**4. Case Studies**

4.1 9–14 January 1982 CAO

*4.1.1 Case Overview*

During 9–14 January 1982, a CAO impacted a widespread area of central and eastern North America, leading to significant socioeconomic impacts. According to NOAA NCEI, the 9–14 January 1982 CAO (hereafter referred to as the January 1982 CAO) contributed to an estimated cost of 1.7 billion dollars, after consumer price index adjustment, and contributed to 85 deaths (<https://www.ncdc.noaa.gov/billions/>). The January 1982 CAO qualifies as a CAO for the Northeast, Central, South, and Southeast regions (see Fig. 2.1 for map of regions) and a CAO that is linked to a cold pool associated with a TPV for all of these regions as well. Given the widespread and significant socioeconomic impacts of the January 1982 CAO, this CAO and its linkages to the cold pool associated with a TPV will be examined in detail.

*4.1.2 TPV and Cold Pool Track and Intensity*

Figure 4.1 shows the tracks of the TPV and cold pool of interest for the January 1982 CAO. The TPV forms over Baffin Bay at 0600 UTC 15 December 1981. The TPV moves over Greenland between 17 and 21 December 1981 and then spends much of its lifetime meandering over northern and central Canada during late December 1981 and early January 1982. The TPV moves equatorward into the U.S. and then eastward across the U.S. during 8–11 January 1982, before moving quickly northeastward toward southern Greenland through 12 January 1982. The TPV undergoes lysis off the southeast coast of Greenland at 0000 UTC 13 January 1982. Overall, the TPV has a long lifetime of ~29 days. The cold pool forms over the northern coast of Greenland at 1800 UTC 20 December 1981, several days after the TPV forms (Fig. 4.1). The cold pool has a similar track to the TPV throughout much of its lifetime, spending much of its lifetime meandering over northern and central Canada before moving equatorward into the U.S. between 8 and 10 January 1982. However, unlike the TPV, the cold pool meanders over eastern North America during 11–12 January 1982 before finally moving off the east coast of North America. The cold pool undergoes lysis at 1800 UTC 13 January 1982 east of Newfoundland. Overall, the cold pool also has a long lifetime (~24 days). The large spatial overlap and temporal coincidence of the TPV and cold pool tracks suggest that the TPV and cold pool are dynamically linked. Furthermore, the movement of the TPV and cold pool over and near the U.S. during 9–11 January 1982, when the January 1982 CAO is occurring, suggests that these features play an important role in the development of the January 1982 CAO.

To understand the intensity evolution of the TPV and cold pool during their lifetimes, Fig. 4.2 shows a time series of the minimum DT potential temperature of the TPV and the minimum 1000–500-hPa thickness of the cold pool. In general, ignoring rapid fluctuations of TPV intensity as evident by large spikes in the time series of the minimum DT potential temperature of the TPV (e.g. during 29 December 1981), which may be numerical artifacts, the TPV gradually intensifies during late December 1981, with the minimum DT potential temperature of the TPV decreasing from ~290 K on 15 December 1981 to close to 260 K on 28 December 1981. During early January 1982, the minimum DT potential temperature of the TPV is very low (~255 K) and decreases to ~249 K on 9 January 1982. However, during 10–12 January 1982, the TPV rapidly weakens as the minimum DT potential temperature of the TPV increases to ~284 K on 13 January 1982. The overall gradual intensification of the TPV suggests that longwave radiative cooling likely contributes importantly to the intensification of the TPV, especially given that Cavallo and Hakim (2009, 2010, 2012, 2013) have shown that longwave radiative cooling can play a critical role in the intensification of TPVs. In addition, as will be shown in section 4.1.3, an extratropical cyclone (EC) undergoes rapid cyclogenesis off the east coast of North America during 10–11 January 1982, and latent heat release associated with this EC may contribute to the weakening of the TPV during this time period.

About two days after the cold pool forms, similarly as the TPV, the cold pool gradually intensifies during late December 1981 and early January 1982, with the minimum 1000–500-hPa thickness of the cold pool decreasing from ~488 dam on 20 December 1981 to ~463 dam on 7 January 1982 (Fig. 4.2). Between 7 and 9 January 1982, the cold pool slowly weakens as it moves toward southern Canada and into the U.S., but rapidly weakens as the TPV does during 10–13 January 1982. Longwave radiative cooling (e.g., Curry 1983; Emanuel 2008) may contribute to the gradual intensification of the cold pool during late December 1981 and early January 1982. Some of the weakening of the cold pool during 7–13 January 1982 may be related to sensible heat fluxes from the surface, especially as the cold pool moves over the relatively warmer land of the U.S compared to that of Canada, considering that there may be less snow cover and a thinner snowpack over the U.S. compared to Canada. Latent heat release and warm air advection associated with the aforementioned EC that will be discussed in section 4.1.3 may also contribute to the weakening of the cold pool during 10–13 January 1982. Overall, the similar patterns in intensity changes for the TPV and cold pool further suggest that the TPV and cold pool may be dynamically linked.

*4.1.3 Synoptic Evolution of TPV, Cold Pool, and CAO*

Ridge amplification over the eastern North Pacific and western North America likely plays an important role in the equatorward transport of the TPV and cold pool. At 0000 UTC 5 January 1982, an EC is intensifying over the central North Pacific in the left exit region of the North Pacific jet stream, downstream of a short-wave trough (Figs. 4.3a,b). Afterward, through 0000 UTC 8 January 1982, the EC moves northeastward toward the Gulf of Alaska and the aforementioned trough over the central North Pacific moves eastward and becomes negatively tilted, with a corridor of precipitable water values in excess of 25 mm becoming established from the subtropics poleward toward the Gulf of Alaska (Figs. 4.3c–h). Warm air advection, as suggested by the nearly perpendicular orientation of the SLP contours to the 1000–500-hPa thickness contours downstream of the EC during 5–8 January 1982 (Figs. 4.3b,d,f,h), likely supports the ridge amplification that occurs downstream of the EC over the eastern North Pacific into western North America (Figs. 4.3a,c,e,g). In addition, Figs. 4.4a,b show that at 0000 UTC 6 and 0000 UTC 7 January 1982, associated with the EC over the central North Pacific and with the corridor of precipitable water values in excess of 25 mm is a widespread region of mid-level ascent and concomitantly, diabatically-driven upper-tropospheric divergent outflow and associated negative PV advection by the irrotational wind, which likely also supports the aforementioned ridge amplification. Furthermore, as the ridge comes into closer proximity with the TPV and associated cold pool located over western North America, the thermal gradient over the Gulf of Alaska and southwestern Canada strengthens, supporting the development and intensification of a jet streak over this region during 7–8 January 1982 (Figs. 4.3e–h).

Between 0000 UTC 8 and 0000 UTC 9 January 1982, as the ridge continues to build eastward and comes into closer proximity with the TPV and associated cold pool, the jet streak between the ridge and TPV strengthens as TPV–jet interaction begins to occur (Figs. 4.5a–d). Between 0000 UTC 9 and 0000 UTC 10 January, as TPV–jet interaction occurs, the TPV and associated cold pool move equatorward into the U.S. and a strong surface anticyclone over northwestern North America rapidly strengthens and expands southeastward in the left entrance region of the jet streak into the central U.S., just east of the Rocky Mountains (Figs. 4.5c–f). Strong cold air advection over the Great Plains as suggested by Fig. 4.5f, along with an expected terrain-tied northerly component of motion (e.g., Colle and Mass 1995) on the east side of the Rockies where there is a strong SLP gradient associated with the surface anticyclone likely help allow the cold air associated with the cold pool to spread far away from the core of the cold pool, illustrating that the cold pool associated with the TPV has a geographically widespread impact.

By 0000 UTC 11 January 1982, the TPV and cold pool have moved eastward toward the U.S east coast, and cold air associated with the cold pool has overspread the entire eastern U.S. (Figs. 4.5g,h). As discussed in section 4.1.2, the TPV and cold pool rapidly weaken during 10–11 January 1982. Between 0000 UTC 10 and 0000 UTC 11 January 1982, the TPV becomes deformed in shape as suggested by the TPV becoming less isotropic during a time of TPV–jet interaction when the TPV is located on the cyclonic shear side of the jet streak (Figs. 4.5e,g). Horizontal shear on the cyclonic shear side of the jet streak in combination with confluent flow in the entrance region of the jet streak and diffluent flow in the exit region of the jet streak, as suggested in Figs. 4.5e,g may contribute to the deformation of the TPV, which may contribute to the weakening of the TPV. In addition, rapid cyclogenesis of an EC occurs in the left exit region of the jet streak between 0000 UTC 10 and 0000 UTC 11 January 1982 (Figs. 4.5f,h). Latent heat release associated with widespread ascent in the vicinity of the EC during 0000 UTC 10 and 0000 UTC 11 January 1982 (Figs. 4.6a,b) may also contribute to the weakening of the TPV and cold pool, and warm air advection that is suggested over Labrador and eastern Quebec at 0000 UTC 11 January 1982 (Fig. 4.5h) may further contribute to the weakening of the cold pool.

Overall, the fact that the TPV and cold pool move directly together into the U.S. during the time of the January 1982 CAO suggest again that the TPV and cold pool are dynamically linked and that the TPV and cold pool play an important role in the development of the January 1982 CAO. In section 4.1.4, the evolution of the three-dimensional structure of the TPV and cold pool will be examined to gain more insight on how these features evolve together over time. In addition, the relationship between TPV–jet interaction and the development of the strong surface anticyclone important for the evolution of the CAO will be explored further in section 4.1.5.

*4.1.4 Three-dimensional Structure of TPV and Cold Pool*

Figure 4.7 shows a meridional cross section (AA’) transecting the TPV early in its lifecycle at 1200 UTC 16 December 1981, when the TPV is relatively weak and located over Baffin Bay. Using the position of the 2-PVU surface in the cross section as a proxy for the position of the DT, the TPV only extends downward to ~450 hPa, which is not too far beneath the background DT. Overall, the TPV is a mesoscale feature embedded in a broad region of lowered DT and relatively low DT potential temperature air. Although the TPV has not been objectively identified to be associated with a cold pool at this time, there is a broad region of cold air characterized by 1000–500-hPa thickness values below 500 dam where the TPV and broad region of relatively low DT potential temperature air is located, suggesting that the TPV is still associated with cold air beneath it throughout the troposphere. The cross section indeed shows an upward bowing of isentropes beneath the TPV, with potential temperature values less than 268 K in the lower troposphere beneath the TPV.

As shown in Fig. 4.2, the TPV strengthens throughout late December 1981 and is intense during early January 1982. Also, the cold pool intensifies during late December 1981 and early January 1982 after it forms. Figure 4.8 shows a meridional cross section (BB’) transecting the TPV and cold pool at 1200 UTC 2 January 1982, when both features are intense, highly isotropic in shape, and vertically aligned. Compared to early in its lifecycle at 1200 UTC 16 December 1982 (Fig. 4.7), the cross section in Fig. 4.8 at 1200 UTC 2 January 1982 shows that the TPV has become a much better defined, larger, and deeper feature, with the TPV now extending downward to ~650 hPa. In addition, there is substantial upward bowing of the isentropes within and beneath the TPV throughout the depth of the troposphere, indicative of the well-defined cold pool that is collocated with the TPV and the tropospheric-deep influence of the TPV. Potential temperature values are less than 240 K near the surface beneath the TPV, indicative of the very cold air associated with the cold pool. Overall, as the TPV has become better defined and deeper, there has become a more robust upward bowing of the isentropes beneath the TPV and concomitantly a better defined and intense cold pool, illustrating that the TPV and cold pool are likely dynamically linked. It is also evident in the cross section that near the surface, isentropes spread outward, far away from the center of the TPV and cold pool, indicative of surface-based Arctic air spreading far from the center of the TPV and cold pool. In fact, especially to the south and north of the TPV, there is a very strong near-surface vertical potential temperature gradient, with near-surface PV values in excess of 4 PVU north of the TPV. As shown by Emanuel (2008), Arctic air can be characterized by relatively high values of PV. The strong near-surface vertical potential temperature gradient also suggests that a steep temperature inversion is in place, which is likely related to longwave radiative cooling (e.g., Curry 1983).

Figure 4.9 shows a meridional cross section (CC’) transecting the TPV and cold pool at 0000 UTC 10 January 1982 when both features have moved into the northern U.S. and are contributing to CAO development, as well as when the TPV is interacting with a jet streak. Compared to at 1200 UTC 2 January 1982 (Fig. 4.8), the TPV has is even deeper at 0000 UTC 10 January 1982, now extending downward to ~750 hPa (Fig. 4.9). A PV wall associated with a large horizontal PV gradient is evident throughout much of the troposphere between the TPV and warm air to its south, concomitant with the jet streak. There continues to be a very notable upward bowing of isentropes throughout the troposphere within and beneath the TPV, illustrative of the cold pool associated with the TPV. The surface-based Arctic air continues to spread far away from the center of the TPV and cold pool, with the leading edge of the Arctic air associated with the Arctic front likely being located south of 40°N where there is a strong surface horizontal potential temperature gradient. Furthermore, within the Arctic air, the boundary layer appears well mixed and neutrally stable given the nearly vertical orientation of the isentropes within the boundary layer. A well-mixed boundary layer was also shown by Shapiro et al. (1987; Fig. 1.7 in this thesis) in a cross section transecting the January 1985 “polar vortex” feature (or TPV as will be discussed in section 4.2), and they suggest that this well-mixed boundary layer may result from diabatic heating induced by the flow of Arctic air over the relatively warm land surface. Overall, the cross sections in especially Figs. 4.8 and 4.9 illustrate that the influence of the TPV extends throughout the depth of the troposphere and over a widespread geographical area, and also illustrate that the TPV and associated cold pool plays a crucial role in CAO development.

*4.1.5 Q-vector Diagnosis*

The concomitant occurrence of the interaction of the TPV with a jet streak over western North America and the rapid strengthening and southeastward expansion of the strong surface anticyclone over western North America in the left entrance region of the jet streak during 8–10 January 1982 (Figs. 4.5a–f) suggests that TPV–jet interaction may play an important role in the rapid strengthening and expansion of the strong surface anticyclone. Figure 4.10 shows plots of 600–400-hPa **Q**n and **Q**s and their associated forcings for vertical motion from 1200 UTC 8 to 0000 UTC 10 January 1982. At 1200 UTC 8 January 1982, between the ridge and the TPV, there is a region of divergence of **Q**nand **Q**n forcing for descent collocated with and to the south and east of the ~1047 hPa surface anticyclone over western North America (Fig. 4.10a). The orientation of the Qn vectorsfrom cold to warm air suggests that upper-level frontogenesis may be occurring, which may support a strengthening of the jet streak over western North America and concomitantly a strengthening of the thermally direct ageostrophic circulation in the entrance region of the jet streak. A strengthened thermally direct ageostrophic circulation in the entrance region of the jet streak may be associated with increased descent and thus increased forcing for anticyclogenesis in the left entrance region of the jet streak. There is also a small region of **Q**s forcing for descent over northern Alberta on the downstream side of the ridge and to the south and east of the surface anticyclone (Fig. 4.10b). The position of the regions of **Q**n and **Q**s forcing for descent over and to the south and east of the surface anticyclone suggests that the surface anticyclone may strengthen and expand southeastward.

By 0000 UTC 9 January 1982, the maximum SLP of the surface anticyclone remains ~1047 hPa, but the surface anticyclone has expanded southeastward over the last 12 h (compare Figs. 4.10a,c). As the TPV has come into close proximity with the jet streak between the TPV and ridge (Fig. 4.5c), the magnitude of **Q**n and **Q**n forcing for descent has concomitantly increased between the TPV and the ridge, in the left entrance region of the jet streak, over and to the south and east of the surface anticyclone (Fig. 4.10c). Additionally, the crossing of the geopotential height contours and potential temperature contours in the region of **Q**n forcing for descent, on the cyclonic shear side of the jet streak suggests that geostrophic cold air advection may be occurring. Shapiro (1981) has shown that geostrophic cold air advection in cyclonic shear may provide a favorable environment for upper-level frontogenesis, which is suggested by the orientation of the Qn vectors from cold to warm air over western North America. Furthermore, by comparing Figs. 4.10b,d, the region of **Q**s forcing for descent has increased in size and magnitude between 1200 UTC 8 and 0000 UTC 9 January 1982, with a maximum of **Q**s forcing for descent located upstream of the TPV, over and to the south and east of the surface anticyclone at 0000 UTC 9 January 1982. This region of **Q**s forcing for descent may be related the large curvature of the flow associated with TPV, as noted in the geopotential height contours. The maximum of **Q**s forcing for descent being located upstream of the TPV is also anticipated given that Lang and Martin (2010) have shown that **Q**s forcing for descent is expected to be located upshear of an isolated vertical vorticity maximum. Overall, the position of the regions of **Q**n and **Q**s forcing for descent over and to the south and east of the surface anticyclone suggests that the surface anticyclone may strengthen and continue to build southeastward.

Between 0000 UTC and 1200 UTC 9 January 1982, the surface anticyclone does rapidly strengthen from ~1047 hPa to ~1054 hPa**,** or by ~7 hPa in 12 h, and does build southward and especially eastward (compare Figs. 4.10c,e). At 1200 UTC 9 January 1982, as the TPV is interacting with the jet streak, **Q**n and **Q**s and their associated forcings for descent have continued to increase, with a maximum of both **Q**n and **Q**s forcing for descent located upstream of the TPV, over and to the south and east of the surface anticyclone, which would suggest that the surface anticyclone may strengthen further and build southeastward (Figs. 4.10e,f). Between 1200 UTC 9 and 0000 UTC 10 January 1982, the surface anticyclone indeed strengthens from ~1054 hPa to ~1060 hPa**,** or by ~6 hPa in 12 h, and builds southeastward, and their continues to be similar patterns of **Q**n and **Q**s and their associated forcings for descent at 0000 UTC 10 January 1982 as 12 h earlier (compare Figs. 4.10e,f with Figs. 4.10g,h). Overall, the concomitant occurrence of the increase in magnitude of **Q**n and **Q**s and their associated forcings for descent during TPV–jet interaction with the rapid strengthening and expansion of the surface anticyclone suggests that TPV–jet interaction may play an important role in the rapid strengthening and expansion of the surface anticyclone. Given the important role this surface anticyclone plays in the development of the January 1982 CAO, TPV–jet interaction may play an important role in CAO development.

4.2 19–24 January 1985 CAO

*4.2.1 Case Overview*

During 19–24 January 1985, a CAO impacted a widespread area of central and eastern North America, leading to significant socioeconomic impacts. According to NOAA NCEI, the 19–24 January 1985 CAO (hereafter referred to as the January 1985 CAO) contributed to an estimated cost of 1.9 billion dollars, after consumer price index adjustment, and contributed to 150 deaths (<https://www.ncdc.noaa.gov/billions/>). As with the January 1982 CAO, the January 1985 CAO qualifies as a CAO for the Northeast, Central, South, and Southeast regions (see Fig. 2.1 for map of regions) and a CAO that is linked to a cold pool associated with a TPV for all of these regions as well. Given the widespread and significant socioeconomic impacts of the January 1985 CAO, this CAO and its linkages to the cold pool associated with a TPV will be examined in detail. Furthermore, this research confirms that the “polar vortex” examined by Shapiro et al. (1987) to play an important role in the development of the January 1985 CAO, as discussed in Chapter 1, qualifies as a TPV associated with a cold pool.

*4.2.2 TPV and Cold Pool Track and Intensity*

Figure 4.11 shows the tracks of the cold pool and TPV of interest for the January 1985 CAO. The cold pool is identified to form over the northeastern coast of Baffin Island at 1800 UTC 11 January 1985. The cold pool meanders for several days over northern Canada before being transported equatorward from northern Canada into the U.S. during 17–20 January 1985. The cold pool then moves northeastward toward Labrador during 21–22 January 1985 and meanders over and near Labrador during 22–27 January 1985. The cold pool undergoes lysis over Quebec at 0600 UTC 28 January 1985. In total, the cold pool has a lifetime of ~17 days. The TPV forms at 1800 UTC 16 January 1985 over north central Canada, several days after the cold pool is identified to form (Fig. 4.11). The TPV moves generally westward over the next day before being quickly transported equatorward into southern Canada and then into the central U.S. between 18 and 20 January 1985 together with the cold pool. Similarly to the cold pool, the TPV is quickly transported northeastward toward Labrador during 21–22 January 1985, and then meanders over and near Labrador during 22–27 January 1985. Unlike the cold pool, the TPV then moves over the northwestern North Atlantic, where it meanders for a few more days before undergoing lysis at 1200 UTC 2 February 1985. Overall, the TPV has a lifetime of 17 days. In general, the large spatial overlap and temporal coincidence of the TPV and cold pool tracks over North America suggest that the these features are dynamically linked. Furthermore, the movement of the TPV and cold pool over and near the U.S. during 19–21 January 1985, when the January 1985 CAO is occurring, suggests that these features play an important role in the development of the January 1985 CAO.

Figure 4.12 shows a time series of the minimum DT potential temperature of the TPV and the minimum 1000–500-hPa thickness of the cold pool. At the beginning of its lifecycle, the cold pool is already characterized by a 1000–500-hPa thickness value of ~478 dam. Although the cold pool is objectively identified to form on 11 January 1985, subjectively, the cold pool appears to form several days earlier (not shown). There may have been separate 1000–500-hPa thickness minima embedded within the cold pool that may have contributed to multiple cold pool tracks. In general, the cold pool strengthens during 11–19 January 1985, with the cold pool reaching a minimum 1000–500-hPa thickness value of ~468 dam on 19 January 1985 (Fig. 4.12) over southern Canada (Fig. 4.11). The cold pool then weakens as it moves across the U.S. and then northeastward toward Labrador during 19–22 January 1985. The cold pool then generally intensifies as is meanders over and near Labrador until 27 January 1985, before weakening again. When the TPV forms, it is characterized by an already low minimum DT potential temperature value of ~269 K. As will be shown in section 4.2.3, the TPV forms from a preexisting broad area of low DT potential temperature air. In general, the TPV intensifies until 20 January 1985 as the cold pool intensifies, though reaches a minimum DT potential temperature value of ~259 K on 20 January 1985, when the cold pool is already weakening. Afterward, during 21–29 January 1985, the patterns of intensity change of the TPV are similar to that of the cold pool.

The periods of strengthening of the cold pool and TPV may be related to longwave radiative cooling. Weakening of both the cold pool and TPV during 21–22 January 1985 may be related to latent heat release in association with a strong EC that rapidly intensifies during 21–22 January 1985 over the east coast of North America, as will be discussed in section 4.2.3. In addition, weakening of the cold pool as it moves across the U.S. during 19–21 January 1985 may be related to sensible heat fluxes from the surface as the cold pool moves over the relatively warmer land of the U.S compared to that of Canada, especially considering that there may be less snow cover and a thinner snowpack over the U.S. compared to Canada.

*4.2.3 Synoptic Evolution of TPV, Cold Pool, and CAO*

The equatorward transport of the TPV and cold pool is likely tied to flow amplification over the North Pacific and North Atlantic. Poleward fluxes of warm, moist air associated with a strengthening EC in the left exit region of a strong North Pacific jet stream over the central North Pacific and with an EC moving poleward off the coast of Labrador likely support ridge amplification over the eastern North Pacific and over Greenland, respectively, between 0000 UTC 14 and 0000 UTC 15 January 1985 (Figs. 4.13a–d). Meanwhile, during this same time period, the cold pool over northeastern North America moves southward on the west side of this ridge over Greenland. Collocated with the cold pool is a broad region of low DT potential temperature air that will evolve into the TPV. The EC over the North Pacific moves northeastward to the Gulf of Alaska between 0000 UTC 15 and 0000 UTC 17 January 1982, while the ridge over the eastern North Pacific concomitantly amplifies into western North America (Figs. 4.13c–h). During the same period, a new EC rapidly develops and then occludes as it moves from the northeastern U.S. to southeastern Canada, contributing to ridge amplification over the northwestern North Atlantic (Figs. 4.13c–h). Both ECs are associated with a narrow corridor of relatively high precipitable water air extending from the subtropics toward the high latitudes (e.g., Fig. 4.13f). Figures 4.14a,b shows that at 0000 UTC 15 and 0000 UTC 16 January 1982, associated with the aforementioned ECs and narrow corridors of relatively high precipitable water air are widespread regions of mid-level ascent and concomitantly, diabatically-driven upper-tropospheric divergent outflow and associated negative PV advection by the irrotational wind, which likely contributes to the ridge amplification over the eastern North Pacific and western North America, and over the northwestern North Atlantic. Also, at 0000 UTC 17 January 1985, the broad region of low DT potential temperature air has consolidated into a single TPV that is now tracked (Fig. 4.13g). Although the consolidation of the broad region of low DT theta air into a single TPV has not been examined closely, it may be related to the close approach and merger of separate mid-level vorticity maxima (not shown).

The combination of ridging over western North America and over Greenland and Baffin Bay force the TPV and cold pool to move equatorward over Canada between 0000 UTC 18 and 0000 UTC 19 January 1985 (Figs. 4.15a–d). A jet streak strengthens between the TPV and the ridge over western North America as the TPV moves equatorward and interacts with the jet streak. In addition, a strong surface anticyclone over the Arctic and extending southward into the Canadian Archipelago strengthens and builds southeastward over western Canada in the left entrance region of the jet streak. As the TPV and cold pool move equatorward into the northern U.S. through 0000 UTC 20 January 1985, the strong surface anticyclone strengthens in the left entrance region of the jet streak and rapidly builds southeastward into the central U.S (Figs. 4.15e,f). Strong cold air advection over the central and southern U.S. as implied in Fig. 4.15f along with an expected terrain-tied northerly component of motion (e.g., Colle and Mass 1995) on the east side of the Rockies where there is a strong SLP gradient associated with the surface anticyclone may allow the cold air associated with the cold pool to spread far away from the core of the cold pool. As the TPV and cold pool move eastward over the U.S. through 0000 UTC 21 January 1985 (Figs. 4.15g,h), the cold air associated with the cold pool overspreads the entire eastern U.S. Also, as the TPV has been interacting with a jet streak to its south and east, it has become deformed in shape and weakens between 0000 UTC 20 and 0000 UTC 21 January 1985 (Figs. 4.15e,g), suggesting perhaps that the deformation of the TPV may contribute to the weakening of the TPV. In addition, cyclogenesis of an EC is noted over southeastern Canada at 0000 UTC 21 January 1985 in the left exit region of the aforementioned jet streak (Fig. 4.15h), and this EC undergoes rapid cyclogenesis over the next day (not shown). Latent heat release associated with widespread ascent found in the vicinity of the EC during 0000 UTC 21 and 0000 UTC 22 January 1982 (Figs. 4.16a,b) may be contributing to the weakening of the TPV and cold pool during 21–22 January 1985 shown in Fig. 4.12.

Overall, as in the January 1982 CAO case, the fact that the TPV and cold pool in the January 1985 CAO case move directly together into the U.S. during the time of the January 1985 CAO suggest that the TPV and cold pool are dynamically linked and that the TPV and cold pool play an important role in the development of the January 1985 CAO. To get a better sense of the evolution of the three-dimensional structure of the TPV and cold pool in this case, cross sections of these features will be examined in section 4.2.4. In addition, the relationship between TPV–jet interaction and the development of the strong surface anticyclone important for the evolution of the January 1985 CAO will be explored further in section 4.2.5.

*4.2.4 Three-dimensional Structure of TPV and Cold Pool*

Figure 4.17 shows a meridional cross section (DD’) transecting the cold pool early in its lifecycle at 0000 UTC 13 January 1985, a time during which the TPV has not yet been identified. At this time, there is broad region of low DT potential temperature air with individual embedded DT potential temperature minima that is collocated with the broad cold pool. The cross section shows that this broad region of low DT potential temperature air corresponds to a broad region of depressed DT extending downward below 400 hPa, with embedded mesoscale undulations of the DT. There is a broad region of upward bowing of isentropes in the cross section beneath the broad region of lowered DT, indicative of the broad cold pool, with potential temperature values near the surface as low as 240 K illustrating the very cold air associated with the cold pool.

Figure 4.18 shows a meridional cross section (EE’) transecting the TPV and cold pool at 0000 UTC 17 January 1985, shortly after the TPV has been first identified. Since the broad region of low DT potential temperature air has consolidated into a TPV, there is a more focused region of lowered DT in the cross section in Fig. 4.18 at 0000 UTC 17 January 1985 compared to in the cross section in Fig. 4.17 at 0000 UTC 13 January 1985, with the DT now extending downward below 500 hPa where the TPV is located. There is also a more robust upward bowing of the isentropes beneath the TPV in the cross section in Fig. 4.18 at 0000 UTC 17 January 1985 compared to beneath the broad region of low DT potential temperature air in the cross section in Fig. 4.17 at 0000 UTC 13 January 1985, corresponding to a more intense cold pool at 0000 UTC 17 January 1985 (Fig. 4.18) compared to at 0000 UTC 13 January 1985 (Fig. 4.17).

Figure 4.19 shows a meridional cross section (FF’) transecting the TPV and cold pool at 0000 UTC 20 January 1985, when both features have moved into the northern U.S. and are contributing to CAO development. The TPV has intensified and now extends downward to ~750 hPa. The robust upward bowing of the isentropes within and beneath the TPV illustrates the impressive cold pool collocated with the TPV. In addition, the TPV is interacting with a jet streak to its south, and a PV wall similar to that in the cross section in Fig. 4.9 for the January 1982 CAO case extends throughout much of the troposphere between the TPV and warm air to the south, concomitant with the jet streak. Near the surface, the isentropes spread horizontally outward, far away from the core of the TPV and cold pool, with the Arctic front associated with the leading edge of Arctic air likely located south of 37.5°N where there is a strong surface horizontal potential temperature gradient. As discussed by Shapiro et al. (1987) for a cross section transecting the same TPV and cold pool (Fig. 1.7 in this thesis), the cross section in Fig. 4.19 shows that within the Arctic air, the boundary layer appears well-mixed and neutrally stable, which, as suggested by Shapiro et al. (1987), may result from diabatic heating induced by the flow of Arctic air over the relatively warm land surface. Overall, the cross section in Fig. 4.19 illustrates that the TPV has a tropospheric deep influence and impacts a widespread geographical area, playing an important role in the development of the January 1985 CAO.

*4.2.5 Q-vector Diagnosis*

The concomitant occurrence of the interaction of the TPV with a jet streak over western North America and the strengthening and rapid southeastward expansion of the strong surface anticyclone over western North America in the left entrance region of the jet streak during 18–20 January 1985 (Figs. 4.15a–f) suggests that TPV–jet interaction may play an important role in the strengthening and expansion of the strong surface anticyclone. Figure 4.20 shows plots of 600–400-hPa **Q**n and **Q**s and their associated forcings for vertical motion from 1200 UTC 18 to 0000 20 January 1985. At 1200 UTC 18 January 1985, a region of divergence of **Q**n and **Q**n forcing for descent is located to the northwest of the TPV, over and to the southeast of the surface anticyclone over northwestern Canada (Fig. 4.20a). The orientation of the Qn vectors from cold to warm air in this region suggests that upper-level frontogenesis may be occurring, which would support the strengthening of the jet streak over western North America that occurs between 0000 UTC 18 and 0000 UTC 19 January 1985 (Figs. 4.15a–d). The strengthened jet streak may be associated with a strengthened thermally direct ageostrophic circulation in the entrance region of the jet streak and concomitantly increased descent and thus increased forcing for anticyclogenesis in the left entrance region of the jet streak. Also at 1200 UTC 18 January 1985, over northwestern North America, the Qs vectors are weak in magnitude and there is not an organized region of **Q**s forcing for descent (Fig. 4.20b). The region of **Q**n forcing for descent over and to the southeast of the surface anticyclone suggests that the surface anticyclone may strengthen and expand southeastward.

By 0000 UTC 19 January 1985, while the surface anticyclone has not strengthened, it has expanded southeastward over western Canada since 1200 UTC 18 January 1985 (compare Figs. 4.20a,c). As the TPV is interacting with the strengthening jet steak over western Canada (Fig. 4.15c), the magnitude of **Q**n has increased and the region of **Q**n forcing for descent has expanded in areal coverage just northwest of the location of the TPV, in the left entrance region of the jet streak, and over and to the south and east of the surface anticyclone (Fig. 4.20c). In addition, given the crossing of the geopotential height contours and potential temperature contours in and near the region of **Q**n forcing for descent, on the cyclonic shear side of the jet streak, there is suggestion that geostrophic cold air advection is occurring, providing additional evidence that upper-level frontogenesis may be occurring. The region of **Q**n forcing for descent over and to the south and east of the surface anticyclone suggests that the surface anticyclone may strengthen and expand southeastward. Also, there is now an organized region of **Q**s forcing for descent positioned west of the TPV and to the south and east of the surface anticyclone (Fig. 4.20d), which would at least suggest the surface anticyclone may build southeastward.

Between 0000 UTC and 1200 UTC 19 January 1985, the surface anticyclone does strengthen from ~1046 hPa to ~1051 hPa, or by ~5 hPa in 12 h, and expands southeastward (compare Figs. 4.20c,e). There continues to be a broad region of **Q**n forcing for descent just west of the TPV, over and to the southeast of the surface anticyclone (Fig. 4.20e), as well as a smaller and weaker area of **Q**s forcing for descent west of the TPV and to the south and east of the surface anticyclone (Fig. 4.20f), suggesting that the surface anticyclone may strengthen further and build southeastward. The surface anticyclone does strengthen between 1200 UTC 19 and 0000 UTC 20 January 1985 from ~1051 hPa to ~1054 hPa, or by about 3 hPa in 12 h, and continues to build southeastward, and there continues to be generally similar patterns of **Q**n and **Q**s and their associated forcings for descent near the TPV at 0000 UTC 20 January 1985 as 12 h earlier (compare Figs. 4.20e,f with Figs. 4.20g,h). In general, the concomitant occurrence of the increase in magnitude of **Q**n and increase in areal coverage of the **Q**n forcing for descent during TPV–jet interaction with the strengthening and expansion of the surface anticyclone suggests that TPV–jet interaction may play an important role in the strengthening and expansion of the surface anticyclone. Given the important role this surface anticyclone plays in the development of the January 1985 CAO, TPV–jet interaction may play an important role in CAO development.

4.3 Summary

Both the January 1982 CAO and January 1985 CAO are linked to a cold pool associated with a TPV. Ridge amplification appears to play an important role in the equatorward transport of the TPV and cold pool in each case. Ridge amplification in both cases is likely related to ECs and their associated poleward fluxes of warm, moist air from lower latitudes to high latitudes, as well as diabatically driven upper-tropospheric outflow associated with widespread mid-level ascent in the vicinity of the ECs. Thus, precursor disturbances like ECs may play a critical role in leading to ridge amplification necessary for the equatorward transport of TPVs and cold pools.

The large spatial overlap and temporal coincidence of the TPV and cold pool throughout much of their lifetimes in each case suggest that the TPV and cold pool are dynamically linked in each case and demonstrates that the influence of TPVs can extend throughout the depth of the troposphere and over a widespread geographical area. Indeed, in each case, cross sections transecting the TPV and cold pool illustrate that the TPV has a tropospheric deep impact, and as the TPV becomes a better defined and more robust feature, the cold pool does as well, demonstrating a dynamical linkage between the TPV and cold pool. In general, it is anticipated that as a TPV becomes stronger and better defined, the TPV will become associated with a greater degree of upward bowing of isentropes within and beneath the TPV such that a cold pool associated with the TPV will concomitantly become stronger and better defined. The cross sections also illustrate that near the surface, the Arctic air associated with the TPV and cold pool spread far away from the core of the TPV and cold pool, demonstrating that the TPV and cold pool have a geographically widespread impact in each case. Although the TPV and cold pool generally overlap spatially to a large degree in each case, their centers are not always completely vertically aligned. Changes in the tilt of TPVs and cold pools and the impact of changes in the tilt of TPVs and cold pools on the structure of these features and CAO development should be further explored. Also, processes hypothesized to play a role in intensity changes of the TPV and cold pool in each case such as longwave radiative cooling, latent heating, and sensible heating from the surface should be quantified in the future in order to better understand their impact on TPV and cold pool intensity in each case.

In addition, in each case, as ridge amplification occurs over western North America and as the TPV is transported equatorward, the TPV interacts with a strengthening jet streak located between the ridge and the TPV. The concomitant occurrence of the increase in magnitude of **Q**n and the increase in magnitude and/or areal coverage of **Q**n forcing for descent during TPV–jet interaction with the strengthening and expansion of a surface anticyclone over western North America in the left entrance region of the jet streak and within the region of **Q**n forcing for descent in each case suggests that TPV–jet interaction may play an important role in the strengthening and expansion of the surface anticyclone in each case. Given that the surface anticyclone helps allow cold air from the cold pool associated with the TPV to spread far away from the core of the cold pool in each case, TPV–jet interaction may play an important role in CAO development in each case. Overall, given that 1) the TPV is associated with a cold pool that moves directly along with the TPV into the U.S. during the time of the CAO in each case and 2) the TPV via TPV–jet interaction may help to strengthen a strong surface anticyclone that helps transport cold air from the cold pool associated with the TPV far away from the core of the cold pool in each case illustrates that TPVs may play a crucial role in CAO development in each case.

Fig. 4.1. Tracks of TPV (red) from 0600 UTC 15 December 1981 to 0000 UTC 13 January 1982 and cold pool (blue) from 1800 UTC 20 December 1981 to 1800 UTC 13 January 1982 for January 1982 CAO case. Stars denote locations of genesis, crosses denote locations of lysis, and red and blue dots represent 0000 UTC positions of TPV and cold pool, respectively, every 48 h. Numbers pointing toward dots represent dates of the 0000 UTC positions of the TPV and cold pool, such that numbers ≥17 correspond to dates in December 1981 and numbers ≤12 correspond to dates in January 1982.

Fig. 4.2. Time series of minimum DT potential temperature (θ) of TPV (K, red) every 6 h from 0600 UTC 15 December 1981 to 0000 UTC 13 January 1982 and minimum 1000–500-hPa thickness of cold pool (dam, blue) every 6 h from 1800 UTC 20 December 1981 to 1800 UTC 13 January 1982 for January 1982 CAO case.

Fig. 4.3. DT (2-PVU surface) potential temperature (K, shaded), wind speed (black contours every 10 m s−1, beginning at 50 m s−1), and wind (m s−1, flags and barbs) at (a) 0000 UTC 5 January, (c) 0000 UTC 6 January, (e) 0000 UTC 7 January, and (g) 0000 UTC 8 January 1982; 250-hPa wind speed (m s−1,shaded), 1000–500-hPa thickness (dashed red and blue contours every 10 dam, contoured red for values >540 dam and blue otherwise), SLP (black contours every 8 hPa), and precipitable water (mm, shaded) at (b) 0000 UTC 5 January, (d) 0000 UTC 6 January, (f) 0000 UTC 7 January, and (h) 0000 UTC 8 January 1982. Green line and dot represent track and position of TPV, respectively, and yellow line and dot represent track and position of cold pool, respectively.

Fig. 4.4. Precipitable water (mm, shaded), 600–400-hPa ascent (red contours every 2.5 × 10−3 hPa s−1), and 300–200-hPa PV (PVU, gray) and negative PV advection by the irrotational wind (PVU day−1, shaded) at (a) 0000 UTC 6 January and (b) 0000 UTC 7 January 1982.

Fig. 4.5. DT (2-PVU surface) potential temperature (K, shaded), wind speed (black contours every 10 m s−1, beginning at 50 m s−1), and wind (m s−1, flags and barbs) at (a) 0000 UTC 8 January, (c) 0000 UTC 9 January, (e) 0000 UTC 10 January, and (g) 0000 UTC 11 January 1982; 250-hPa wind speed (m s−1,shaded), 1000–500-hPa thickness (dashed red and blue contours every 10 dam, contoured red for values >540 dam and blue otherwise), SLP (black contours every 8 hPa), and precipitable water (mm, shaded) at (b) 0000 UTC 8 January, (d) 0000 UTC 9 January, (f) 0000 UTC 10 January, and (h) 0000 UTC 11 January 1982. Green line and dot represent track and position of TPV, respectively, and yellow line and dot represent track and position of cold pool, respectively.

Fig. 4.6. Precipitable water (mm, shaded), 600–400-hPa ascent (red contours every 2.5 × 10−3 hPa s−1), and 300–200-hPa PV (PVU, gray) and negative PV advection by the irrotational wind (PVU day−1, shaded) at (a) 0000 UTC 10 January and (b) 0000 UTC 11 January 1982.

Fig. 4.7. (left) Cross section along line AA’ of PV (PVU, shading), potential temperature (K, black), and wind speed (dashed white contours every 10 m s−1, beginning at 50 m s−1); (upper-right) DT (2-PVU surface) potential temperature (K, shaded), wind speed (black contours every 10 m s−1, beginning at 50 m s−1), and wind (m s−1, flags and barbs); and (lower-right) 1000–500-hPa thickness (dam, shading) at 1200 UTC 16 December 1981. Green line in upper-right and lower-right panels represents transect of cross section AA’. Label “TPV” represents location of TPV.

Fig. 4.8. As in Fig. 4.7, but for cross section along line BB’ at 1200 UTC 2 January 1982. Label “CP” represents location of cold pool.

Fig. 4.9. As in Figs. 4.7 and 4.8, but for cross section along line CC’ at 0000 UTC 10 January 1982.

Fig. 4.10. SLP (blue contours every 4 hPa, beginning at 1040 hPa) and 600­–400-hPa **Q**n (K m−1 s−1, vectors), **Q**n forcing for vertical motion (10−17 Pa−1 s−3, shaded), geopotential height (dark gray contours every 10 dam), and potential temperature (dashed red contours every 5°C) at (a) 1200 UTC 8 January, (c) 0000 UTC 9 January, (e) 1200 UTC 9 January, and (g) 0000 UTC 10 January 1982; SLP (blue contours every 4 hPa, beginning at 1040 hPa) and 600­–400-hPa **Q**s (K m−1 s−1, vectors), **Q**s forcing for vertical motion (10−17 Pa−1 s−3, shaded), geopotential height (dark gray contours every 10 dam), and potential temperature (dashed red contours every 5°C) at (b) 1200 UTC 8 January, (d) 0000 UTC 9 January, (f) 1200 UTC 9 January, and (h) 0000 UTC 10 January 1982. Green line and dot represent track and position of TPV, respectively.

Fig. 4.11. Tracks of TPV (red) from 1800 UTC 16 January to 1200 UTC 2 February 1985 and cold pool (blue) from 1800 UTC 11 January to 0600 UTC 28 January 1985 for January 1985 CAO case. Stars denote locations of genesis, crosses denote locations of lysis, and red and blue dots represent 0000 UTC positions of TPV and cold pool, respectively, every 48 h. Numbers pointing toward dots represent dates of the 0000 UTC positions of the TPV and cold pool, such that numbers ≥13 correspond to dates in January 1985 and the number “2” corresponds to 2 February 1985.

Fig. 4.12. Time series of minimum DT potential temperature (θ) of TPV (K, red) every 6 h from 1800 UTC 16 January to 1200 UTC 2 February 1985 and minimum 1000–500-hPa thickness of cold pool (dam, blue) every 6 h from 1800 UTC 11 January to 0600 UTC 28 January 1985 for January 1985 CAO case.

Fig. 4.13. DT (2-PVU surface) potential temperature (K, shaded), wind speed (black contours every 10 m s−1, beginning at 50 m s−1), and wind (m s−1, flags and barbs) at (a) 0000 UTC 14 January, (c) 0000 UTC 15 January, (e) 0000 UTC 16 January, and (g) 0000 UTC 17 January 1985; 250-hPa wind speed (m s−1,shaded), 1000–500-hPa thickness (dashed red and blue contours every 10 dam, contoured red for values >540 dam and blue otherwise), SLP (black contours every 8 hPa), and precipitable water (mm, shaded) at (b) 0000 UTC 14 January, (d) 0000 UTC 15 January, (f) 0000 UTC 16 January, and (h) 0000 UTC 17 January 1985. Green line and dot represent track and position of TPV, respectively, and yellow line and dot represent track and position of cold pool, respectively.

Fig. 4.14. Precipitable water (mm, shaded), 600–400-hPa ascent (red contours every 2.5 × 10−3 hPa s−1), and 300–200-hPa PV (PVU, gray) and negative PV advection by the irrotational wind (PVU day−1, shaded) at (a) 0000 UTC 15 January and (b) 0000 UTC 16 January 1985.

Fig. 4.15. DT (2-PVU surface) potential temperature (K, shaded), wind speed (black contours every 10 m s−1, beginning at 50 m s−1), and wind (m s−1, flags and barbs) at (a) 0000 UTC 18 January, (c) 0000 UTC 19 January, (e) 0000 UTC 20 January, and (g) 0000 UTC 21 January 1985; 250-hPa wind speed (m s−1,shaded), 1000–500-hPa thickness (dashed red and blue contours every 10 dam, contoured red for values >540 dam and blue otherwise), SLP (black contours every 8 hPa), and precipitable water (mm, shaded) at (b) 0000 UTC 18 January, (d) 0000 UTC 19 January, (f) 0000 UTC 20 January, and (h) 0000 UTC 21 January 1985. Green line and dot represent track and position of TPV, respectively, and yellow line and dot represent track and position of cold pool, respectively.

Fig. 4.16. Precipitable water (mm, shaded), 600–400-hPa ascent (red contours every 2.5 × 10−3 hPa s−1), and 300–200-hPa PV (PVU, gray) and negative PV advection by the irrotational wind (PVU day−1, shaded) at (a) 0000 UTC 15 January and (b) 0000 UTC 16 January 1985.

Fig. 4.17. As in Fig. 4.7, but for cross section along line DD’ at 0000 UTC 13 January 1985.

Fig. 4.18. As in Fig. 4.7, but for cross section along line EE’ at 0000 UTC 17 January 1985.

Fig. 4.19. As in Fig. 4.7, but for cross section along line FF’ at 0000 UTC 20 January 1985.

Fig. 4.20. SLP (blue contours every 4 hPa, beginning at 1032 hPa) and 600­–400-hPa **Q**n (K m−1 s−1, vectors), **Q**n forcing for vertical motion (10−17 Pa−1 s−3, shaded), geopotential height (dark gray contours every 10 dam), and potential temperature (dashed red contours every 5°C) at (a) 1200 UTC 18 January, (c) 0000 UTC 19 January, (e) 1200 UTC 19 January, and (g) 0000 UTC 20 January 1985; SLP (blue contours every 4 hPa, beginning at 1032 hPa) and 600­–400-hPa **Q**s (K m−1 s−1, vectors), **Q**s forcing for vertical motion (10−17 Pa−1 s−3, shaded), geopotential height (dark gray contours every 10 dam), and potential temperature (dashed red contours every 5°C) at (b) 1200 UTC 18 January, (d) 0000 UTC 19 January, (f) 1200 UTC 19 January, and (h) 0000 UTC 20 January 1985. Green line and dot represent track and position of TPV, respectively.