A planetary-scale, 30–90-day oscillation in zonal wind, surface pressure, and precipitation that propagates slowly eastward.

<u>Wavelength</u> = 12,000–20,000 km <u>Speed</u> =  $\sim 5 \text{ m s}^{-1}$ 



**Figure 3.** Time-space power spectra of (a) 850 hPa zonal wind (NCEP/NCAR reanalysis) and (b) precipitation [*Xie and Arkin*, 1997] for 1979 through 1998, averaged over  $20^{\circ}N-20^{\circ}S$  and  $60^{\circ}-180^{\circ}E$ . Positive (negative) periods correspond to eastward (westward) propagating power. Data resolutions for the spectra are pentad in time and  $10^{\circ}$  in longitude.

#### Zhang (2005)

# The dominant mode of *intraseasonal* convective variability in the tropics.



Schreck et al. (2013)

# The dominant mode of *intraseasonal* convective variability in the tropics.

#### The convective signal of the MJO is strongest over the Indian and Pacific oceans.

The MJO is strongest in the Southern Hemisphere summer.

The convective signature is largest 10–15° north/south of the Equator.



# with the MJO (above), ER waves (upper left), and Kelvin waves (lower left)

Schreck et al. (2013)

# The dominant mode of *intraseasonal* convective variability in the tropics.



#### How was the MJO discovered?

In the late 1960s, new computer power at NCAR allowed scientists to look for patterns in weather observations.

They could only look for patterns in time (at a given location), and not in space (at a given time).











1) **Surface pressure** oscillates with a period of 40–50 days.



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3) The signal was limited to the deep tropics.

4) There was little signal in the meridional winds and in the zonal winds in the mid-troposphere.

#### How did they interpret their findings?



1) The MJO is a region of low-level convergence and, thus, convection.

2) **Pressure** is low in the regions of strong convection.

3) The convection only propagates eastward.

4) The circulation associated with the MJO circumnavigates the globe in ~40 days, but the convection weakens as it moves into the Western Hemisphere.

#### Schematic of MJO structure



As the convection approaches, easterly (westerly) trade winds are enhanced at the Equator at low levels (aloft).

This enhanced flow creates counter-rotating vortices to the north and south = shear vorticity!

**Anticyclones** (a) low levels **Cyclones** (a) upper levels

#### Schematic of MJO structure



Associated with, and behind the convection, are strong westerly (easterly) winds at low levels (aloft).

**Twin cyclones** (a) low levels **Anticyclones** (a) upper levels

This structure is reminiscent of a *Rossby wave* to the west of the convection, and of a *Kelvin wave* to the east of the convection.





**Zhang** (2005)





OLR anomalies: 7.5°S - 2.5°S



Wave Type	Color	Propagation Direction	Wavenumber	Period (days)
Madden–Julian Oscillation (MJO)	Blue	Eastward	0–9	30–96
Kelvin Waves	Green	Eastward	1–14	2.5–17
Equatorial Rossby (ER) Waves	Black	Westward	1–10	9–72



OLR anomalies: 7.5°S - 2.5°S 30-Nov-2012 to 22-Feb-2013 + 21-day Fourier Projection 30-Nov 21-Dec 11-Jan 1-Feb 22-Feb Fourier Projection 15-Mar 60E 120E 180 120W 60W 0 0 Obs: W/m^2 -84 -72 -60 -48 -36 -24 36 48 60 72 84 Sum of Waves: W/m^2 -18 -12 -6 0 6 12 18

MJO (blue, CINT=12); ER (black, CINT=12); Kelvin (green, CINT=12)









FIG. 2. Anomalous OLR and circulation from ERA-15 reanalysis on day 0 associated with a  $-40 \text{ W} \text{ m}^{-2}$  perturbation in MJO-filtered OLR at the equator, 155°E for the period 1979–93, all seasons included; (a) 850 and (b) 200 hPa. Dark (light) shading denotes OLR anomalies less than  $-32 \text{ W} \text{ m}^{-2}$  ( $-16 \text{ W} \text{ m}^{-2}$ ). Streamfunction contour interval is (a)  $4 \times 10^5 \text{ m}^2 \text{ s}^{-1}$  and (b)  $10 \times 10^5 \text{ m}^2 \text{ s}^{-1}$ . Locally statistically significant wind vectors at the 95% level are shown. The largest vectors are about 2 m s<sup>-1</sup> in (a) and around 5 m s<sup>-1</sup> in (b).

Kiladis et al. (2005)

Zonal Wind along Equator



Kiladis et al. (2005)