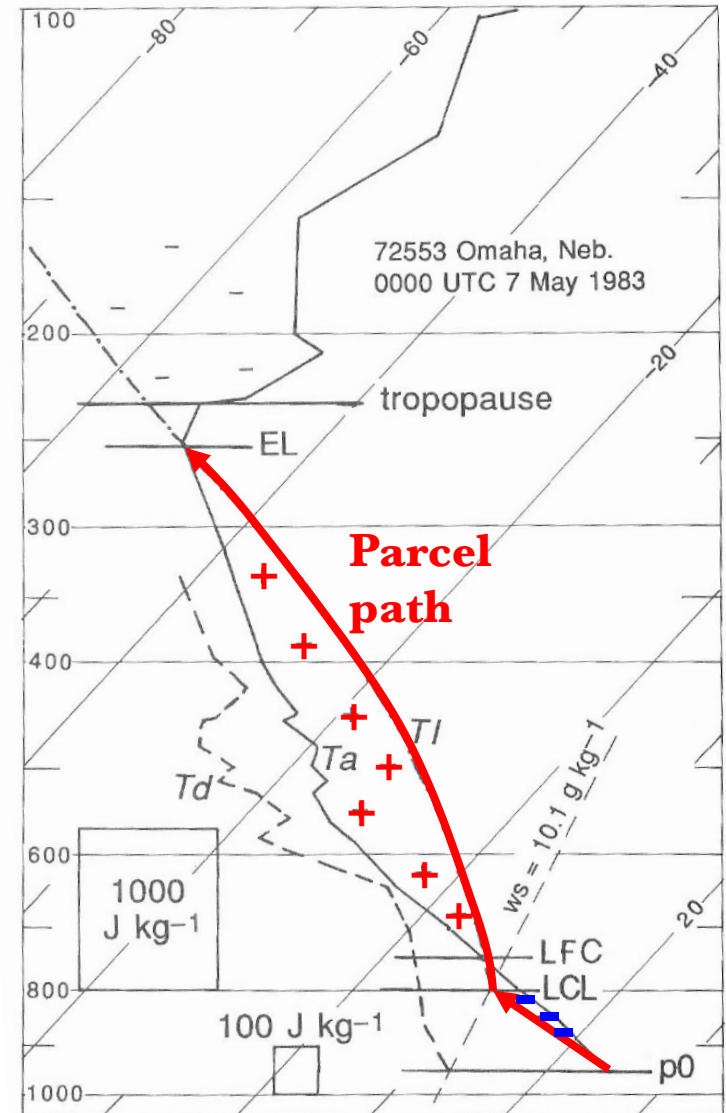
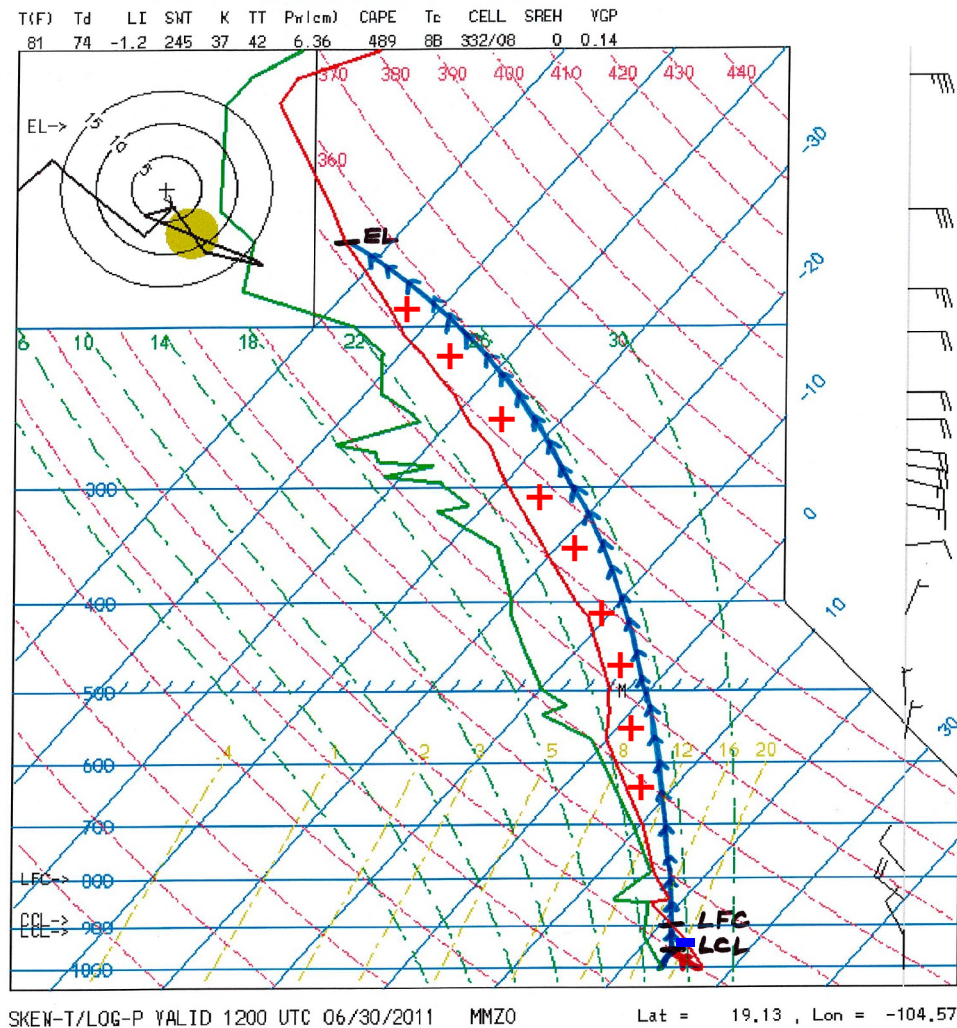


Midlatitude versus tropical convection: Skew T-log p diagrams



Lapse rates and stability

$$\gamma < \Gamma_m$$

Absolutely stable

$$\gamma = \Gamma_m$$

Saturated neutral

$$\Gamma_m < \gamma < \Gamma_d$$

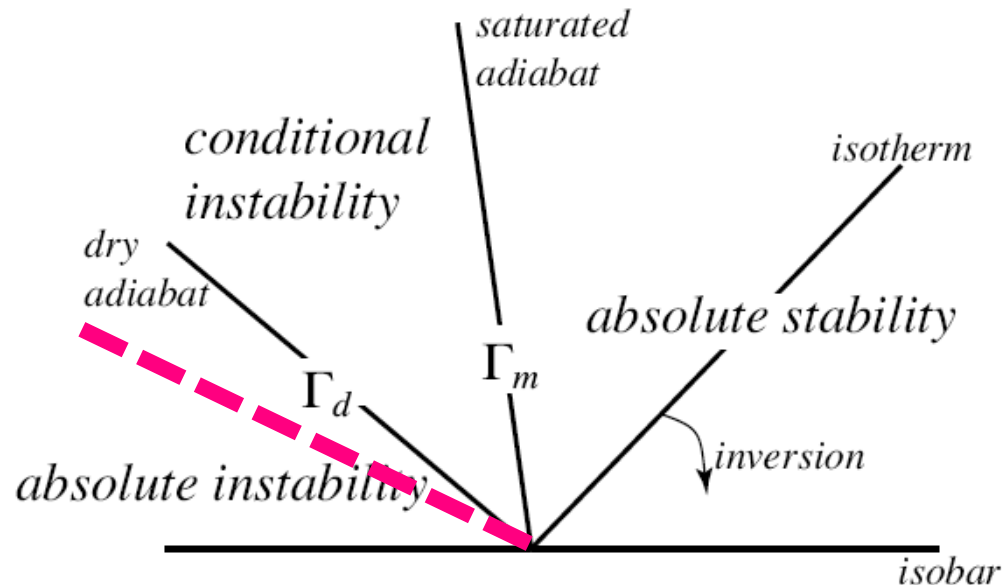
Conditionally unstable

$$\gamma = \Gamma_d$$

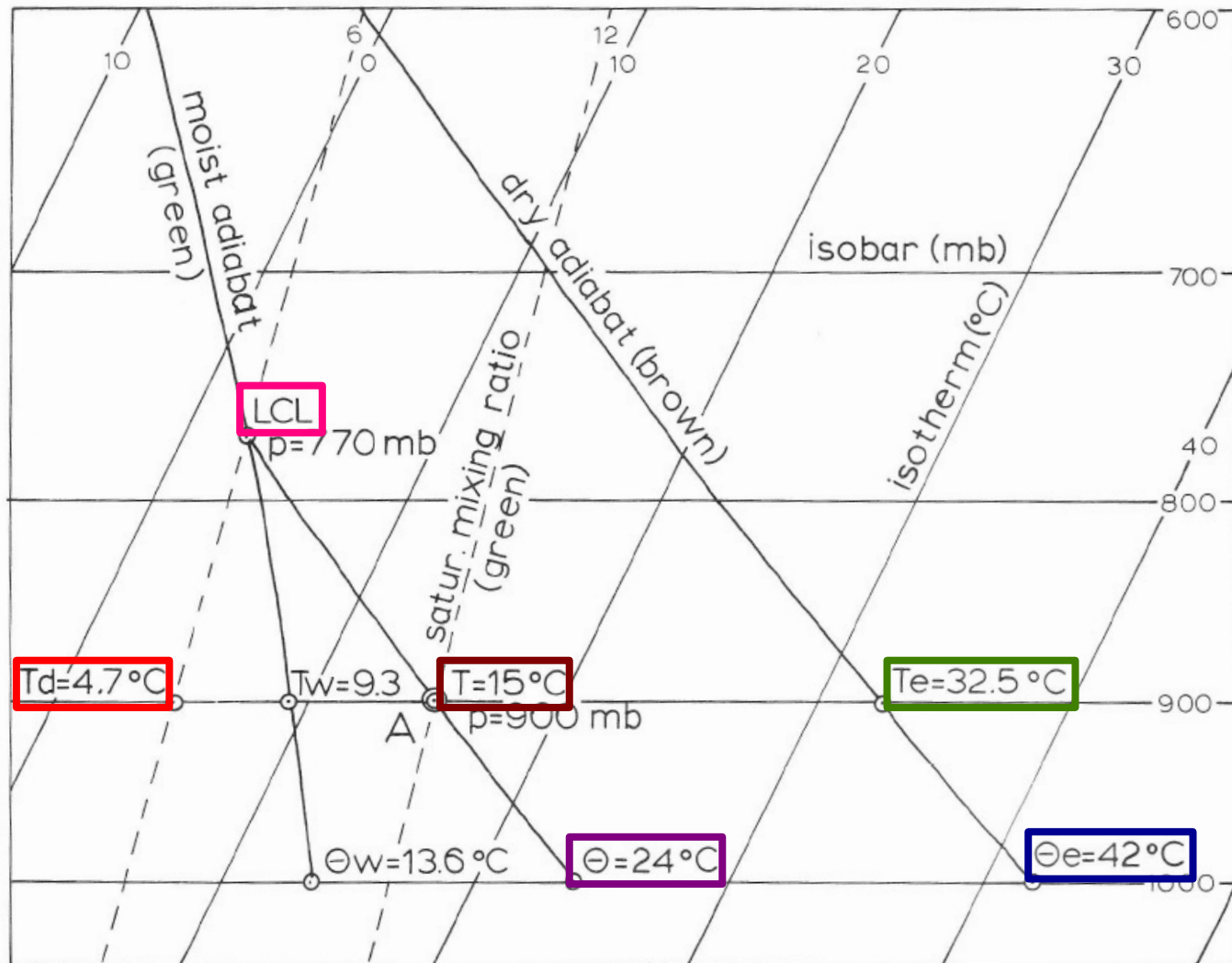
Dry neutral

$$\gamma > \Gamma_d$$

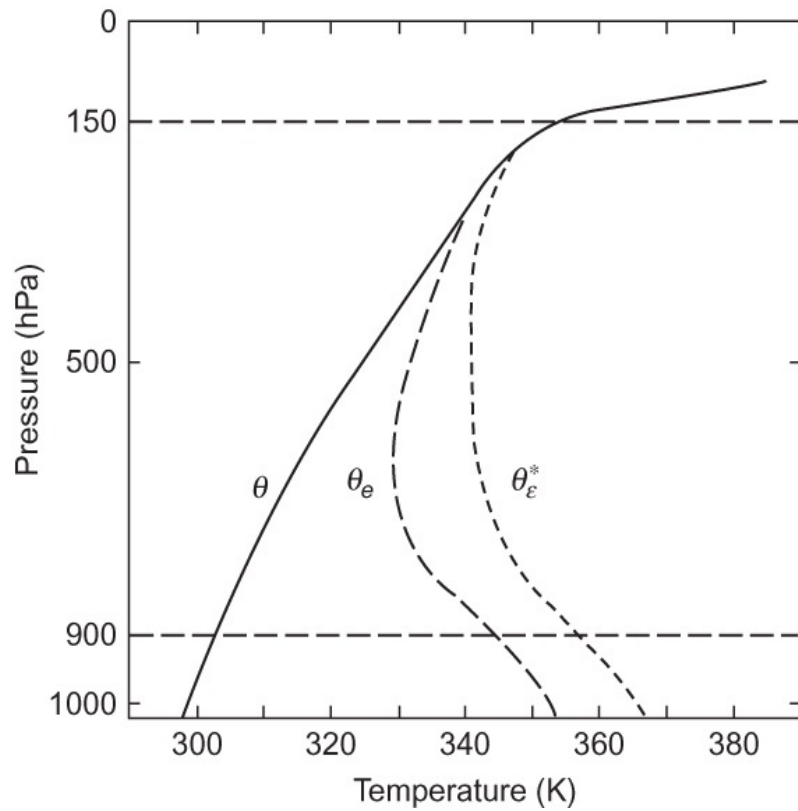
Absolutely unstable



Finding Θ s on a skew T-log p diagram

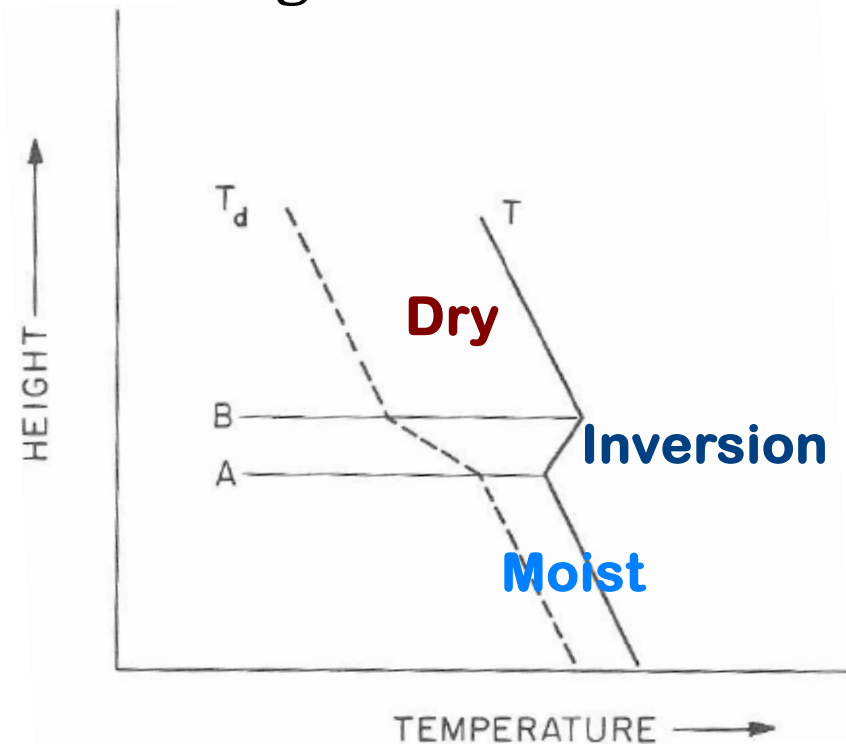


Typical tropical thermodynamic profiles



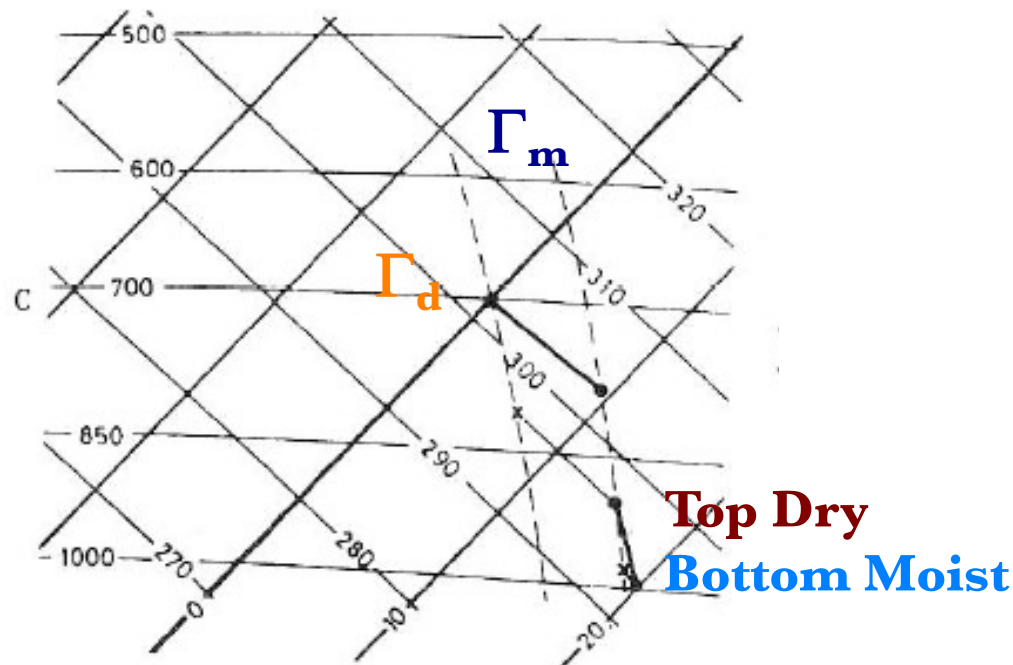
Potential instability

- Potential instability is related to the lifting of layers, instead of individual air parcels, and the vertical stratification of moisture.
- The dew point decreases rapidly with height above the inversion layer AB.
- Suppose layer AB is lifted. The air at A will reach its LCL almost immediately and cool moist adiabatically above that, while the air at B will need to cool dry adiabatically though a deep layer before saturation.

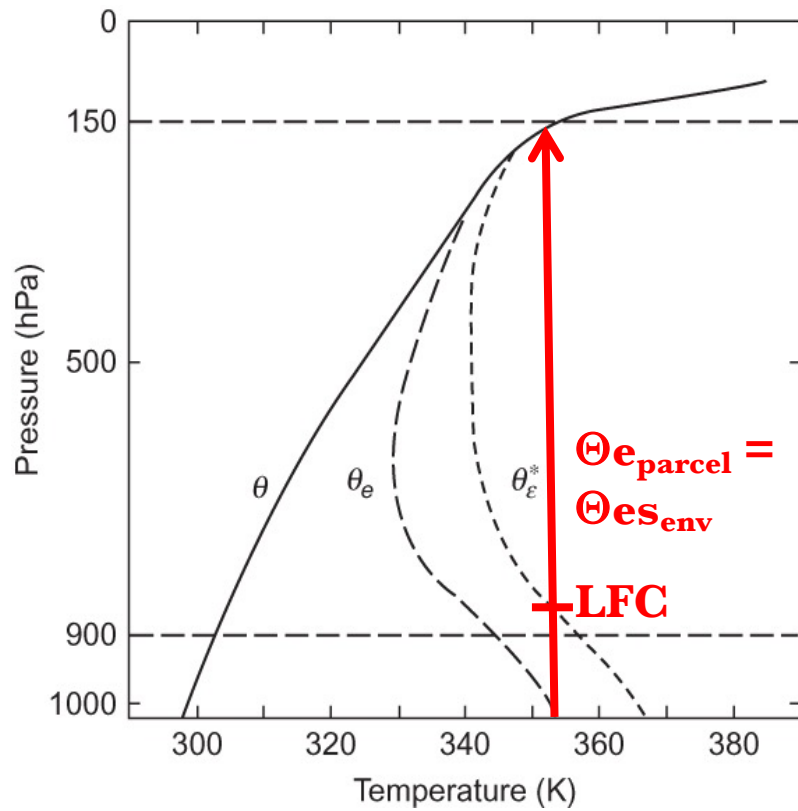


Potential instability

- Therefore, the top part of the layer cools much more rapidly than the bottom part & the lapse rate destabilizes!
- Sufficient lifting may cause the layer to become conditionally unstable, even if the entire sounding is absolutely stable to begin with!

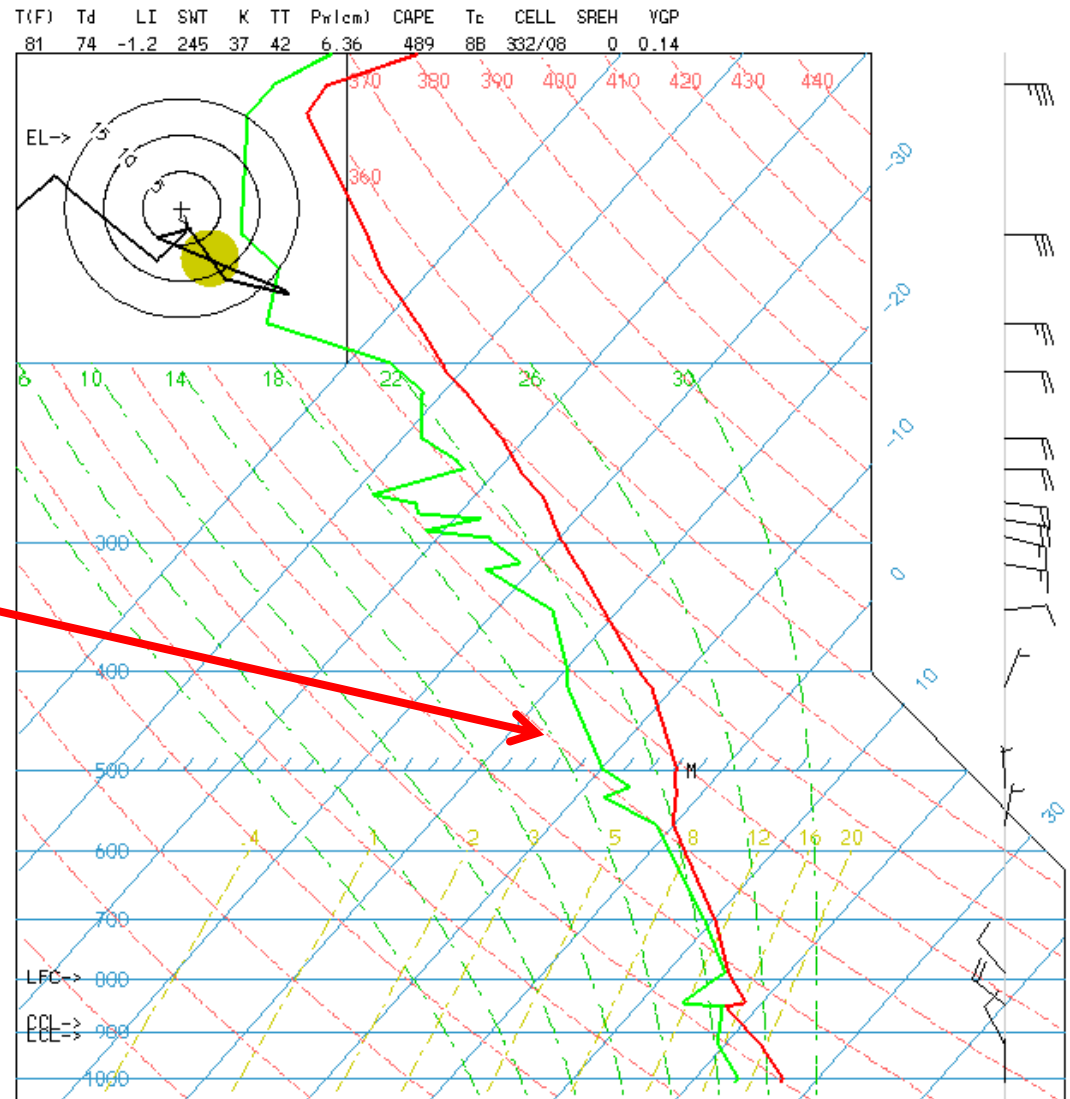


Typical tropical thermodynamic profiles



The LFC, Θ_e , and Θ_{es}

Skew-T/Log-P diagrams created from rawinsonde data



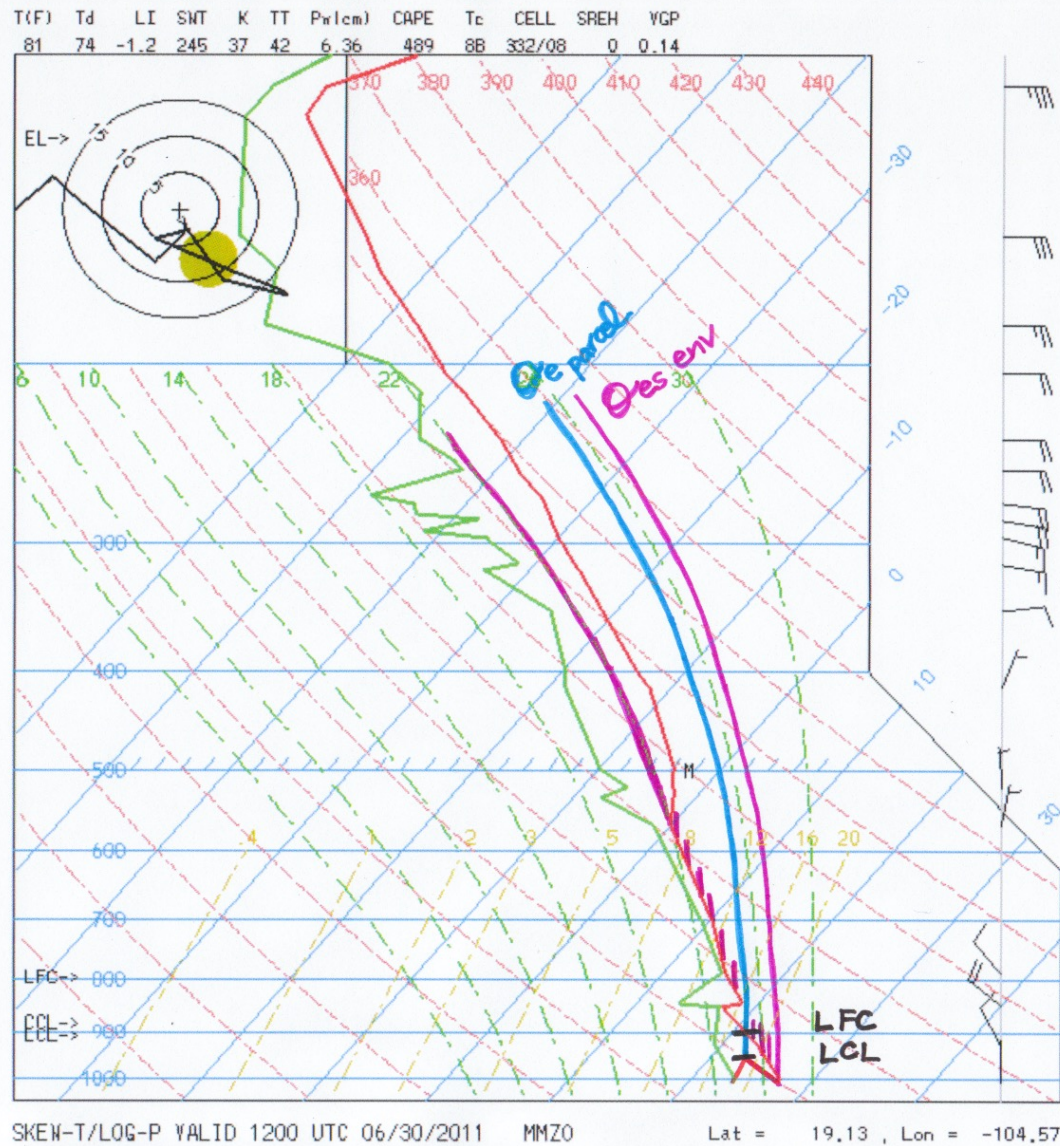
SKEN-T/LOG-P VALID 1200 UTC 06/30/2011

MMZO

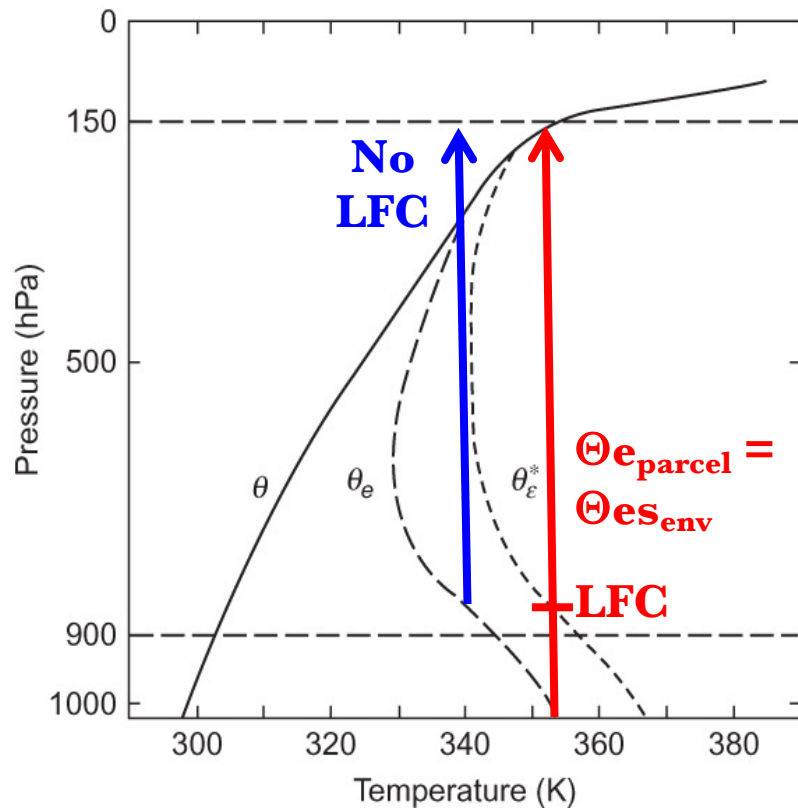
Lat = 19.13 , Lon = -104.57

Can we see that
 $\Theta_{e_{\text{parcel}}} = \Theta_{e_{\text{env}}}$
 at the LFC from a
 “normal”
 sounding?

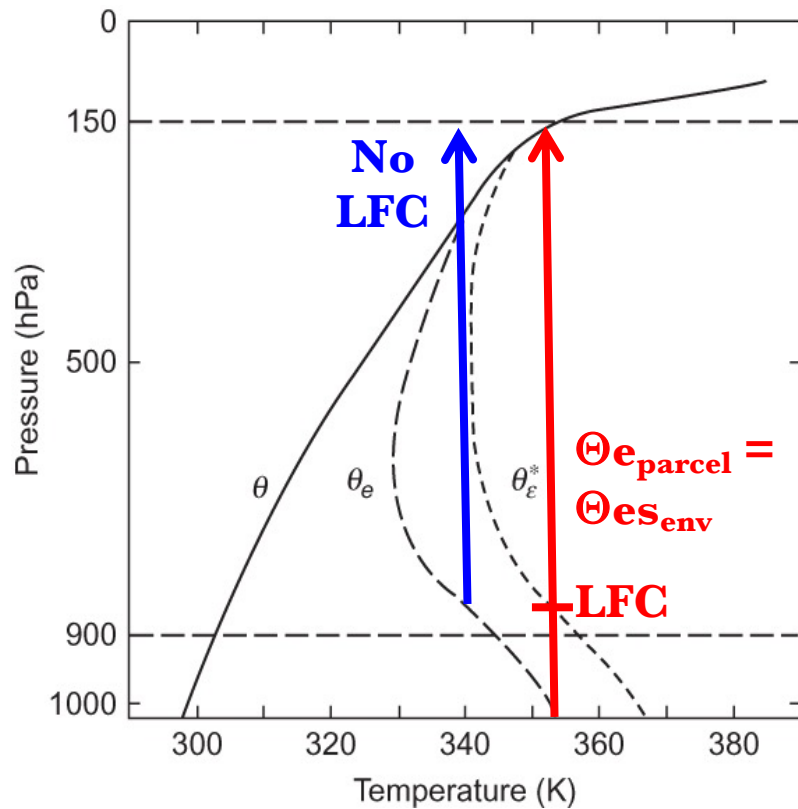
The LFC, Θ_e , and Θ_{es}



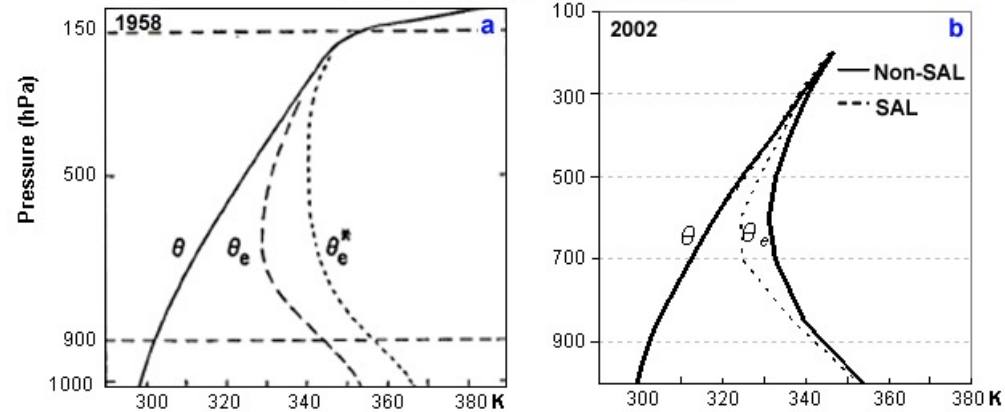
Typical tropical thermodynamic profiles



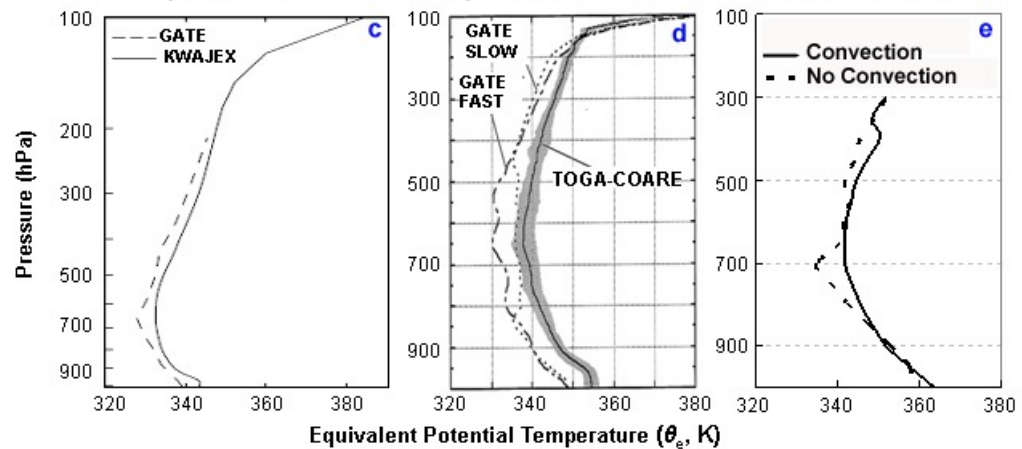
Typical tropical thermodynamic profiles



Average Profiles of Equivalent Potential Temperature and Potential Temperature
Caribbean and Tropical North Atlantic



Equatorial West Pacific and Tropical East Atlantic



a. Ooyama (1969)

d. LeMone et al. (1998)

b. Data courtesy of Jason Dunion

e. Data from Grossman and Durran (1984)

c. Cetrone and Houze (2006)

How convection changes the thermodynamic profile

Miller and Betts (1977)

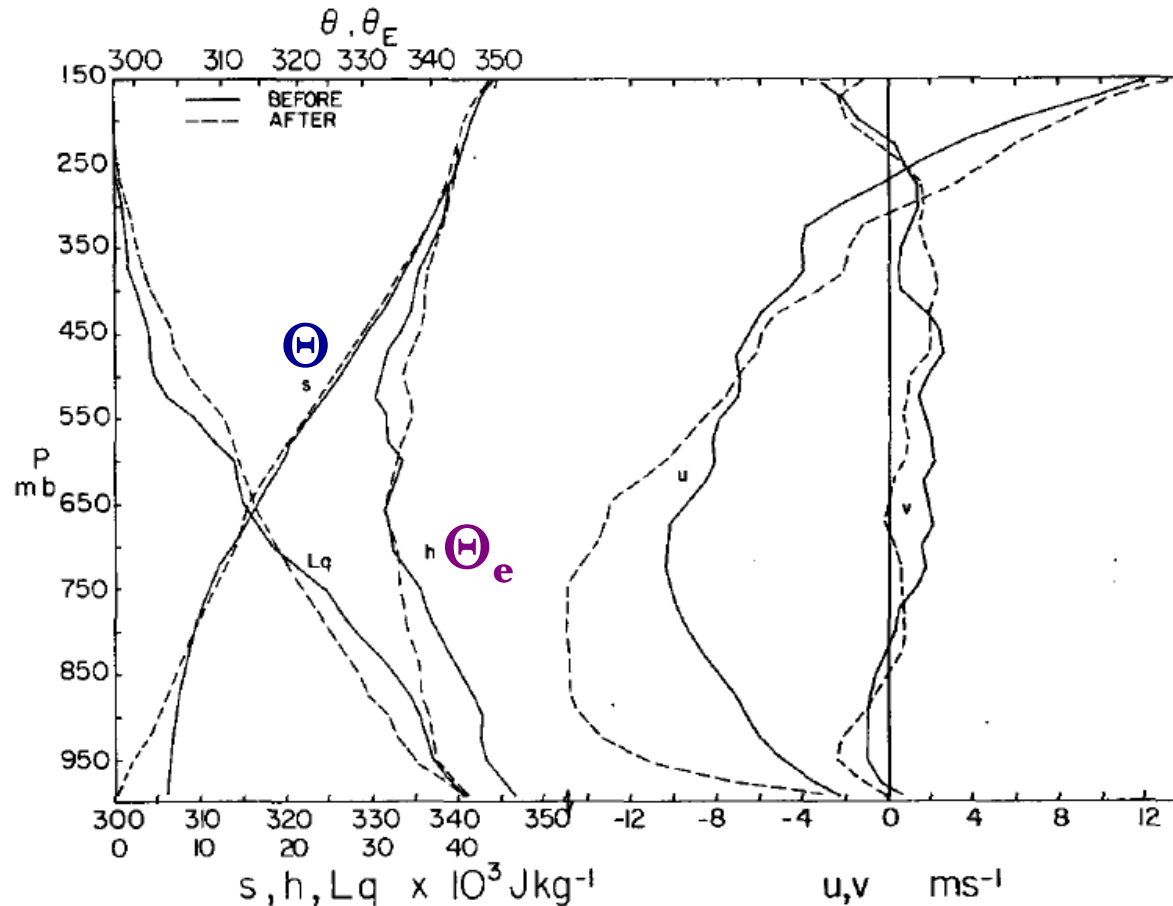


FIG. 1. Mean tropospheric profiles of static energy s , moist static energy h , water vapor (as latent energy Lq) and wind components u , v before and after the passage of a traveling convective storm over the rawinsonde site. The upper scale shows potential temperatures θ , θ_E corresponding (with slight approximation) to s and h . The averages include all soundings in Table 1 except 203 and 204.

How convection changes the thermodynamic profile

Miller and Betts (1977)

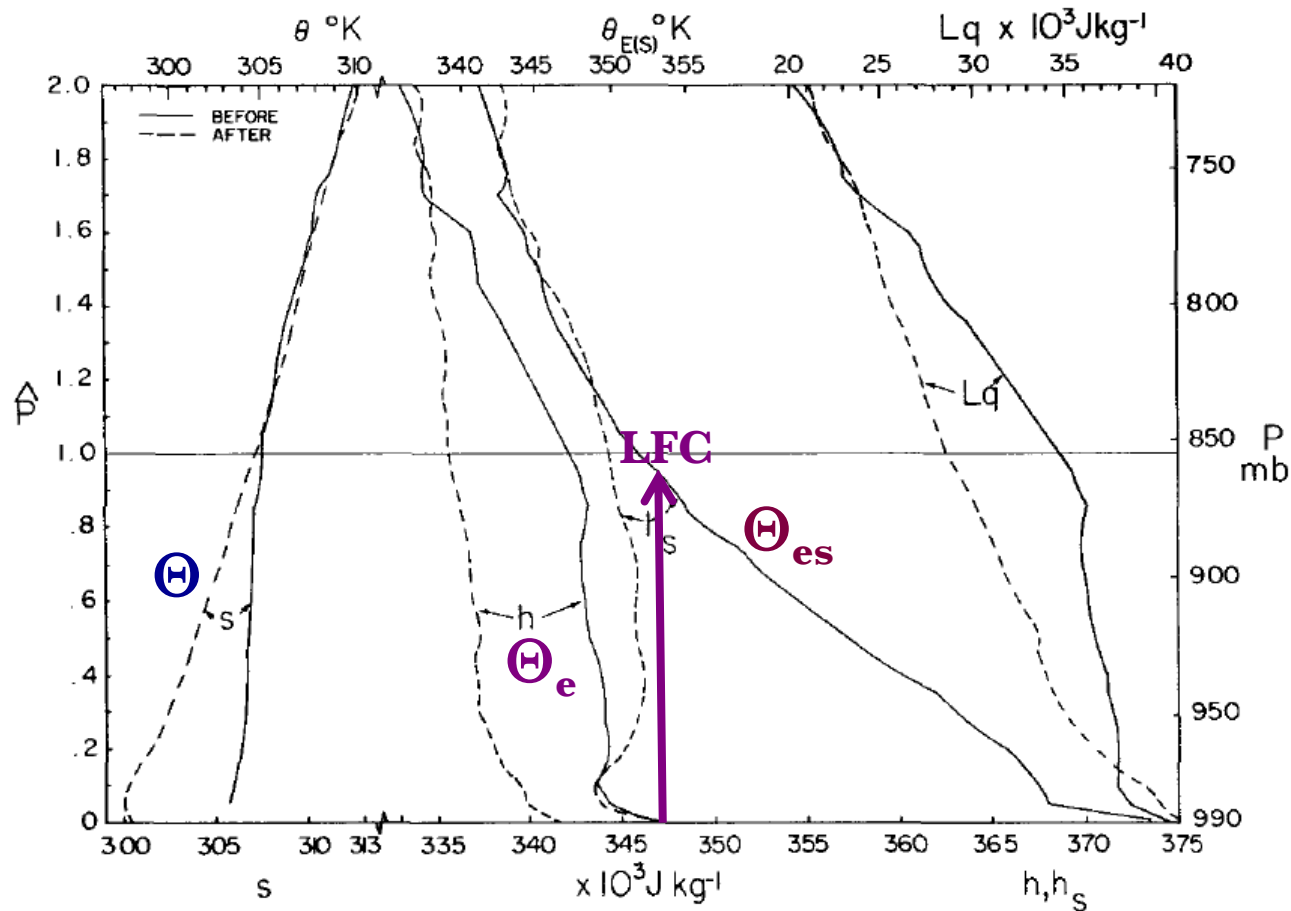


FIG. 4. Mean thermodynamic profiles before and after the passage of a traveling storm; $\hat{p}=1$ corresponds to the mean cloud base before the storm. Potential temperature (θ , θ_E , θ_{ES}) corresponding to static energies s , h , h_s are shown on the upper scale. The two averages include all soundings in Table 1.