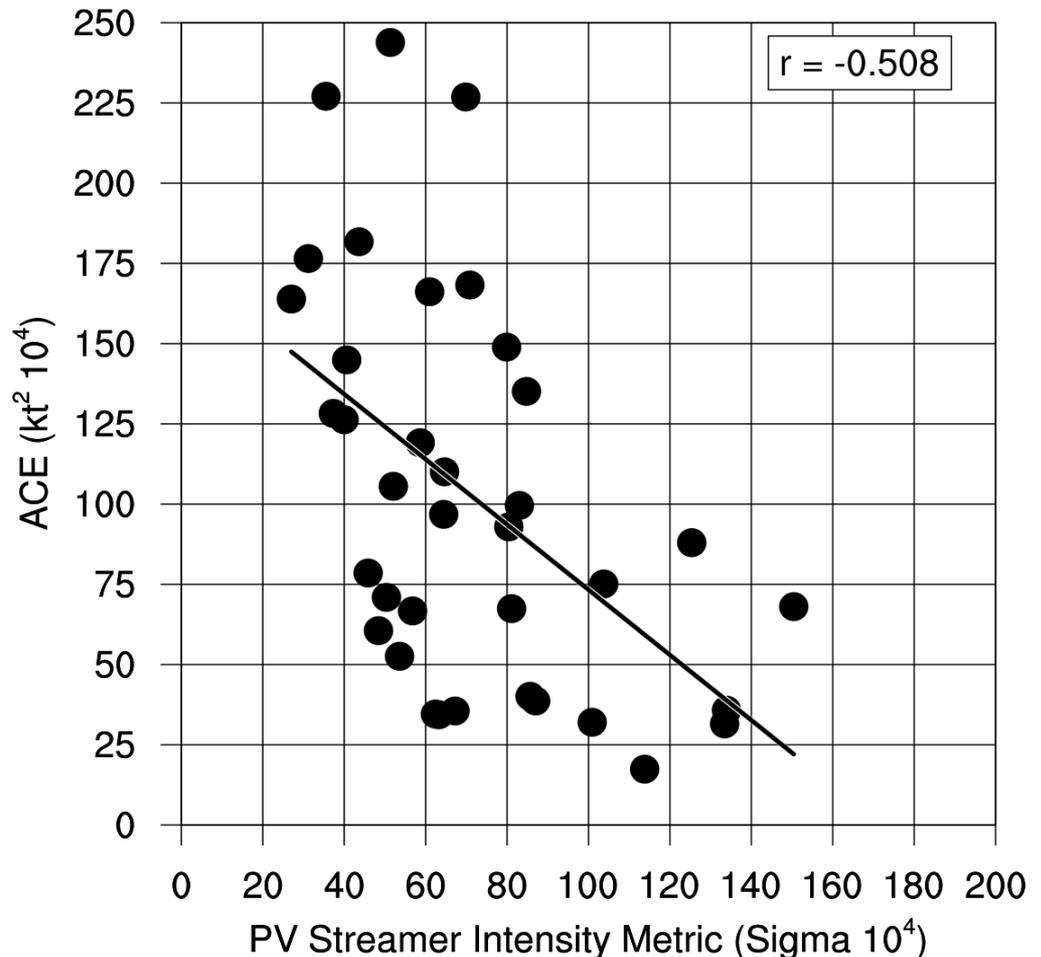
$$\text{Seasonal Intensity Metric} = \iint PV_{std_anom} dAdt$$

June–November 1979–2015

- Combining these variables produces the highest negative correlation between PV streamer activity and ACE

- Caveats

- Not a perfectly linear relationship
- Other factors may be responsible for correlation
 - ENSO
 - SSTA



Results: PV Streamer Intensity Metric vs. ACE

- Putting these variables together over the course of a season

- PV Streamer Occurrence
- PV Streamer Area
- PV Streamer Intensity

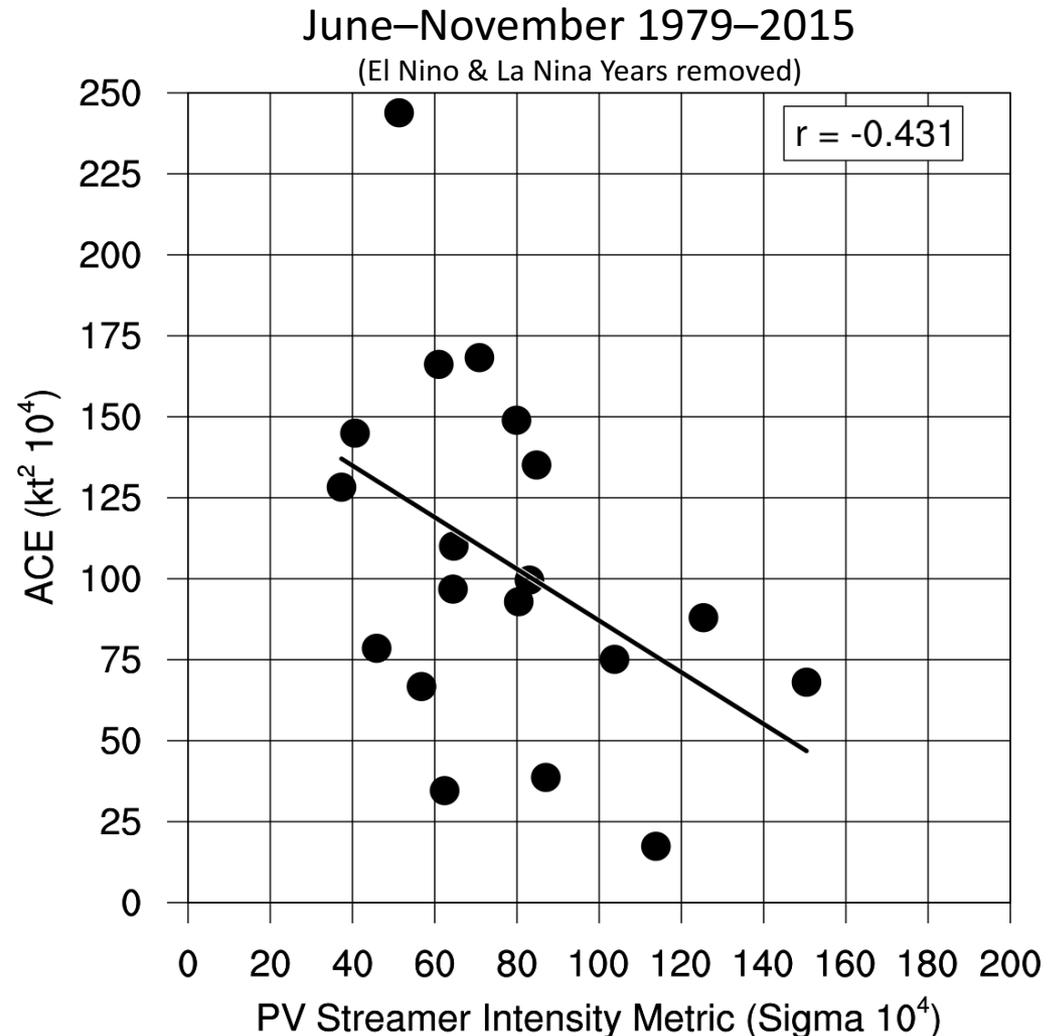
$$\text{Seasonal Intensity Metric} = \iint PV_{std_anom} dAdt$$

- Combining these variables produces the highest negative correlation between PV streamer activity and ACE

- Caveats

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 - SSTA

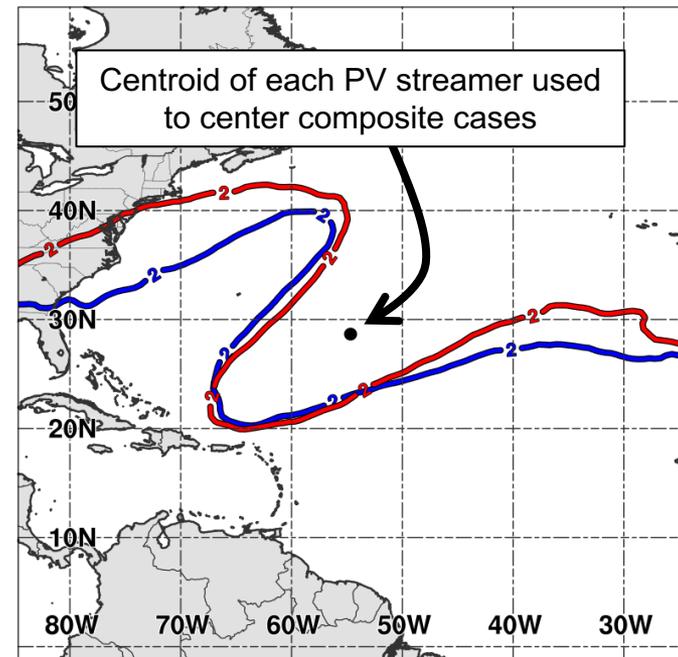
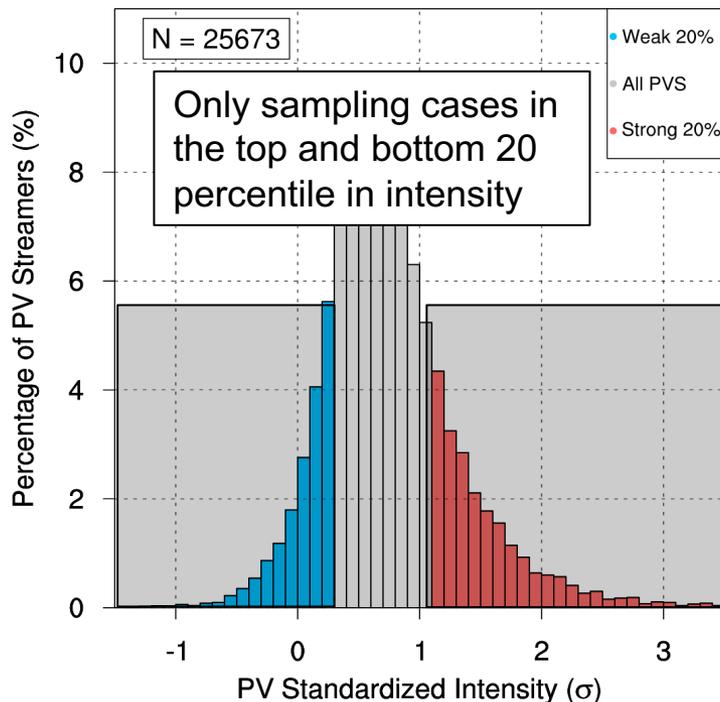
Proven correlation but can we now explain causation?



Results: Composite Differences

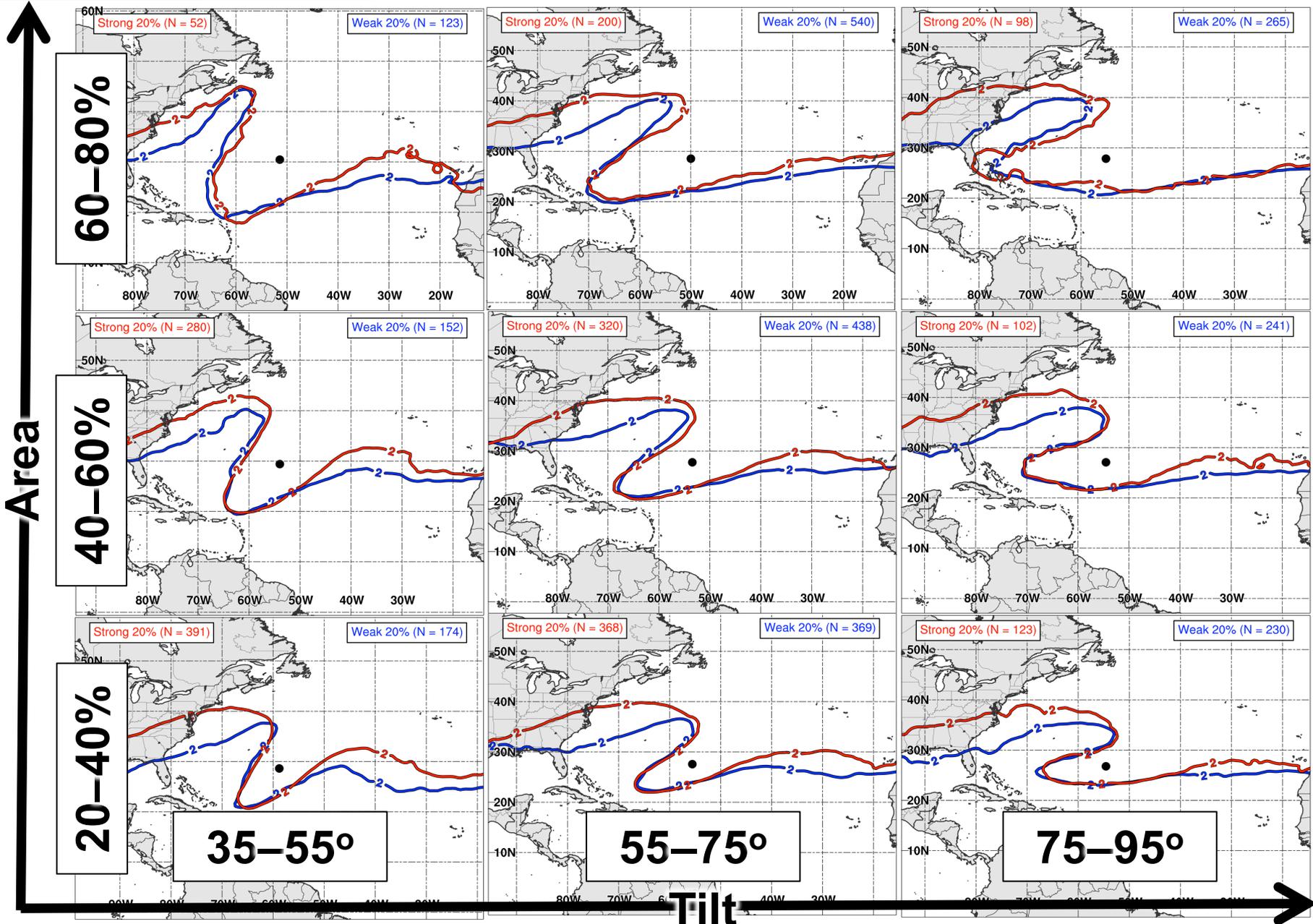
Results: PV Streamer Composite Differences

- There are a number of ways to composite PV streamers
 - Similar intensities (using top and bottom 20 percentile categories)
 - Similar areas (using percentile areas)
 - Similar tilts (using tilt degree thresholds)
- Emphasis of this study is on PV streamer intensity differences
 - Compare strong and weak PV streamers using similar areas and tilts
 - A “composite matrix” can be produced by organizing composites by tilt and area

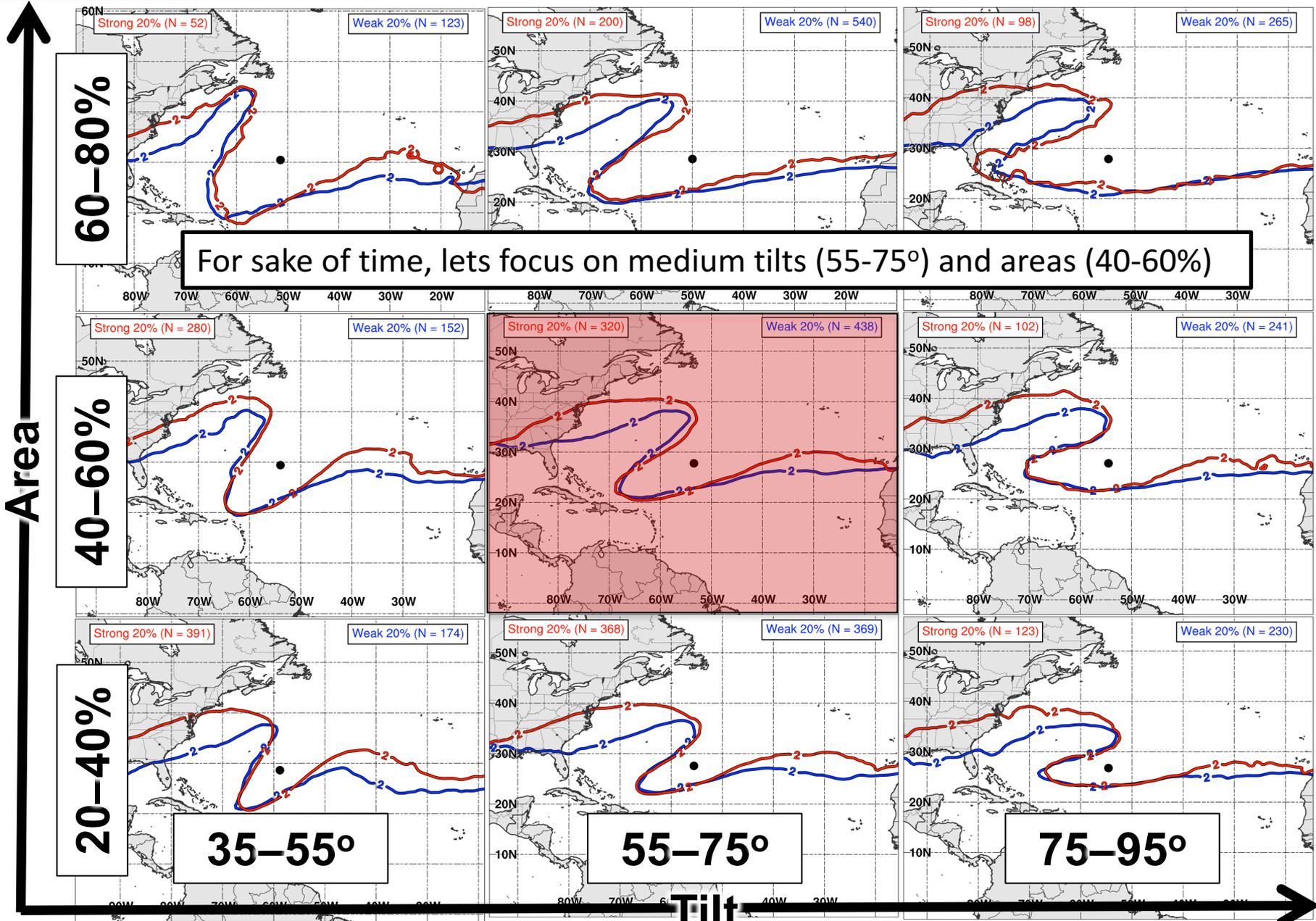


Average 350-K 2-PVU contour for weak and strong PV streamers (blue and red contours respectively)

Results: PV Streamer Composite Differences



Results: PV Streamer Composite Differences



Results: PV Streamer Composite Differences

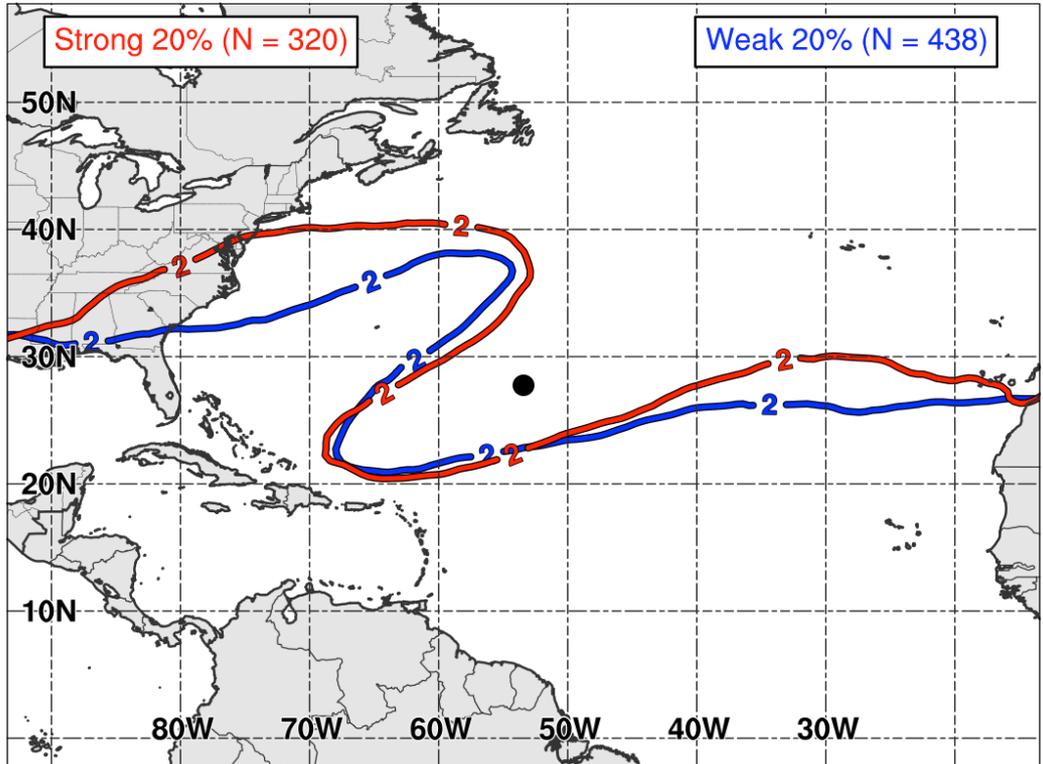
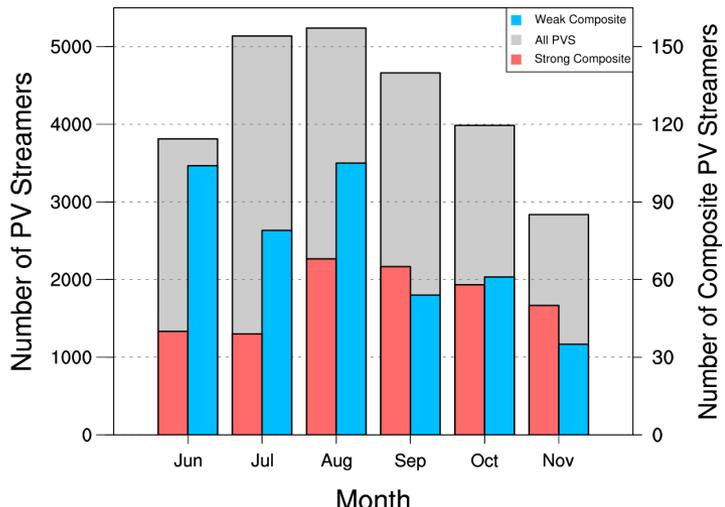
- Composite differences created between strong & weak PV streamers
 - Variable anomalies are normalized to allow for comparison of PV streamers in different locations or different times for each case

$$\Delta x_{strong-weak} = \left\{ \frac{\bar{x}_i^{Composite} - \bar{x}_i^{Mean}}{\sigma_{x_i}} \right\}_{strong} - \left\{ \frac{\bar{x}_i^{Composite} - \bar{x}_i^{Mean}}{\sigma_{x_i}} \right\}_{weak}$$

$N_{strong} = 320$
 $N_{weak} = 438$

Average 350-K 2-PVU contour for weak and strong PV streamers (blue and red contours respectively)

- Statistical significance will be assessed using bootstrap resampling (useful when comparing different sample sizes)

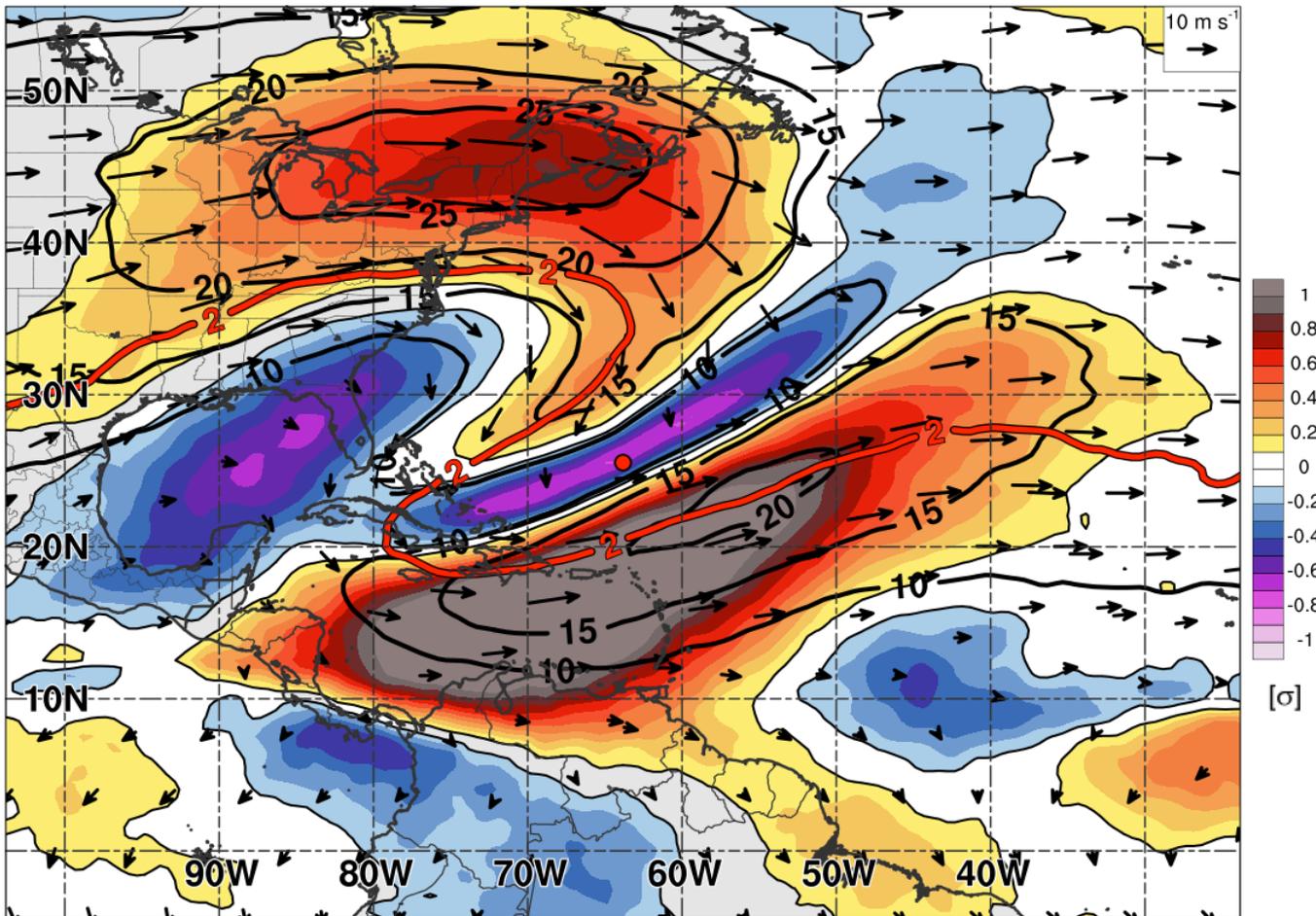


Results: Composite Differences - VWS

- Vertical Wind Shear (VWS)
 - 200–850-hPa

$$\left\{ \frac{\bar{x}_i^{Composite} - \bar{x}_i^{Mean}}{\sigma_{x_i}} \right\}_{strong} \quad N_{strong} = 320$$

❖ Enhanced shear primarily equatorward of strong PV streamer > 1σ



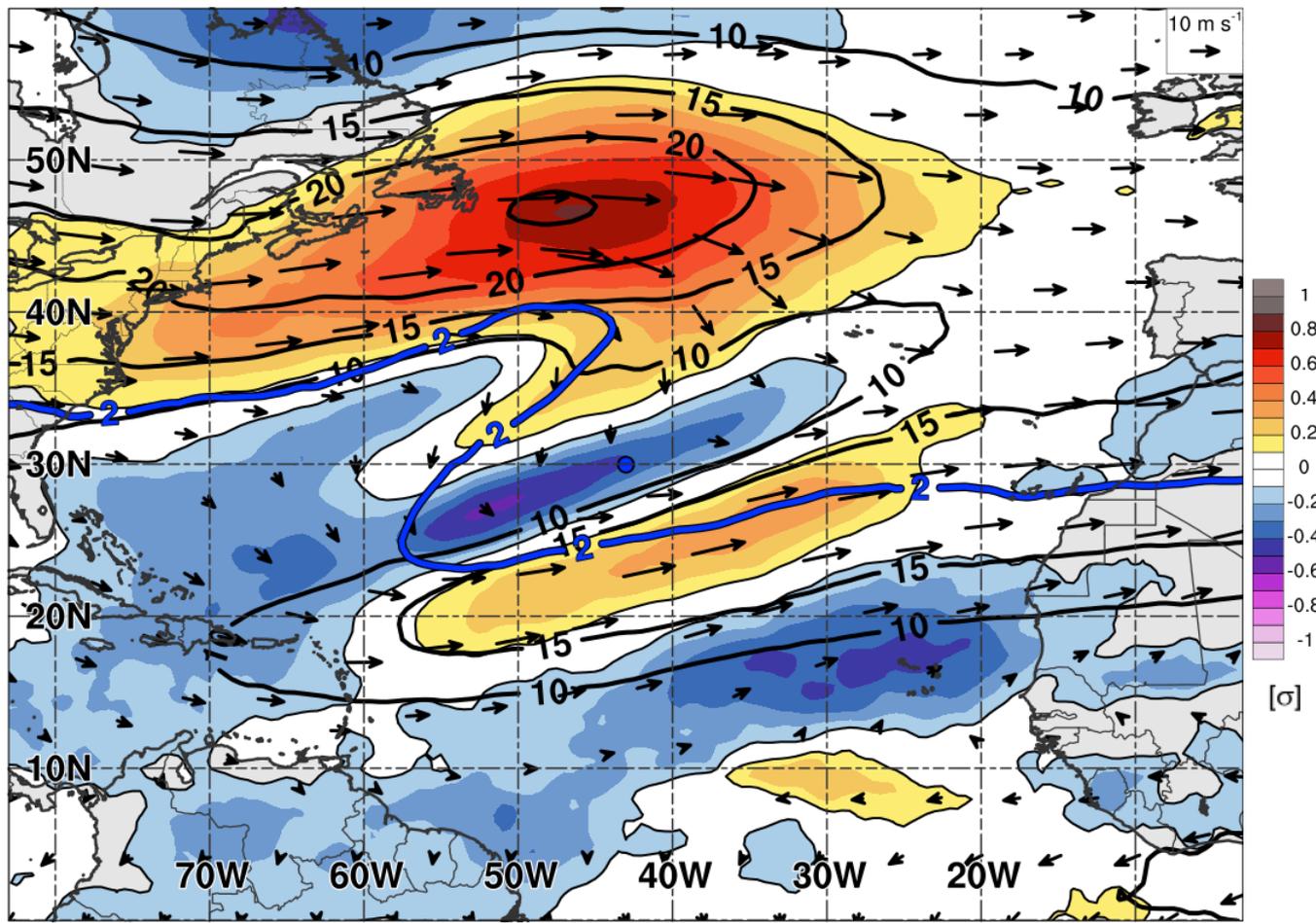
Normalized VWS Differences (shaded, σ), Mean magnitude VWS (black contours, m s⁻¹) and direction (vectors), Mean 350-K 2-PVU contour of strong PV streamers (red line)

Results: Composite Differences - VWS

- Vertical Wind Shear (VWS)
 - 200–850-hPa

$$\left\{ \frac{\bar{x}_i^{Composite} - \bar{x}_i^{Mean}}{\sigma_{x_i}} \right\}_{weak} \quad N_{weak} = 438$$

❖ Only small corridor of enhanced shear equatorward of PV streamer
 $< 0.5\sigma$



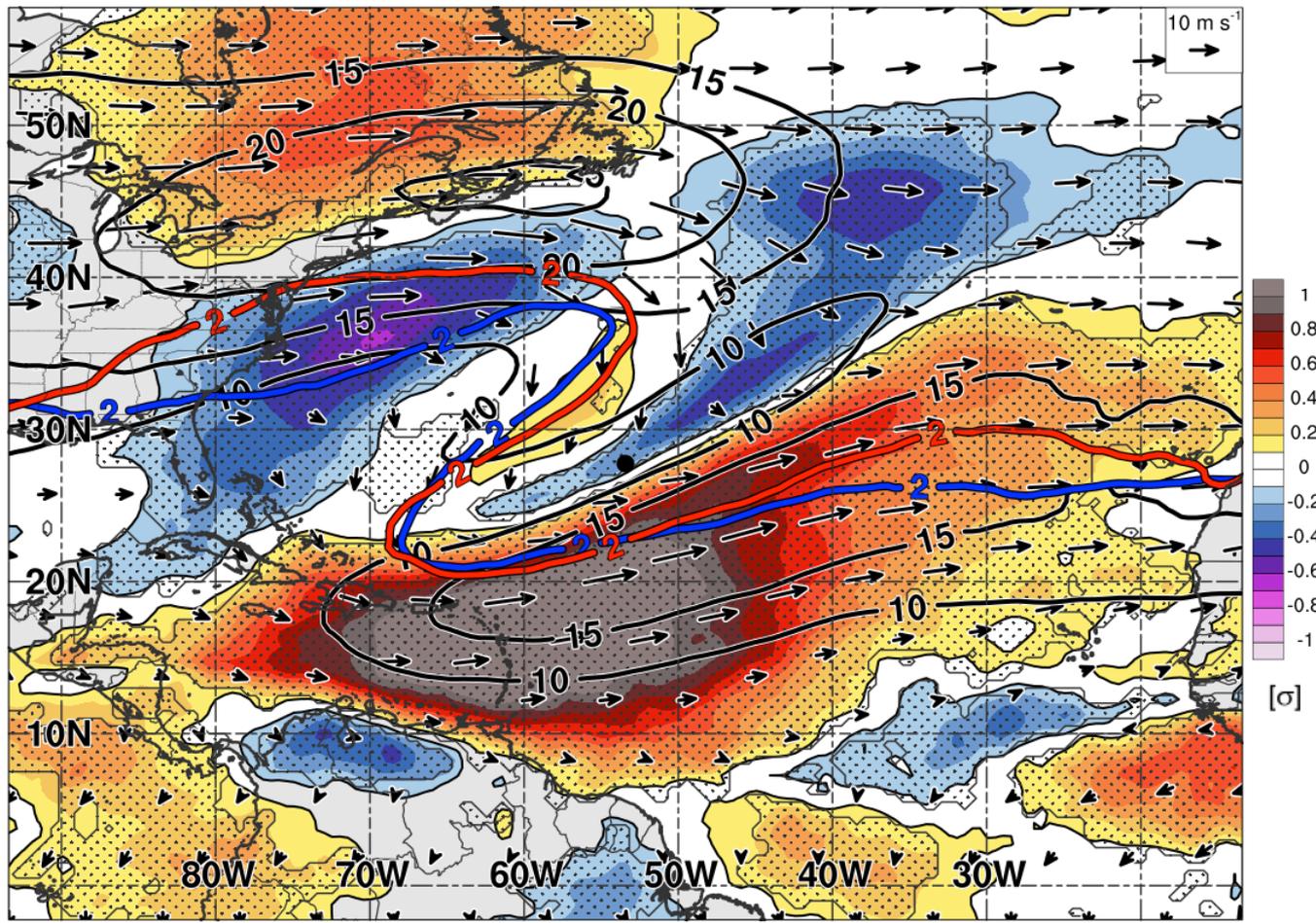
Normalized VWS Differences (shaded, σ), Mean magnitude VWS (black contours, $m\ s^{-1}$) and direction (vectors), Mean 350-K 2-PVU contour of weak PV streamers (blue line)

Results: Composite Differences - VWS

- Vertical Wind Shear (VWS)
 - 200–850-hPa

$$\Delta x_{strong-weak}$$

❖ Strong PV streamers have a much larger and more intense corridor of VWS



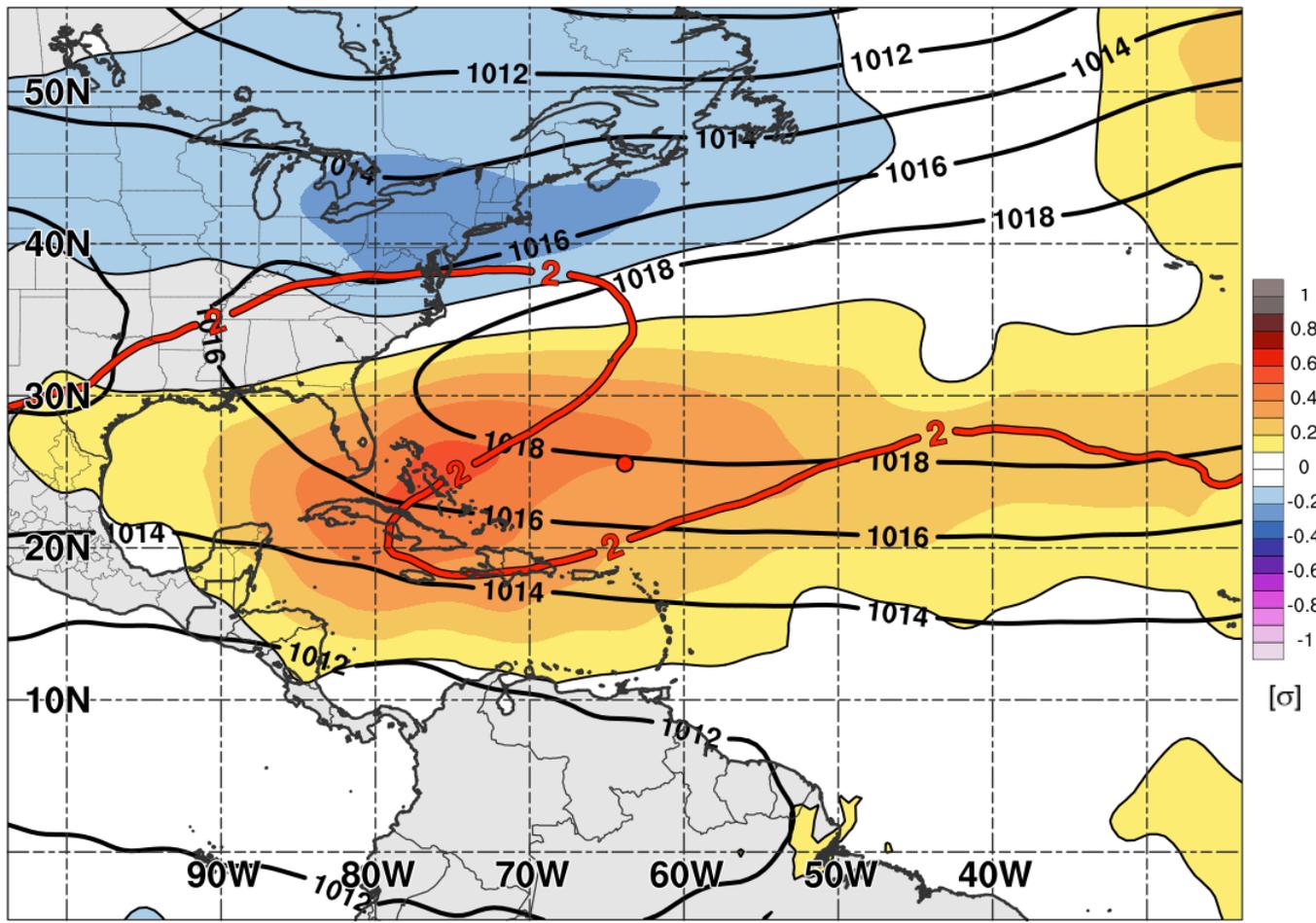
Normalized VWS Differences (shaded, σ), Mean magnitude VWS (black contours, $m s^{-1}$) and direction (vectors), Mean 350-K 2-PVU contour of strong & weak PV streamers (blue and red lines respectively), hatched areas indicate statistical significance to the 99% confidence interval

Results: Composite Differences - SLP

- Sea Level Pressure (SLP)

$$\left\{ \frac{\bar{x}_i^{Composite} - \bar{x}_i^{Mean}}{\sigma_{x_i}} \right\}_{strong} \quad N_{strong} = 320$$

Enhanced SLP over strong PV streamer trough axis equatorward of strong PV streamer



Normalized SLP Differences (shaded, σ), Mean magnitude SLP (black contours, hPa)
 Mean 350-K 2-PVU contour of strong PV streamers (red line)

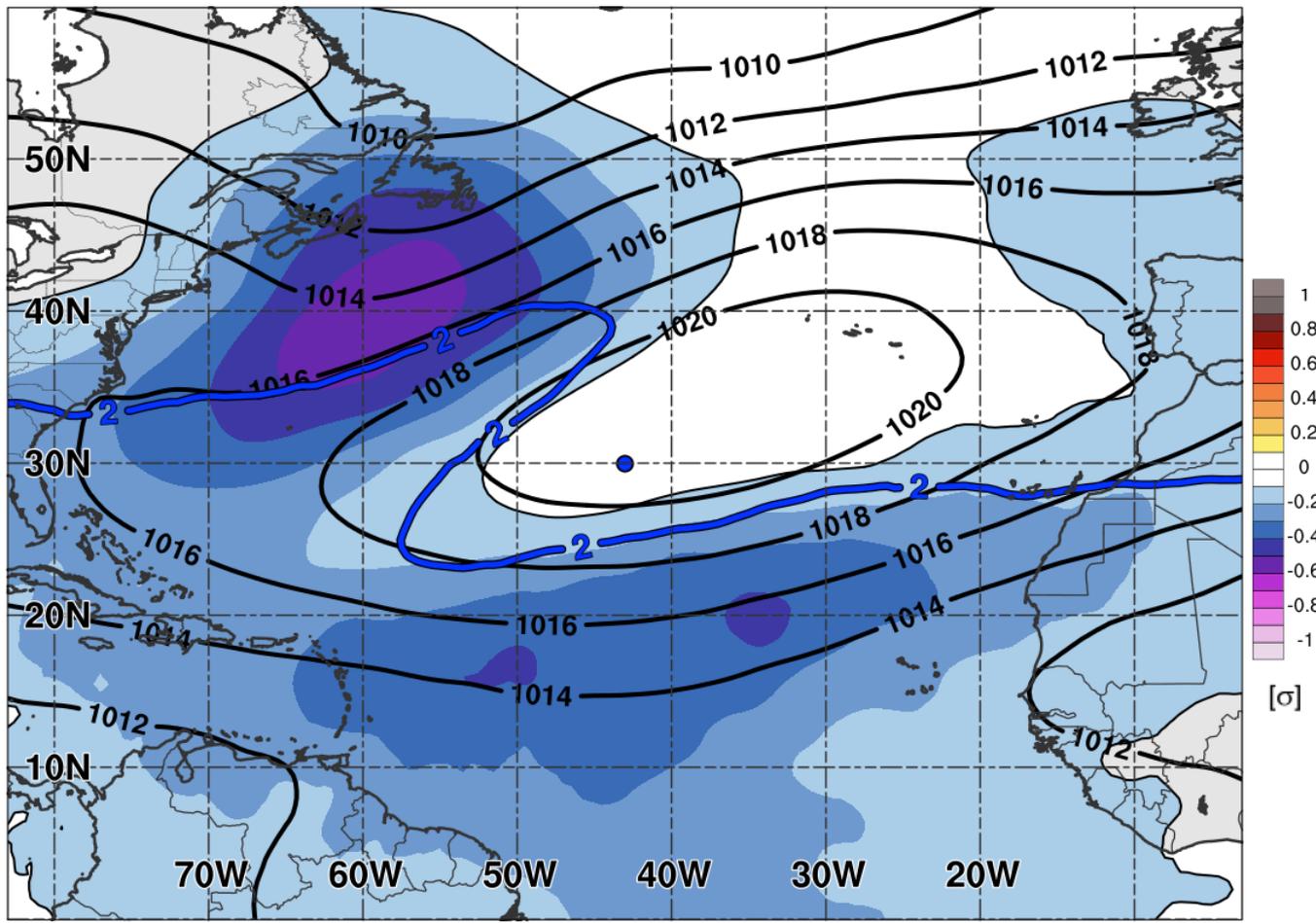
Results: Composite Differences - SLP

- Sea Level Pressure (SLP)

$$\left\{ \frac{\bar{x}_i^{Composite} - \bar{x}_i^{Mean}}{\sigma_{x_i}} \right\}_{weak}$$

$N_{weak} = 438$

❖ Reduced SLP surrounding weak PV streamer trough axis



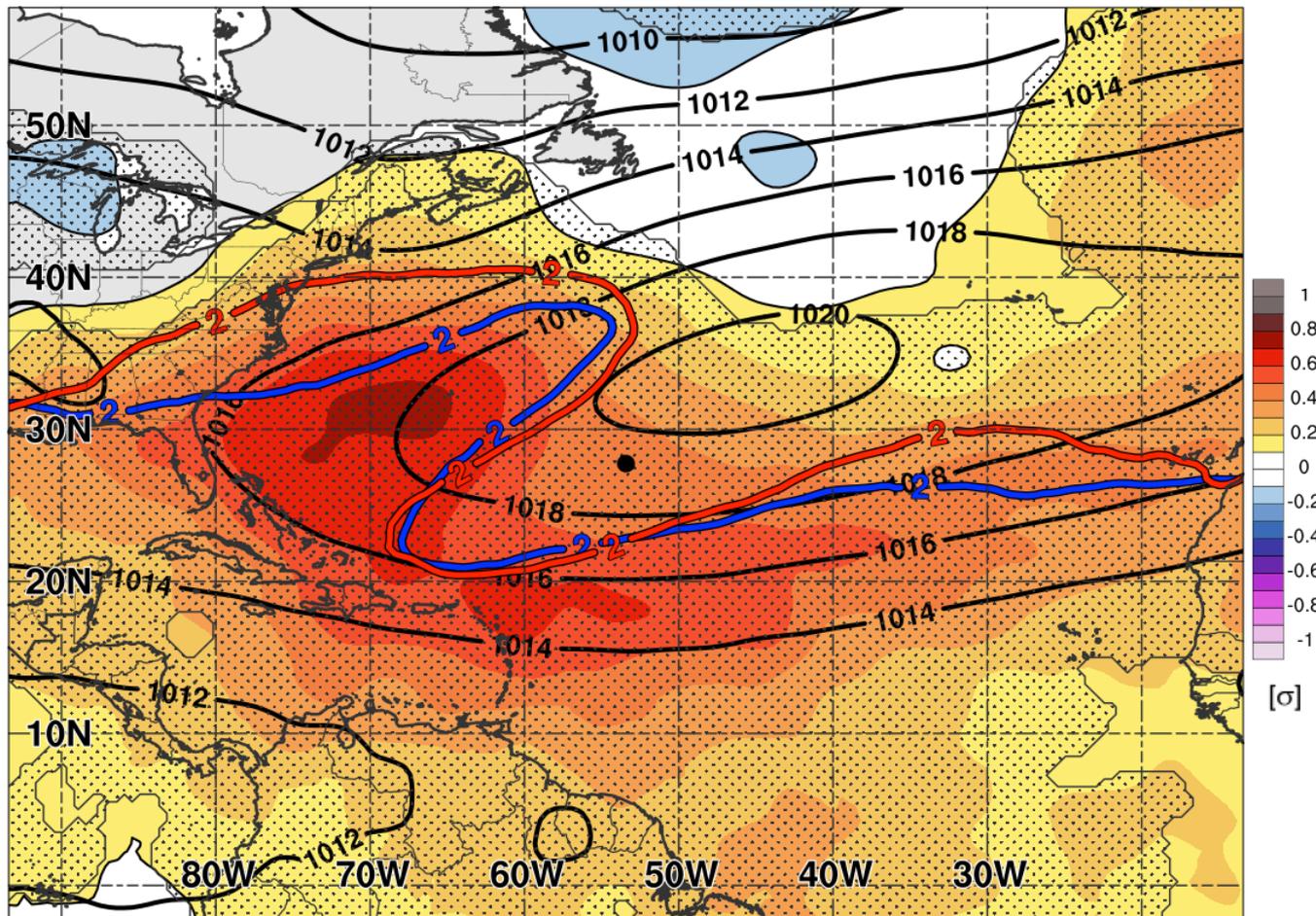
Normalized SLP Differences (shaded, σ), Mean magnitude SLP (black contours, hPa) and Mean 350-K 2-PVU contour of weak PV streamers (blue line)

Results: Composite Differences - SLP

- Sea Level Pressure (SLP)

$$\Delta x_{strong-weak}$$

❖ Near basin wide increased in SLP, especially upstream of PV streamer trough axis



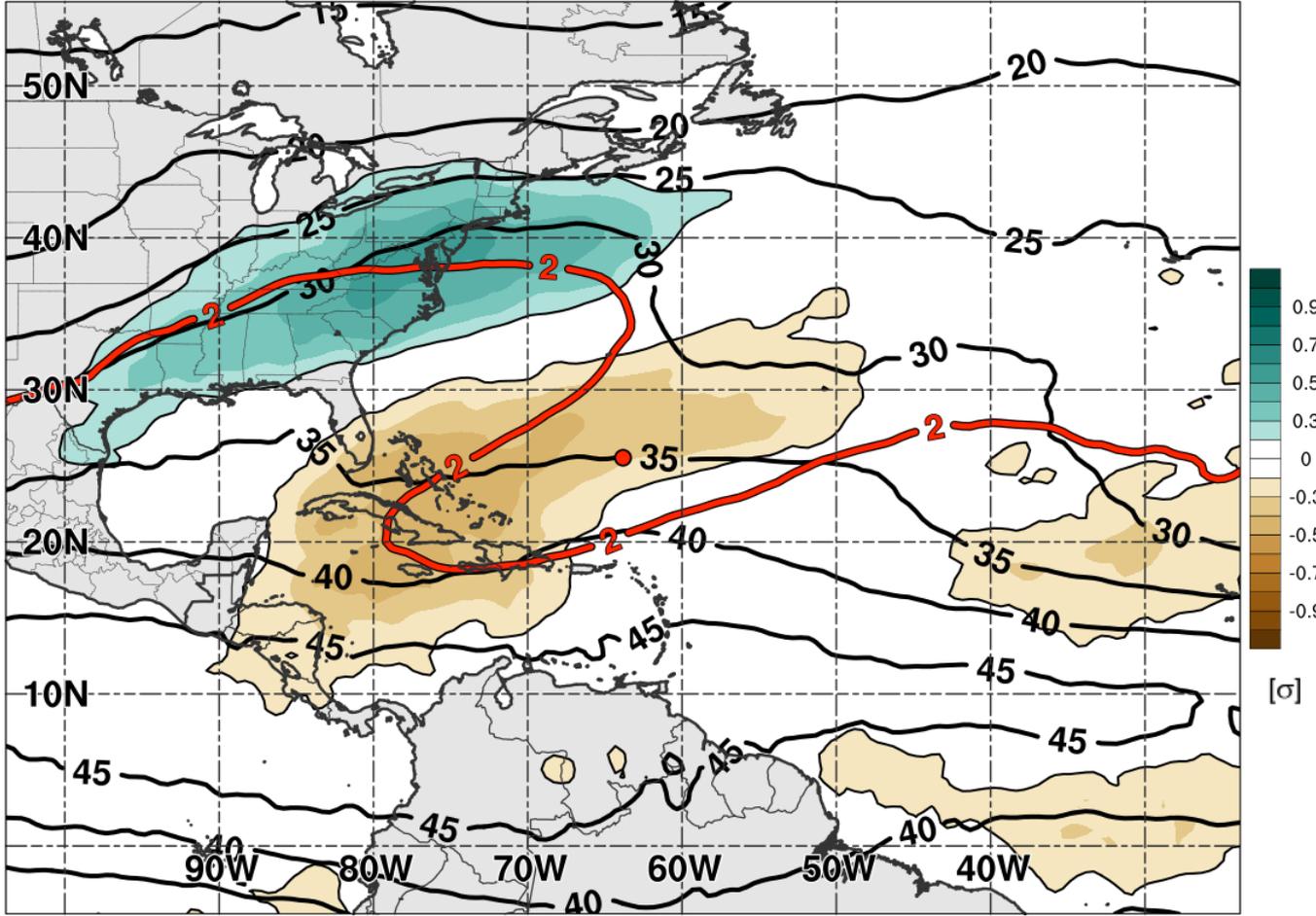
Normalized SLP Differences (shaded, σ), Mean magnitude SLP (black contours, hPa)
Mean 350-K 2-PVU contour of strong and weak PV streamers (red and blue lines respectively)
hatched areas indicate statistical significance to the 99% confidence interval

Results: Composite Differences - PW

- Precipitable Water (PW)

$$\left\{ \frac{\bar{x}_i^{Composite} - \bar{x}_i^{Mean}}{\sigma_{x_i}} \right\}_{strong} \quad N_{strong} = 320$$

- ❖ Couplet of increased moisture upstream with decreased moisture along strong PV streamer trough axis



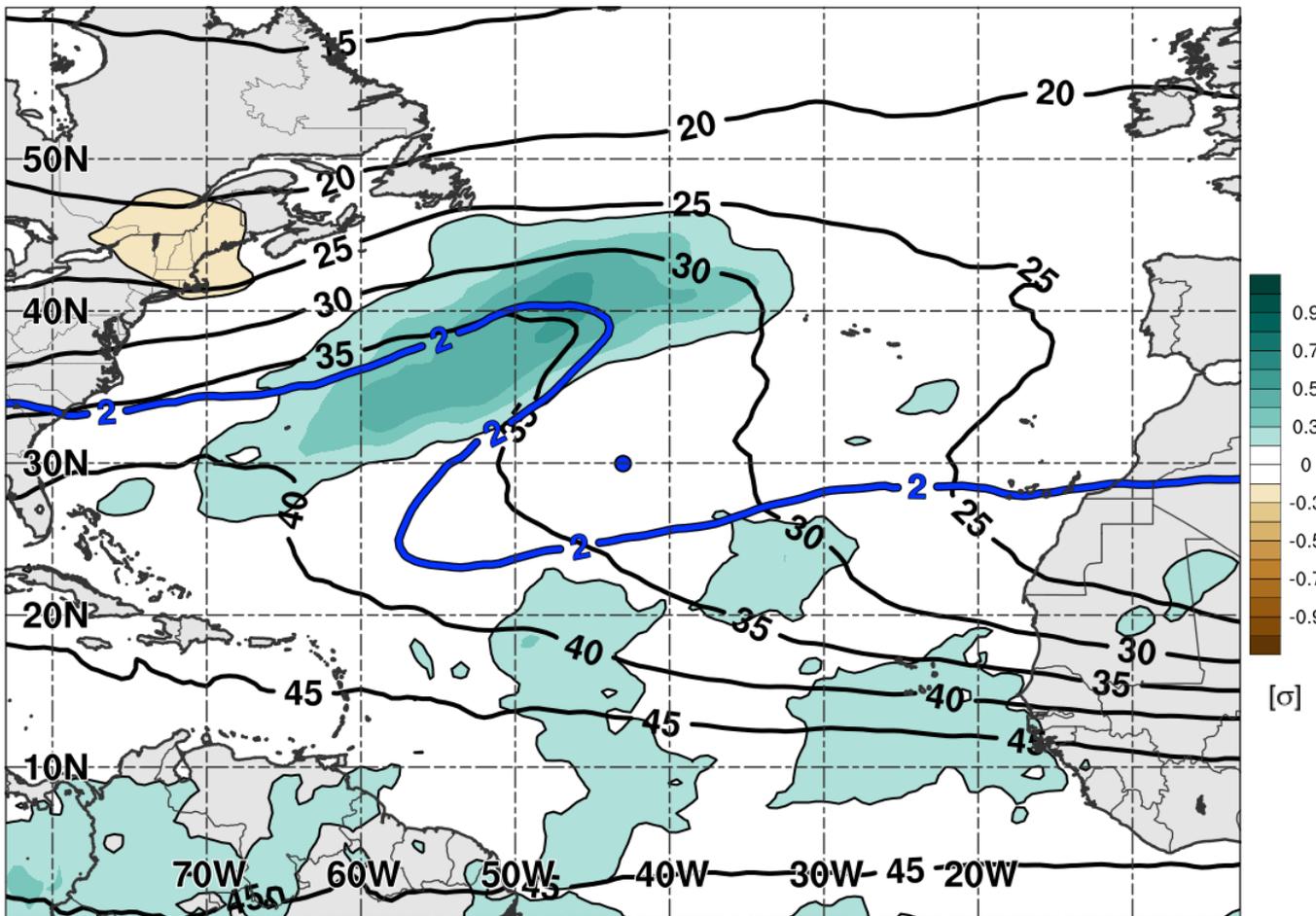
Normalized PW Differences (shaded, σ), Mean magnitude PW (black contours, mm)
 Mean 350-K 2-PVU contour of strong PV streamers (red line)

Results: Composite Differences - PW

- Precipitable Water (PW)

$$\left\{ \frac{\bar{x}_i^{Composite} - \bar{x}_i^{Mean}}{\sigma_{x_i}} \right\}_{weak} \quad N_{weak} = 438$$

- ❖ Only increased moisture upstream of weak PV streamer



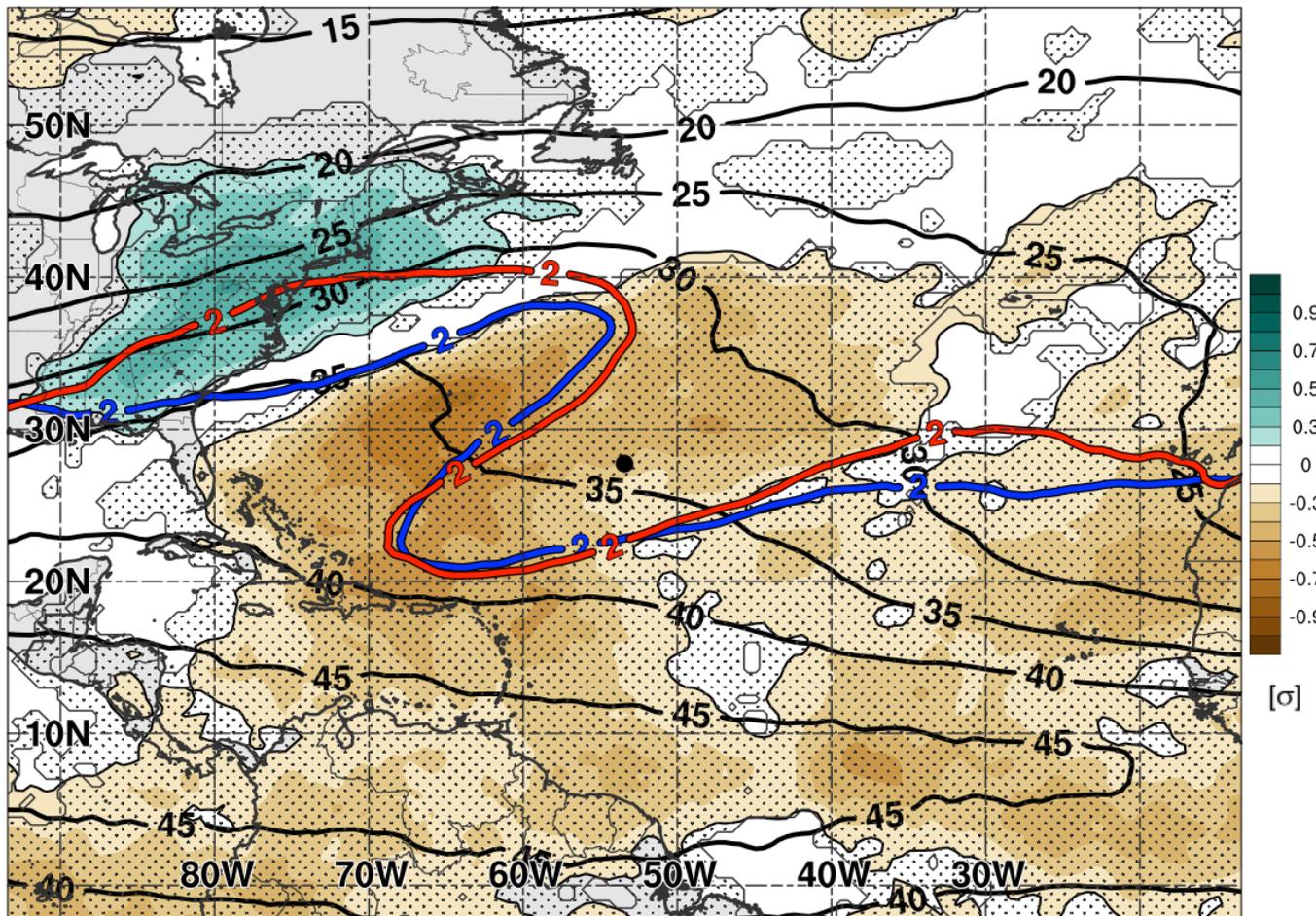
Normalized PW Differences (shaded, σ), Mean magnitude PW (black contours, mm)
 Mean 350-K 2-PVU contour weak PV streamers (blue line)

Results: Composite Differences - PW

- Precipitable Water (PW)

$$\Delta x_{strong-weak}$$

❖ Large region of negative PW anomalies surrounding PV streamer trough axis

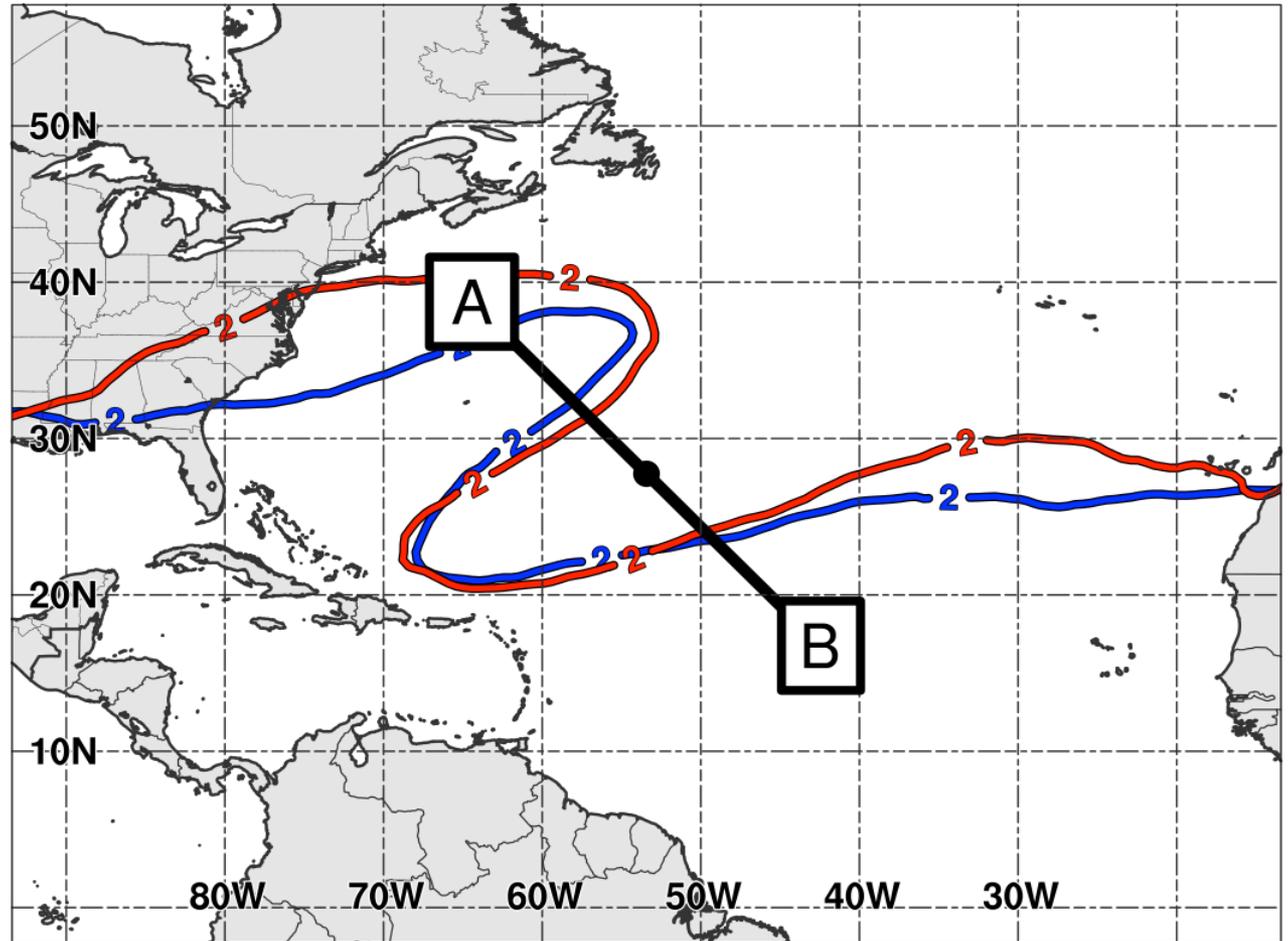


Normalized PW Differences (shaded, σ), Mean magnitude PW (black contours, mm)
Mean 350-K 2-PVU contour of strong and weak PV streamers (red and blue lines respectively)
hatched areas indicate statistical significance to the 99% confidence interval

Results: Composite Differences - PW

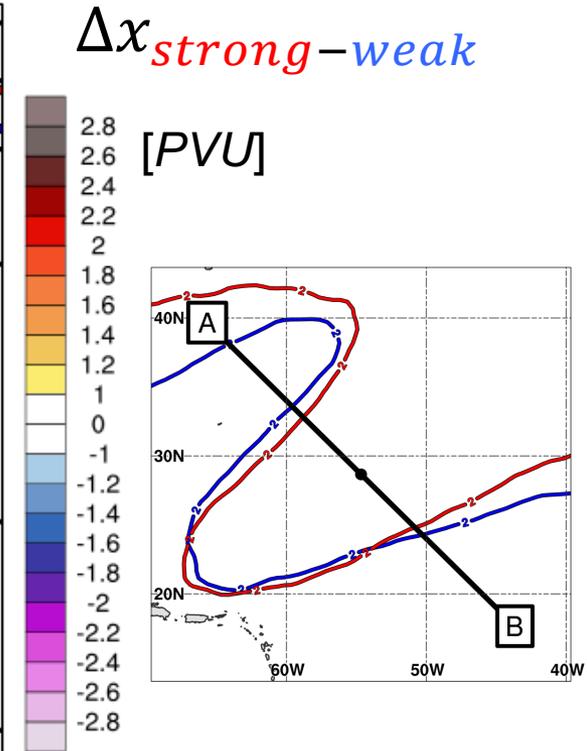
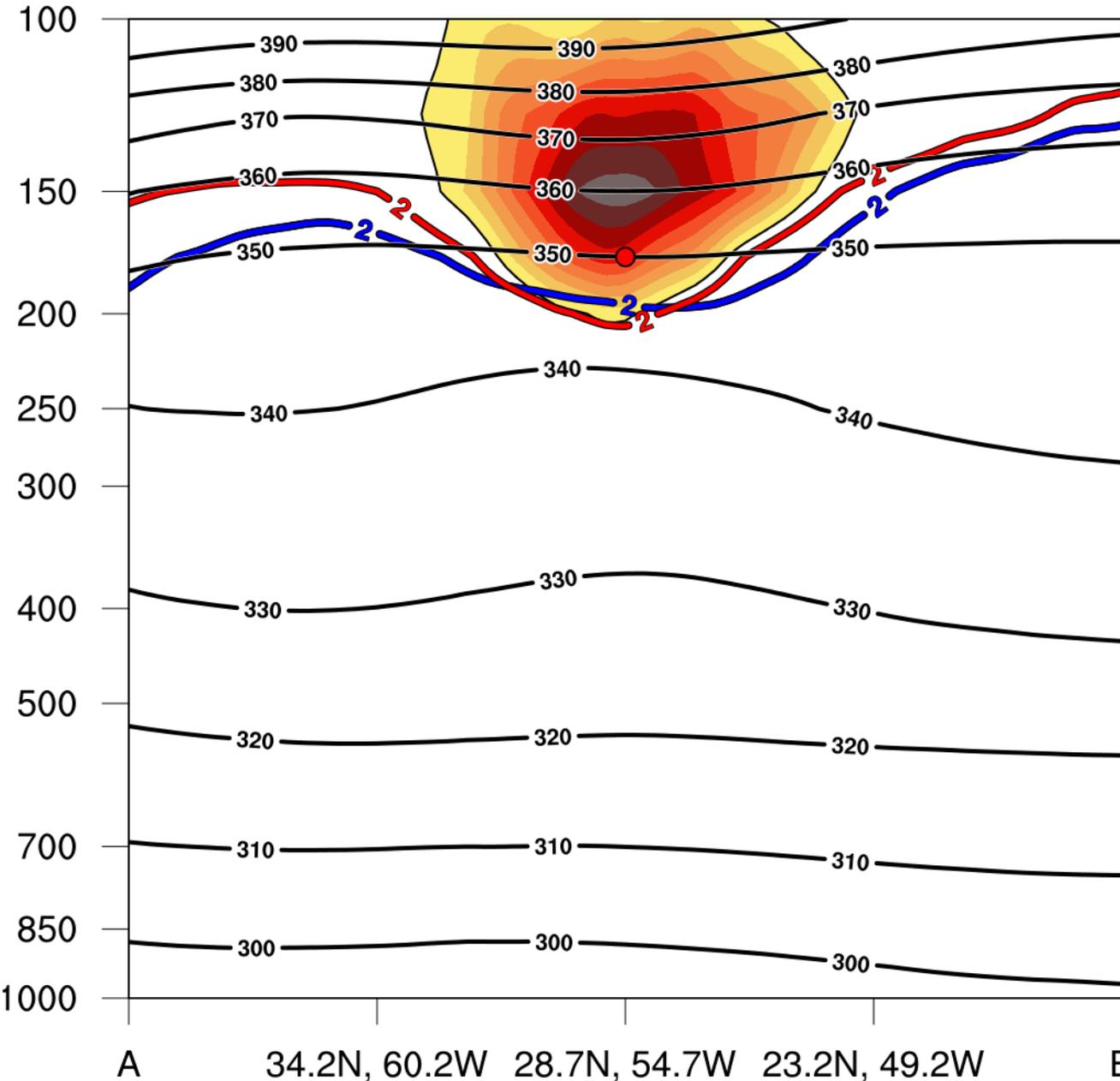
- Lets take a cross-section through the PV streamer trough

$$\Delta x_{\text{strong-weak}}$$

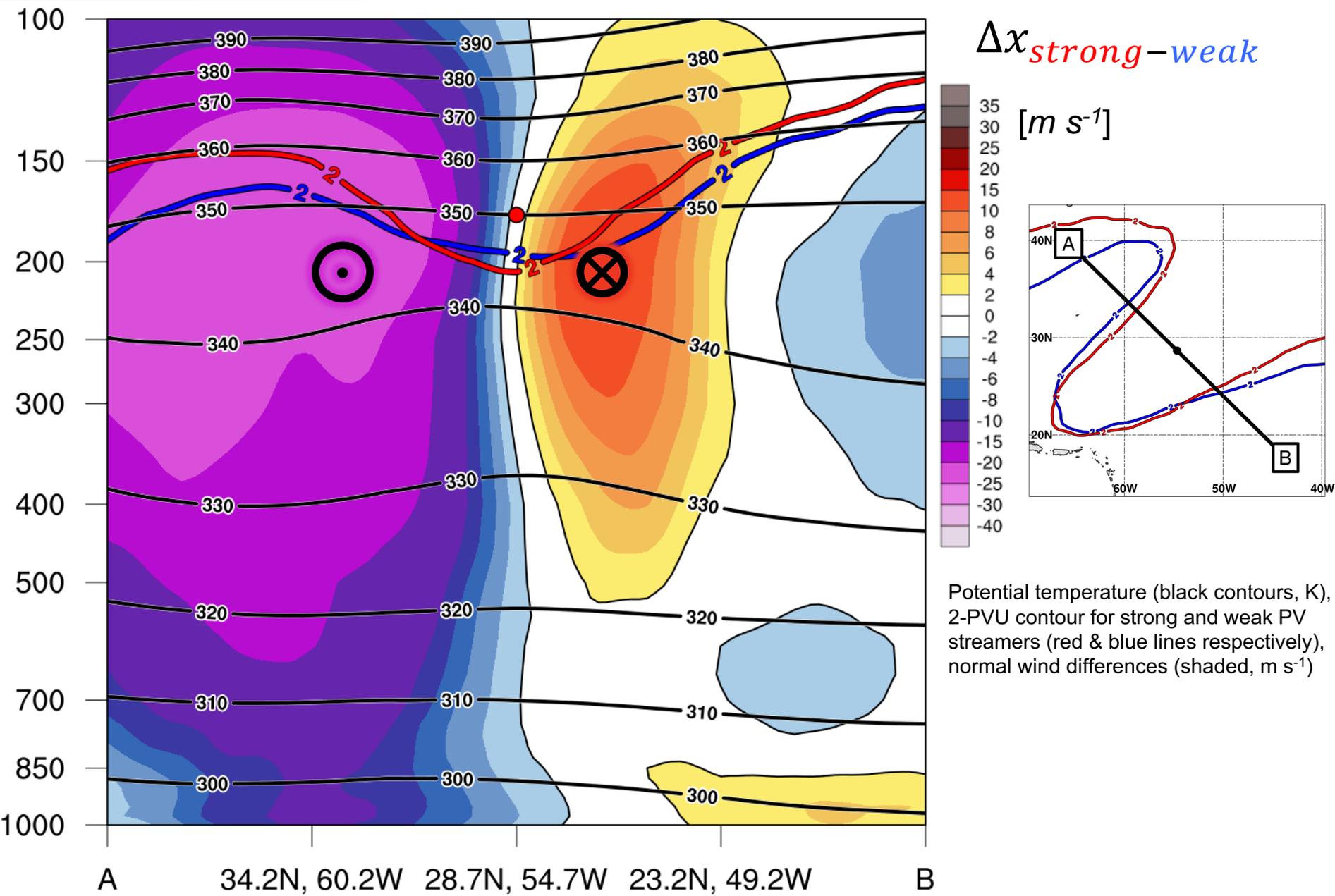


Mean 350-K 2-PVU contour of strong and weak PV streamers (red and blue lines respectively)

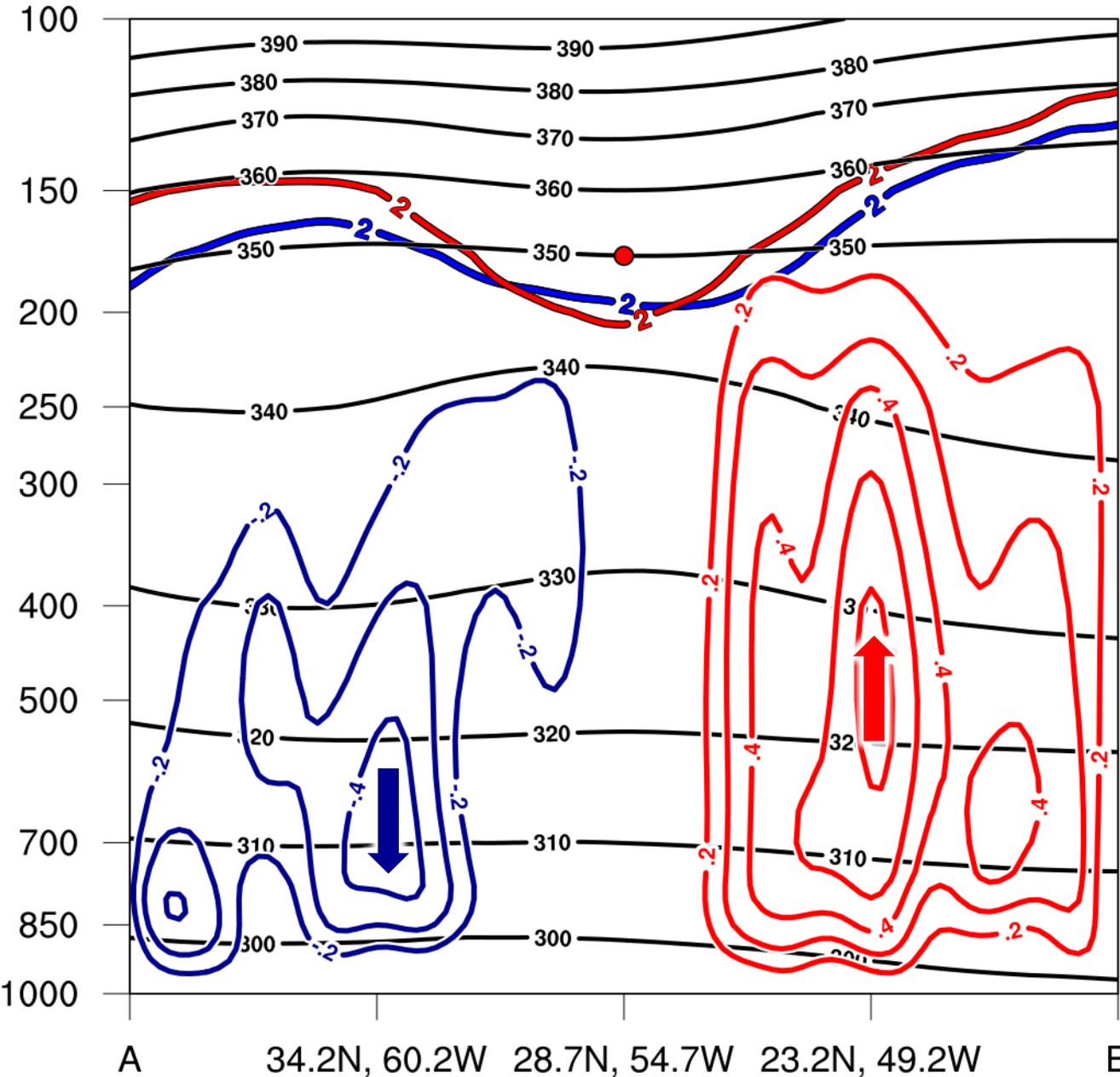
Composite Environment Differences: PV



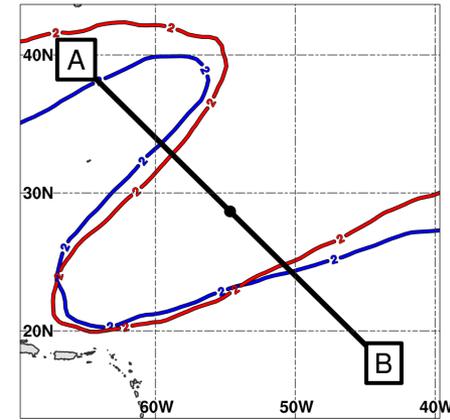
Composite Environment Differences: Wind



Composite Environment Differences: W

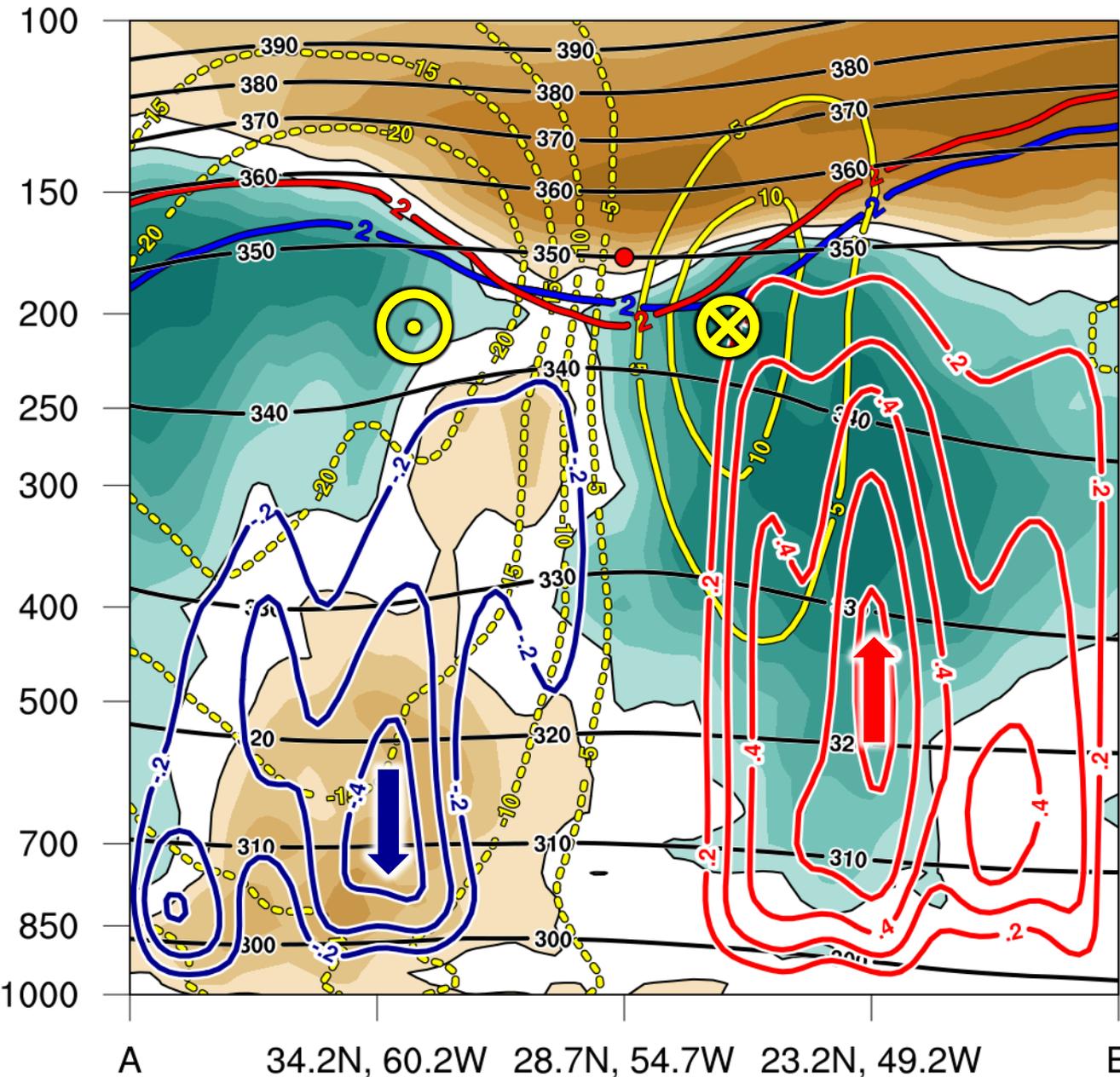


$\Delta x_{strong-weak}$



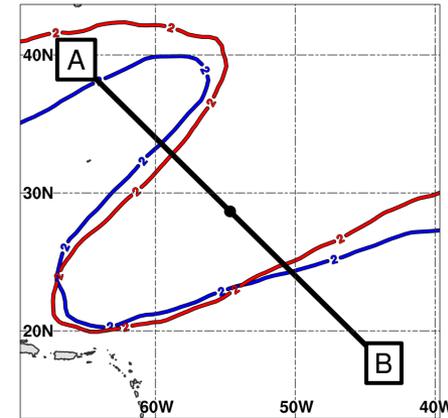
Potential temperature (black contours, K),
 2-PVU contour for strong and weak PV
 streamers (red & blue lines respectively),
 upward and downward vertical motion
 differences (red & blue contours, 10^{-3} hPa s $^{-1}$)

Composite Environment Differences: RH



$\Delta x_{\text{strong-weak}}$

[%]



Potential temperature (black contours, K),
 2-PVU contour for strong and weak PV
 streamers (red & blue lines respectively),
 upward and downward vertical motion
 differences (red & blue contours,
 $10^{-3} \text{ hPa s}^{-1}$), relative humidity
 anomaly differences (shaded, %), normal wind
 differences (yellow contours, m s^{-1})

Concluding Summary

Concluding Summary: Part 1

- A 1979–2015 climatology of PV streamers on the 350-K surface is created by adapting previous techniques
 - During the TC season (June–November) in Atlantic basin (10-100°W)
- PV streamer climatology in the Atlantic basin
 - Highest frequency equatorward of 200-hPa jet
 - Notable shifts occur both from month to month and year to year
 - PV streamers drive the dynamical portion of the TUTT (i.e., MOT) where climatological westerlies occur
- PV streamer intensity and area compared to TC activity
 - Strong and large PV streamers are correlated with lower TC activity
 - Weak PV streamers are correlated with higher TC activity
 - Most obvious differences shown when comparing top and bottom 8 ACE years
 - PV streamer activity metric (combining amount, size, and intensity) exhibits greatest negative correlation with TC activity.

Concluding Summary: Part 2

- Composite differences of PV streamers
 - An effective comparison needs to compare different intensities that are similar in area and similar in tilt
 - Top and bottom 20 percentile in intensity are compared to 55-75° tilts and 40-60 percentile areas
- Vertical Wind Shear
 - Strong PV streamers have much larger and stronger corridors of shear downstream of their trough axis
- Sea Level Pressure
 - Strong PV streamers have higher SLP in and around their trough axis
- Precipitable Water
 - Strong PV streamers have lower PW in their trough axis, but higher PW upstream of the trough axis in the upstream ridge
- Cross Section
 - Enhanced shear from stronger tropopause based winds of a strong PV anomaly
 - Enhanced anomalies related to stronger vertical motion upstream and downstream of strong PV streamers

Final Thoughts and Future Work

- Seasonal PV streamer activity represents an important extratropical impact on TC activity
 - While not fully independent of other factors (ENSO, SSTs, AEW activity) it may still add predictability when combined with these well known seasonal prediction factors
- Future Work
 - Assessing the seasonal predictability of PV streamer activity
 - Understanding upstream precursors lead to strong vs. weak PV streamers (in progress)
 - Real time PV streamer identification, including assessment of relevant variables
 - Area, Intensity, Tilt
 - May provide additional clues into how these individual events are likely to influence TCs that occur nearby

Questions? ppapin@albany.edu

Extra Slides

Results: Reference Websites for more Figures

Yearly PV Streamers 1979-2015

http://www.atmos.albany.edu/student/ppapin/lb13_img/phd/pvs_year.html

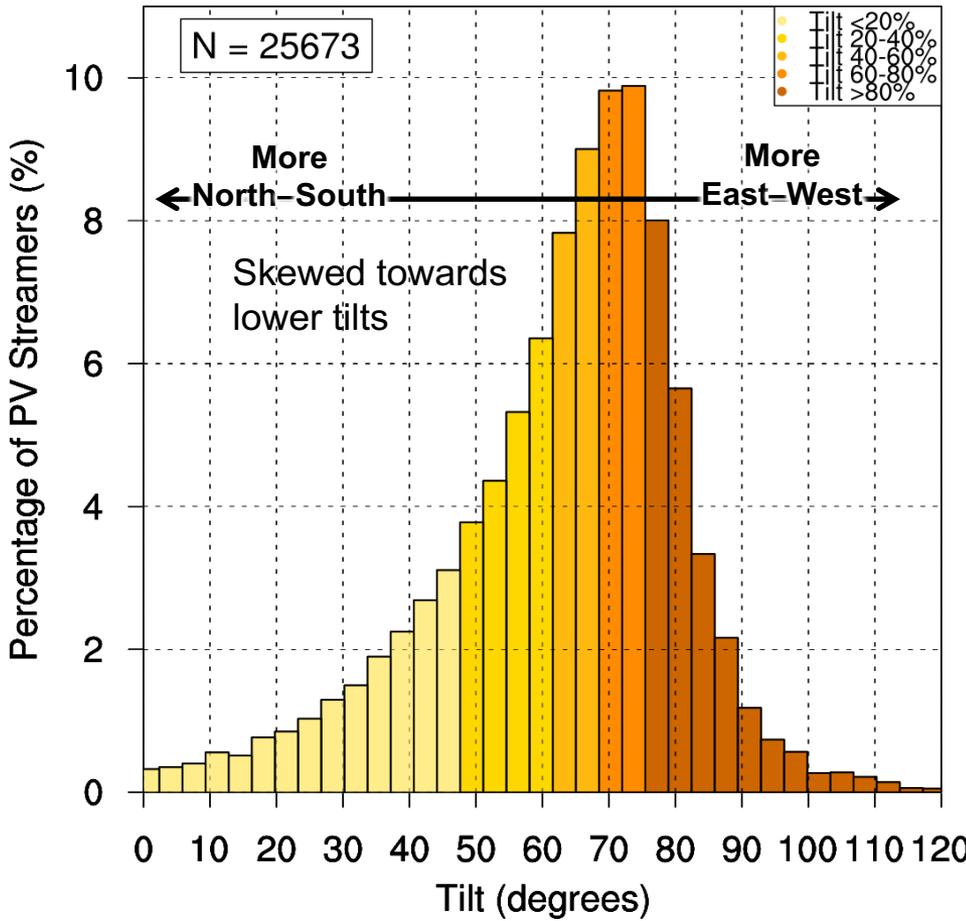
All Composite Results of PV Streamers

http://www.atmos.albany.edu/student/ppapin/lb13_img/phd/pvs_composite.html

Results: Variability of PV Streamer Tilt

- Tilt of PV streamers (degree tilt relative to a meridian)
 - 1 Jun–30 Nov between 100–10°W
 - Mean: 59.5°
 - Stdev: +/- 26.1°

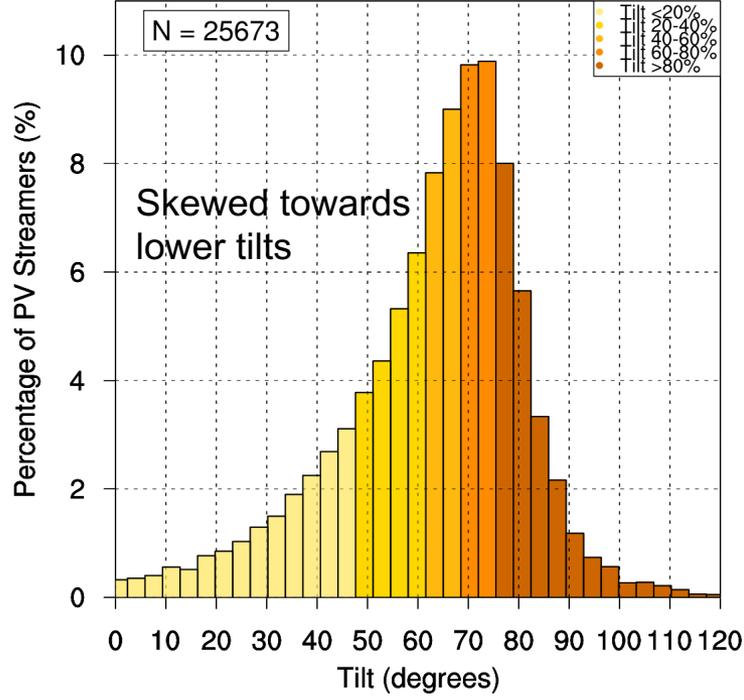
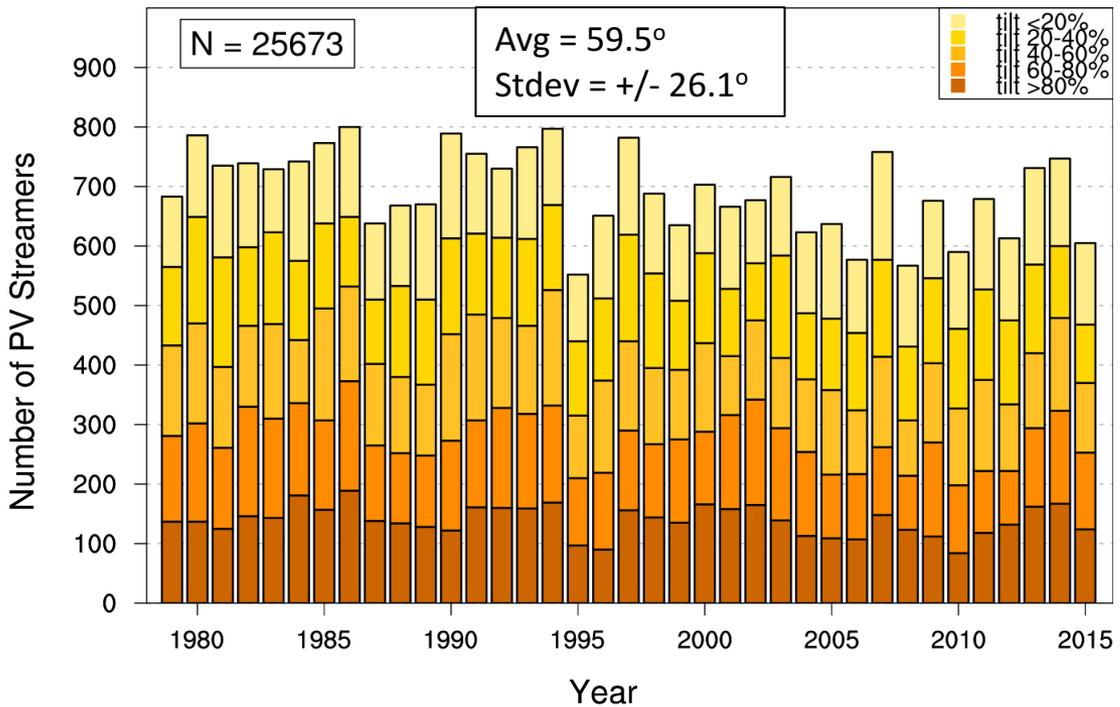
Because of the highly skewed distribution, it is useful to use percentiles when describing tilt values.



PDF of PV Streamer Tilt

Results: Variability of PV Streamer tilt

- Interseasonal variability in PV streamer area
 - 1 Jun–30 Nov between 100–10°W



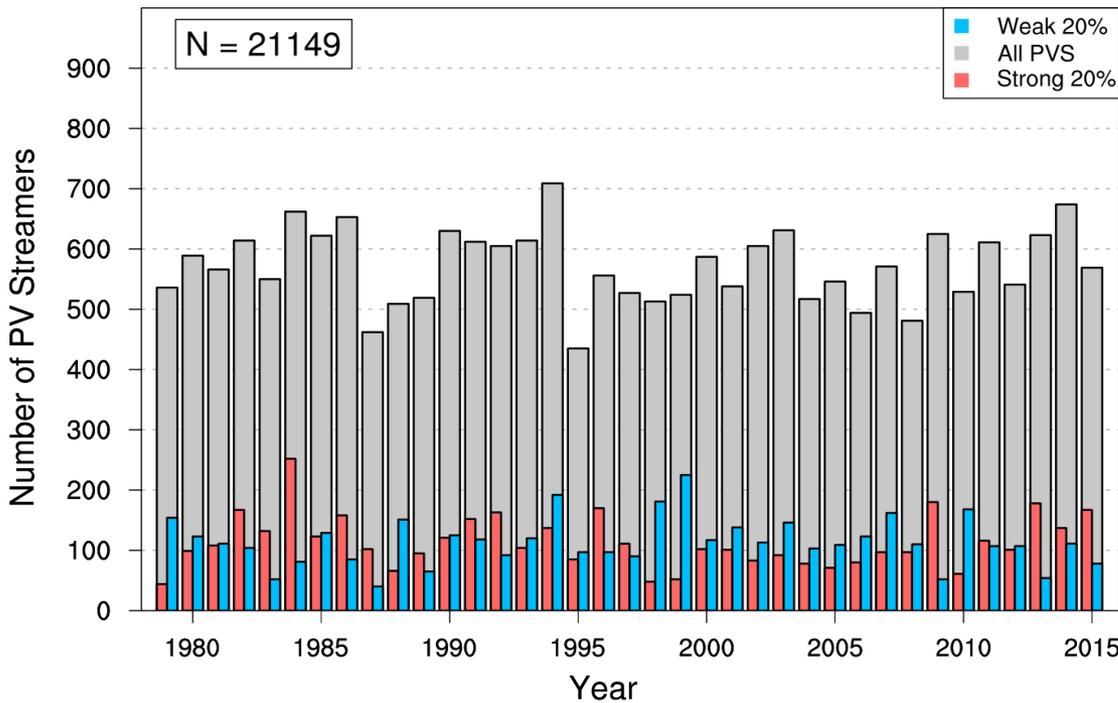
Yearly Distribution of PV Streamer Count by Tilt

PDF of PV Streamer Tilt

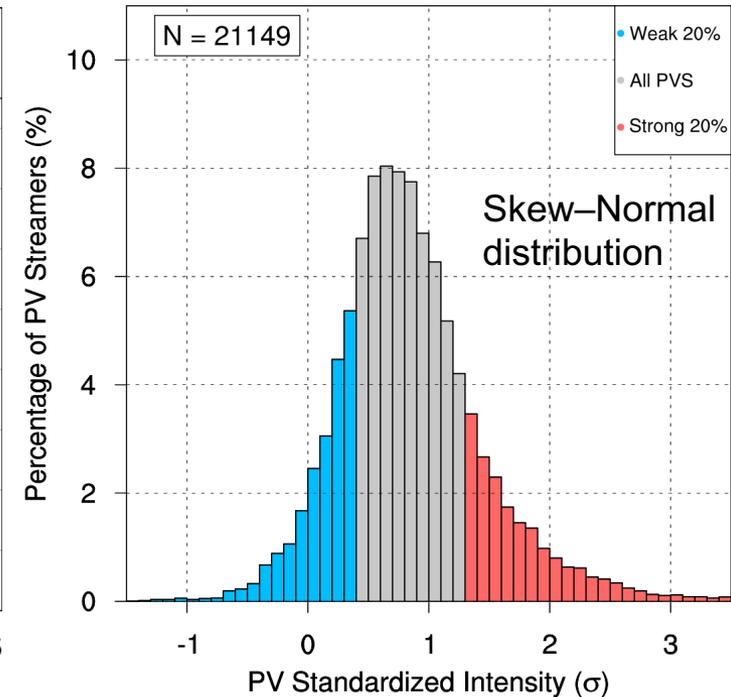
✧ No obvious tilt trends over the CFSR period

ERA Interim: PV Streamer Intensity

- Interseasonal variability in the number of PV streamers
 - 1 Jun–30 Nov between 100–10°W



Yearly Distribution of PV Streamer Count

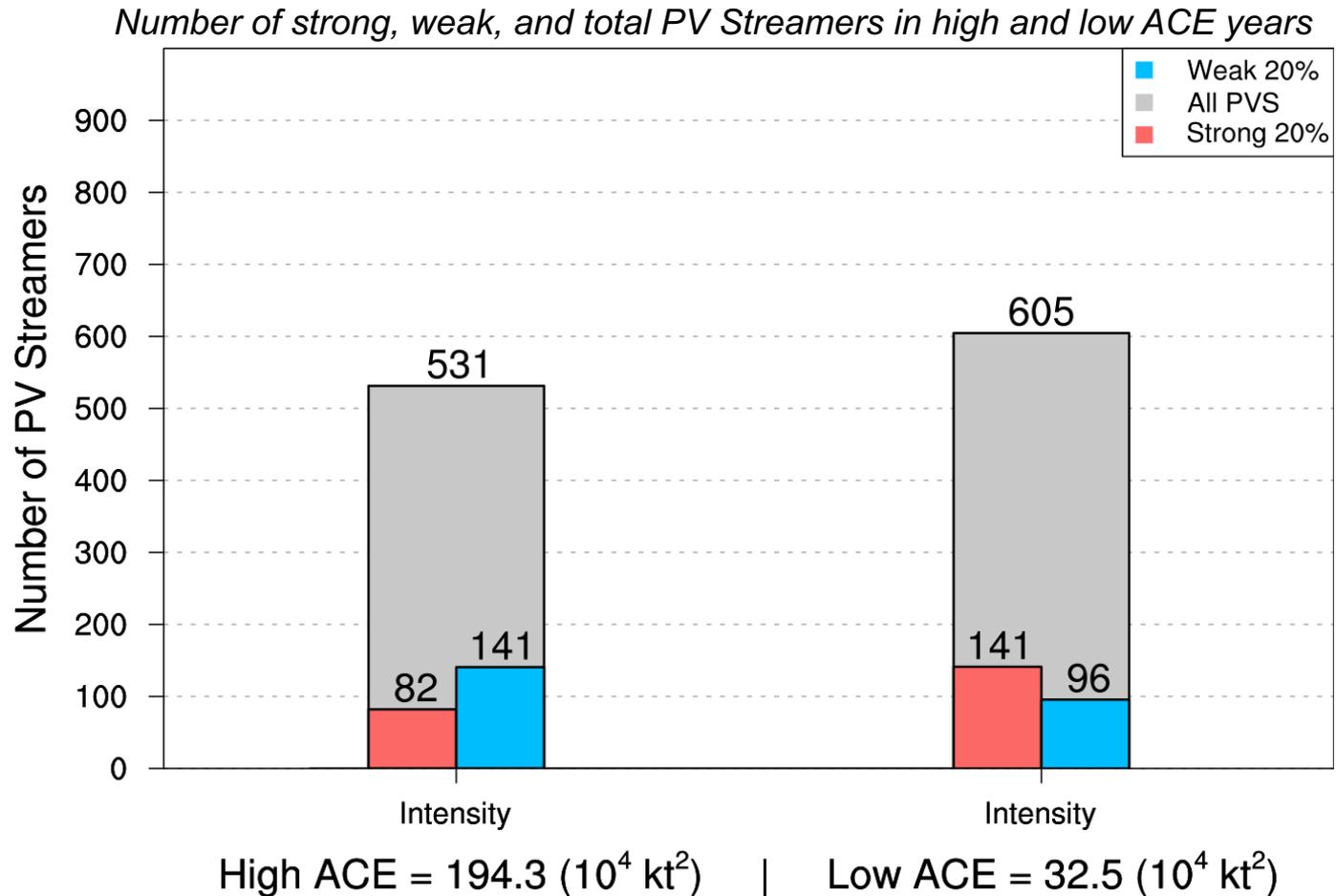


PDF of PV Streamer Intensity

- ✧ ERA-Interim does not display the same long-term intensity trends that were observed in the CFSR.

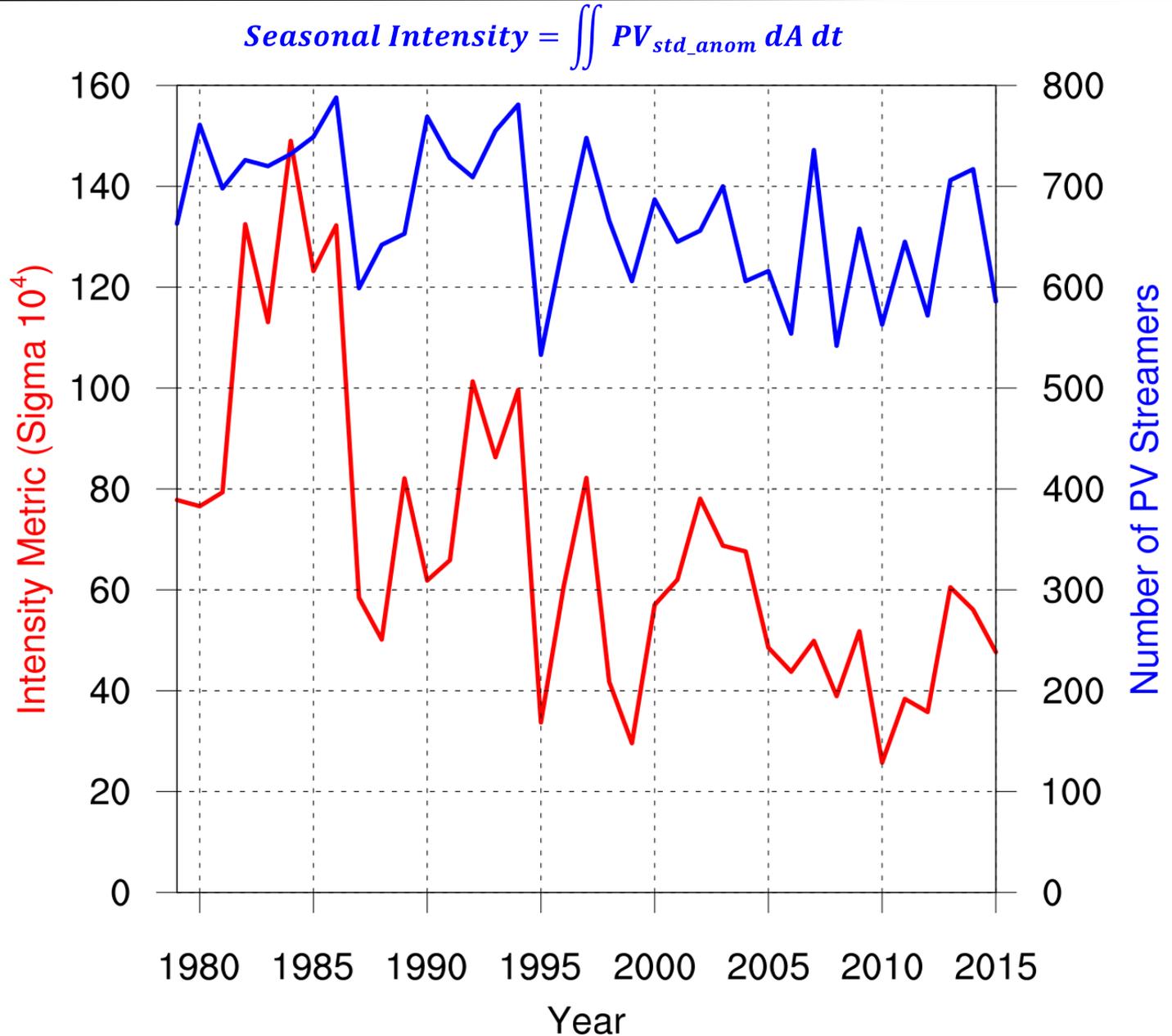
ERA Interim: High ACE vs. Low ACE

Top 8 High ACE Years	Bottom 8 Low ACE Years
2005	1983
1995	1982
2004	1994
1998	1987
1999	2013
2003	1991
1996	1986
2010	1993



Time Series: Intensity Metric vs. PVS Count

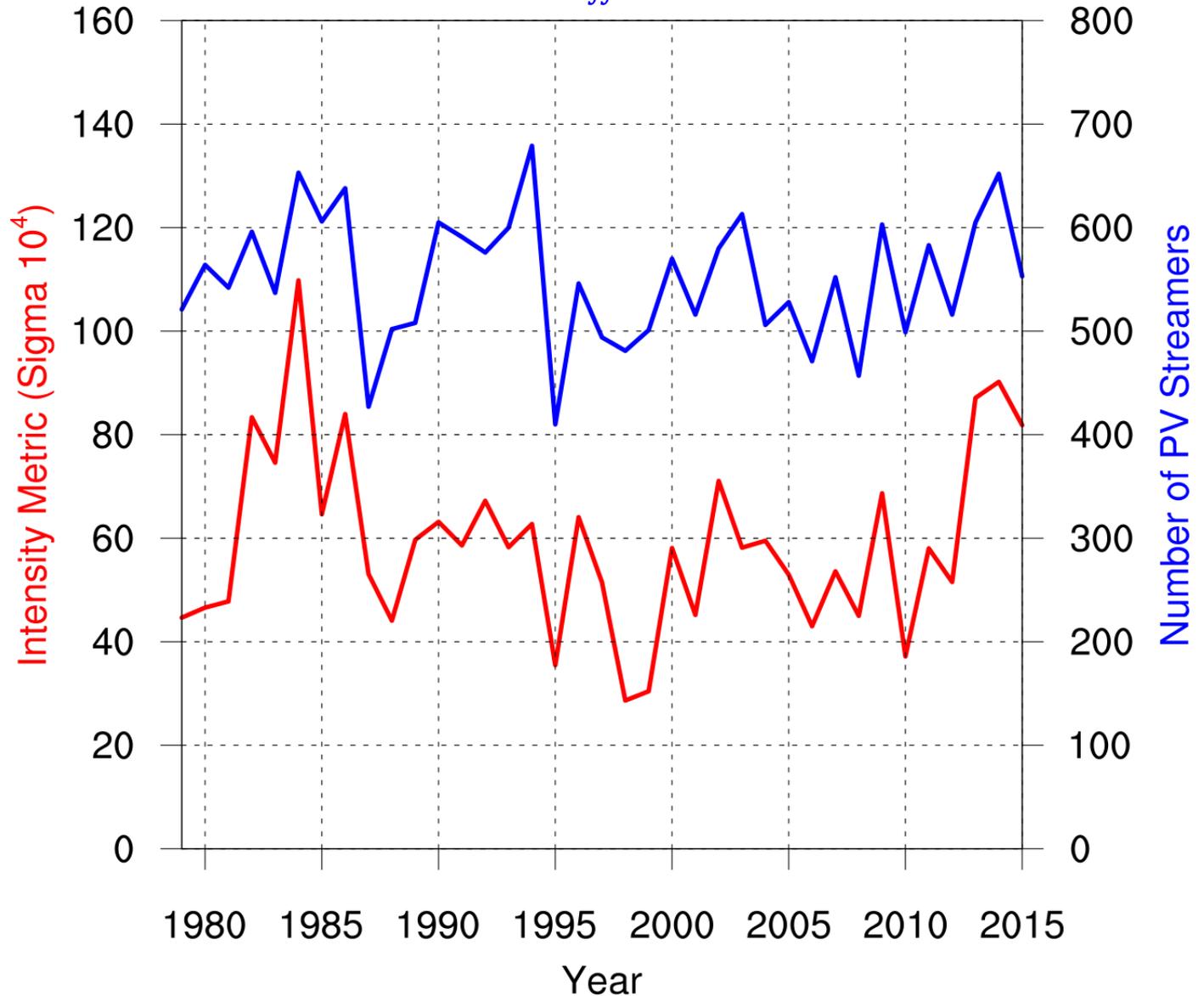
CFSR



Time Series: Intensity Metric vs. PVS Count

ERAI

$$\text{Seasonal Intensity} = \iint PV_{std_anom} dA dt$$

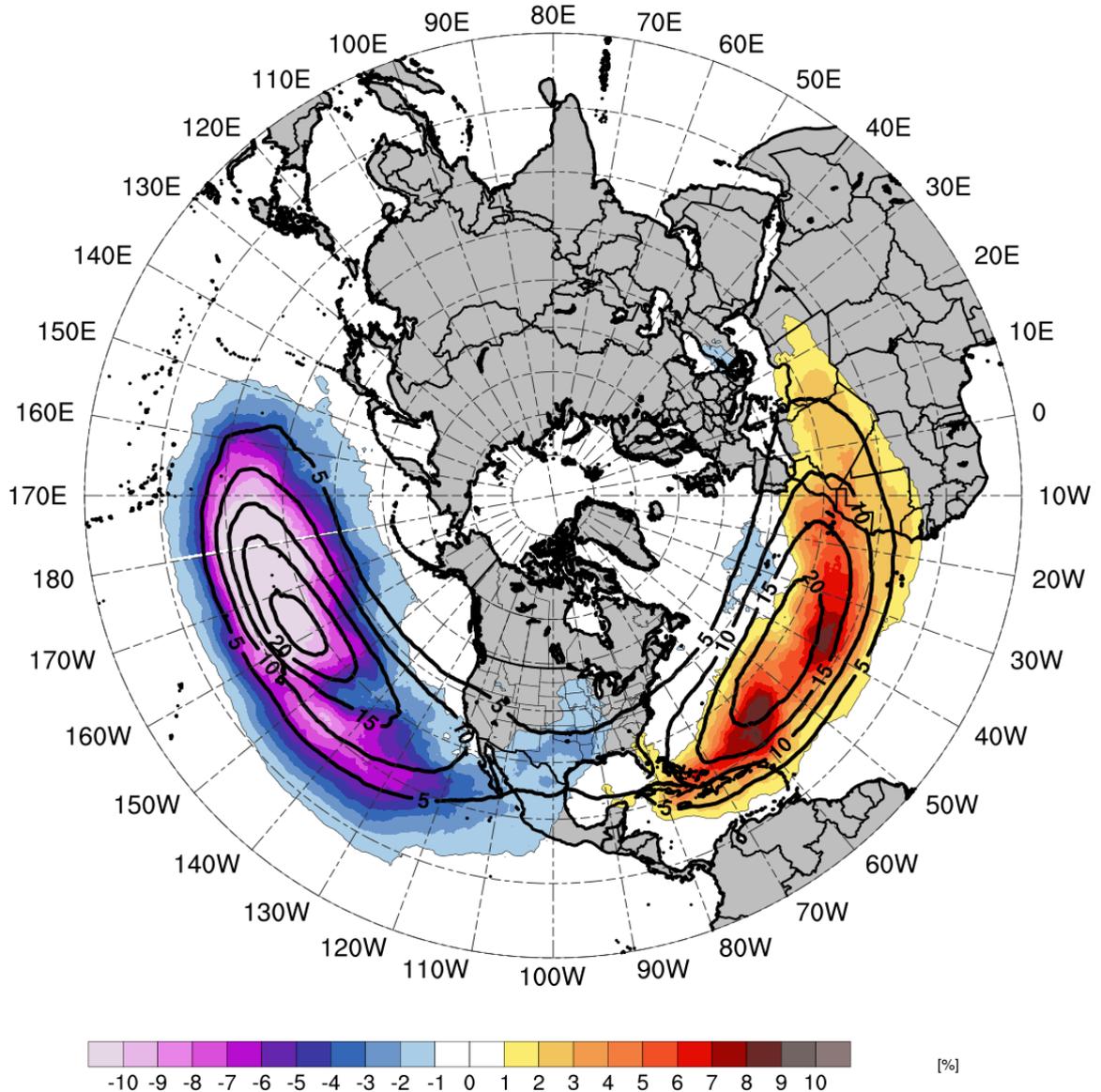


PV Streamer Frequency Anomalies vs. ENSO

CFSR

El Nino – La Nina

**Percentage change in
PV Streamer frequency
relative to climatology**

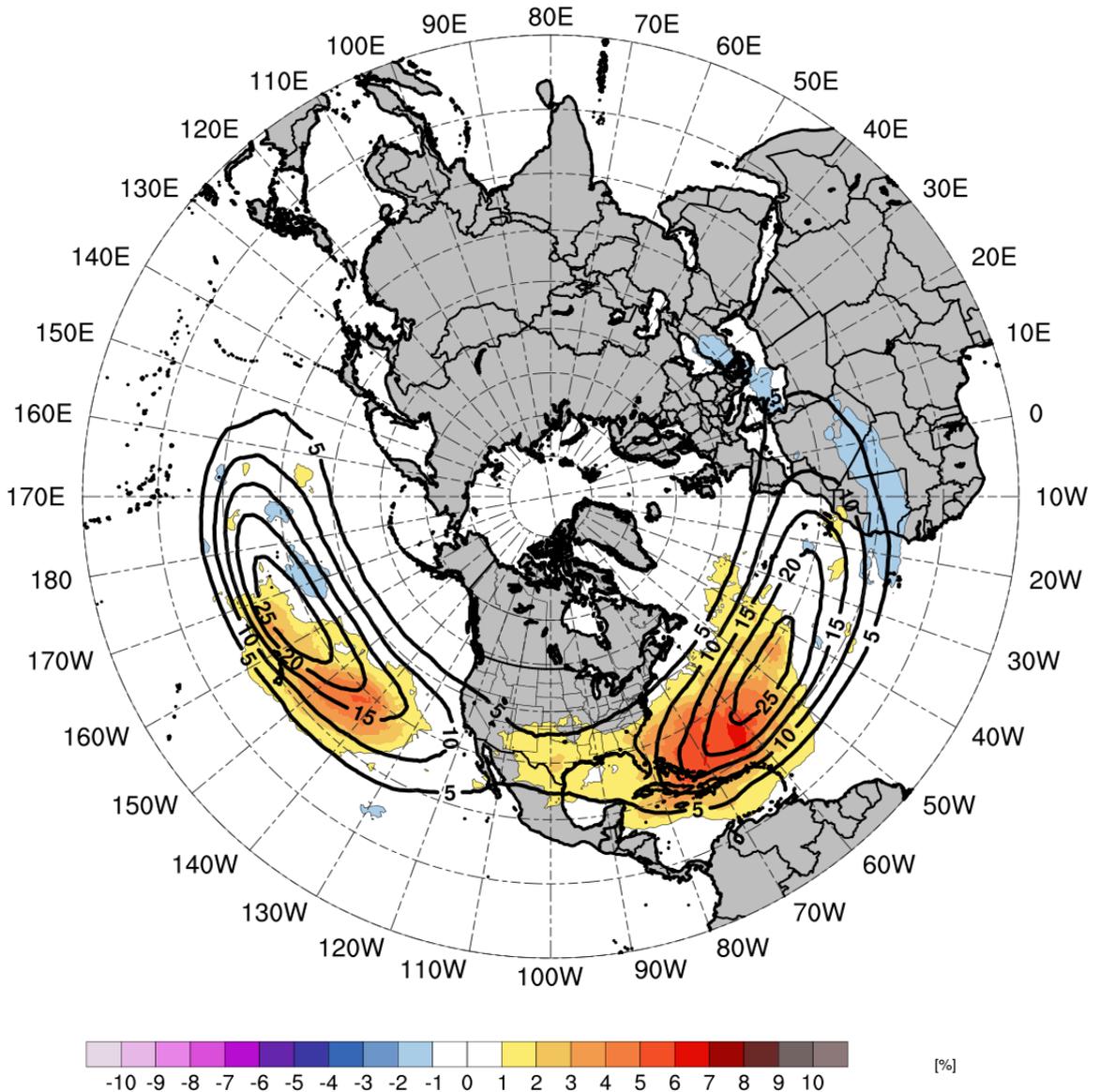


PV Streamer Frequency Anomalies vs. NAO

CFSR

Positive – Negative NAO

**Percentage change in
PV Streamer frequency
relative to climatology**

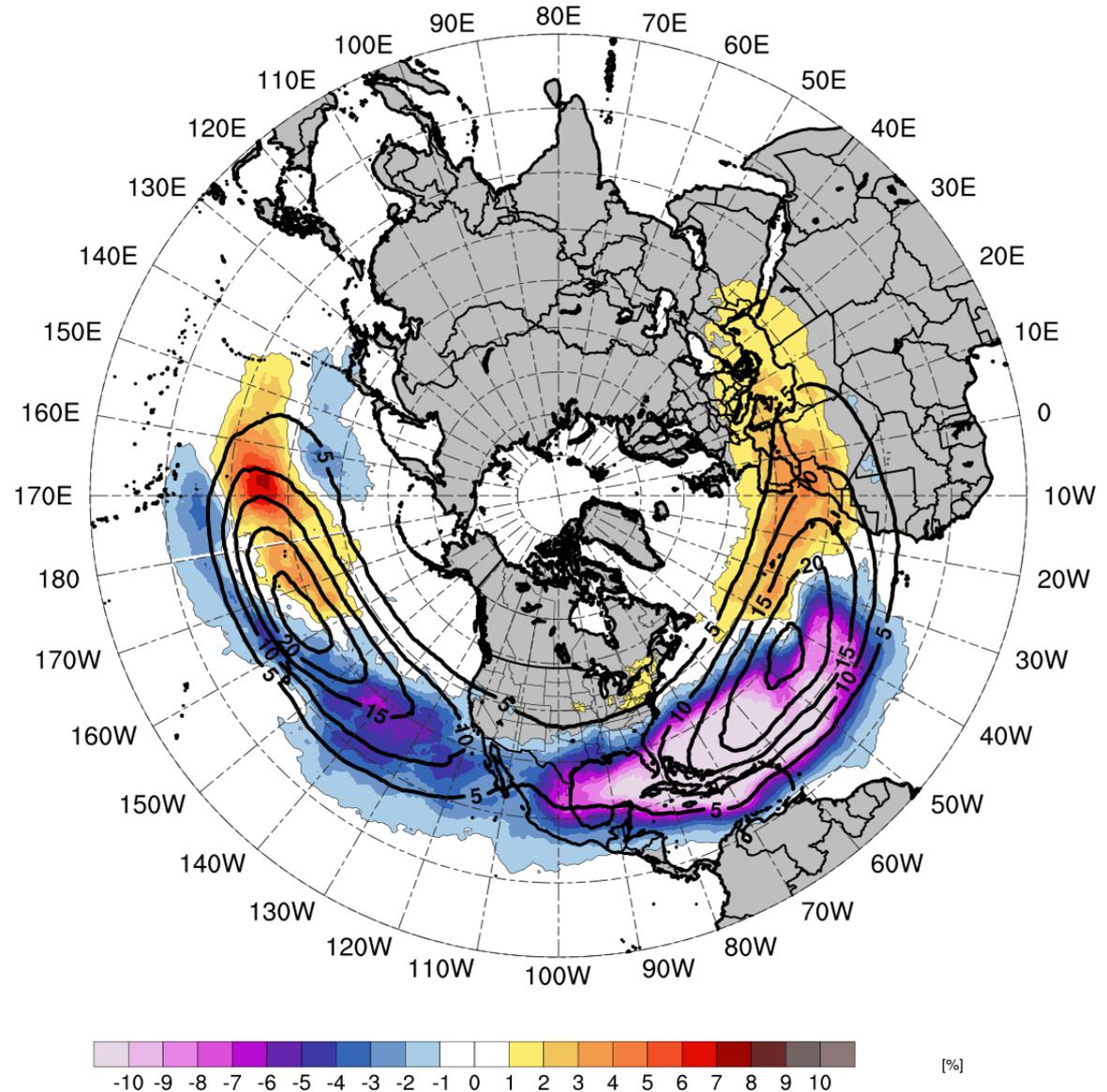


PV Streamer Frequency Anomalies vs. SSTAs

CFSR

Positive – Negative SSTAs

**Percentage change in
PV Streamer frequency
relative to climatology**



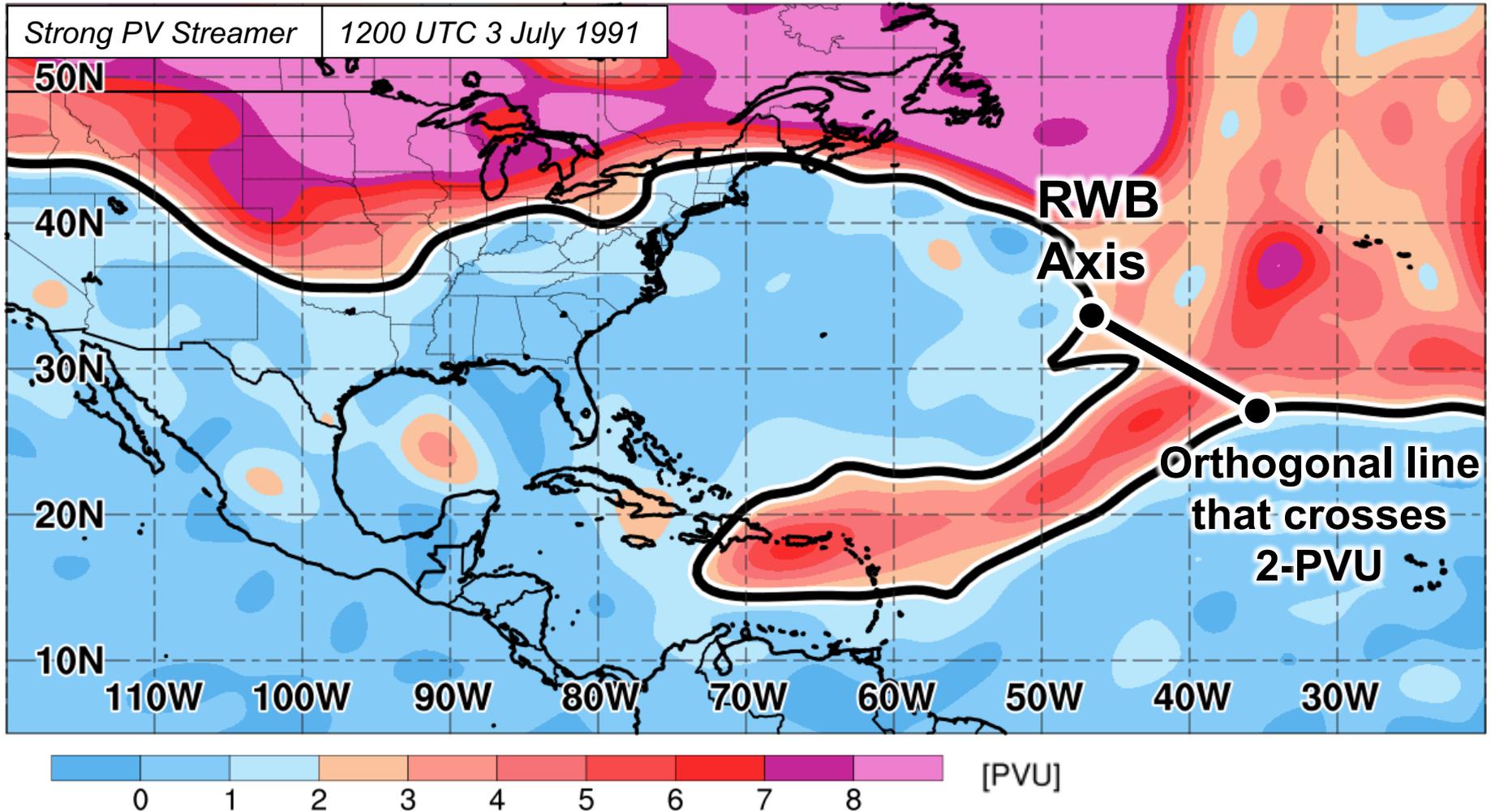
[%]

PV Streamer Identification: Example

Mean and Standard Deviations are derived from a 1979-2009 CFSR climatology

$$\bullet \text{ PV}_{\text{std_anom}} = (\text{PV} - \text{PV}_{\text{mean}}) / \text{PV}_{\text{stdev}}$$

350-K PV (shaded, PVU), and 2-PVU contour (black contour)

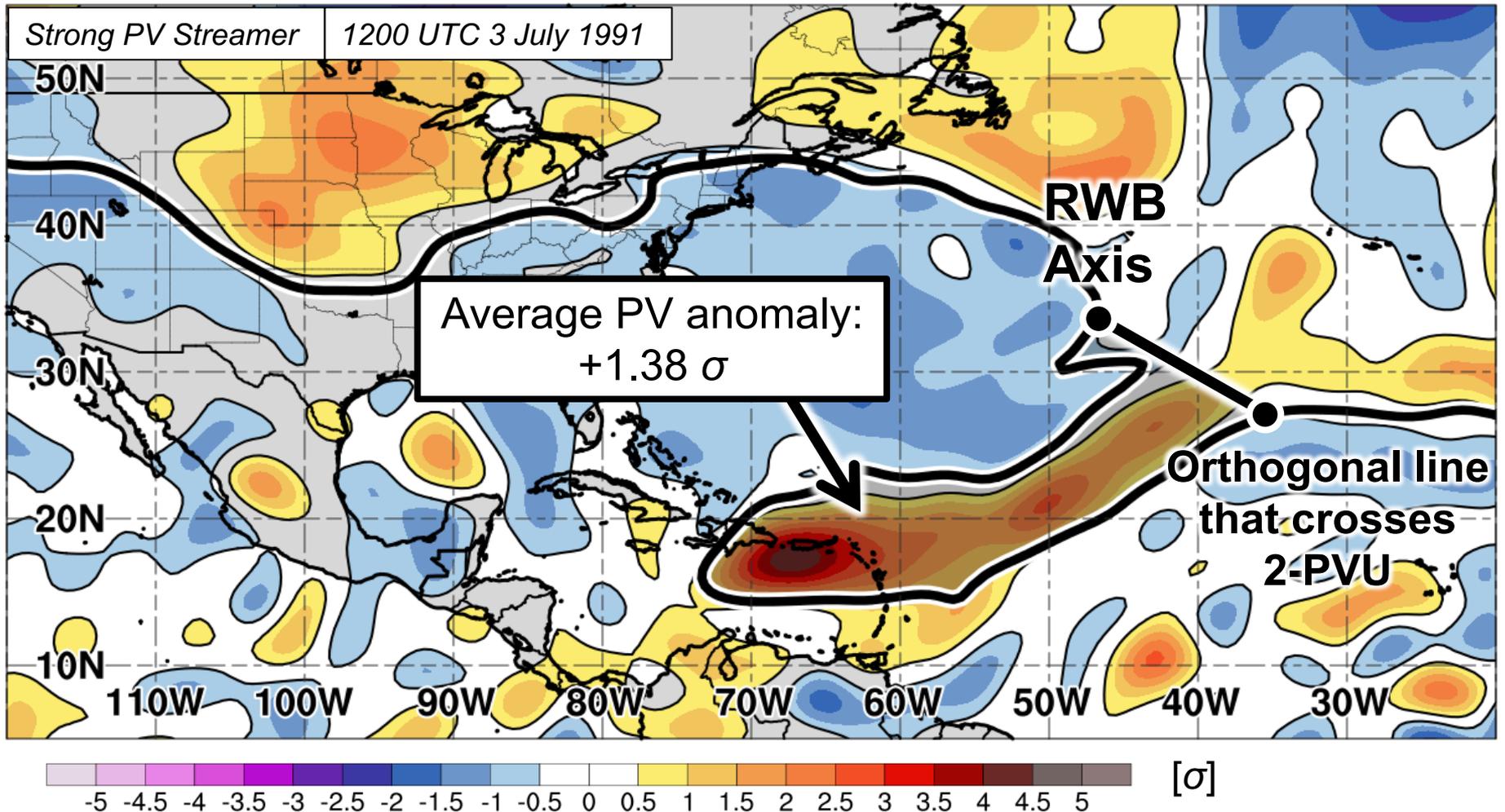


PV Streamer Identification: Example

Mean and Standard Deviations are derived from a 1979-2009 CFSR climatology

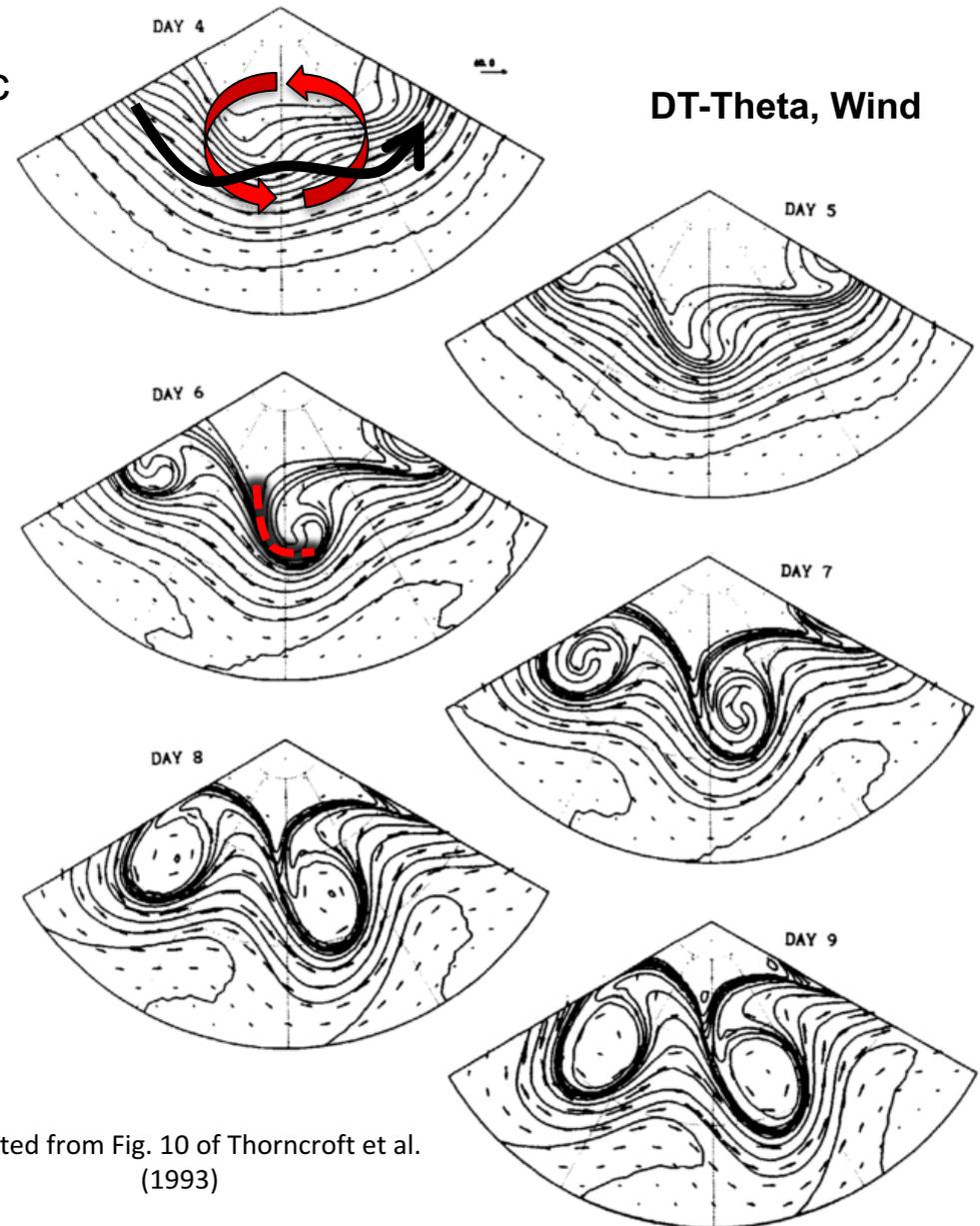
$$\bullet \text{ PV}_{\text{std_anom}} = (\text{PV} - \text{PV}_{\text{mean}}) / \text{PV}_{\text{stdev}}$$

350-K Standardized PV Anomaly (shaded, Sigma), and 2-PVU contour (black contour)



Literature Review: Rossby Wave Breaking

- RWB Manifests as two characteristic baroclinic wave lifecycles
 - ✧ Anticyclonic Wave Breaking (LC1, AWB)
 - ✧ Cyclonic Wave Breaking (LC2, CWB)
- Anticyclonic meridional shear found equatorward of the waveguide
 - ✧ Thin positively tilted PV streamer
- Cyclonic meridional shear found poleward of the waveguide
 - ✧ Thick negatively tilted PV streamer



Adapted from Fig. 10 of Thorncroft et al.
(1993)

PV Streamer Identification

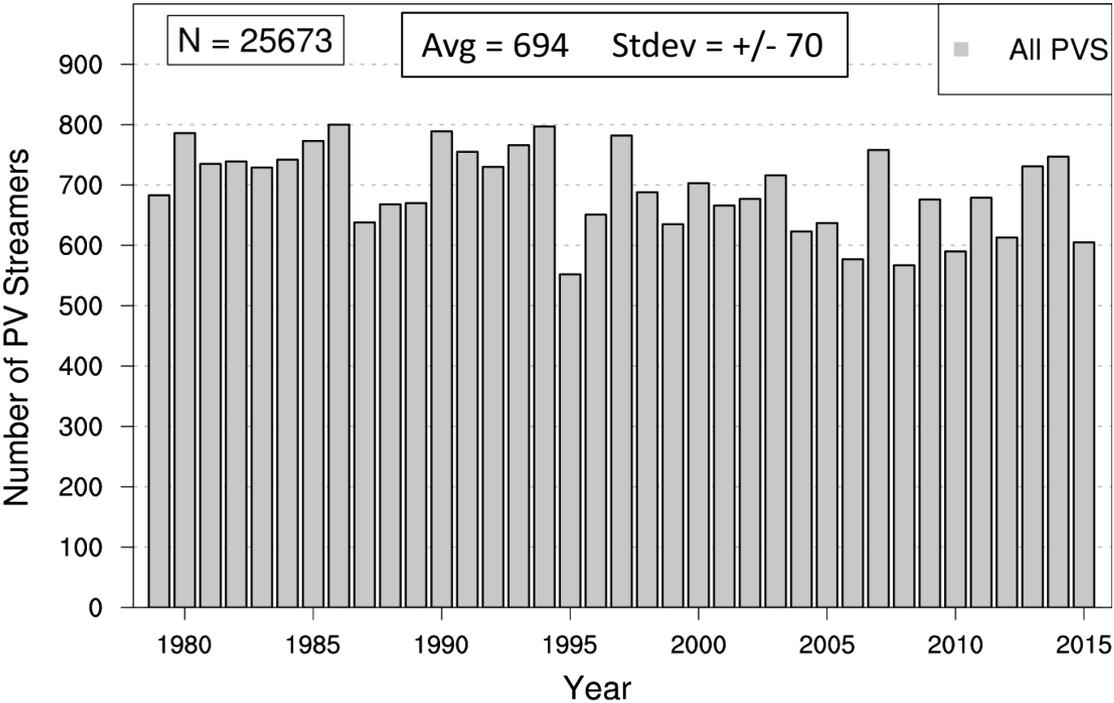
- PV streamers are identified from June–November 1979–2015 using the 0.5° NCEP CFSR (Saha et al. 2010).
- A new PV streamer algorithm is created that combines previous methodologies
 - Postel and Hitchman (1999)
 - ✧ Identifies locations where RWB occurs (meridional gradient reversal in PV)
 - Wernli and Sprenger (2007)
 - ✧ Identifies elongated filaments of high PV air using width and perimeter of PV streamer
- Identification of PV streamers occurs on a isentropic surface that approximates the location of the subtropical tropopause
 - 350-K surface using the 2-PVU contour as the dynamical tropopause
 - PV streamer intensity is calculated as a standardized PV anomaly of the area encompassed by the PV streamer

$$PV_{\text{std_anom}} = (PV - PV_{\text{mean}}) / PV_{\text{stdev}}$$

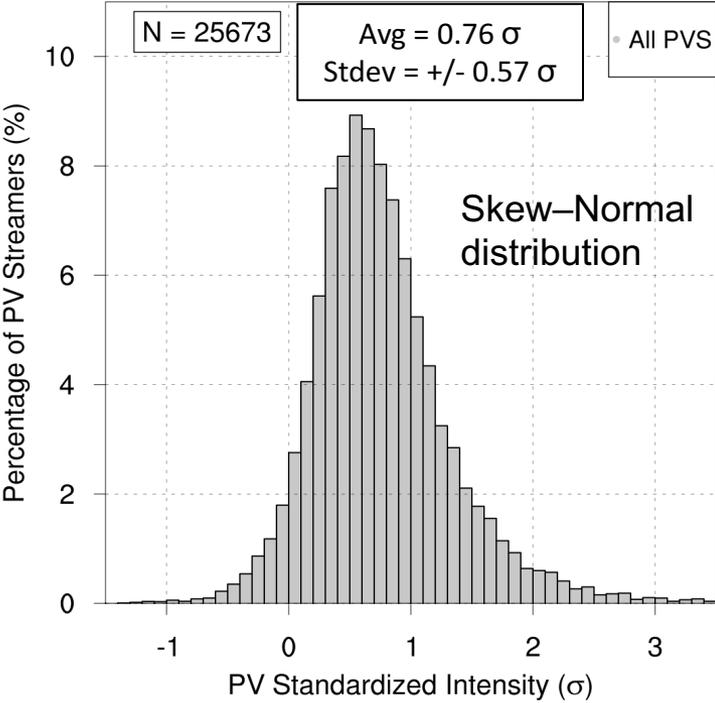
Mean and Standard Deviations are
derived from a 1979-2009 CFSR
climatology

Results: PV Streamer Variability 1979–2015

- Interseasonal variability in the number of PV streamers
 - 1 Jun–30 Nov between 100–10°W



Yearly Distribution of PV Streamer Count



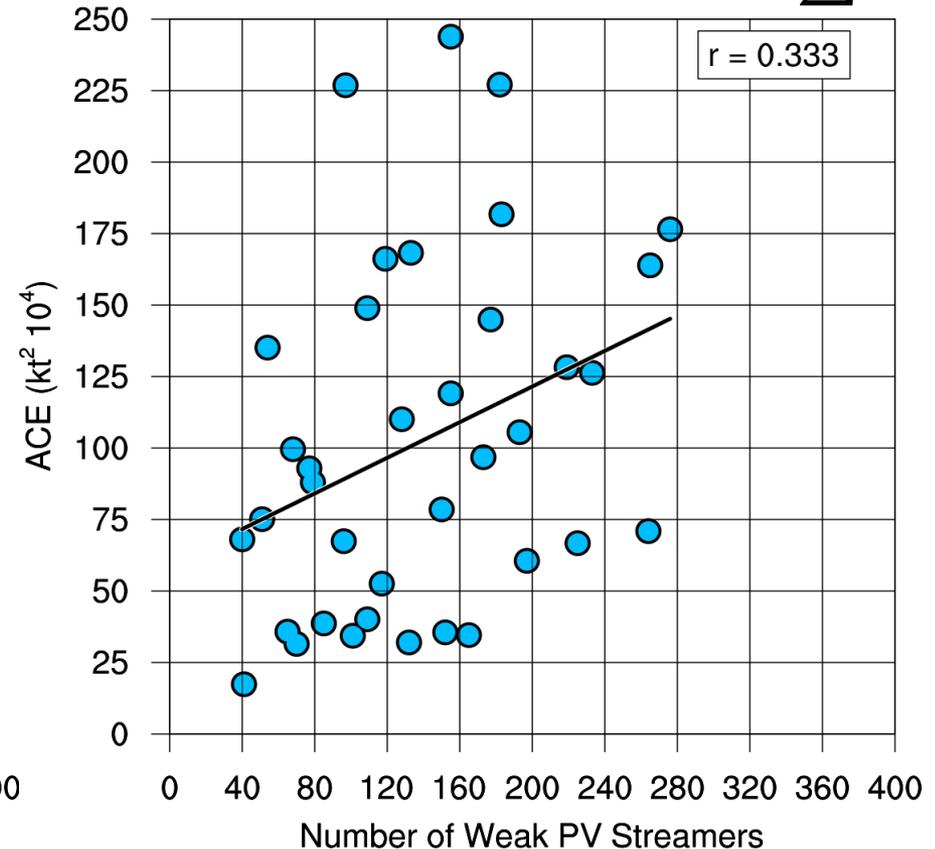
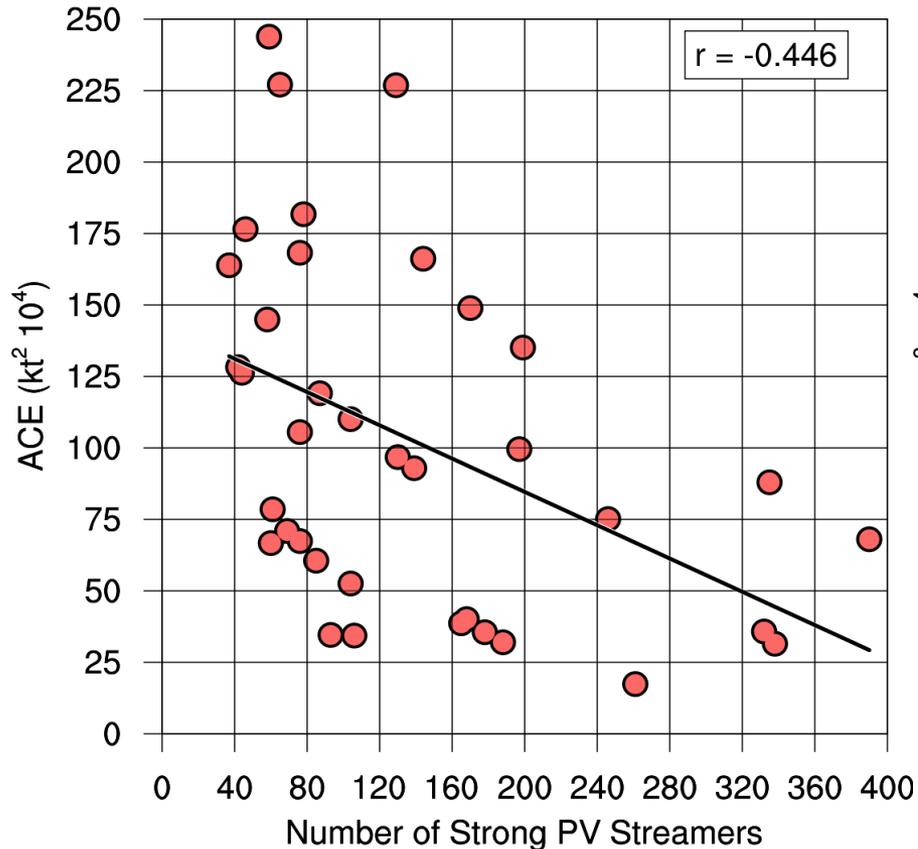
PDF of PV Streamer Intensity

- ✧ CFSR depicts a trend towards a decrease in the number of strong PV streamers and increase in the number of weak PV streamers with time
 - ✧ Questions remain if this is a real trend or is dataset dependent

Climatological Results: PV Streamer Intensity

- How TC activity related to differences in PV streamer intensity

➤ Seasonal TC activity measured by accumulated cyclone energy (ACE) $ACE = 10^{-4} \sum v_{max}^2$



- ✧ Strong PV streamers are negatively correlated with ACE
- ✧ Weak PV streamers are positively correlated with ACE

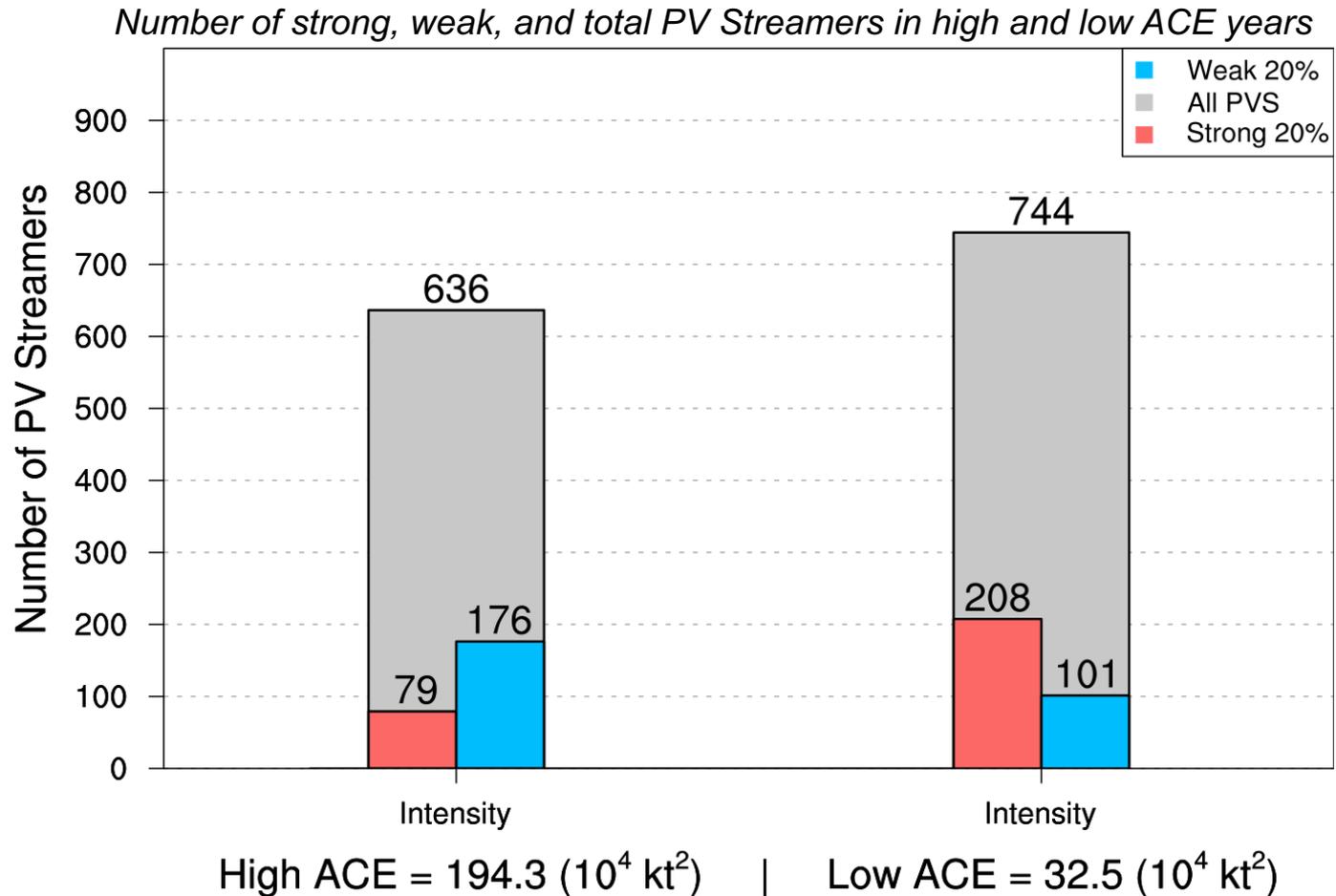
To further illustrate TC activity relationship to PV streamer intensity, let's compare the most and least active TC seasons (using ACE).

Results: TC Activity – High vs. Low ACE

- Active TC seasons (High ACE) feature:

- ✧ Fewer PV streamers (-14.5%)
- ✧ Much Fewer Strong PV streamers (-62.0%)
- ✧ Many More Weak PV streamers (+74.2%)

Top 8 High ACE Years	Bottom 8 Low ACE Years
2005	1983
1995	1982
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Climatological Results: High ACE vs. Low ACE

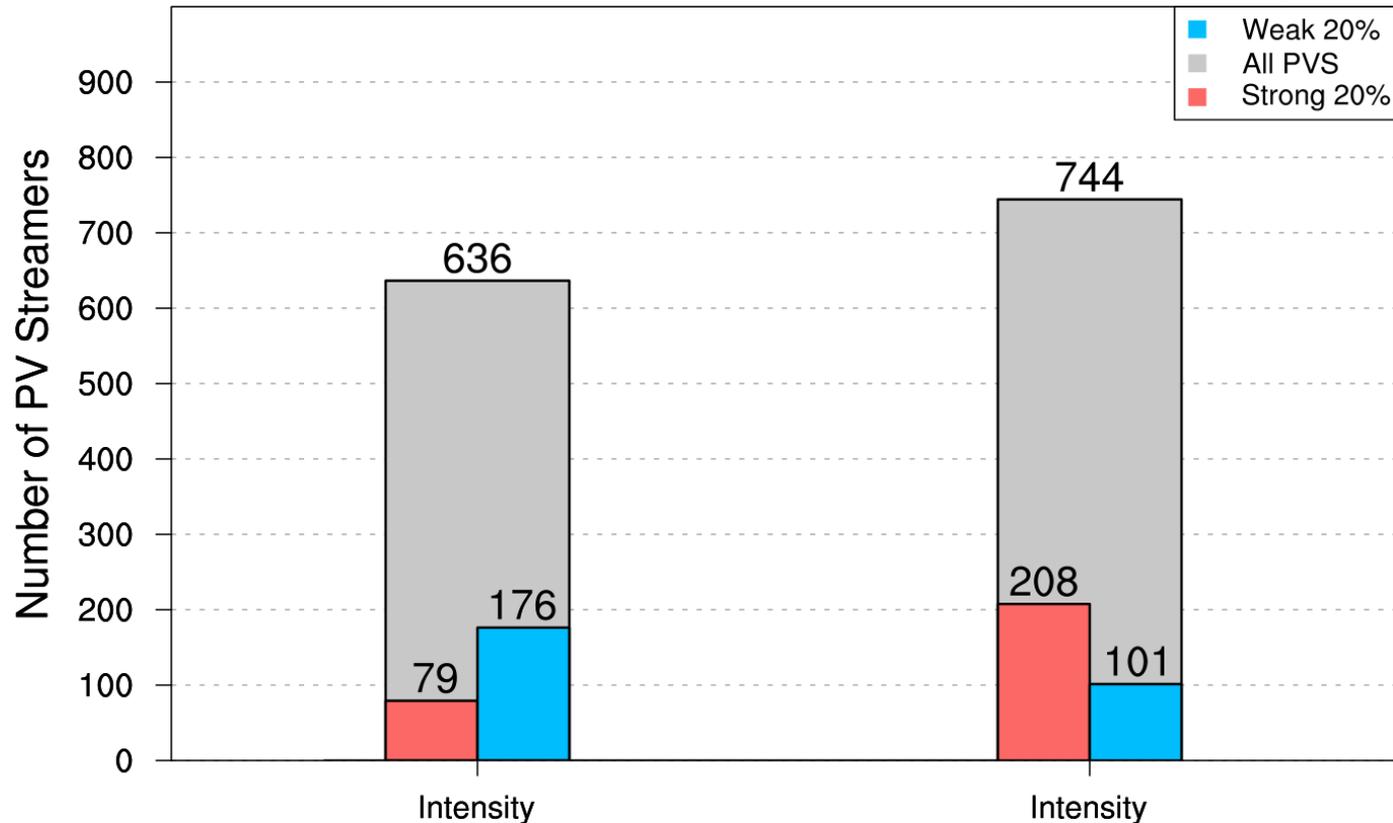
- Active TC seasons (High ACE) feature:

- ✧ Fewer PV streamers (-14.5%)
- ✧ Much Fewer Strong PV streamers (-62.0%)
- ✧ Many More Weak PV streamers (+74.2%)

Why are weak PV streamers more common in high ACE years while strong PV streamers are more common in low ACE years?

Top 8 High ACE Years	Bottom 8 Low ACE Years
2005	1983
1995	1982
2004	1994
1998	1987
1999	2013
2003	1991
1996	1986
2010	1993

Number of strong, weak, and total PV Streamers in high and low ACE years



High ACE = 194.3 (10^4 kt²) | Low ACE = 32.5 (10^4 kt²)

Questions? ppapin@albany.edu

- A 1979–2015 climatology of PV streamers on the 350-K surface is created by adapting previous techniques
 - During the TC season (June–November) in Atlantic basin (10-100°W)
- PV streamer climatology in the Atlantic basin
 - Highest frequency equatorward of 200-hPa jet.
 - Contribute to the formation of the time-mean Mid-Ocean trough (i.e., TUTT)
- PV streamer intensity and area compared to TC activity
 - Strong PV streamers are correlated with lower TC activity
 - Weak PV streamers are correlated with higher TC activity
- Composite Differences of strong minus weak PV streamers:
 - Strong PV streamers exhibit larger and more intense VWS corridors
 - Strong PV streamers exhibit drier air upstream of the trough
 - Linked to stronger upstream subsidence and northeasterly flow