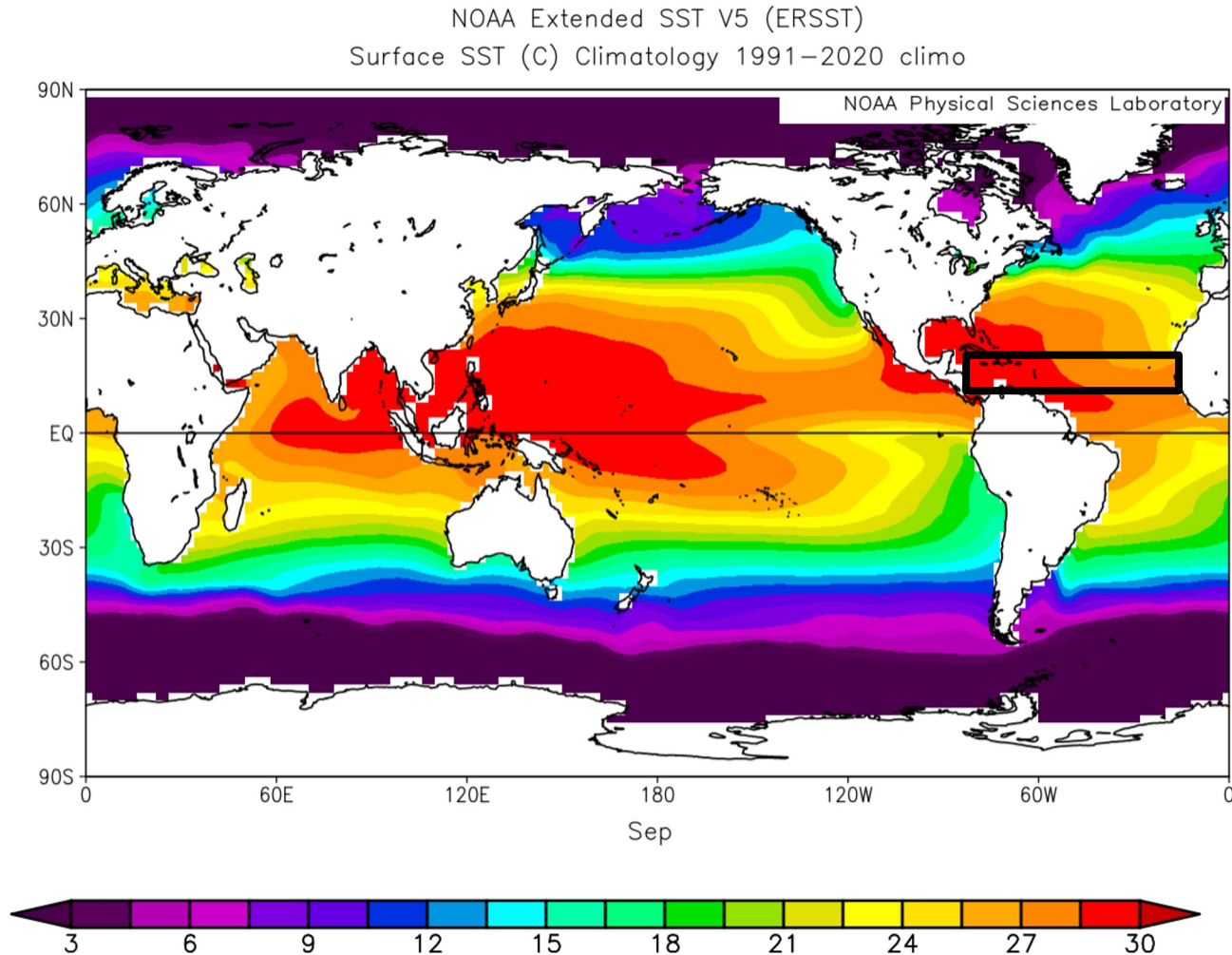


# Necessary conditions for TC development:

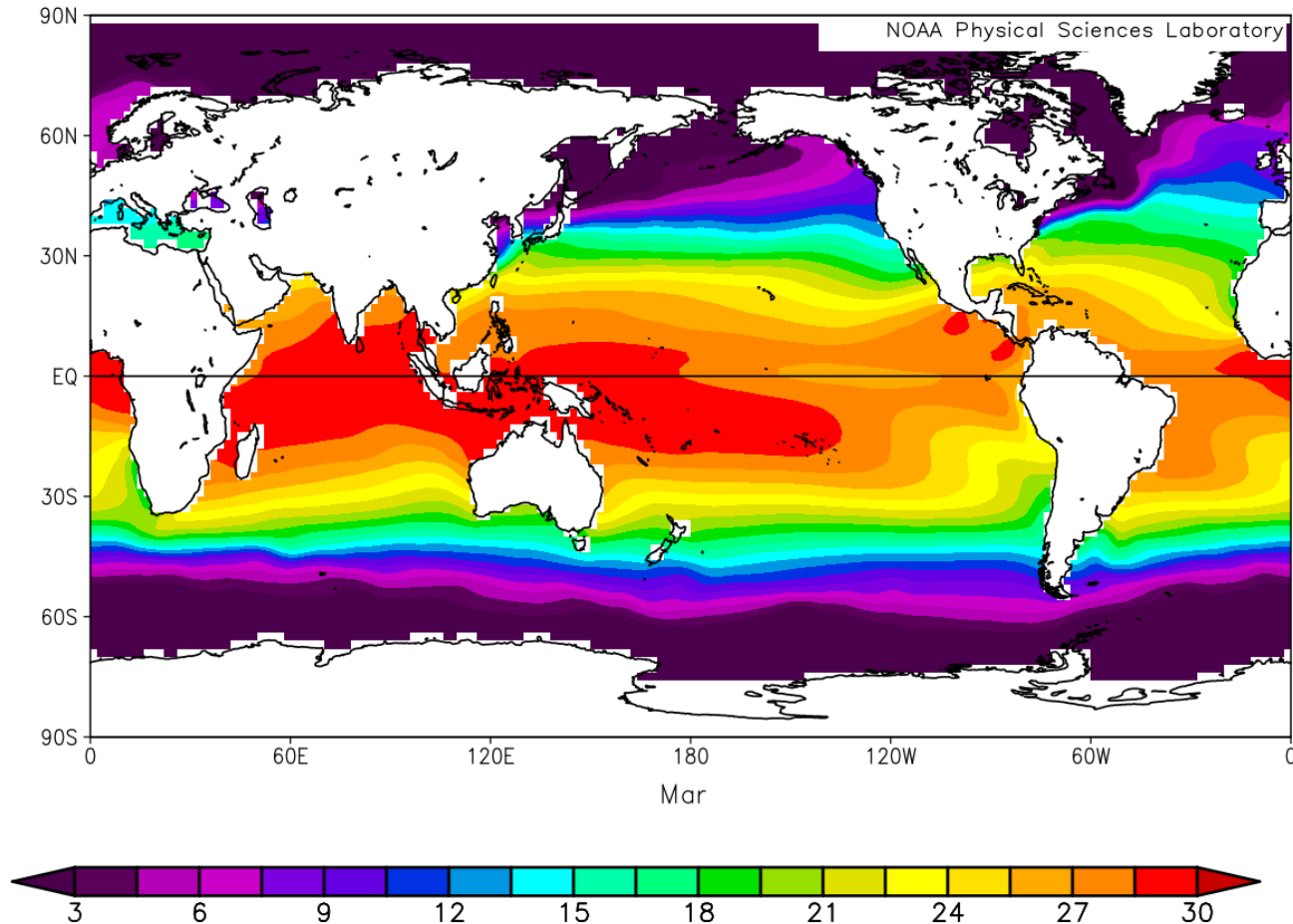
## 1) Warm sea surface temperatures ( $> \sim 26^{\circ}\text{C}$ )



# Necessary conditions for TC development:

## 1) Warm sea surface temperatures ( $> \sim 26^{\circ}\text{C}$ )

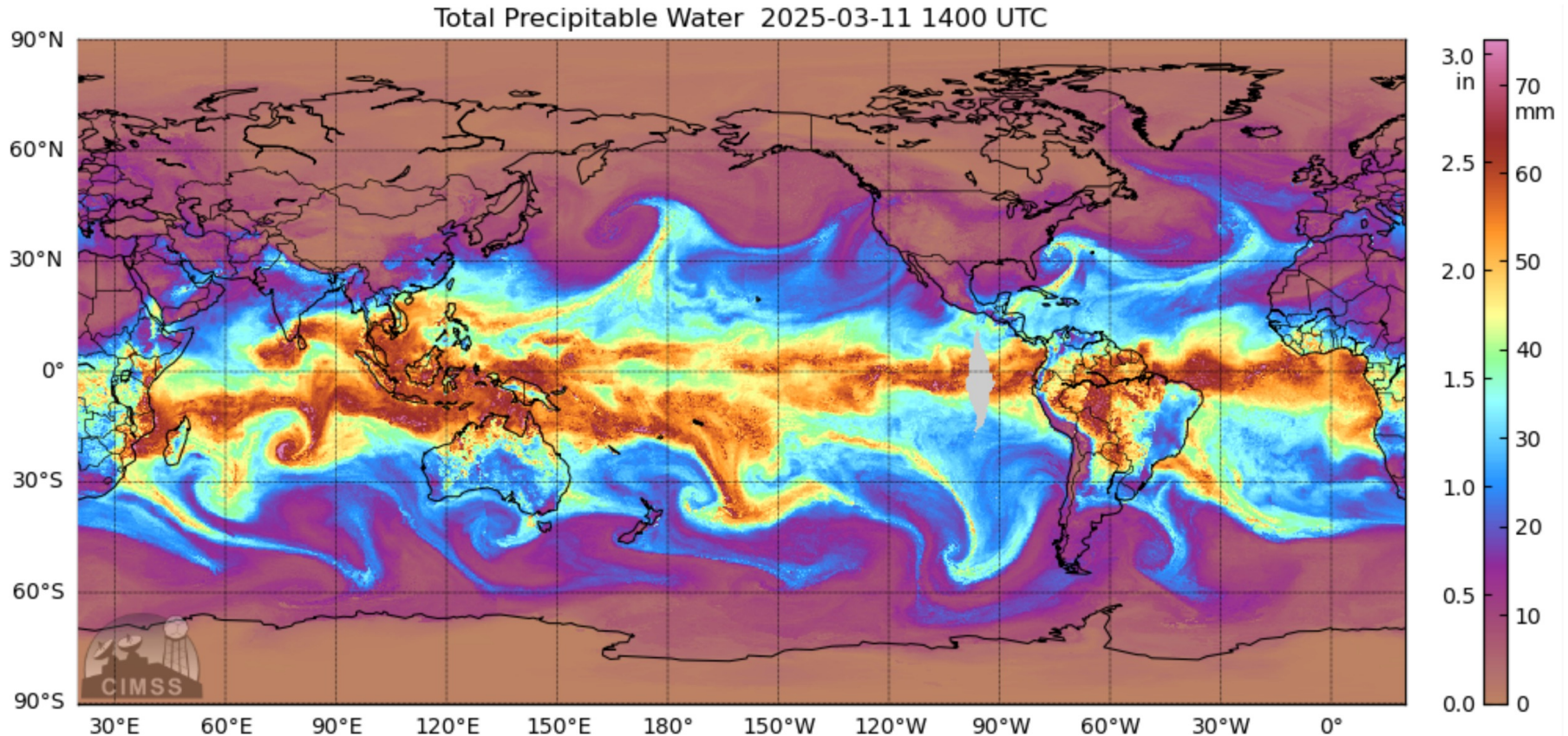
NOAA Extended SST V5 (ERSST)  
Surface SST ( $^{\circ}\text{C}$ ) Climatology 1991–2020 climo





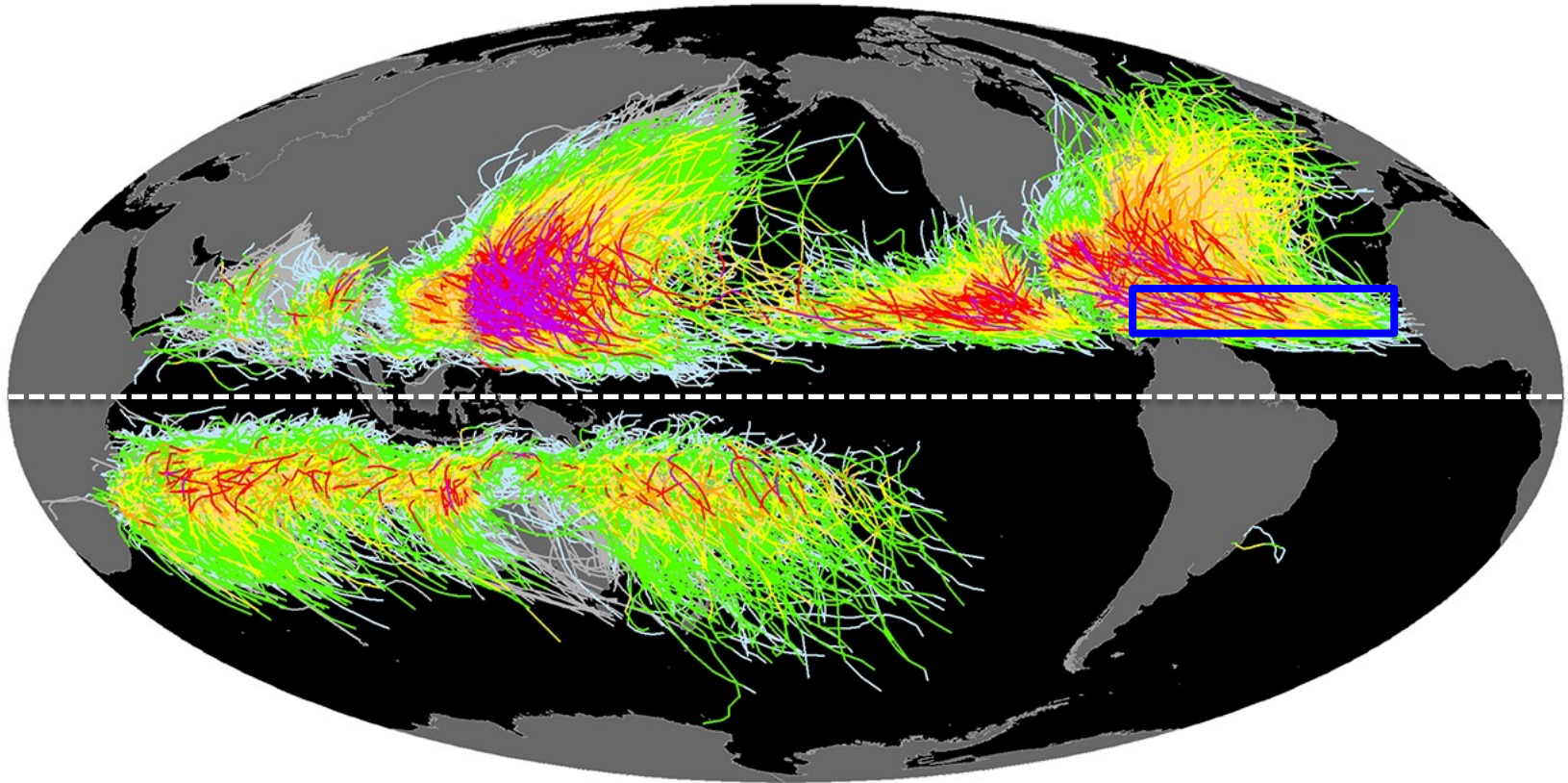
# Necessary conditions for TC development:

## 1a) Warm SSTs ( $> \sim 26^{\circ}\text{C}$ ) & a moist atmosphere



# Necessary conditions for TC development:

## 2) Off equatorial location (Coriolis effect)



The International Best Track Archive for Climate Stewardship (IBTrACS) stores global tropical cyclone information.

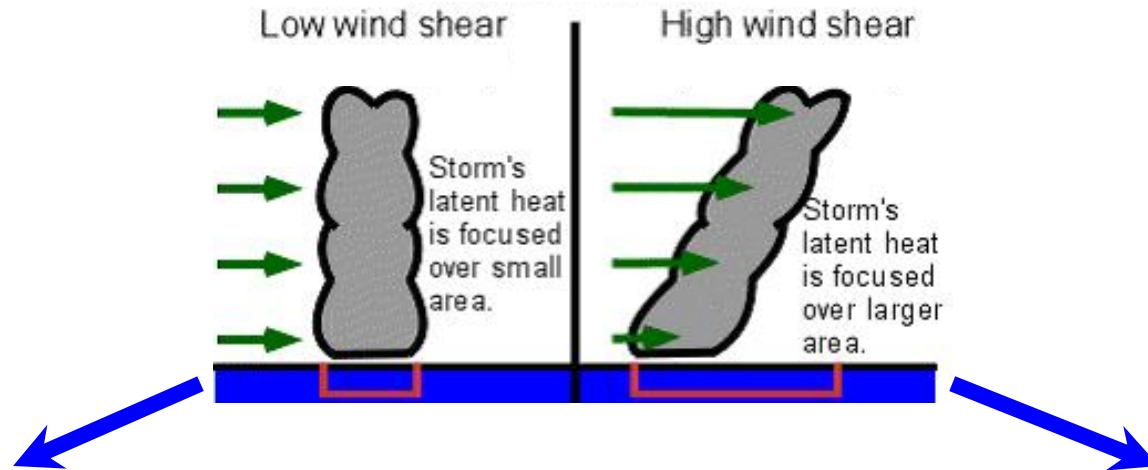
### Saffir-Simpson Hurricane Wind Scale

Intensity Missing	—	Category 1	—
Tropical Depression	—	Category 2	—
Tropical Storm	—	Category 3	—
		Category 4	—
		Category 5	—

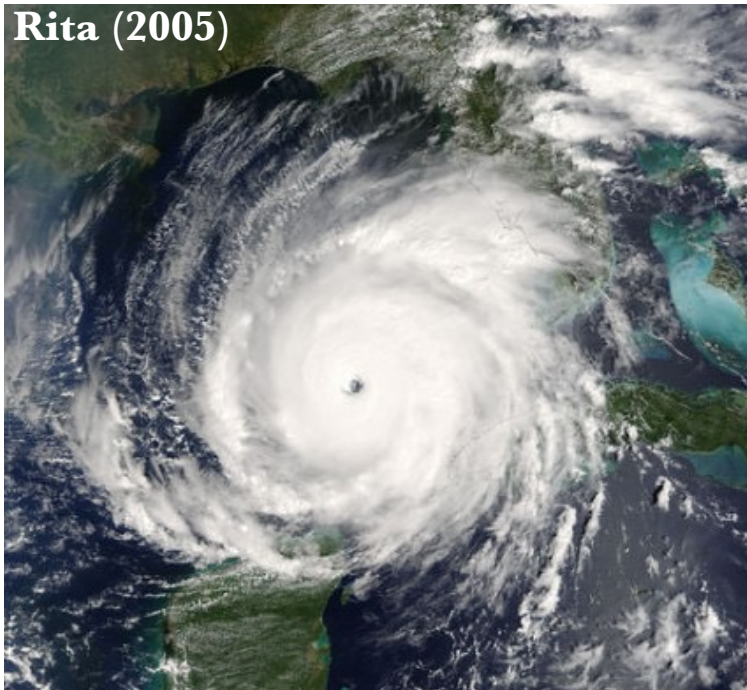


# Necessary conditions for TC development:

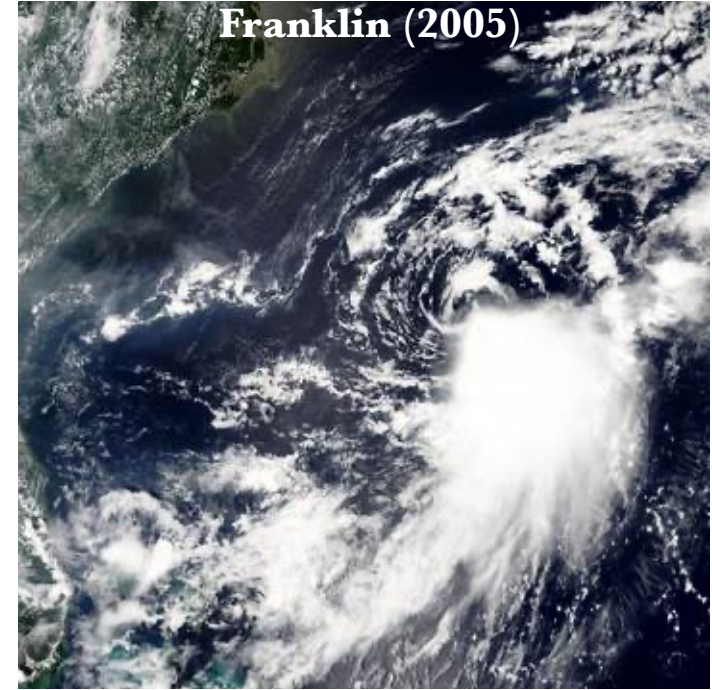
## 3) Relatively “weak” vertical wind shear



**Rita (2005)**

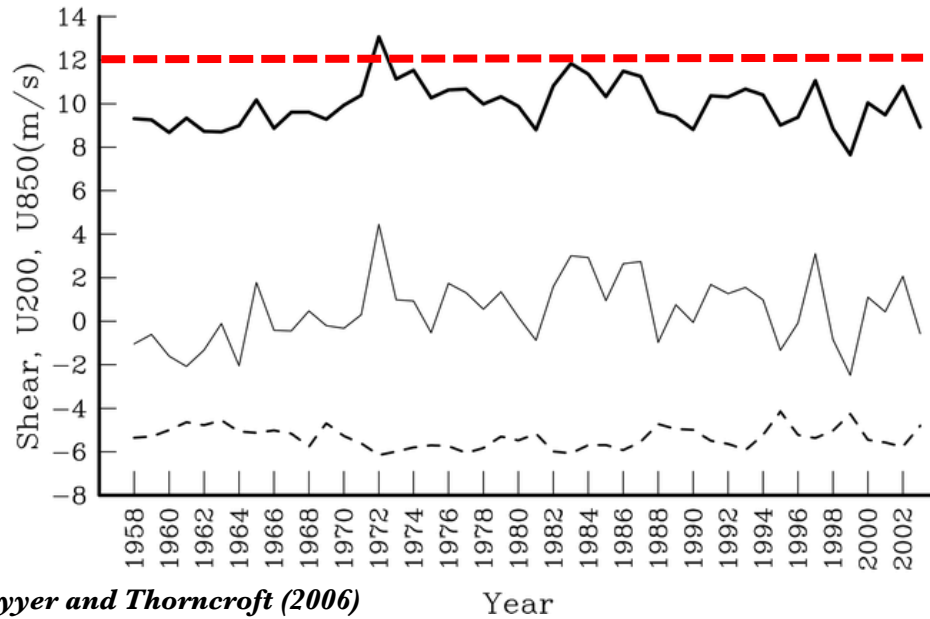


**Franklin (2005)**



# Necessary conditions for TC development:

## 3) Relatively “weak” vertical wind shear



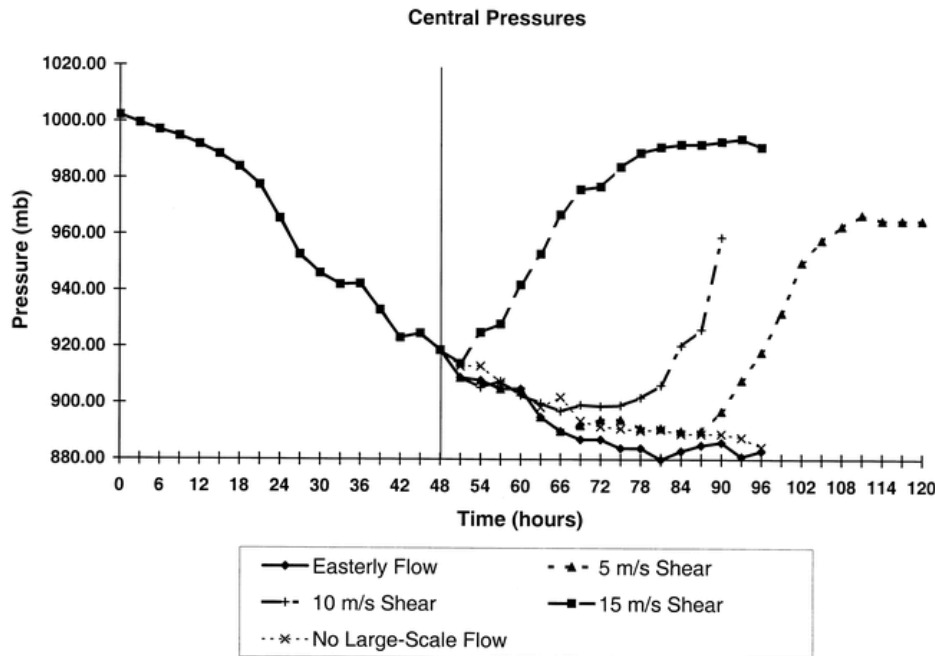
Vertical wind shear  
**Atlantic main  
development region (MDR)  
(left; bold solid line)**

*Aiyyer and Thorncroft (2006)*

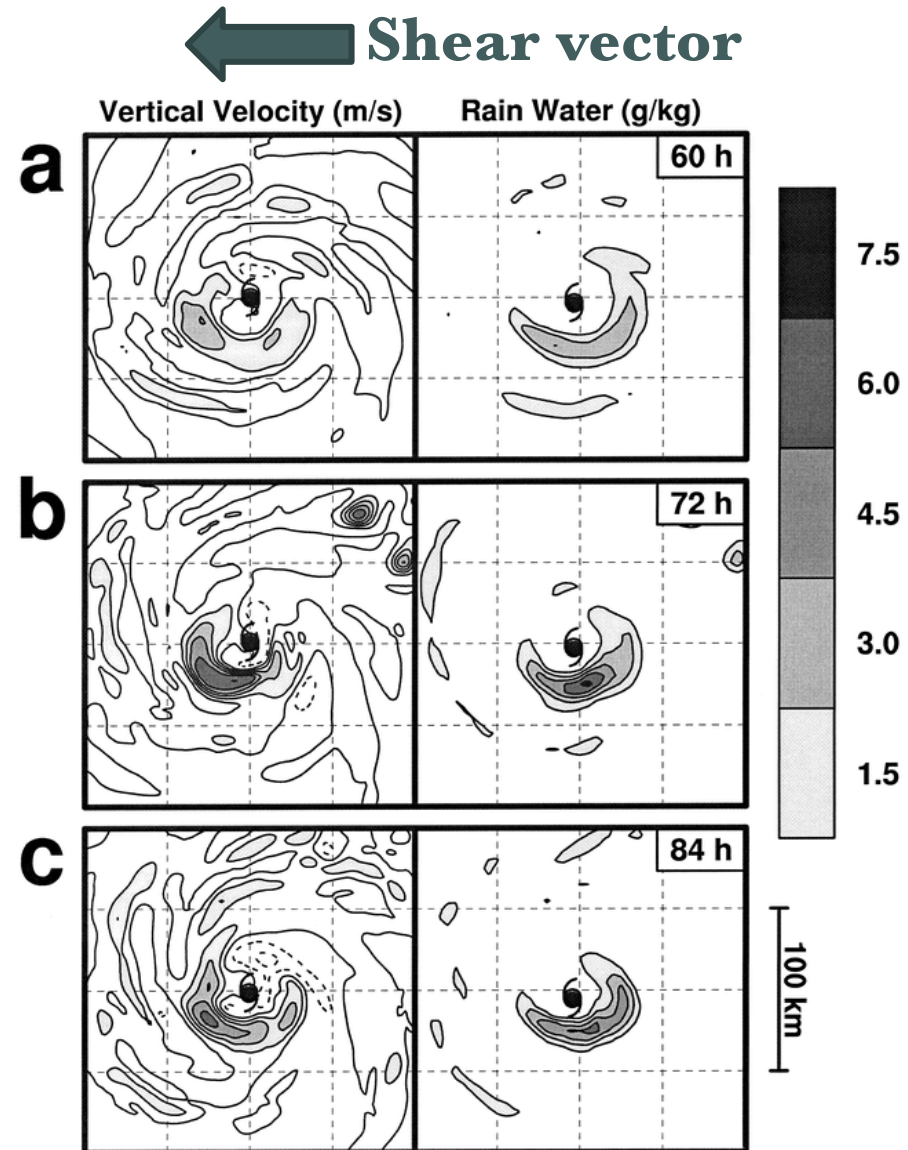


# Necessary conditions for TC development:

## 3) Relatively “weak” vertical wind shear

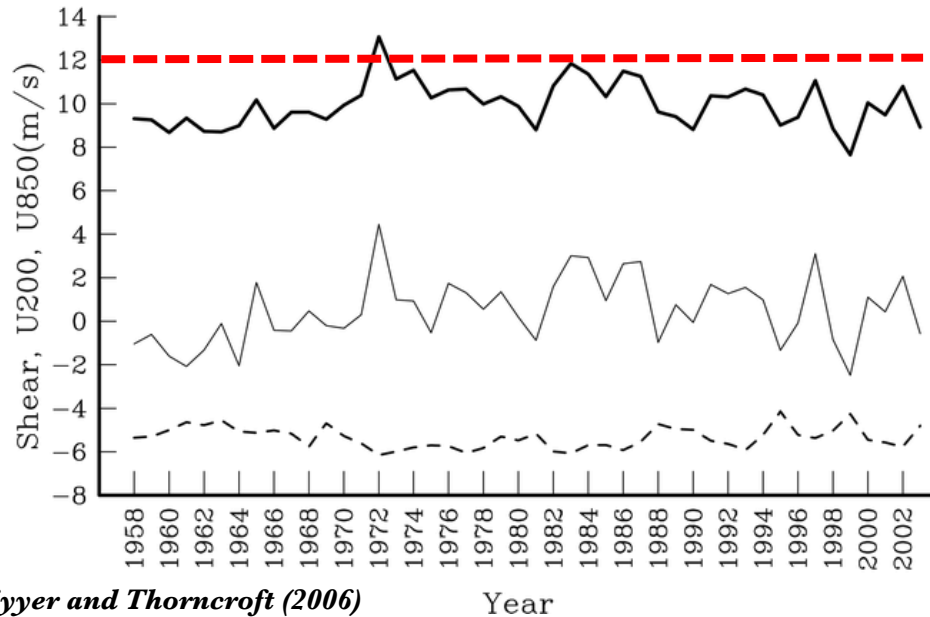


**5 m s<sup>-1</sup> of vertical shear is enough to generate significant asymmetries and weaken a TC after one to two days**



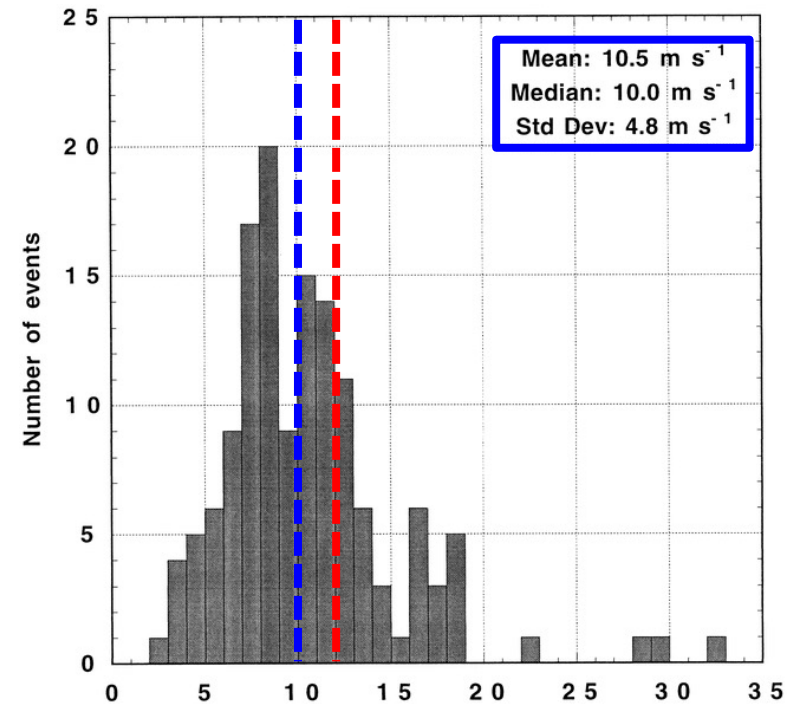
# Necessary conditions for TC development:

## 3) Relatively “weak” vertical wind shear



Vertical wind shear  
At time of genesis for 139  
Atlantic basin tropical  
depressions  
(ATOLL = ~900 hPa)

Vertical wind shear  
Atlantic main  
development region (MDR)  
(left; bold solid line)

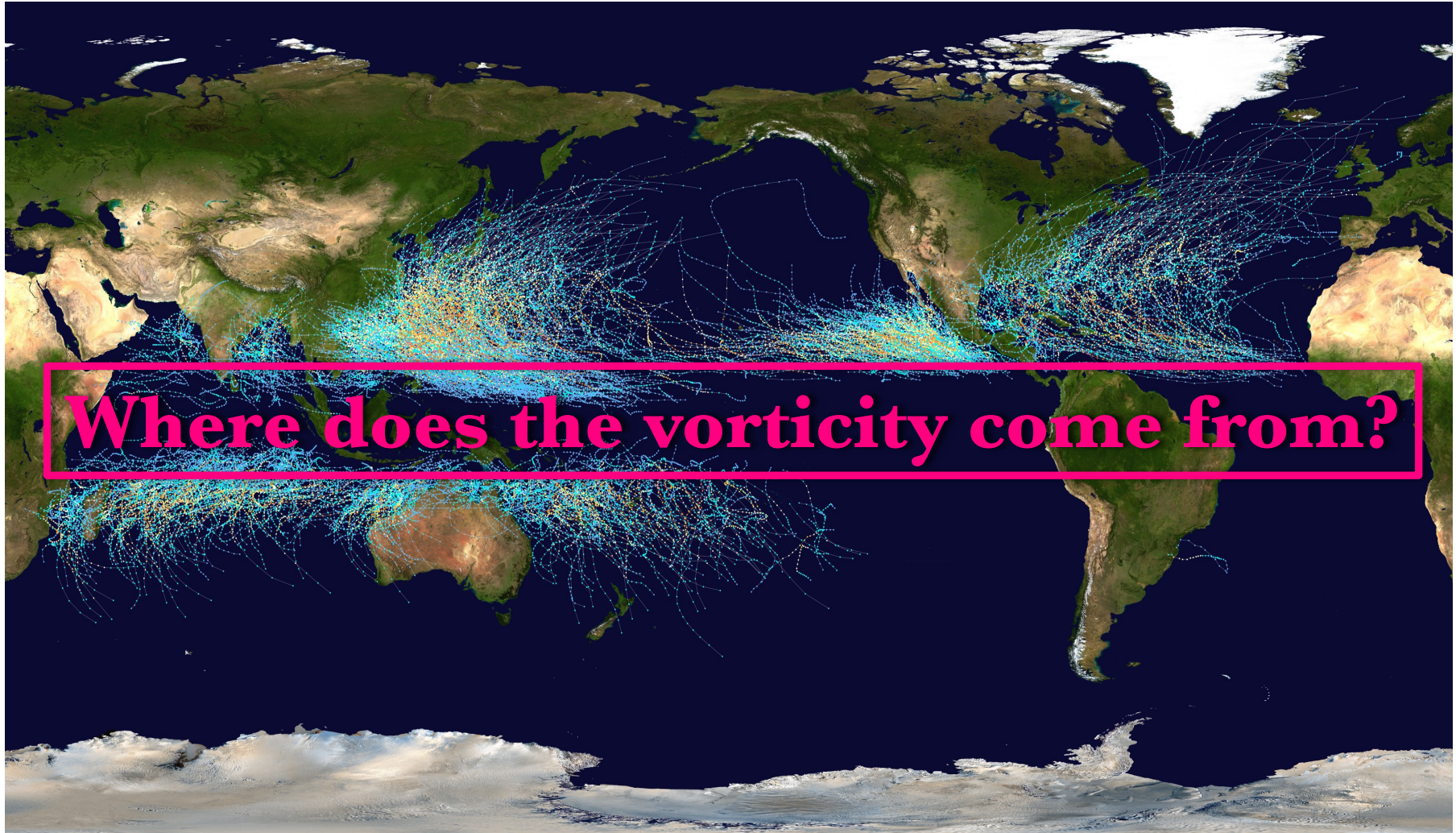


*Bracken and Bosart (2000)*



# Necessary conditions for TC development:

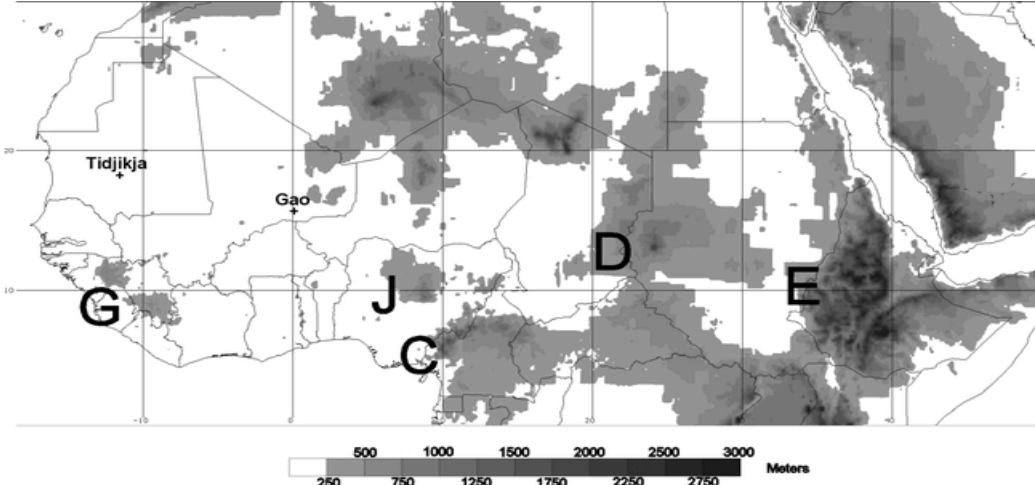
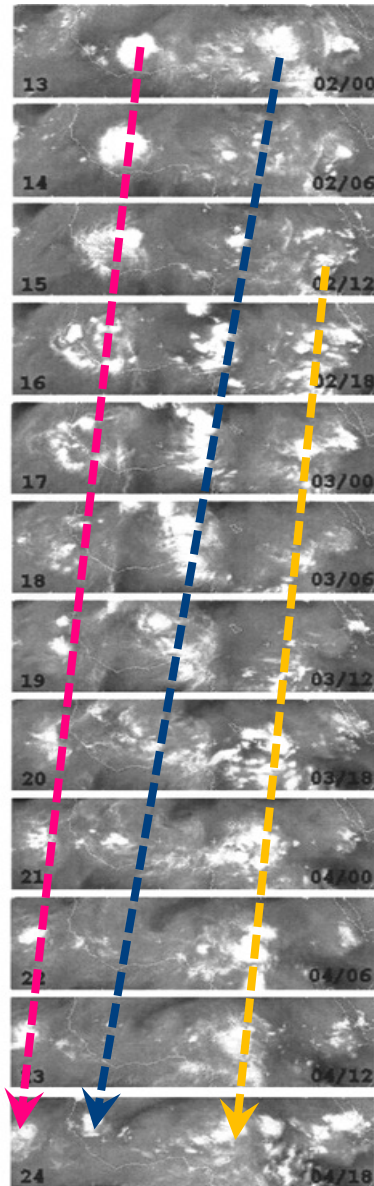
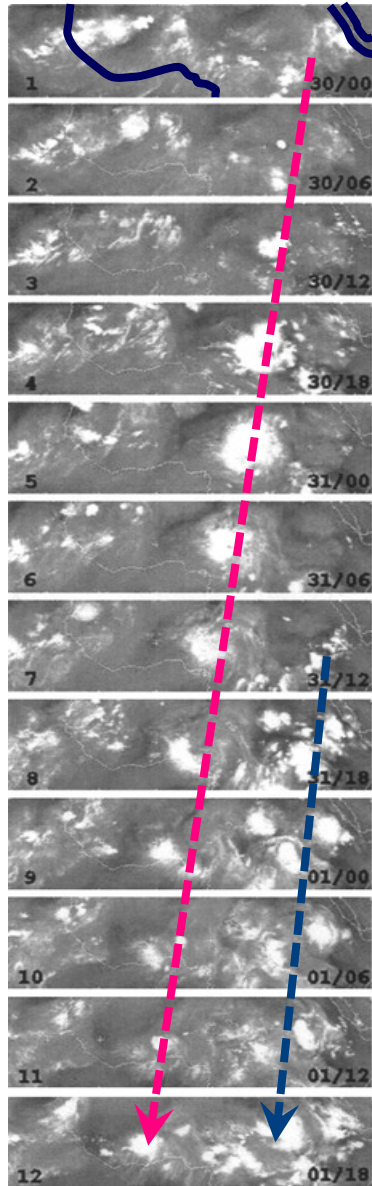
## 4) Pre-existing finite amplitude cyclonic vorticity



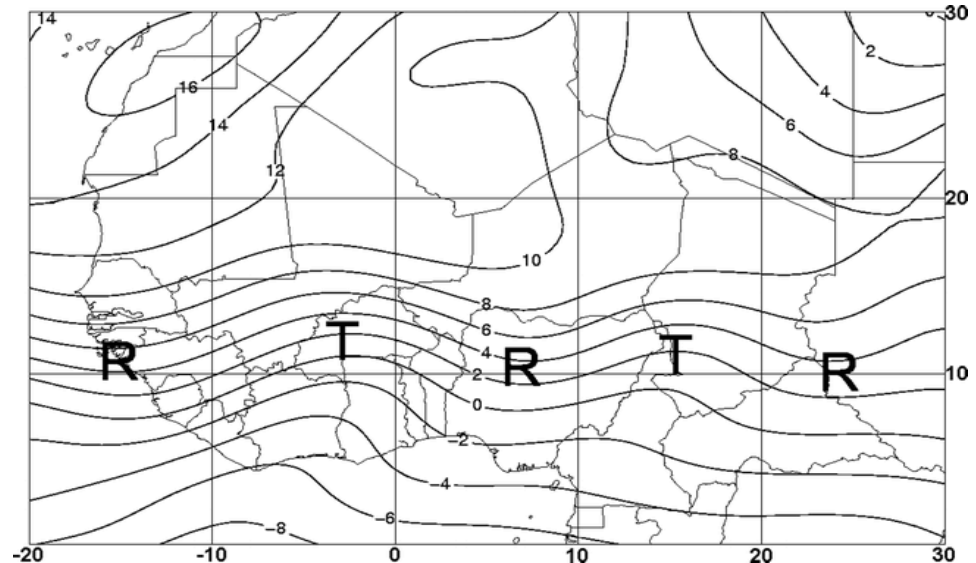


# Necessary conditions for TC development:

## 4) Pre-existing finite amplitude cyclonic vorticity



## African easterly waves & PV



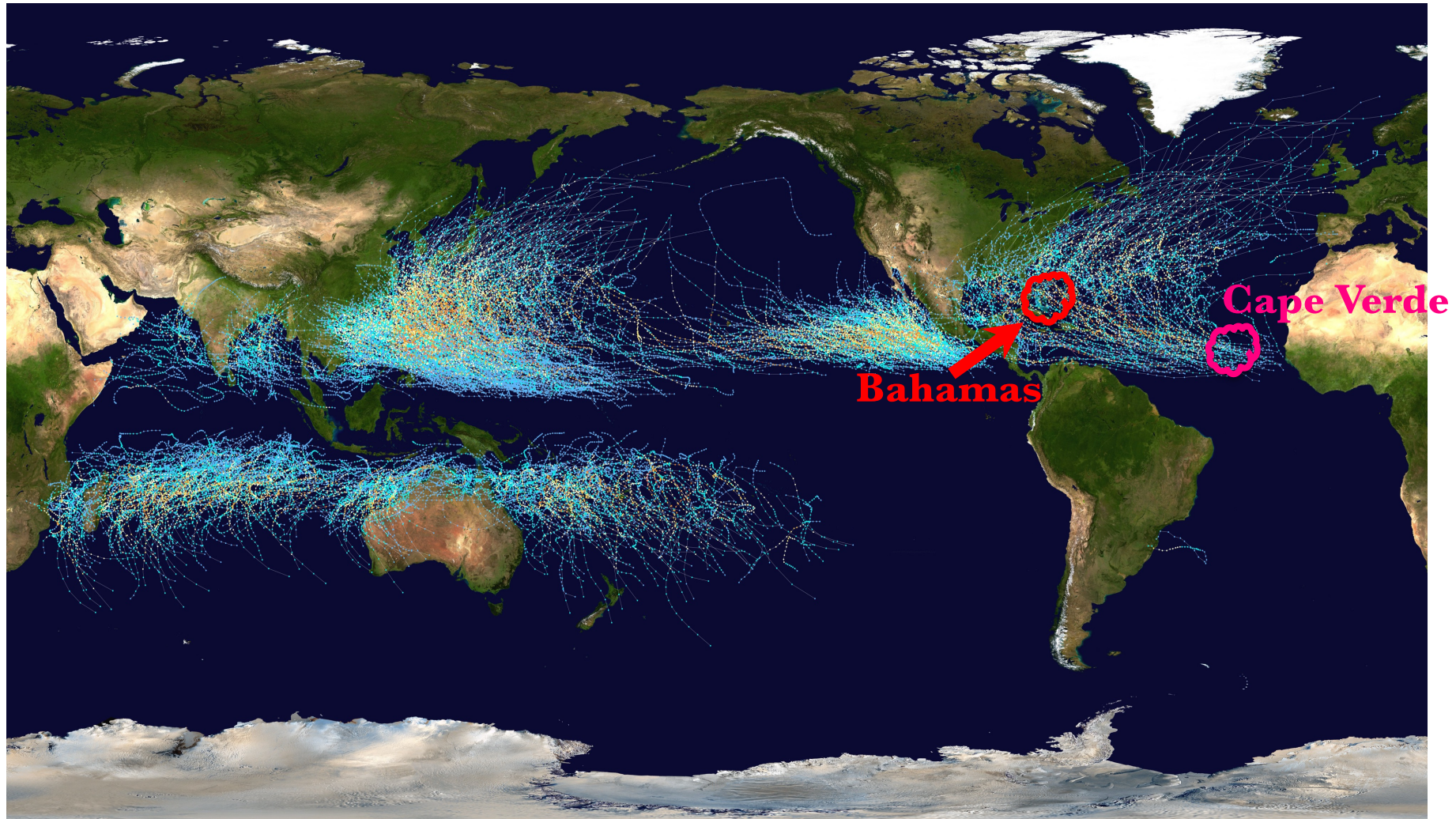
Berry and Thorncroft (2005)

Berry et al. (2007)



# Necessary conditions for TC development:

## 4) Pre-existing finite amplitude cyclonic vorticity

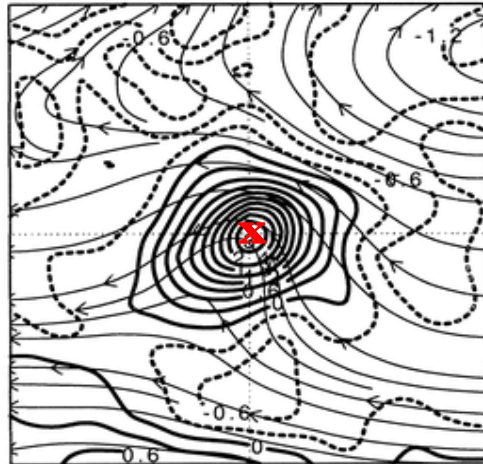




# Necessary conditions for TC development:

## 4) Pre-existing finite amplitude cyclonic vorticity

**Bahamas**

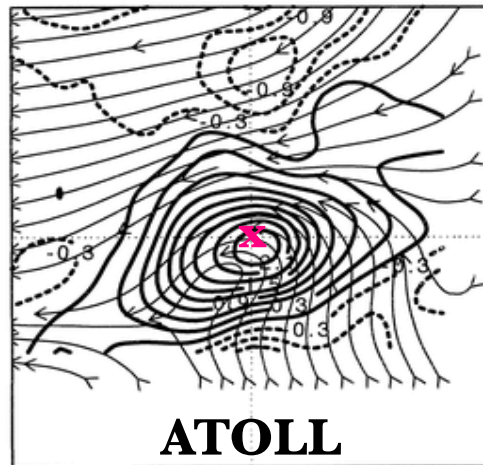


a. ATOLL level Bahamas cases

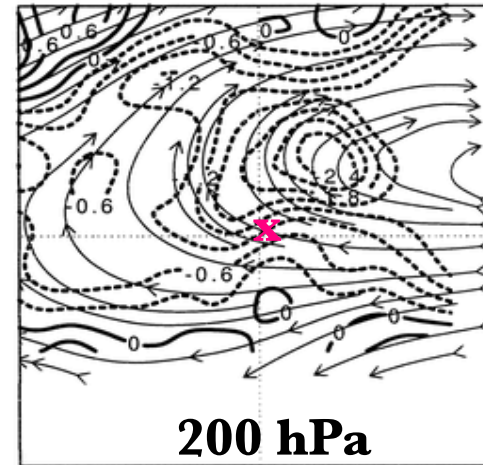


b. 200 hPa Bahamas cases

**Cape Verde**



c. ATOLL level Cape Verde cases



d. 200 hPa Cape Verde cases

*Bracken and Bosart (2000)*

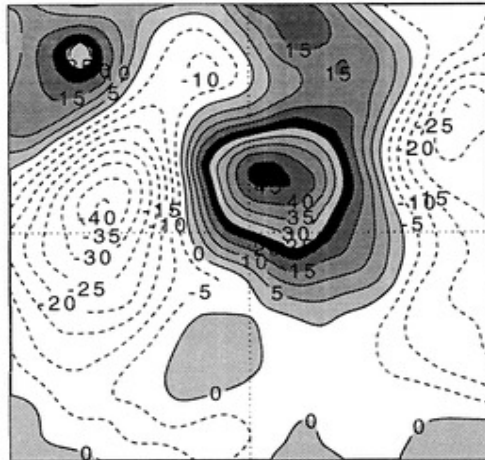
Composite streamlines (thin solid) and relative vorticity (every  $0.3 \times 10^{-5} \text{ s}^{-1}$ , cyclonic solid and anticyclonic dashed)



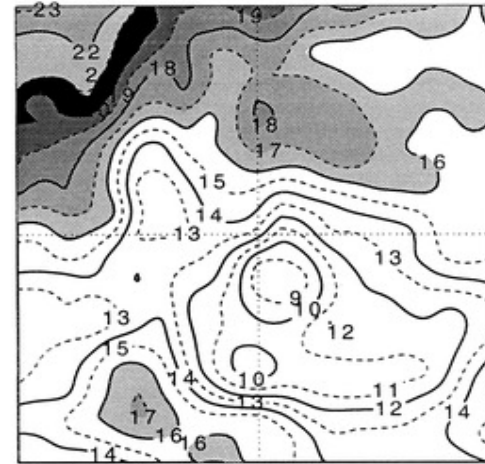
# Necessary conditions for TC development:

## 4) Pre-existing finite amplitude cyclonic vorticity

**Bahamas**

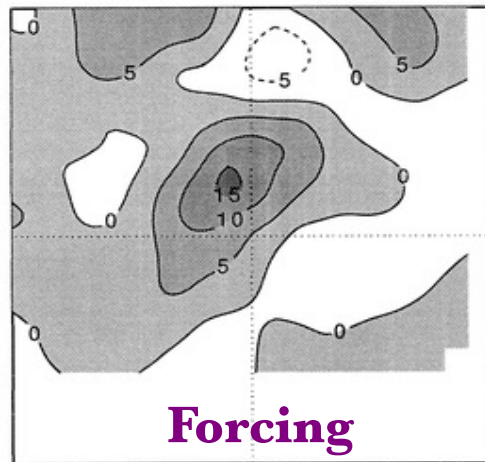


a. Forcing Bahamas cases

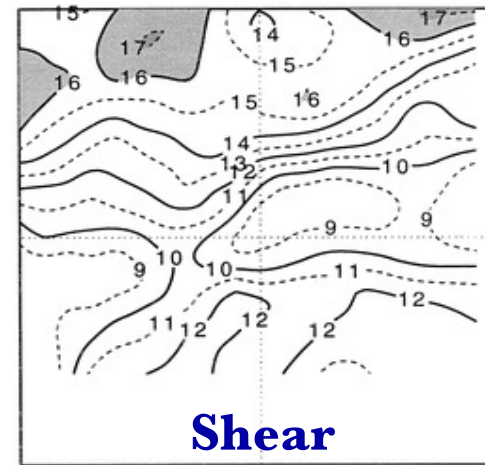


b. Shear Bahamas cases

**Cape Verde**



c. Forcing Cape Verde cases

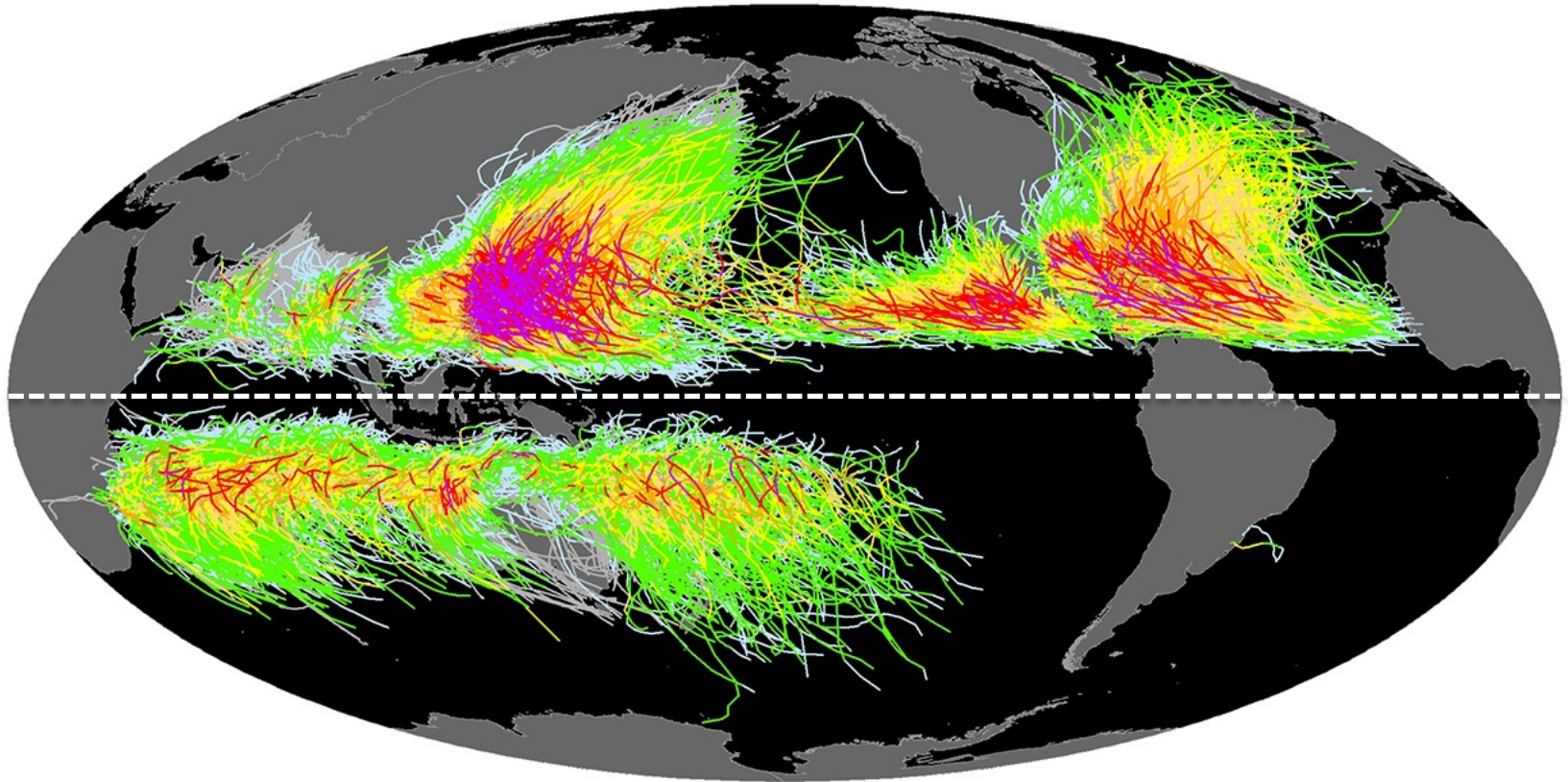


d. Shear Cape Verde cases

*Bracken and Bosart (2000)*

**Sutcliffe–Trenberth forcing for vertical motion**  
**and ATOLL–200-hPa vertical shear**

# Necessary conditions for TC development: Putting it all together!



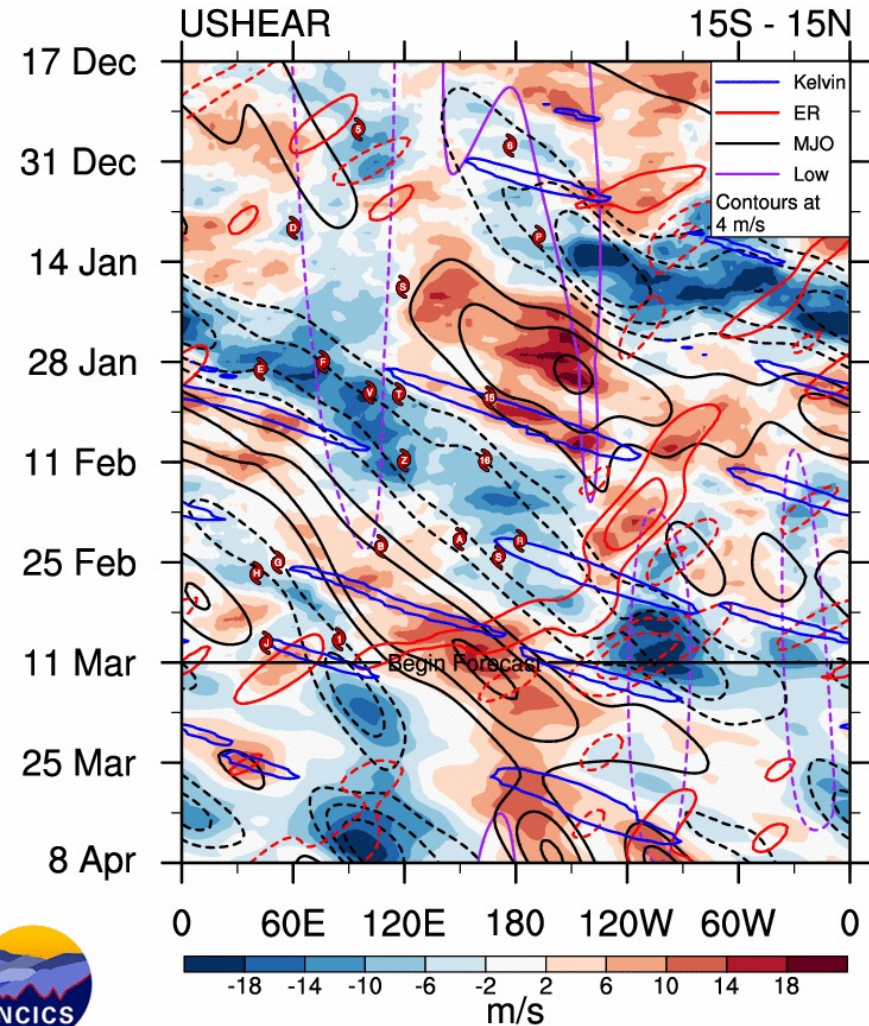
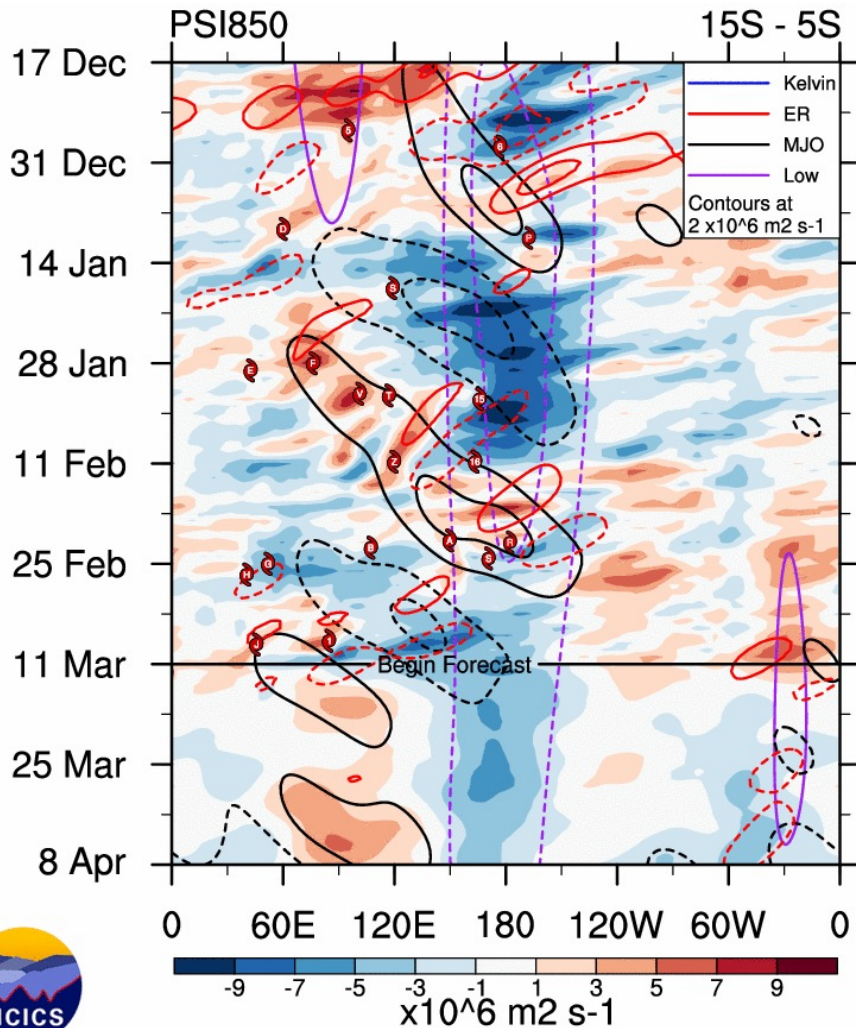
The International Best Track Archive for Climate Stewardship (IBTrACS) stores global tropical cyclone information.

## Saffir-Simpson Hurricane Wind Scale

Intensity Missing	—	Category 1	—
Tropical Depression	—	Category 2	—
Tropical Storm	—	Category 3	—
		Category 4	—
		Category 5	—

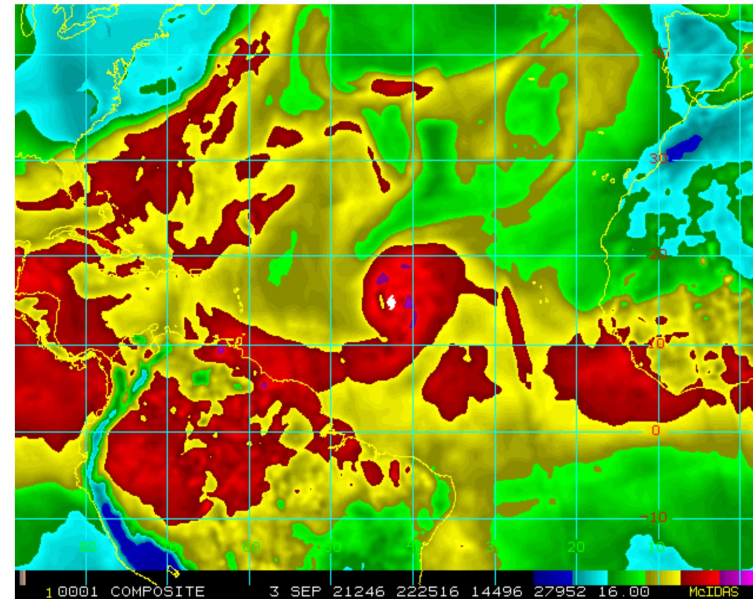
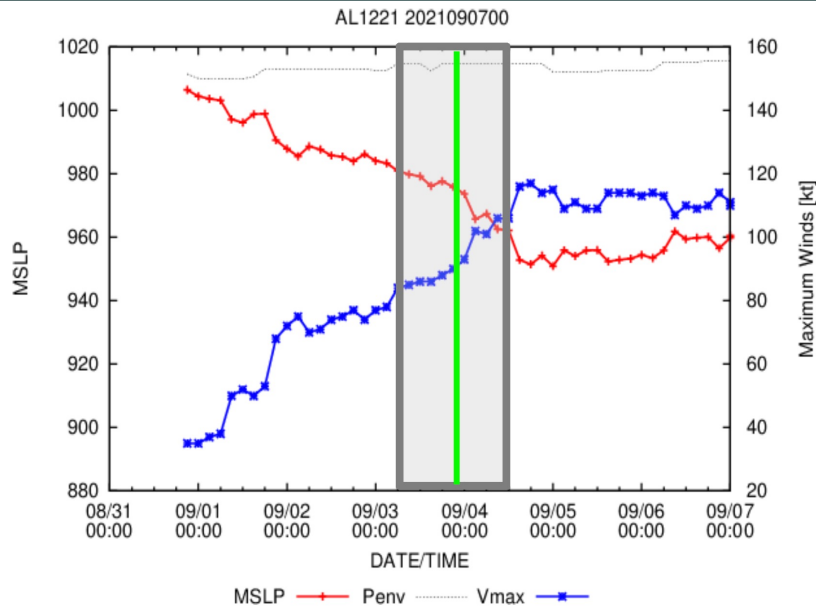


# Necessary conditions for TC development: Putting it all together with Carl's Hovmöllers!

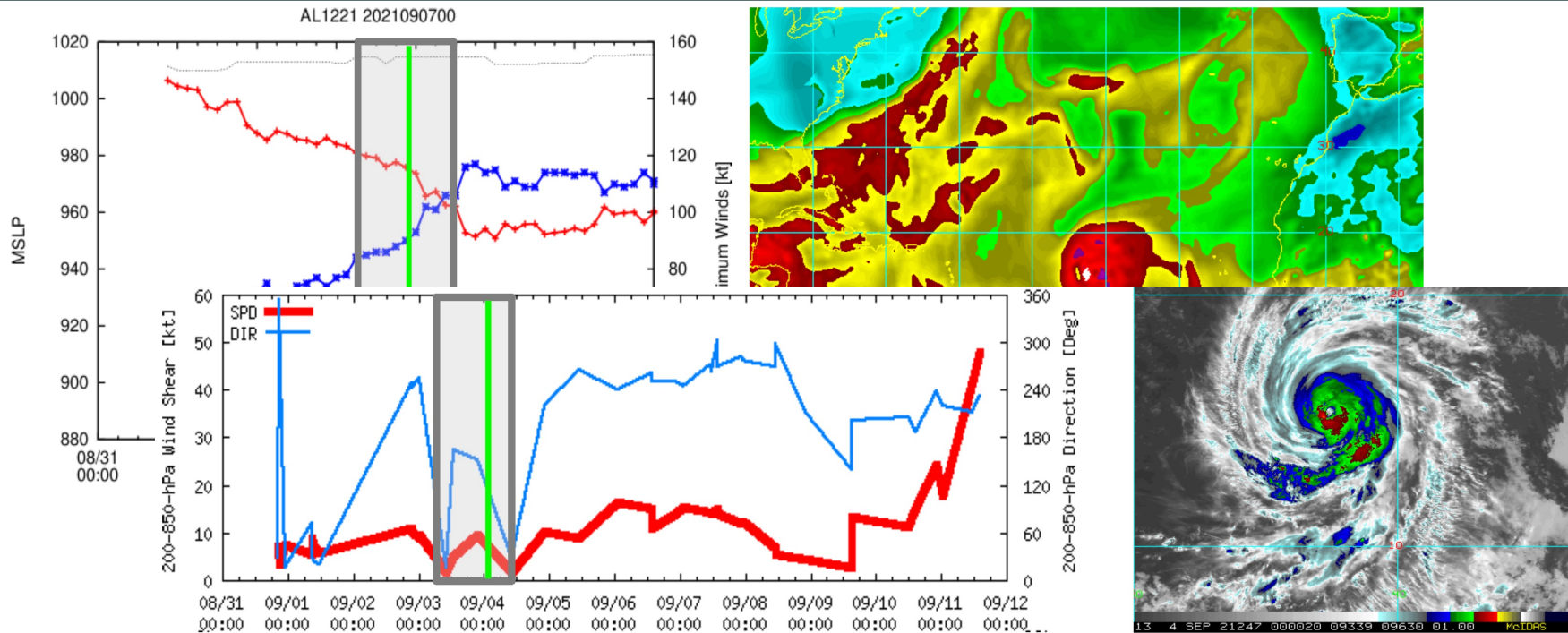




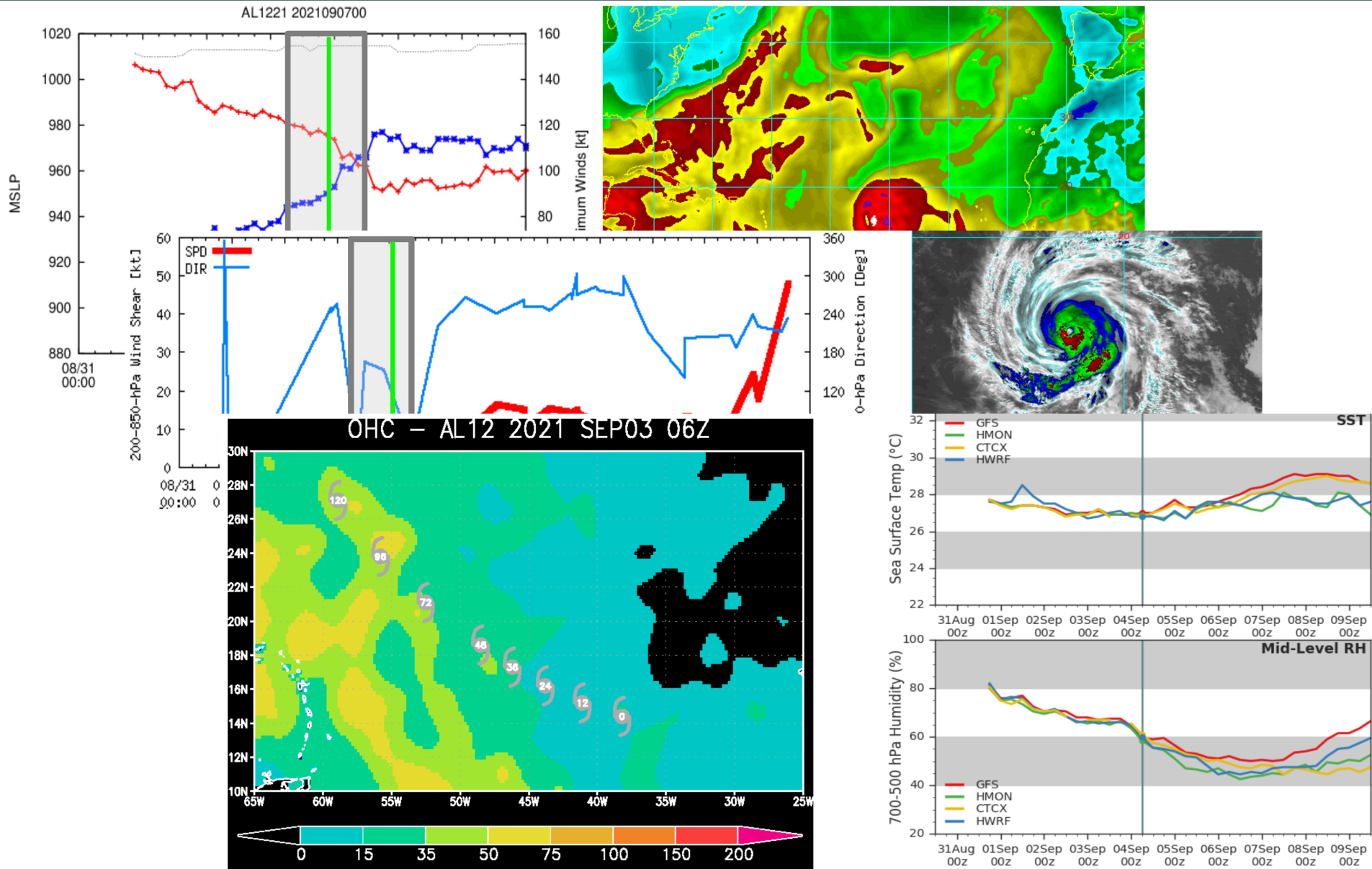
# Necessary conditions for TC development: Putting it all together with CIRA plots!



# Necessary conditions for TC development: Putting it all together with CIRA plots!

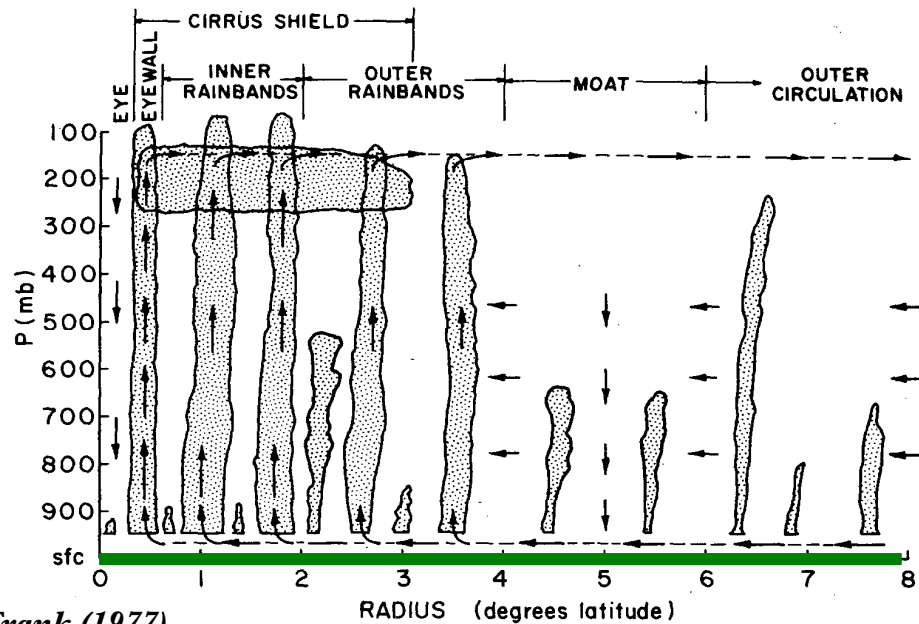
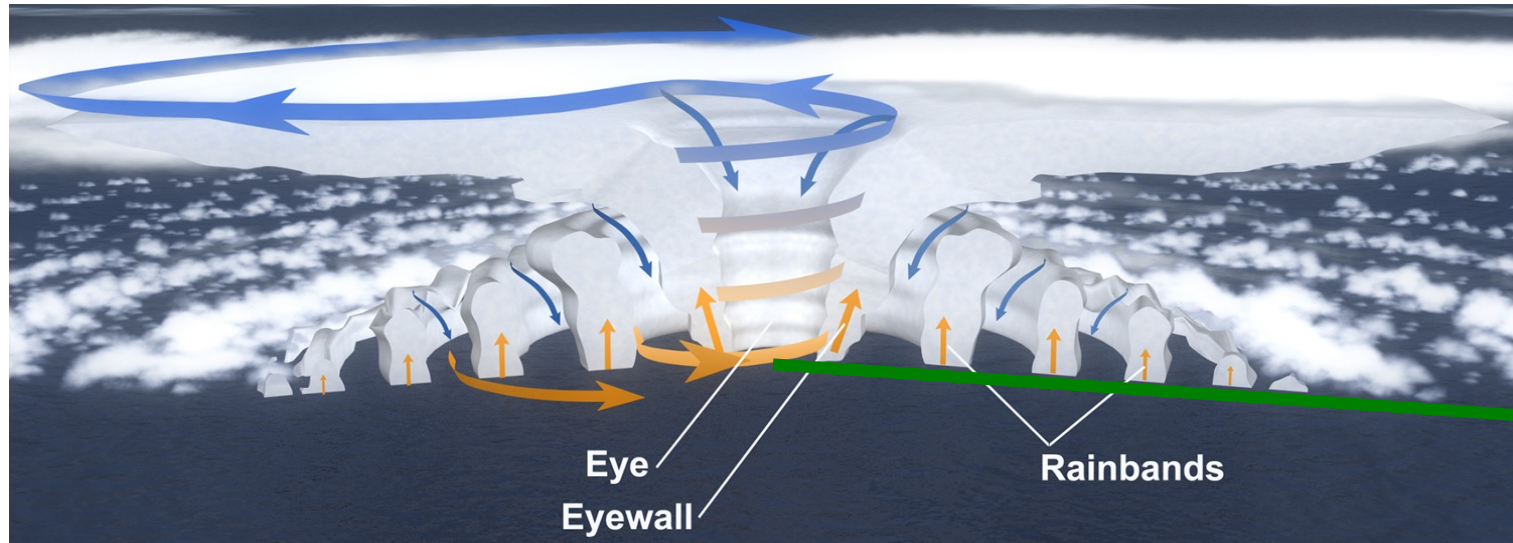


# Necessary conditions for TC development: Putting it all together with CIRA plots!



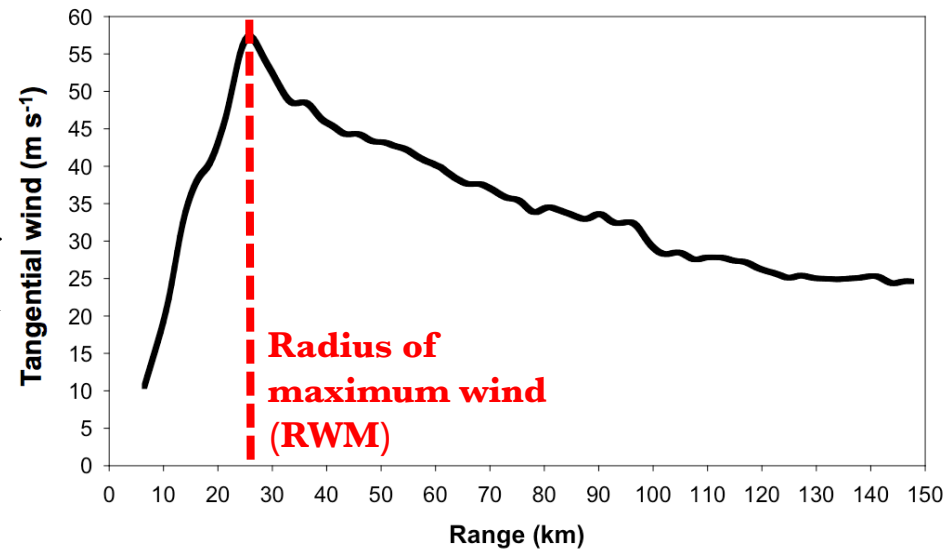
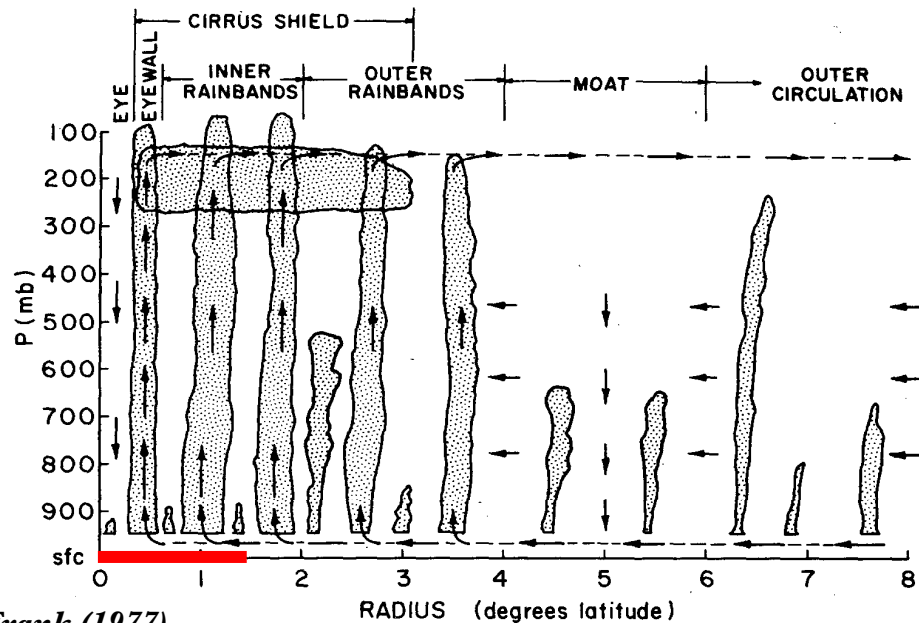
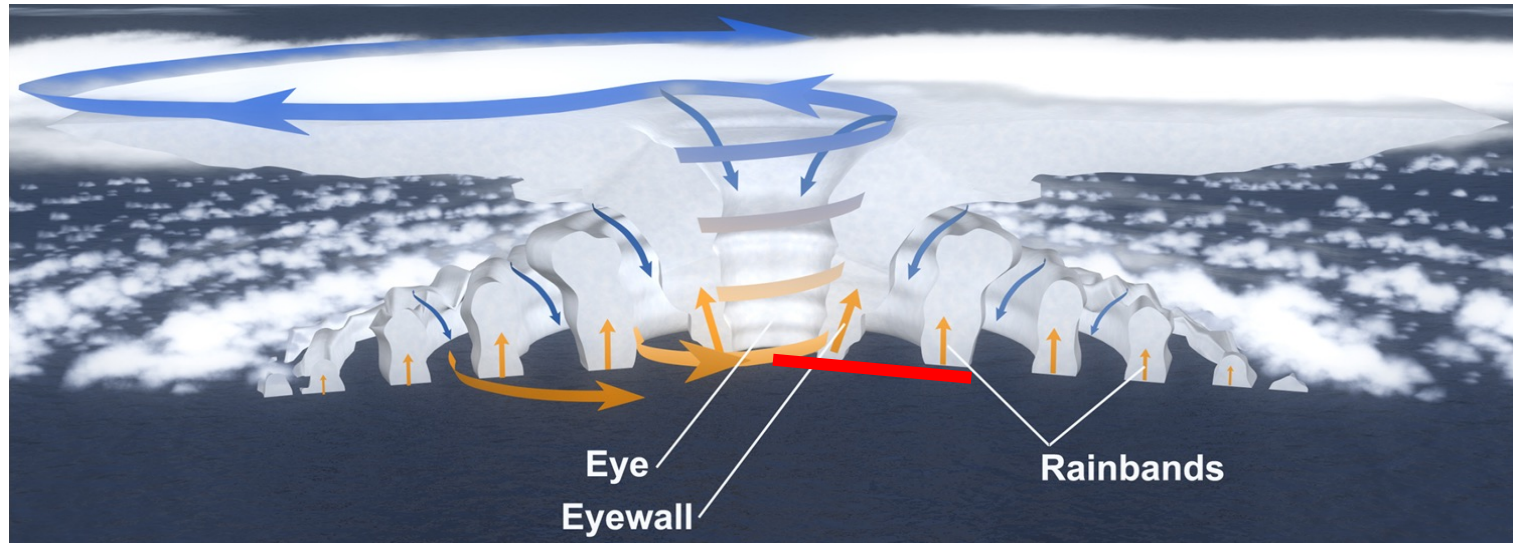


# Tropical cyclone structure: Primary and secondary circulations



Frank (1977)

# Tropical cyclone structure: Primary and secondary circulations





# Tropical cyclone structure: Primary (tangential) circulation

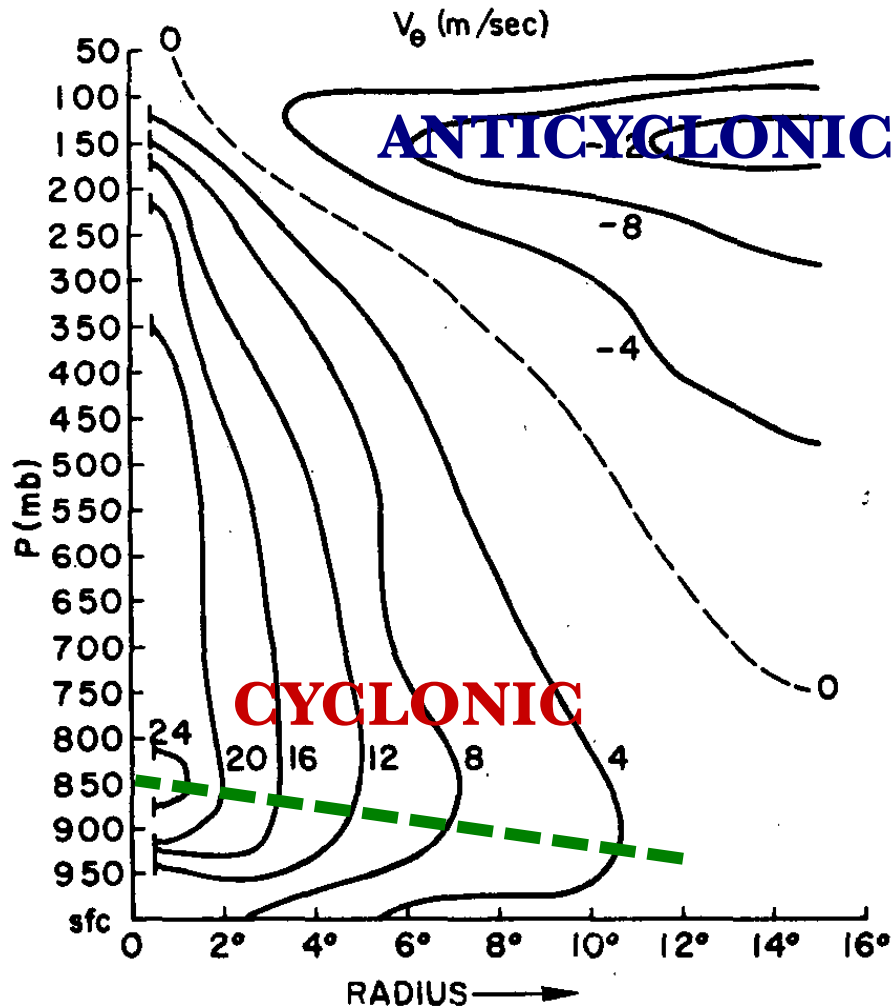


FIG. 9. Two-dimensional cross section of  $V_\theta$  ( $\text{m s}^{-1}$ ) in stationary (NAT) coordinates. Positive numbers denote cyclonic flow.

Frank (1977)

The positive, cyclonic circulation is HUGE ( $\sim 10^\circ$ ), but clouds only extend out to a radius of  $\sim 3^\circ$ – $5^\circ$

Negative, anticyclonic aloft from  $2^\circ$  outward

Winds are maximized at the top of the boundary layer

# Tropical cyclone structure: Temperature anomaly

Latent heat release makes tropical  
cyclones **WARM CORE**

The **maximum temperature  
anomaly** is in the upper troposphere

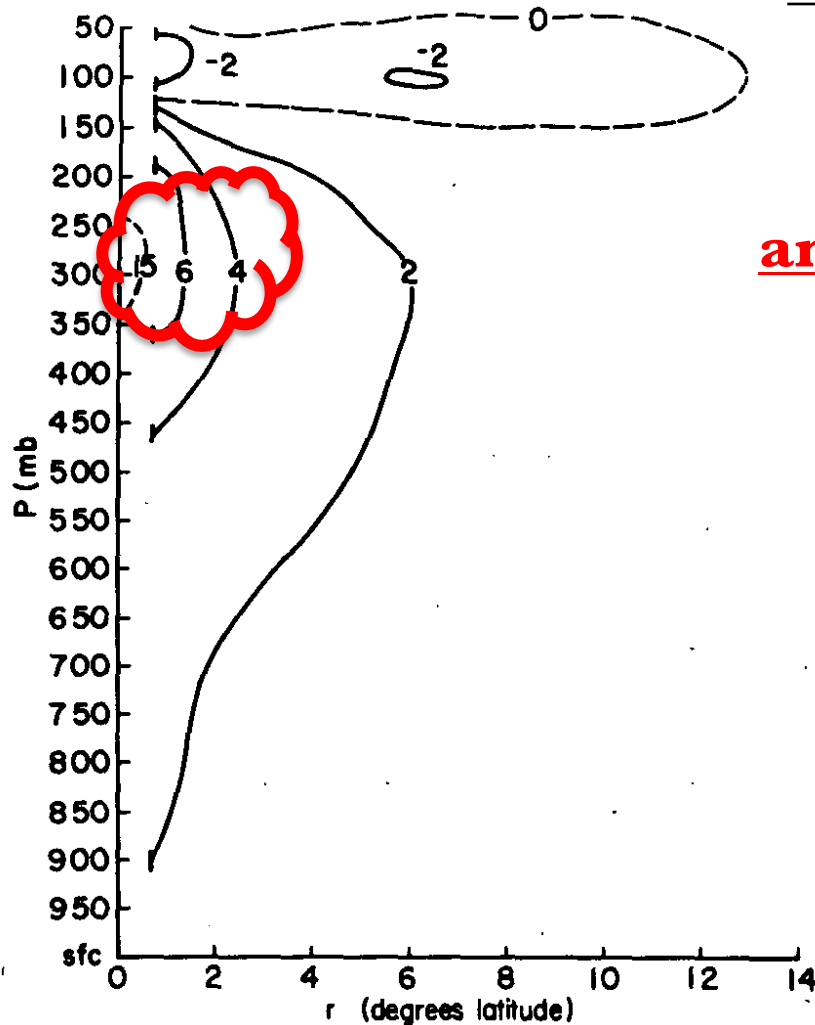
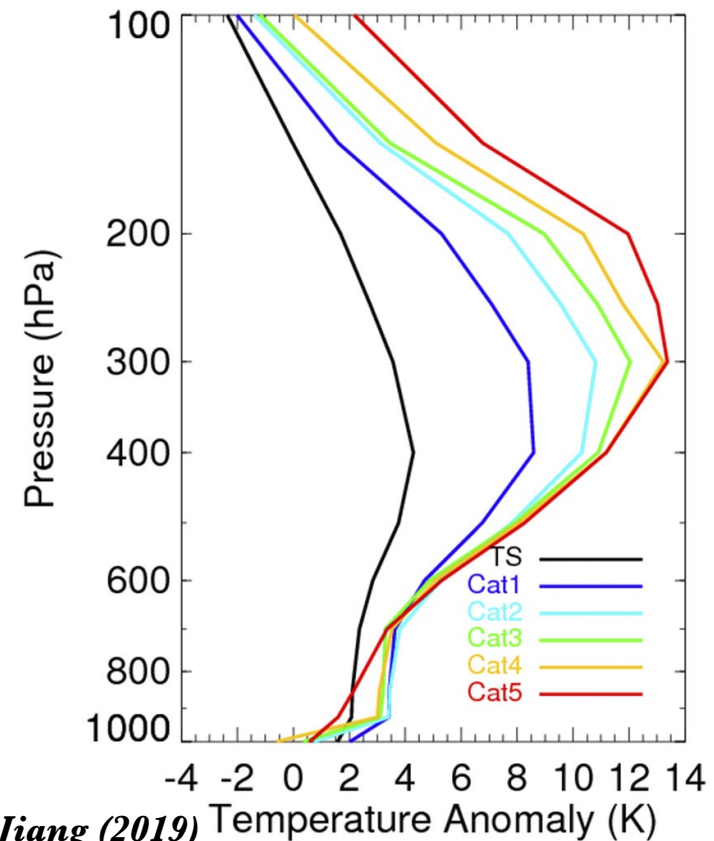


FIG. 3. Temperature anomaly ( $T - \bar{T}_{r=14^\circ}$ ) for mean steady-state typhoon. Units:  $^\circ\text{C}$ .

*Frank (1977)*



*Wang & Jiang (2019)*



# Tropical cyclone structure:

## Temperature anomaly: Warm core

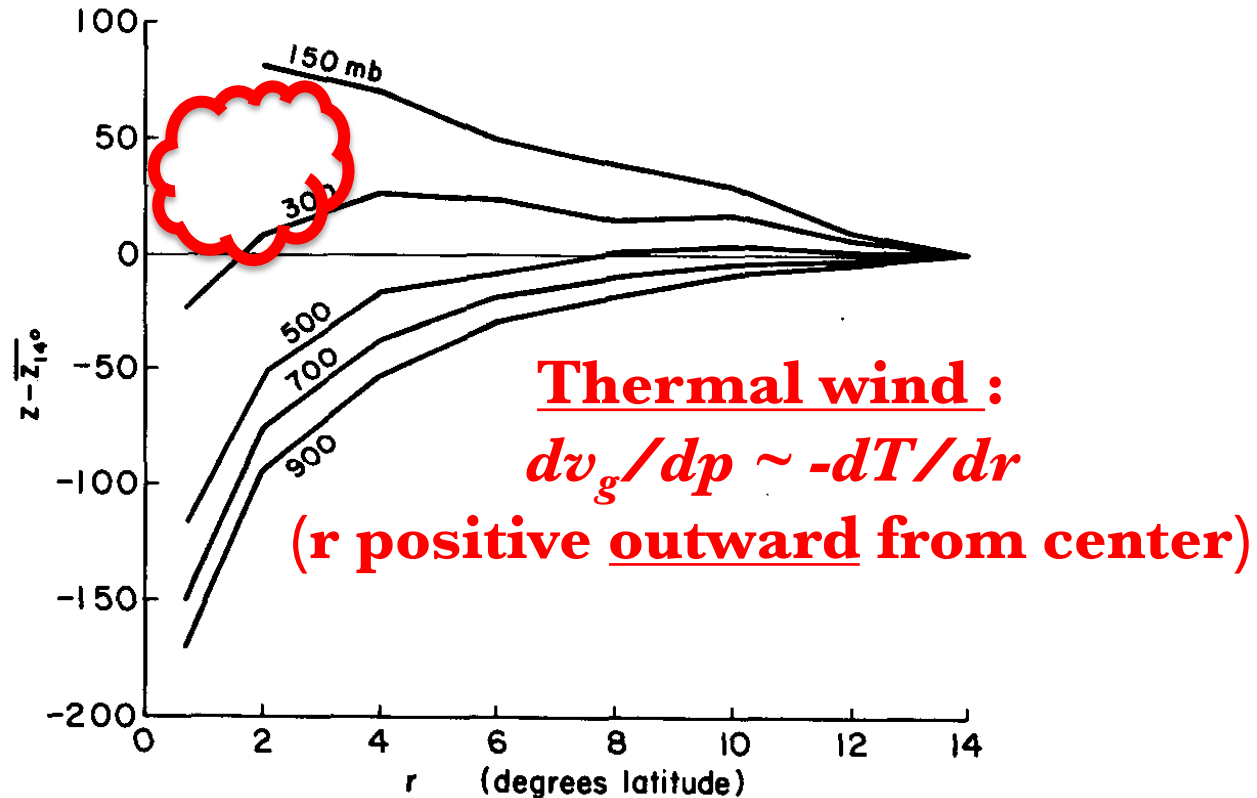


FIG. 4.  $D$  values ( $Z - \bar{Z}_{r=14^\circ}$ ) for mean steady-state typhoon at selected levels. Units: m.

**\* WARM CORE \***

**A low pressure system with the strongest height gradients and, thus, winds, at low levels**

# Tropical cyclone structure: Secondary (radial) circulation

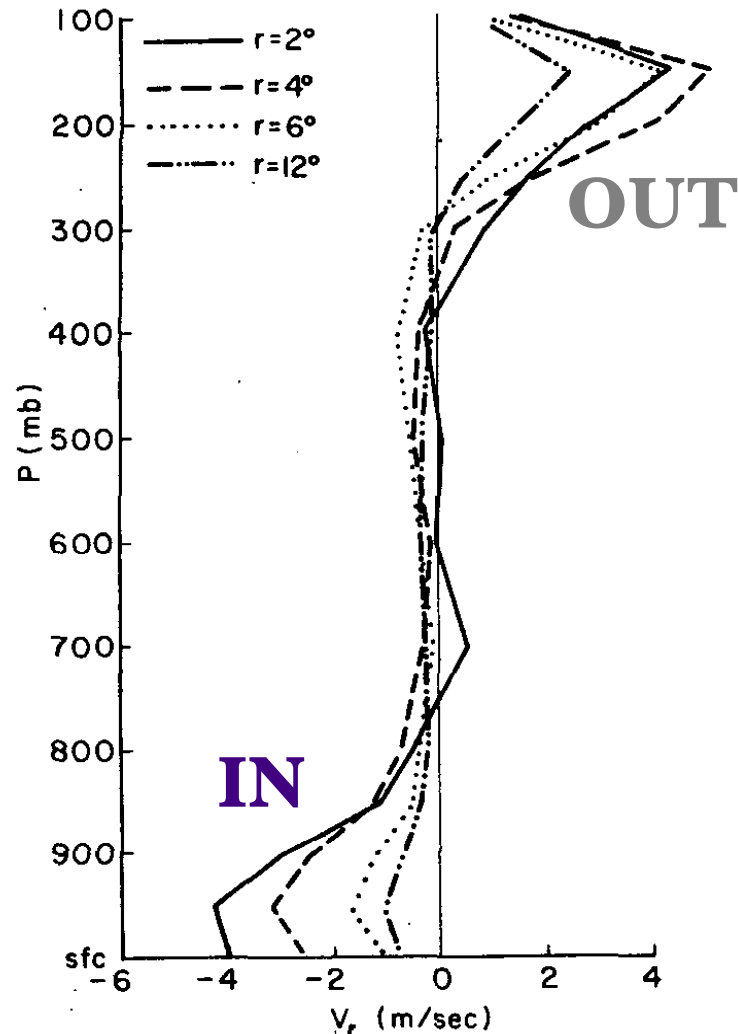


FIG. 12. Two-dimensional vertical profile of radial winds  $V_r$  ( $\text{m s}^{-1}$ ) at  $r=2^\circ$ ,  $4^\circ$ ,  $6^\circ$  and  $12^\circ$ .

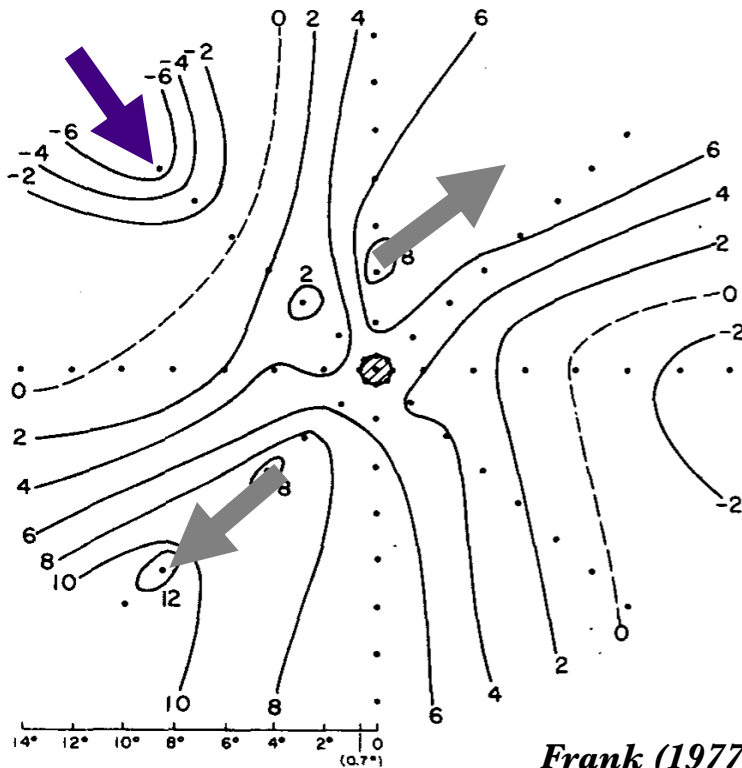
Frank (1977)

The maximum radial INFLOW occurs at ~950 hPa inside a radius of  $2^\circ$  from the center

The maximum OUTFLOW occurs at 150 hPa at  $4^\circ$  from the center

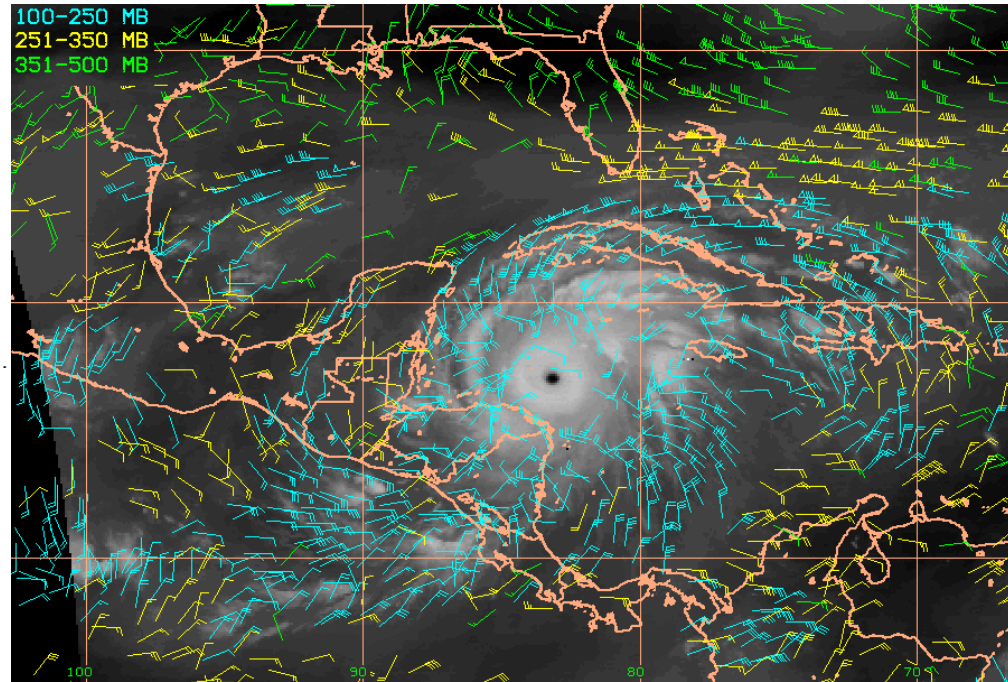


# Tropical cyclone structure: Outflow jets



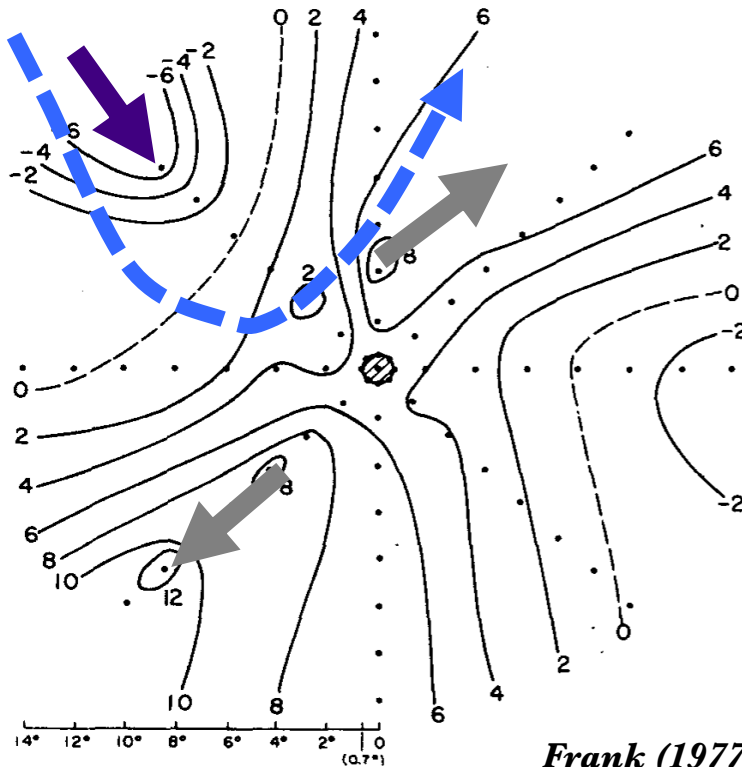
*Frank (1977)*

FIG. 15. Plan view of  $V_r$  ( $\text{m s}^{-1}$ ) in NAT coordinates at 150 mb.



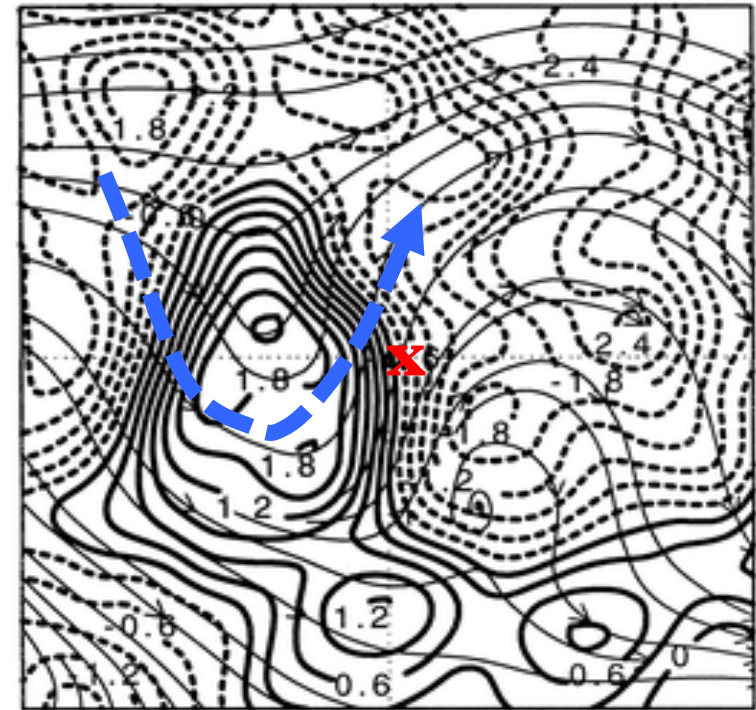
**The upper-level OUTFLOW is concentrated in jets northeast and southwest of the center.**

# Tropical cyclone structure: Outflow jets



*Frank (1977)*

FIG. 15. Plan view of  $V_r$  (m s<sup>-1</sup>) in NAT coordinates at 150 mb.



b. 200 hPa Bahamas cases

*Bracken and Bosart (2000)*

**The upper-level OUTFLOW is concentrated in jets northeast and southwest of the center.**

**In the composite, there is also evidence of a trough to the northwest of the TC like in Bracken and Bosart (2000).**



# Tropical cyclone structure: Convergence and divergence

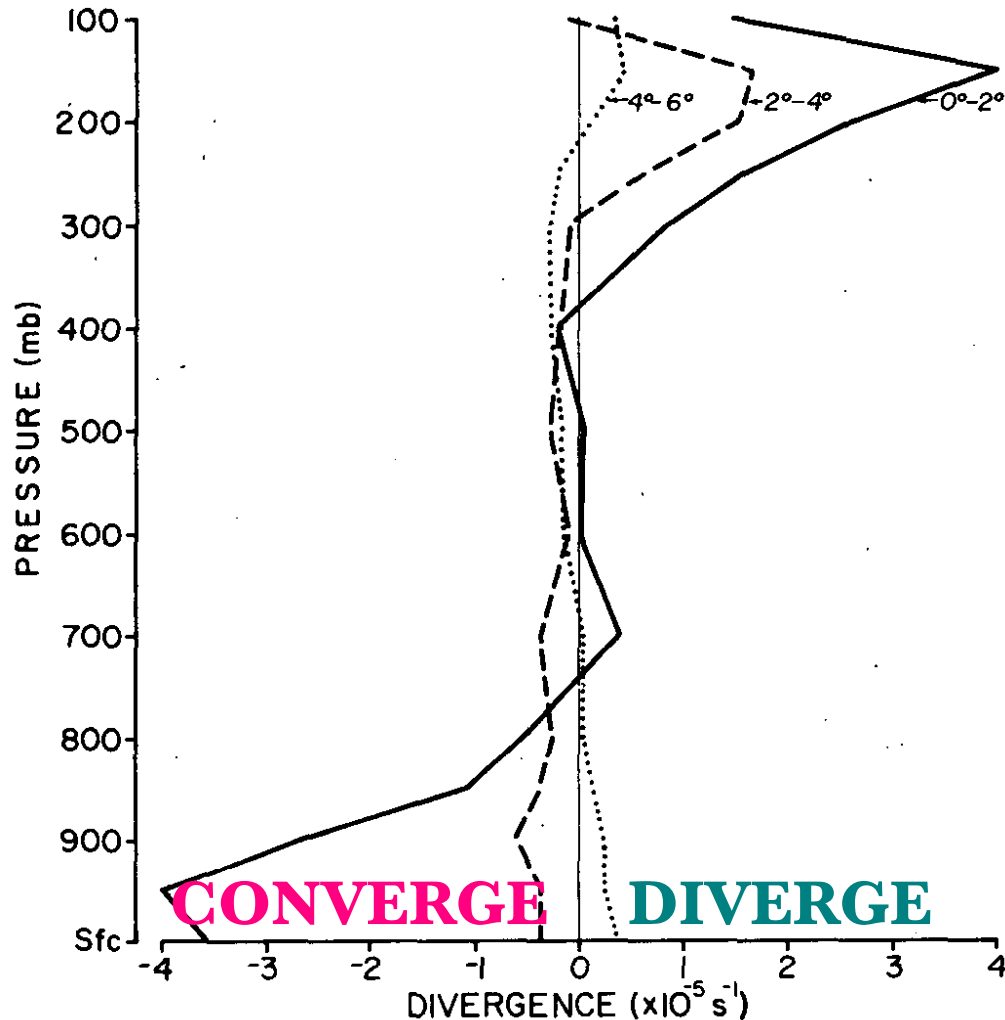


FIG. 18. Divergence for mean steady-state typhoon for the 0-2°, 2-4° and 4-6° regions.

Strong convergence is confined to the lowest levels and inner most radii

The 4°-6° ring shows convergence over divergence indicative of subsidence

# Tropical cyclone structure: Vertical motion

Weak **SUBSIDENCE**  
occurs 4°–6° from the  
center, at the edge of  
the cirrus cloud shield

Mean **UPWARD**  
motion is observed  
inside 4° through the  
depth of the  
troposphere

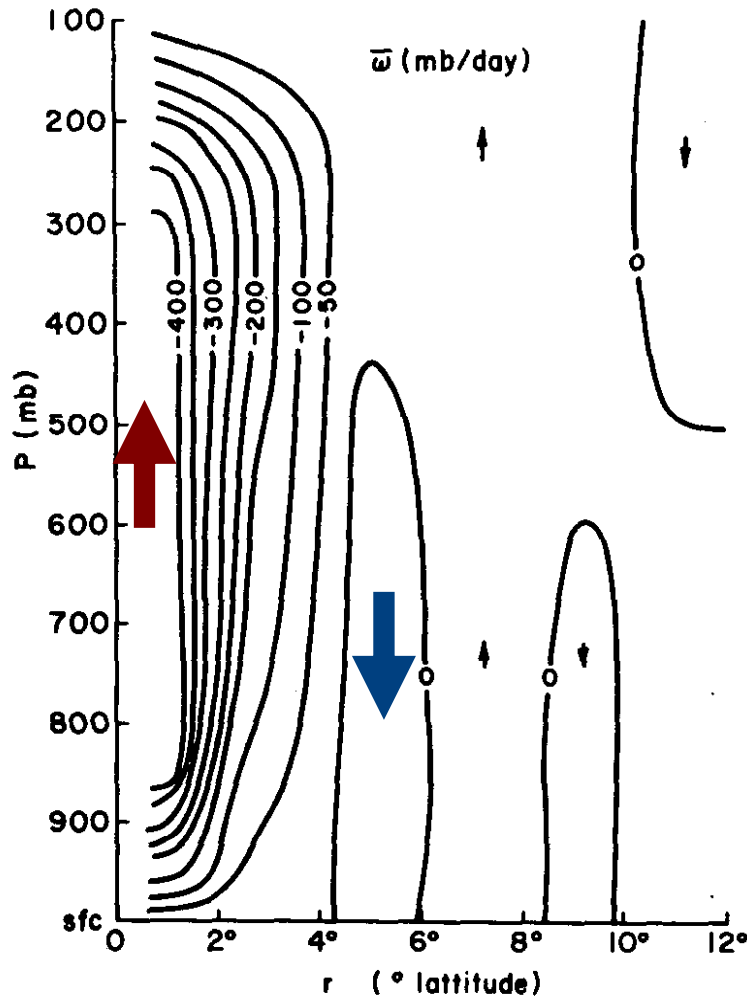


FIG. 19. Vertical motion (mb day<sup>-1</sup>) from 0.7–12°  
for mean steady-state typhoon.